

**A review of the principles of turbidity measurement**

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Abstract:	<p>Turbidity of water due to the presence suspended sediment is measured and interpreted in a variety of ways, which can lead to the misinterpretation of data. This paper re-examines the physics of light-scattering in water, and exposes the extent to which the reporting of turbidity data is inconsistent. It is proposed that the cause of this inconsistency is the fact that the accepted turbidity standards USEPA Method 180.1, ISO 7027 and GLI Method 2 are mutually inconsistent, as these standards give rise to a large number of measurement units that are not based on the optical properties of light absorption and scattering by suspensions in water, but by the arbitrary definition of the degree of turbidity being due to a concentration of formazin or other similar polymer-based calibration standard. It is then proposed that all turbidity-measuring devices should be calibrated with precise optical attenuators such as ND filters. Such calibration would allow for the definition of a beam attenuation coefficient (BAC) for every turbidity measuring instrument which would be cross-comparable with any other instrument calibrated in the same way. The units for turbidity measurements should be based on attenuation and reported as dB m<sup>-1</sup>. It is also proposed that a new standard should be drafted according to this attenuation-based method, and this new standard should also define the nomenclature for reporting data collected at any specific scattering angle in terms of an attenuation in dB m<sup>-1</sup>. The importance of multi-parameter turbidity measurements for the improvement of the quality of turbidity data, and the application of parameter-rich data sets to new methods of sediment characterization are discussed. It is suggested that more research into multi-parameter turbidity measurements is needed, as these new methods will facilitate an increase in parity between turbidity and suspended sediment concentration (SSC), a relationship that is subjective.</p>

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## 1 Abstract

2 Turbidity of water due to the presence suspended sediment is measured and interpreted in a variety of  
3 ways, which can lead to the misinterpretation of data. This paper re-examines the physics of light-  
4 scattering in water, and exposes the extent to which the reporting of turbidity data is inconsistent. It is  
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20 as these new methods will facilitate an increase in parity between turbidity and suspended sediment  
21 concentration (SSC), a relationship that is subjective.

## 22 Key words

23 Turbidity; Suspended Sediment; Sediment Transport

## 24 Introduction

25 The term “turbidity” is used widely throughout the physical sciences, and is interpreted in different  
26 ways in different contexts. It is commonly used to describe the optical clarity of a fluid (for example,  
27 the atmosphere), but for the purposes of this paper it refers to another common usage of the term  
28 which is the optical clarity of water. The presence of suspended particulates, dissolved inorganic  
29 chemical species, organic matter content and temperature can all affect the turbidity of a body of  
30 water. Investigators from different fields (waste water treatment; drinking water quality; forestry;  
31 civil engineering, aquaculture and ecology), and from the sub-disciplines within physical geography  
32 (fluvial; marine; glacial; coastal and estuarial) use turbidity measurement as a surrogate relative  
33 indicator of some other physical property, typically suspended sediment concentration (SSC) or total  
34 suspended solids (TSS). The amount of literature available on the subject of water turbidity is large,  
35 and a number of reviews have already been undertaken by investigators from some of the sub-  
36 disciplinary groups (Bilotta & Brazier 2008; Davies-Colley & Smith 2001; Kerr 1995; Ziegler 2003).  
37 There is however, some disagreement about what turbidity actually *means*, partly due to the  
38 different sub-disciplinary contexts in which the term is used, and partly because of the way in which  
39 the various measurement standards are assumed to be based on a correct *a priori* understanding of  
40 the physical processes of light-scattering and absorption.

41 Why is turbidity measurement important? The answer to this question depends on the  
42 perspective of the investigator. Some researchers are purely interested in the effect that the  
43 attenuation of light has on, for example, aquatic ecosystems, so that knowledge of the mass  
44 concentration of the suspended particles is not always the primary concern. In this case other  
45 parameters of interest include the reduction of visual range in water (affecting the ability of  
46 predators to hunt), and the amount of light available for photosynthesis (Bilotta & Brazier 2008).  
47 Other investigators are concerned directly with the study of sediment-transport processes, in which  
48 case knowledge of the mass concentration of the suspended particles and other parameters such as

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3 49 the particle-size distribution (PSD) is highly desirable for a number of reasons. Turbidity  
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5 50 measurement is important in this context, as although the turbidity measurement itself is heavily  
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7 51 biased by the PSD (Gippel 1989), it is not specifically designed to provide detailed information about  
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9 52 the PSD. For example, knowledge of particle size is important as the transport of fine sediment  
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11 53 derived from different land uses through catchments will impact directly on ecosystem services,  
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13 54 such as the provision of drinking water. Fine sediment delivery into river systems is also known to  
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15 55 cause problems such as irritation to fish gills whilst it is in suspension (Davies-Colley & Smith 2001).  
16  
17 56 Bilotta & Brazier (2008) summarize the effects of what they refer to as suspended solids (SS) on  
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19 57 periphyton and macrophytes, invertebrates and salmonid fish species. The displacement of many  
20  
21 58 fish species can often be due to an increase in turbidity caused by the cumulative effects of fine  
22  
23 59 sediment introduced into the riparian environment as a direct result of human activities such as  
24  
25 60 deforestation (Kerr 1995), or by natural events such as sediment-transport by stormwater runoff.  
26  
27 61 The use of turbidity measurement as a surrogate indicator for parameters such as suspended  
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29 62 sediment concentration has been explored by many researchers, as reviewed by Ziegler (2003). It  
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31 63 has been shown that the particle size distribution (PSD) of a homogenous sediment can vary  
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33 64 temporally from its source (e.g. hillslope runoff) as it is transported through a catchment into a  
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35 65 stream, due to a variation in the relative proportion of aggregates (flocs) present in the measured  
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37 66 flux (Slattery & Burt 1997). Therefore knowledge of how the PSD varies dynamically in this fluvial  
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39 67 context due to a variability in the degree of flocculation (DOF) is important for the study of the  
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41 68 transport processes of both sediment and organic species in flocs (Williams et al. 2007). There is  
42  
43 69 clearly some variation in the importance given to the parameters of turbidity by the different sub-  
44  
45 70 disciplinary groups, and so the aim of this paper is to evaluate how relevant turbidity measurement  
46  
47 71 is to the study of sediment-transport processes specifically, and to propose methods for the  
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49 72 improvement of the measurement and reporting of turbidity in a general context. The steps required  
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51 73 to achieve this evaluation are given by the following list of objectives:  
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3 74 1. To analyse critically the measurement methodologies described in the literature  
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5 75 including any inconsistencies in nomenclature of measurement principles.  
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7 76 2. To review briefly the physics of light absorption and scattering processes in water in  
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9 77 order to provide an underpinning for the discussion of the definition of terms  
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11 78 according to various investigators from different sub-disciplinary groups.  
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13 79 3. To present a critique of the measurement units, calibration methods and standards  
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15 80 applicable to the measurement of turbidity, SSC and TSS, and to examine of the  
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17 81 origins of the relationship between turbidity measurements and the implied  
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19 82 properties of suspended sediment. This step is vital because the cross-comparability  
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21 83 of turbidity data obtained in the field is often invalid due to a widespread reliance on  
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23 84 the assumed integrity of Formazin calibration methods.  
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25 85 4. To propose, based on objective 3, that a new turbidity instrumentation standard is  
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27 86 required, and to describe its fundamental content.  
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## 32 87 **Turbidity measurement principles and nomenclature**

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35 88 The measurement of turbidity is split into two basic methodologies: *turbidimetry*, in which the  
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37 89 degree of *transmission* of light is determined, and *nephelometry*, in which the degree of light-  
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39 90 scattering is evaluated (see reviews by Ziegler, 2003 and Lawler, 2005). This division has its roots in  
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41 91 the mathematical descriptions employed to model the various phenomena. In the case of  
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43 92 turbidimetry, the appropriate theories are due to Beer (1852) and Lambert (1760) ; as for  
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45 93 nephelometry, many theories and models have been developed to describe a range of scattering  
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47 94 processes, and these models are mostly derived from Mie theory (Mie 1908). Nephelometry itself is  
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49 95 sub-divided into three further categories which are *forward-scattering*, *side-scattering* and *back-*  
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51 96 *scattering*. Side-scattering is generally accepted to be a measurement angle of 90° to the incident  
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53 97 beam, although the existing standards impose different upper and lower bounds on that value  
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55 98 (Table 3). Forward-scattering ( $0^\circ < \vartheta < 90^\circ$ ) and Back-scattering ( $90^\circ < \vartheta < 180^\circ$ , often referred to as  
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3 99 *optical back-scattering* or *OBS*) however, do not have a well-defined relative measurement angle.  
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5 100 Different instruments employ different measurement angles, and these values are not always  
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7 101 reported.  
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10 [Insert Figure 1.]  
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13 103 Before continuing with the discussion another ambiguity in terminology must be addressed. The  
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15 104 definition of the scattering angle in terms of where the 0° position is located spatially also varies  
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17 105 throughout the literature (Table 1). For example in some cases a forward-scattering angle is stated,  
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19 106 which *implies* that the transmitted (direct) beam is located at 0° (Agrawal et al., 2008 and Jansson,  
20  
21 107 1992). Contradictory to this position, Bilro et al. (2010) define the transmitted beam as being located  
22  
23 108 at the 180° position. In one instance two contradictory diagrams are presented in the same paper  
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25 109 (Sadar 2004, pp.8-9), and in many other cases the scattering-regime nomenclature is not associated  
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27 110 with a specific scattering angle (e.g. Fugate & Friedrichs, 2002).  
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31 111 The interpretation that is adopted throughout this paper is that the scattering-angle is  
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33 112 specified in terms of a detector placed at a position with respect to the incident beam *after* a  
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35 113 physical interaction has occurred in the sample, i.e. the direct beam detector is placed at the 0°  
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37 114 position (denoting “pure” attenuation measurement), forward-scattering detectors are placed  
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39 115 anywhere from  $0^\circ < \vartheta < 90^\circ$ , a side-scattering detector is placed at exactly 90°, and back-scattering  
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41 116 detectors are placed at  $90^\circ < \vartheta \leq 180^\circ$ .  
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45 117 [Insert Table 1]  
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## 118 The physics of light absorption and scattering through turbid water

### 119 A brief review of optical theories

120 To understand the physics of light scattering by particles suspended in water, it is necessary to have  
121 some knowledge of the mathematical models employed to describe the various absorption and  
122 scattering processes. Fundamental theory and mathematical model development are continually  
123 progressing in this area, but the basic points of interest pertinent to the understanding of turbidity in  
124 water for the practical investigator are summarised in this section. Three main theories are  
125 discussed: Rayleigh theory, Mie theory and geometric optics. Also discussed are two theories that  
126 can be considered as approximations to Mie theory for specific conditions. These are the *Fraunhofer*  
127 *diffraction* theory (FDT) and the *Anomalous diffraction* theory (ADT) of Van De Hulst (1957). The  
128 reason that these two theories are considered here is that they both yield computationally fast  
129 algorithms that are utilised by laser-based particle-sizing instruments. These instruments are used  
130 widely in suspended particle analysis (organic and inorganic) both *in situ* and off-line in laboratories,  
131 and are extensively employed for suspended sediment characterization.

132

### 133 Rayleigh and Mie scattering

134 The third Baron Rayleigh formulated his scattering theory to account for the blue colour of the sky  
135 (Strutt 1871). Rayleigh scattering involves particles that are much smaller than the wavelength of the  
136 incident light, and are also defined as being *optically soft* – meaning that the particles are limited to  
137 having a refractive index very close to 1 (air molecules in the case of Rayleigh's model). Rayleigh  
138 demonstrated that scattering from small particles is strongly wavelength dependent in favour of the  
139 shorter wavelengths and is spatially isometric (i.e. scattered equally in all directions), hence the blue  
140 colour of the sky. He determined that this blue colour is predominant because the scattered light  
141 intensity is inversely proportional to the fourth power of the incident light wavelength, i.e. the



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3 142 shorter wavelengths of light (e.g. blue end of the visible spectrum) are scattered more readily than  
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5 143 the longer wavelengths of light (e.g. red end of the visible spectrum).  
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8 144 Gustav Mie originally developed his theory to explain the colouration of metals in the  
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10 145 colloidal state (Mie 1908). Mie theory successfully explains the dominance of forward scattering  
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12 146 where particles are of a similar size to or larger than the incident wavelength of light, unlike the case  
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14 147 of isotropic scattering of light by much smaller particles as in Rayleigh scattering.  
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17 148 In order to get some sense of the particle size ranges that are applicable to the different  
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19 149 scattering regimes it is first necessary to define the *dimensionless size parameter*  $x$ ,  
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$$x = \frac{2\pi r}{\lambda} \quad (1)$$
  
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26 151 where  $r$  is the spherical particle radius [m] and  $\lambda$  is the wavelength of the incident light [m]. Figure 2  
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28 152 shows how the forward-lobed nature of a set of light intensity distribution functions develops as  $x$   
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30 153 increases from 0.1 to 10. These spatial intensity distribution functions are also known as *scattering*  
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32 154 *phase functions*, which are calculated using Mie theory.  
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35 155 [insert Figure 2.]  
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## 38 156 Geometric optics

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41 157 Geometric optics, otherwise known as ray optics, describes the light traversing a medium in terms of  
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43 158 a straight path (hence “ray”). It explains *refraction*, in which there is a change in direction of a light  
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45 159 ray at the interface between two regions with differing refractive indices. It also accounts for  
46  
47 160 *reflection* and *absorption*, and is best applied in situations where the wavelength of light is much less  
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49 161 than the size of the scattering particle. Figure 3 depicts a simplified diagram of scattering and  
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51 162 absorption processes of a particle suspended in water as viewed from the perspective of ray optics.  
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[insert Figure 3.]

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6 164 Fraunhofer diffraction theory (FDT)

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9 165 Fraunhofer diffraction occurs at small angles to the forward-scattered beam, i.e.  $<30^\circ$ . Under these  
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11 166 conditions of wavelength and scattering angle, FDT is a useful approximation to Mie theory, and is  
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13 167 popular due to the relative simplicity of its algorithms. Due to the wavelength and particle size  
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15 168 restrictions FDT cannot be applied to sub-micron sized particles. For example, the smallest sized  
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17 169 sediment particle that could exhibit Fraunhofer diffraction when illuminated by a beam of red light  
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19 170 (wavelength 630 nm) would be  $6.3 \mu\text{m}$ , i.e. well above the sub-micron size limit.

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26 172 Anomalous diffraction theory (ADT)

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29 173 ADT (Van De Hulst 1957) is a computationally efficient method by which the scattering from small  
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31 174 particles can be modelled. The caveat is that the particles must be *optically soft* as in Rayleigh  
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33 175 scattering (i.e. they must have a refractive index close to 1), and they must also have a large size  
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35 176 parameter  $x \gg 1$ .

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41 178 The single scattering albedo

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43 179 The single scattering albedo, denoted  $\omega$ , is a useful unitless quantity defined as the ratio of  
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45 180 scattering efficiency to total extinction efficiency. If the attenuation observed by a detector placed in  
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47 181 the "direct beam" configuration as in Figure 1 was due entirely to absorption, then  $\omega = 0$ . When the  
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49 182 observed attenuation is due to scattering processes alone, then  $\omega = 1$ . The scattering albedo is useful  
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51 183 when describing the particle size range that can be effectively modelled by the various regimes  
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53 184 (Rayleigh, Mie etc.). A graph of scattering albedo ( $\omega$ ) versus size parameter ( $x$ ) is presented by  
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55 185 Moosmüller & Arnott (2009, Fig.1, p.1031), which shows the particle size ranges covered by Rayleigh  
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3 186 and Mie theory for particles with a refractive index of 1.55 (similar to that of silica). On this graph,  
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5 187 the approximate scattering-model regime boundaries are observed, as shown in Figure 4. The *large*  
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7 188 *particle limit* of Mie theory is also shown, and the size parameter at which Mie theory converges  
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9 189 with this limit is the point at which geometric optics (not shown on the graph) becomes an  
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11 190 alternative scattering model (at  $x \approx 2000$ ).

### 191 **Light absorption and scattering by suspensions in water**

192 In the terminology of physical optics absorption is a *non-parametric* process, i.e. one that is  
193 inherently lossy – meaning that energy is dissipated in the absorbing medium. The *parametric*  
194 processes that are to be considered do not involve any imparting of energy to the physical system  
195 through which the radiation is traversing, i.e. the wavelength of the scattered light is not altered  
196 (elastic scattering). The pertinence of these (and other) theories to the study of suspended particles  
197 in general, and suspended sediment specifically, must be considered. Rayleigh theory is applicable to  
198 small, non-absorbing (dielectric) spherical particles. Mie theory is the most ubiquitous of the models  
199 that is applied to the study of light scattering by suspensions in water. It represents a general  
200 solution to scattering from absorbing or non-absorbing spherical particles, with no limits on particle  
201 size. Rayleigh theory is less complex to apply than Mie theory, but is limited to small particles. The  
202 dimensionless size parameter  $x$  (Equation 1) for the scattering regimes, and the equivalent  
203 approximate particle size ranges are:

204  $x \ll 1$  Rayleigh scattering (*2 nm to 75 nm*)

205  $x \cong 1$  Mie scattering (*20 nm to 765  $\mu\text{m}$* )

206  $x \gg 1$  Geometric optics (*>200  $\mu\text{m}$* )

207 The graph of wavelength vs. particle diameter (Figure 4) shows the accepted boundaries  
208 between the various scattering regimes, as adapted from Lelli (2014) and confirmed by Moosmüller

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3 209 & Arnott (2009). Also plotted on the graph are the clastic sediment size ranges that are of interest in  
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5 210 this paper.  
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8 211 **[insert Figure 4.]**  
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11 212 Interpretation of this plot must however be considered carefully, as the data it represents are  
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13 213 limited to a *single scattering event* from a *purely spherical* particle. The regime boundaries located at  
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15 214  $x=0.02$ ,  $x=0.2$  and  $x=2000$  (Lelli 2014 and Moosmüller & Arnott (2009) are not strict demarcation  
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17 215 lines (i.e. Mie theory includes Rayleigh theory as  $x \rightarrow 0$ ), but are there to suggest the generally  
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19 216 accepted view of where the various models are used with respect to particle size parameter  $x$ .  
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21 217 These boundaries should be considered to be somewhat blurred when applied to multiple-scattering  
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23 218 from non-homogenous suspended sediment particles. Considerable model development is needed  
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25 219 to account for scattering from large, non-spherical sediment particles. This work will lead to a  
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27 220 redefinition of the scattering regime boundaries as depicted in Figure 4, with new models specific to  
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29 221 suspended sediment being represented on the graph. There would also be one omission from the  
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31 222 graph, namely Rayleigh scattering. As far as light scattering from suspended sediment is concerned,  
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33 223 this theory has no application due to the restrictions in particle size (i.e. very small:  $< 76.4$  nm) and  
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35 224 refractive index (i.e.  $n \approx 1$ ). Although Mie theory is limited to small, spherical particles only, it has  
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37 225 many extensions that describe much more complex scattering regimes (including multiple-scattering  
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39 226 and scattering from small non-spherical particles), and also simpler scattering regimes such as FDT  
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41 227 (valid for particle diameter  $d \geq 10 \lambda$ , and scattering angle  $\vartheta \leq 30^\circ$ ). Other theories such as ADT which  
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43 228 as with Rayleigh theory was originally designed for optically soft particles (but in this case with a  
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45 229 large  $x$  value), are also adaptable to cope with higher refractive indices and non-spherical particles  
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47 230 (Liu et al. 1998).  
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53 231 There is clearly a need to find a light-scattering model framework that is consistent with  
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55 232 both small and large particle scattering, and which is also extensible to many-particle analysis. In the  
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57 233 case of back-scattering from suspended sediment it has been shown that the reflectivity of the  
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3 234 sediment also has a direct effect on the scattered light intensity (Sutherland et al. 2000), suggesting  
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5 235 that geometric optics may play a part in future model development. Without a comprehensive  
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7 236 understanding of the complex manner by which particle size, shape and concentration affect the  
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9 237 absorption and scattering of light, it will not be possible to interpret what a turbidity measurement  
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11 238 actually means.

### 14 239 **The definition of the beam attenuation coefficient.**

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17 240 The attenuation coefficient  $\Sigma$  is commonly referred to as the *beam attenuation coefficient* (BAC) in  
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19 241 the turbidity literature, but these two quantities are defined in different ways by different authors. It  
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21 242 is important that the ambiguities in both the definition and application of the BAC as a method for  
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23 243 comparing turbidity data obtained by different methods are appreciated, as these ambiguities can  
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25 244 lead to the misinterpretation of that data. The following discussion focusses on how the *a priori*  $\Sigma$  is  
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27 245 defined, and then leads on to a definition of the BAC as an expression of  $\Sigma$  in terms of observable  
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29 246 quantities, i.e. a *measured* attenuation and the *optical path-length* of the measurement instrument.

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### 34 248 **The attenuation coefficient $\Sigma$**

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39 249 Light is absorbed by water and this absorption is a function of the wavelength of the incident light  
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41 250 (Figure 5). The strongest absorption occurs at a wavelength of  $\lambda = 417.5\text{nm}$  (Pope & Fry 1997) which  
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43 251 gives a maximum reduction in transmitted light intensity of 0.05% over a distance of 0.1 m, which is  
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45 252 the typical limit to the optical path length of existing turbidity instruments. As this is the worst-case  
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47 253 scenario, the absorption of light by water is considered to be negligible in the context of turbidity  
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49 254 measurement.

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52 255 **[insert Figure 5.]**

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56 256 Light is also absorbed by any other material that may be suspended in the water. In order to  
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58 257 determine practically a value for absorption it is necessary to measure the amount of light  
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3 258 transmitted through a given sample of water. This is termed the transmittance,  $T$ , which is defined  
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5 259 as the ratio of the transmitted light intensity  $I$  to the light source intensity  $I_0$ , and has units of  $\text{Wm}^{-2}$ .  
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7 260 The transmittance is also related to the optical depth (Equation 2),  $\tau$  (effectively the opacity of the  
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9 261 medium), and the *absorbance*,  $A$ :

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$$T = \frac{I}{I_0} = e^{-\tau} = 10^{-A} \quad (2)$$

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16 263 A quantitative measure of the optical depth  $\tau$  can be expressed in terms of the natural logarithm of  
17  
18 264 the transmittance or in terms of the absorbance (Equation 3). This in turn leads to a definition of  
19  
20 265 absorbance with units of the Neper (Equation 4), or in terms of the base-ten logarithm (Equation 5)  
21  
22 266 yielding a decibel quantity.

23  
24  
25 267 
$$\tau = -\ln(T) = A \ln(10) \quad (3)$$

26  
27  
28 268 
$$A = \frac{\ln(T)}{\ln(10)} = -\log_{10}(T) \quad (4)$$

29  
30  
31  
32 269 
$$A = -10 \log_{10}(T) \quad (5)$$

33  
34  
35 270 This definition of absorbance as a logarithmic function of transmittance is useful as it facilitates a  
36  
37 271 linear relationship with the optical path-length. When a linear relationship between *transmittance*  
38  
39 272 and *path-length* is established it then becomes theoretically easier to relate the *absorbance* to the  
40  
41 273 *concentration* of a suspension, which will consequently itself be a linear function.

42  
43  
44 274 The *a posteriori* description of the attenuation of light through a homogeneous medium is credited  
45  
46 275 to Bouguer (1729) and is also associated with Lambert. It has been called Bouguer's law, Lambert's  
47  
48 276 law (Lambert 1760) and the Bouguer-Lambert law. It states that the attenuation is proportional to  
49  
50 277 the distance travelled through the absorbing medium. The extension to this law which includes a  
51  
52 278 term for the concentration of absorbers is known as Beer's law, or more ubiquitously as the Beer-  
53  
54 279 Lambert law (Equation 6 and Equation 7), which states that the attenuation is proportional to the  
55  
56 280 concentration of the absorbers (Beer 1852).  
57  
58  
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1  
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3 281 The Beer-Lambert law allows the absorbance to be stated under *ideal conditions*, including the  
4  
5 282 assumption that there are *no scattering processes* occurring in the sample, and that the attenuation  
6  
7 283 is linear along the light path. This law enables the absorbance to be directly related to the  
8  
9 284 concentration of absorbers,  $c$ , and the path length  $l$  (Equation 6). Equation 7 expresses the same  
10  
11 285 quantity as a transmittance:

12  
13  
14  
15 286 
$$A = \varepsilon c l \quad (6) \quad T = e^{-\varepsilon c l} \quad (7)$$

16  
17 287 where  $\varepsilon$  is the absorptivity [ $\text{m}^2$ , or  $\text{m}^2 \text{kg}^{-1}$ ] of the absorbers in suspension, and is a constant  
18  
19 288 dependent on the physical properties of the absorbers (i.e. dielectric properties). When defined in  
20  
21 289 these terms, the *attenuation coefficient*  $\Sigma$  can be stated as the product of the absorptivity and the  
22  
23 290 concentration of the absorbers:

24  
25  
26  
27 291 
$$\Sigma = \varepsilon c \quad (8)$$

28  
29  
30 292 Substituting Equation 8 into Equation 6 gives the absorbance in terms of the attenuation coefficient:

31  
32  
33 293 
$$A = \Sigma l \quad (9)$$

34  
35  
36 294 The attenuation coefficient can be expressed in Napierian terms or as a decadic quantity (i.e. in  
37  
38 295 decibels). The measured *luminance* ( $\text{Cd m}^{-2}$ ) represents the *power* delivered by the transmitted light  
39  
40 296 beam per unit area. In electronic design it is more common to use decadic terminology to specify  
41  
42 297 measurement instrument parameters such as those used for the determination of light attenuation.  
43  
44 298 If Equation 7 is substituted into Equation 5, then the absorbance can alternatively be stated in  
45  
46 299 decibels (Equation 10 and Equation 11).

47  
48  
49  
50 300 
$$A = 10 \Sigma l \quad (10) \quad A = 10 \varepsilon c l \quad (11)$$

51  
52  
53 301 It is worth noting that the absorbance  $A$  is a dimensionless parameter, and the attenuation  
54  
55 302 coefficient  $\Sigma$  has units of reciprocal length ( $\text{m}^{-1}$ ). However, the absorptivity  $\varepsilon$  may have different units  
56  
57 303 depending on the context in which the concentration  $c$  is expressed (Equation 11). For example, in  
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60

1  
2  
3 304 the case where the concentration is simply the number of absorbers  $N$  per unit volume, then the  
4  
5 305 units of concentration are reciprocal volume, i.e.  $\text{m}^{-3}$  or  $\text{l}^{-1}$ . Therefore, absorptivity  $\epsilon$  in this instance  
6  
7 306 has units of  $\text{m}^2$ . In the case of suspended sediment, the absorptivity  $\epsilon$  would have units of  $\text{m}^2 \text{kg}^{-1}$ . It  
8  
9 307 is important to recognise the units stated for absorptivity, as other nomenclature could potentially  
10  
11 308 refer to the same physical quantity. For example, the *mass attenuation coefficient* used in chemistry  
12  
13 309 also has units of  $\text{m}^2 \text{kg}^{-1}$ . Hence it is prudent to examine the mathematical definition being used  
14  
15 310 within a given text to determine what physical quantity is actually being discussed, and not to rely on  
16  
17 311 the accuracy of the nomenclature at all. Another example of ambiguous nomenclature is highlighted  
18  
19 312 by Fig.4, which shows the graph of the light absorption spectrum of water. The range of this function  
20  
21 313 is referred to as the *absorption coefficient*, and as it has units of reciprocal length ( $\text{m}^{-1}$ ) it is  
22  
23 314 equivalent to the  $\Sigma$  of this discussion (i.e. the *attenuation coefficient*). This multiplicity of  
24  
25 315 measurement units has the potential to cause confusion, since the absorption coefficient has *the*  
26  
27 316 *same units* as the attenuation coefficient  $\Sigma$ . This is an important point as absorption is *not the same*  
28  
29 317 as attenuation. Attenuation is the *end result* of the effects of the *physical properties* of the medium  
30  
31 318 on the *propagation* of the light waves, and represents a loss of measureable light intensity. Any  
32  
33 319 measured attenuation cannot be presumed to be due to absorption alone (Figure 3). Scattering of  
34  
35 320 light can occur in all directions, and reflection and refraction of light can also distort any attenuation  
36  
37 321 measurement. For example, Gumprecht & Sliepcevich (1953) suggested that *forward scattering* can  
38  
39 322 distort a true attenuation measurement by *adding* to the transmitted light intensity observed by a  
40  
41 323 detector. This forward-scattering component is referred to as the *extinction coefficient* by Clifford et  
42  
43 324 al. (1995, p.774), who describe it as "*the re-formation of light after scattering behind the particle*",  
44  
45 325 and attribute this effect to the presence of suspended particles of diameter less than approximately  
46  
47 326  $4 \mu\text{m}$ .  
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3 328 BAC – the beam attenuation coefficient  
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6 329 The attenuation coefficient  $\Sigma$  is defined for *ideal* conditions, i.e. situations in which the attenuation  
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8 330 of light obeys the Beer-Lambert law and is thus concerned with absorption only, although some  
9  
10 331 definitions of BAC include a term for light-scattering (Kirk 1985). However, light-absorption cannot  
11  
12 332 be measured directly; only the *attenuation* of a light source can be determined by direct  
13  
14 333 measurement of light transmitted through a sample. As this attenuation could be affected by other  
15  
16 334 processes besides absorption (e.g. scattering), the absorption itself is not directly observable. The  
17  
18 335 absorption and scattering processes that occur within the sample do not have any bearing on how a  
19  
20 336 transmitted light intensity is measured at a given angle with respect to the incident beam, as the  
21  
22 337 only available parameters are the measurement angle  $\vartheta$ , and  $I / I_0$  for each  $\vartheta$ . It is crucial that the  
23  
24 338 BAC is accepted only as a measurement of light *attenuation*, and it cannot by itself be used to infer  
25  
26 339 any *a priori* mechanism of absorption or scattering. It is however conceptually convenient to  
27  
28 340 consider the definition of the BAC as being based purely on the effects of absorption alone (i.e. the  
29  
30 341 ideal conditions of the Beer-Lambert law). The measurement of transmissivity and hence the  
31  
32 342 attenuation of light due to the turbidity of water is referred to in the literature as *turbidimetry* or  
33  
34 343 *transmissometry*. The class of device for performing this measurement is consequently termed a  
35  
36 344 *turbidimeter* or a *transmissometer*.  
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45 346 A practical definition of the BAC  
46

47 347 Many devices exist for the measurement of optical transmissivity in water, and in this sense the  
48  
49 348 word “transmissivity” is synonymous with *attenuation* and refers to the measurement of  $I / I_0$  at an  
50  
51 349 angle  $\vartheta$  of  $0^\circ$  with respect to  $I_0$ , i.e. the “direct beam” (Figure 1). This measurement leads to the  
52  
53 350 derivation of the BAC by application of Equation 4, such that the BAC in decibels per metre ( $\text{dB m}^{-1}$ )  
54  
55 351 can be stated as  
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3 352 
$$\text{BAC} = \frac{-10 \log T}{l} \quad (12)$$
  
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5

6 353 where  $l$  is the optical path length (m) as determined by the particular instrument used for the  
7  
8 354 measurement.  
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11 355

## 16 356 **Turbidity measurement units, calibration methods and standards**

### 19 357 **A summary of the major turbidity standards**

22 358 The following three standards are in common use throughout the sub-disciplines of water quality  
23  
24 359 assessment. Although other standards do exist, these three are the most commonly cited by  
25  
26 360 researchers into the properties of natural waters. The summaries of these standards are presented  
27  
28 361 in order to highlight some of the technical imprecision inherent in their measurement  
29  
30 362 methodologies.  
31  
32

#### 34 363 **US EPA Method 180.1**

36  
37 364 This standard has been in use in various revisions since the early 1970s. The most recent revision  
38  
39 365 being 2.0 (US EPA 1993), which states that it is applicable to the measurement of turbidity in  
40  
41 366 "*drinking, ground, surface, and saline waters, domestic and industrial wastes*" (US EPA 1993, p.1).  
42  
43 367 The standard employs the comparison between the light scattered by the test sample to the light  
44  
45 368 scattered by a "*standard reference suspension*" (US EPA 1993, p.1). This reference suspension  
46  
47 369 consists of a defined mixture of two chemicals, hydrazine sulphate and hexamethylenetetramine, to  
48  
49 370 produce a "*stock standard suspension*" known as Formazin (US EPA 1993, p.3). A *primary standard*  
50  
51 371 *suspension* is then created by diluting 10mL of stock standard in 100mL of reagent water. This  
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53 372 concentration is defined as having a turbidity of 40 Nephelometric Turbidity Units (NTU). Another  
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3 373 acceptable commercially available primary standard based on styrene divinylbenzene polymer is also  
4  
5 374 stated.

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7  
8 375 The instrumentation parameters for the measurement of scattered light by this standard are the use  
9  
10 376 of a tungsten light source with a colour temperature from 2200-3000K, and a beam path-length of  
11  
12 377 not greater than 0.1 m. The detector response should peak at 400-600 nm, and the measurement  
13  
14 378 angle should be  $90^\circ \pm 30^\circ$ . Note that this is a very broad range of light wavelengths and scattering  
15  
16 379 angles which encompass forward-, side- and back-scattering geometries.

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21  
22 381 **ISO 7027**

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25  
26 382 This standard has been in effect in Europe since 1994. It relies in part on the use of light scattering  
27  
28 383 and attenuation by standard suspensions for comparison with the same measurements in a test  
29  
30 384 sample, as with EPA Method 180.1. A notable difference between the two standards is that ISO 7027  
31  
32 385 dictates the use of near infrared light ( $\lambda = 860$  nm) for all measurements. The standard suggests that  
33  
34 386 at wavelengths greater than 800nm the interferences caused by natural colouration of the water  
35  
36 387 (e.g. by dissolved humic substances) can be significantly reduced, an effect which has been observed  
37  
38 388 by Hongve & Akesson (1998).

39  
40  
41  
42 389 In addition to the measurement of diffuse radiation (i.e. nephelometry) expressed in Formazin  
43  
44 390 Nephelometric Units (FNU – in the range 0-40), the standard also defines a method for the  
45  
46 391 “*measurement of the attenuation of a radiant flux, more applicable to highly turbid waters (for*  
47  
48 392 *example waste or polluted waters)*” (ISO 1999). This measurement is expressed in Formazin  
49  
50 393 Attenuation Units (FAU), in the range 40-4000 FAU.

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53 394  
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## 395 GLI Method 2

396 This method is explicitly for the determination of turbidity in drinking water. It is a nephelometric  
397 *and* attenuation-based *ratio-metric* method based on infrared light of 860 nm wavelength, in  
398 common with ISO 7027. The use of dual-beam instruments that have two light sources and two  
399 detectors is specified. Each light source is pulsed sequentially, and for each measurement phase a  
400 90° *active* intensity and a 0° *reference* intensity measurement is acquired (Figure 6). A ratio-based  
401 algorithm is then used to calculate an NTU value based on the four data points (i.e. two 0° and two  
402 90° measurements). The accepted reason for employing this method is that it improves instrument  
403 stability due to interferences caused by the degradation of the light source, the fouling of sensor  
404 windows, and the effects of water colouration. It must be noted that the ratio algorithm is not  
405 defined in the standard, which implies that the implementation is left to the instrument designer  
406 (the topic of ratio methods is considered in greater detail later). As in the previously discussed  
407 standards, formazin suspensions are used for calibration. This is an example of a *multiple parameter*  
408 measurement method.

409 [insert Figure 6.]

## 410 A summary of turbidity measurement units

411 The U.S. Geological Survey has summarized currently used turbidity units and their associated  
412 standards as reproduced in Table 2 (USGS 2013), with amendments for the scattering angle  
413 convention in use throughout this paper .

414 [Insert Table 2]

415 Most of the material reviewed for this paper pertains to measurements taken by turbidity  
416 instruments that comply with either USEPA Method 180.1 or ISO 7027, and hence the measurement  
417 units that are most commonly encountered in the literature are NTU, FNU (specifically for drinking-  
418 water assessment) and FAU (specifically for waste-water assessment). The USGS considers these

1  
2  
3 419 units to be the ones that are most commonly applied to submersible turbidimeters. The other units  
4  
5 420 listed in Table 2 are rarely encountered in the turbidity literature. In addition to the USGS website,  
6  
7 421 another useful summary containing greater detail regarding the applications of the different  
8  
9 422 turbidimeter designs is presented by Sadar (2004). A more concise summary of the standards  
10  
11 423 discussed in this paper is presented by (Ziegler 2003), and this summary is reproduced here (Table 3)  
12  
13  
14 424 as it provides pertinent and useful aid to the context of this discussion.

15  
16  
17 425 **[Insert Table 3]**

### 18 19 20 426 **The problem with formazin**

21  
22 427 Formazin is useful as a turbidity standard as it can be reproducibly prepared from raw materials to  
23  
24 428 within  $\pm 1\%$ , and comprises a wide range of particle shapes and sizes ranging from 0.1  $\mu\text{m}$  to 10  $\mu\text{m}$   
25  
26 429 (Buzoianu 2000). However, it also has a number of drawbacks as highlighted by Buzoianu (2000):

- 27  
28  
29 430
- 30 • The preparation temperature affects the resulting PSD.
  - 31 431 • Formazin is carcinogenic.
  - 32 432 • Formazin primary standards do not usually state the concentration uncertainty.
  - 33 433 • The stability of formazin standards decreases as the concentration decreases (Table 4). The  
34 434 dilution ratio can be very high which leads to high uncertainty at low concentrations. This  
35 435 necessitates the use of secondary standards with longer shelf lives, and these standards can  
36 436 have poor repeatability of preparation, they are not formazin (eg latex), and they have  
37 437 different (narrow) PSDs. Hence, the use of secondary standards produces more variation in  
38 438 the response of different measurement instruments to the same nominal turbidity level.

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48 439 **[Insert Table 4]**

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51 440 It is a key fact that all of the units described in the previous section (Table 2 and Table 3) are derived  
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53 441 from a *chemical concentration level* of formazin or a secondary polymer-based standard. By this  
54  
55 442 methodology an increase in concentration is defined as an increase in turbidity. *There is no defined*  
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3 443 *relationship between the stated turbidity and the measured light intensity.* The word  
4  
5 444 “concentration” has effectively been replaced by “turbidity” in the definition of these measurement  
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7 445 units. For example section 7.3 of US EPA Method 180.1 states “Primary calibration standards: Mix  
8  
9 446 and dilute 10.00 mL of stock standard suspension (Section 7.2) to 100 mL with reagent water. **The**  
10  
11 447 **turbidity of this suspension is defined as 40 NTU.** For other values, mix and dilute portions of this  
12  
13 448 suspension as required.”

14  
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16  
17 449 This definition is a serious issue as “turbidity” in these standard techniques no longer refers to an  
18  
19 450 optical property of water, but rather a chemical concentration of what is in terms of particle  
20  
21 451 classification an unknown distribution of both particle sizes and particle shapes. As the particle-size  
22  
23 452 distribution (PSD) is not known, it is therefore not repeatable between measurements due to factors  
24  
25 453 such as chemical degradation and flocculation during storage of the “stock standards”. Also, the fact  
26  
27 454 that it is deemed acceptable to use secondary standards that will not have the exact same optical  
28  
29 455 response as formazin (Sethi et al. 1997, p.110) suggests a flaw in the methodology at its root, as  
30  
31 456 these “stock standards” are clearly not consistent nor are they traceable.

32  
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34  
35 457 The sphericity of the suspended formazin particles is also not quantified. Sadar (1999) states when  
36  
37 458 describing formazin “*the polymer in solution consists of random shapes and sizes.*” Both PSD (Baker  
38  
39 459 & Lavelle 1984, Ziegler 2003) and sphericity (Gibbs 1978) have been shown to have a significant  
40  
41 460 effect on the light-scattering characteristics of a suspension. Referring back to Figure 2, the  
42  
43 461 dimensionless size parameter  $x$  has a large effect on the scattering phase function. For example,  
44  
45 462 nephelometric instruments are most sensitive to particles of  $<1 \mu\text{m}$  diameter as in this size-range  
46  
47 463 there is a significant amount of side-scattering, yet the standards do not state the PSD limits  
48  
49 464 required for reference solutions.

50  
51  
52  
53 465 It has been demonstrated that different instruments measure different turbidity values when  
54  
55 466 calibrated with the same primary standard, due to the differences in instrument design (Buzoianu  
56  
57 467 2000). This is a situation that can occur even when the different instruments are made to comply  
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3 468 with the same measurement standard (e.g. EPA Method 180.1), due to the wide design tolerances  
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5 469 (e.g. a measurement angle of  $90^\circ \pm 30^\circ$ ). In view of the large uncertainties in the concentrations (and  
6  
7 470 PSDs) of the calibration standards, augmented by the variation in measurement instrument  
8  
9 471 response, there is then a scenario in which *one* stock standard and *two* different measurement  
10  
11 472 instruments (made to the same or different standards) could potentially give rise to not *two*, but  
12  
13 473 *multiple different* initial calibration results (Figure 7). An inaccurate surrogate model of turbidity has  
14  
15 474 now effectively become synonymous with turbidity itself *by definition* in these standards. This  
16  
17 475 calibration problem has implications for the measurement of turbidity in the field. The cross-  
18  
19 476 comparability of measurements made by different researchers at different sites using different  
20  
21 477 instrumentation is now questionable, even if each researcher has a self-consistent set of repeatable  
22  
23 478 calibration data for their own particular measurement instrument. It is therefore necessary to take a  
24  
25 479 step back and to re-define the chain of measurement at its first and weakest link, which is the  
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27 480 Formazin standard, and to establish a new methodology based purely on the calibration of  
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29 481 measurement instruments to well-defined light intensities at well-defined wavelengths.  
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34 482 [insert Figure 7.]  
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### 38 483 **Towards a new turbidity instrumentation standard**

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42 484 In order to move towards a new standard for the design of turbidity instrumentation it is first  
43  
44 485 necessary to take a step back from the accepted suspension-based calibration methods as  
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46 486 prescribed by the existing standards. The following discussion attempts to clarify the misconceptions  
47  
48 487 associated with the relationship between SSC, TSS and turbidity, and leads on to a proposed  
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50 488 calibration methodology based on the measurement of light-attenuation due to the presence of  
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52 489 optical neutral density (ND) filters in the optical beam path. To complete the new standard, a new  
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54 490 nomenclature based on the BAC is proposed for the reporting of turbidity at multiple scattering  
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56 491 angles and wavelengths of light. To conclude the discussion, some suggestions for the contents of  
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3 492 potential secondary standards (based on the newly proposed instrumentation standard) for  
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5 493 surrogate SSC determination are then outlined briefly.  
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7

8 494 **Suspended sediment concentration (SSC) and total suspended solids (TSS): their**  
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11 495 **relationship with turbidity and the importance of the PSD**  
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13  
14 496 The surrogacy of physical properties for intrinsic optical properties as is the case regarding *chemical*  
15  
16 497 *concentration* becoming a surrogate for *optical turbidity* has raised the possibility of further  
17  
18 498 misinterpretation, due to the undefined PSD of the calibration standards and the inconsistent  
19  
20 499 response of different measurement instruments to the same PSD (Buzoianu 2000). In this section it  
21  
22 500 is necessary to take a step back from turbidity to examine the meanings of the pre-existing  
23  
24 501 terminology for suspensions (of sediment or otherwise) in water. It is important to understand this  
25  
26 502 terminology as the descriptive acronyms actually refer to documented test methods for the  
27  
28 503 determination of sediment concentration and suspended solids concentration. An understanding of  
29  
30 504 these methods will then facilitate a deeper appreciation of the reasons for the conceptual conflation  
31  
32 505 of sediment concentration with turbidity.  
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35  
36 506 The US convention regarding the attribution of documented test methods to the acronyms “SSC”  
37  
38 507 and “TSS” has been adopted in this paper. Regarding this terminology, as with that of turbidity, the  
39  
40 508 differences in use in different disciplinary areas arises again. For example Holliday et al. (2003)  
41  
42 509 suggest TSS to mean “total suspended sediment concentration”, rather than “total suspended  
43  
44 510 solids”, i.e. the acronym SSC may have been a better choice.  
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47  
48 511 The field techniques and laboratory methods for the measurement of SSC and TSS were reviewed by  
49  
50 512 Gray et al. (2000), who cite Method D 3977-97 (ASTM 1998) for SSC and Method 2540 D (APHA  
51  
52 513 1971) for TSS. They describe the two different analytical methods as follows:  
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- 55 514
- 56 • SSC data are produced by measuring the dry weight of all the sediment from a known  
57  
58 515 volume of a water-sediment mixture.  
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3 516           • TSS data are produced by several methods, most of which entail measuring the dry  
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5 517                   weight of sediment from a known volume of a subsample of the original.  
6

7 518 After an analysis of 3235 paired SSC and TSS measurements was performed, it was concluded that  
8  
9 519 SSC was the more reliable methodology (Gray et al. 2000), especially when the amount of sand in a  
10  
11 520 sample exceeds approximately one quarter of the dry sediment mass. The main reason given for this  
12  
13 521 disparity of results is that the SSC analytical method utilises the entire sample (including all sediment  
14  
15 522 present), whereas the TSS methods typically involve the analysis of only a sub-sampled aliquot of the  
16  
17 523 total sample. The decanting and pipetting techniques employed to obtain this aliquot do not capture  
18  
19 524 a complete representation of the sediment population of the original sample. The resulting sub-  
20  
21 525 sample is therefore sediment deficient, particularly of the larger sand-sized sediment fraction. Gray  
22  
23 526 et al. (2000) go on to suggest that the reason for this loss of sediment during TSS analysis arises from  
24  
25 527 the fact that TSS methods were originally designed for analysis of waste-water samples that were to  
26  
27 528 be collected *after* an initial settling phase, hence larger sediment particles were never intended to  
28  
29 529 be part of the analysis. They finally conclude that SSC and TSS analysis of natural water samples are  
30  
31 530 not comparable, and that SSC is the only viable method for the determination of the sediment  
32  
33 531 concentration of natural waters.  
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38 532 In order to relate a subjective turbidity reading to a real physical property such as SSC, a calibration  
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40 533 procedure is typically performed. This relationship between the optical properties of suspended  
41  
42 534 sediment and its mass concentration must therefore be understood, requiring the characterisation  
43  
44 535 of its lithology. The size of the sediment particle is frequently measured either directly (e.g. filtering  
45  
46 536 and sieving), or analytically (by LASER diffraction) in the case of smaller size fractions. LASER-based  
47  
48 537 particle size measurements give a *volume* concentration value, which then requires further  
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50 538 knowledge of the specific density and mineralogy of the sample in order for an estimate of the *mass*  
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52 539 concentration to be obtained. This process is known as *end-member calibration*.  
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3 540 The problem now arises that the detector response has been pre-calibrated to a primary standard,  
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5 541 with arbitrary units for turbidity based on unstable calibration methods. It has already been  
6  
7 542 suggested ([Figure 7](#)) that these units (NTU etc.) are not comparable between calibrations made on  
8  
9 543 instruments constructed to the same standard. It is therefore highly unlikely that calibrations made  
10  
11 544 by *different* instruments (constructed to the same or different standards) can ever be accurately  
12  
13 545 compared due to the invalidity of these extrinsic turbidity units. It is therefore necessary to  
14  
15 546 determine the true instrument response by a different method entirely. Only then can an end-  
16  
17 547 member calibration have any chance of being meaningful.

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21 548 Optical neutral density filters (ND filters) are regularly employed for the calibration of transmission-  
22  
23 549 based optical instruments, but are seldom employed in turbidimetry or nephelometry. These filters  
24  
25 550 provide a consistent optical density (OD) which in turn will attenuate a well-defined percentage of  
26  
27 551 the transmitted light. One such example of an attempt to calibrate a turbidimeter against a known  
28  
29 552 light attenuator is Finlayson (1985). By not only calibrating a turbidimeter against Formazin  
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31 553 suspension, but also against ND filters, Finlayson has devised a method by which direct comparison  
32  
33 554 between attenuation measurements made on the same sample by different devices could  
34  
35 555 potentially be developed. It can be seen that Formazin concentration does not in fact have a linear  
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37 556 relationship to measured light attenuation ([Figure 8](#)). Although the calibration data are sparse in the  
38  
39 557 upper range of the instrument in this case (Finlayson 1985), there is a good fit of the data to a power  
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41 558 law ( $R^2 = 0.9954$ ). The only two useful axes on this graph are “meter reading” and “neutral density  
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43 559 filters”, as these two alone are all that is required to accurately establish the response of the  
44  
45 560 instrument to attenuation ([Figure 9](#)). Only when this detector attenuation curve has been  
46  
47 561 established can further selective end-member calibrations be performed to determine the effect the  
48  
49 562 PSD has on the response of a particular instrument to a given sediment. Each ND filter represents an  
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51 563 optical density,  $d$ , which is directly equivalent to the absorbance  $A$ , as in [Equation 4](#). So in order to  
52  
53 564 calculate the BAC in  $\text{dB m}^{-1}$  for an instrument with path-length  $l$ , the following equation can be  
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55 565 applied ([Equation 13](#)):

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$$BAC = \frac{10 d}{l} = \frac{10 ND}{l} \quad (13)$$
  
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6 567 [insert Figure 8.]  
7

8  
9 568 [insert Figure 9.]  
10

## 11 569 Instrumentation parameters and calibration methods

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16 570 To arrive at a *consistent* methodology for the measurement of turbidity it is necessary to accept that  
17  
18 571 the only quantity that can be readily measured optically in this context is the transmitted light  
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20 572 intensity, and hence *attenuation* with respect to the light source (i.e.  $I/I_0$ ). It is the methodology for  
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22 573 taking this measurement that should be rigorously specified, regardless of the measurement angle  $\vartheta$   
23  
24 574 with respect to  $I_0$ . The implementation section of the standard should address this methodology, and  
25  
26 575 focus purely on the desired response of the instrument to light at defined intensities and  
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28 576 wavelengths. This aspect of work would involve the definition of parameters such as sensor type,  
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30 577 variable intensity light source specification (including coherence and polarization), detector amplifier  
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32  
33 578 gains and ranges, ND filter calibration procedure involving multiple beam paths, beam path-length  
34  
35 579 and collimation arrangements. It is then necessary to decide which instrument parameters (e.g.  $\vartheta$ ,  $\lambda$   
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37 580 and  $l$ ) should be specified as mandatory for all turbidity measuring instruments, and which ones  
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39 581 should be considered as being application-specific.  
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## 45 583 The reporting of turbidity measurement data

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49 584 The standardization of the reporting of turbidity as attenuation data (Ziegler 2003) and the use of a  
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51 585 more descriptive nomenclature is proposed, which will allow for the easy identification of  
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53 586 application-specific data such that incompatible measurements will not be inadvertently compared  
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55 587 to each other. It is suggested that significant progress could be made if the measurement concepts  
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57 588 for turbidimetry and nephelometry were unified, i.e. by treating them *both* as an attenuation  
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3 589 process. The only difference being that for scattered light measurement the effective concentration  
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5 590 of scatterers is *inversely proportional* to the BAC measured at a specific angle to the incident beam.  
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7 591 However, for that to be achieved formulations of the BAC at specific angles must then be defined,  
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9 592 for example  $BAC_0$  for a standard transmissivity measurement and  $BAC_{90}$  for the nephelometric  
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11 593 counterpart at  $90^\circ$ . For the nephelometric case the relationship between the scattered light intensity  
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13 594 and the concentration could be viewed as an *inverse attenuation*, since a higher concentration of  
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15 595 particles will produce stronger scattering (until the concentration is too high, at which point  
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17 596 multiple-scattering and grain-shielding will dominate and interfere with the measurement of the  
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19 597 side-scattered light). Measurement-instrument calibration now becomes somewhat critical, as any  
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21 598 drift in the incident light intensity or the sensor response will affect the sensitivity of the system to  
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23 599 the low light intensities that need to be detected due to side- or back-scattering. This nephelometric  
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25 600  $BAC_{90}$  measurement results in potentially larger percentage errors than those that are likely for  
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27 601 measurements based on  $BAC_0$ , as greater electronic amplification is required to detect the weaker  
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29 602 scattered-light signal which can be inherently noisy. In order to formulate a generic equation for the  
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31 603 BAC as a function of measurement angle it is necessary to include two terms: one for attenuation  
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33 604 and one for scattering. The use of these terms is in no way a new idea (e.g. Kirk 1985), however the  
34  
35 605 interpretation of scattered light intensity as an *inverse absorbance* has not been previously  
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37 606 considered. In this new method the same measurement units could be employed for practical  
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39 607 comparison between data obtained under different conditions using different instruments, so long  
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41 608 as those instruments complied with the same instrumentation standard, and the reporting of said  
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43 609 data is consistent (Ziegler 2003). For example Kirk (1985) suggested using the correct description of  
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45 610 the measurement method, such as "side-scattering", when stating results – or preferably  $BAC_{90}$  in  
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47 611 this case.  
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### 613 **Standards for surrogate SSC determination**

614 Further standards for the determination of surrogate properties such as SSC should refer to  
615 instruments that are specified according to the new instrumentation standard. In order to estimate  
616 SSC accurately, optical instruments must be capable of producing data rich enough to facilitate  
617 suspended sediment characterization. Methods for the determination of the PSD (and other  
618 properties) of a suspended sediment by multi-parameter measurements need to be developed,  
619 which could include the use of LASER diffraction techniques. Other potential methods of sediment  
620 characterization should also be explored more thoroughly.

621

### 622 **Suspended sediment characterization**

623 For a deeper understanding of sediment transport to be realized, it is essential to know how the  
624 different size-classes of sediment respond to different flow conditions, especially the larger sand-  
625 sized particles that can be transiently in suspension long enough to affect turbidity measurements. A  
626 knowledge of sediment particle shape in terms of *sphericity* and *roundness* can also provide an  
627 insight into the distance travelled by sediment particles that have previously been entrained in a  
628 flow of water. There is a clear need therefore to characterize the suspended sediment to determine  
629 the particle sizes present. This characterization can be achieved by traditional gravimetric sampling  
630 methods, but there is an increasing need to gather data for research purposes *in-situ* and quickly. In  
631 some cases, these measurements could be made “off-line” by optical means, which would still be  
632 much faster than can be achieved by gravimetric methods. LASER-based optical measurements are  
633 the most commonly employed for this purpose, although there have been attempts to derive  
634 particle-size information from multi-parameter turbidity measurements. The effect that particle  
635 shape has on such measurements could also be exploited as a characterization technique.

## 636 Measurement ratios and multi-parameter method development

637 The designers of some turbidity meters (i.e. any commercially available instrument that claims  
638 compliance with GLI Method 2) have adopted the use of multi-parameter measurements in order to  
639 improve instrument performance. This innovation has included the measurement of light intensities  
640 at multiple scattering angles, and the use of the ratios of those intensities to infer some of the  
641 physical properties of the scattering suspension, e.g. sphericity (Gibbs 1978), or to negate the effect  
642 of water colour as an interference to the turbidity measurement (Lawler 2005, Lambrou et al. 2009).  
643 An example of another multi-parameter approach to turbidity measurement is presented by Yang &  
644 Hogg (1979), wherein two different wavelengths of light are used to predict the PSD of the scattering  
645 suspension. These and other multi-parameter approaches to turbidity measurement should be the  
646 focus of further research, and will aid the development of new turbidity standards.

## 647 Conclusions

- 648 1. The use of turbidity purely as an indicator of water clarity is entirely acceptable assuming the  
649 development of more consistent standards. The problem is that the existing standards have  
650 introduced a set of measurement units that actually represent a surrogate for turbidity and  
651 therefore cannot be used to describe water clarity.
- 652 2. Simple turbidity measurements when used as a surrogate for suspended sediment  
653 concentration are only viable under highly constrained conditions. Bias toward the fine  
654 sediment fraction is usually considered unimportant, but this is not always the case.
- 655 3. Sand-sized sediment fractions are not consistently accounted for by existing turbidity  
656 measurements, due to their high settling velocities. The SSC method is also required in order  
657 to quantify the sand fraction fully.
- 658 4. The development of new light-scattering models will permit more sophisticated approaches  
659 to turbidity measurement, in particular by the use of parameter-rich data sets obtainable

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3 660 from multi-parameter methods. This approach will facilitate the improvement of turbidity  
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5 661 standards, and could increase the accuracy of large sediment particle detection.  
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7 662 5. A new turbidity instrumentation standard needs to be drafted, based purely on the principle  
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9 663 of attenuation for calibration and reporting purposes. It should specify the reporting of the  
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11 664 BAC in dB m<sup>-1</sup> (or derived units) for a range of measurement angles and wavelengths of light.  
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13 665 This standard should be a root standard from which other secondary standards are derived,  
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15 666 e.g. standards for suspended sediment characterisation or total suspended solids  
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17 667 assessment by optical turbidity measurement.  
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20 668 6. A further standard for suspended sediment determination by simple multi-parameter  
21  
22 669 turbidity measurements needs to be devised (leading on from point 4 above). This standard  
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24 670 should include basic sediment characterisation as an outcome of optical turbidity  
25  
26 671 measurements (e.g. PSD and sphericity).  
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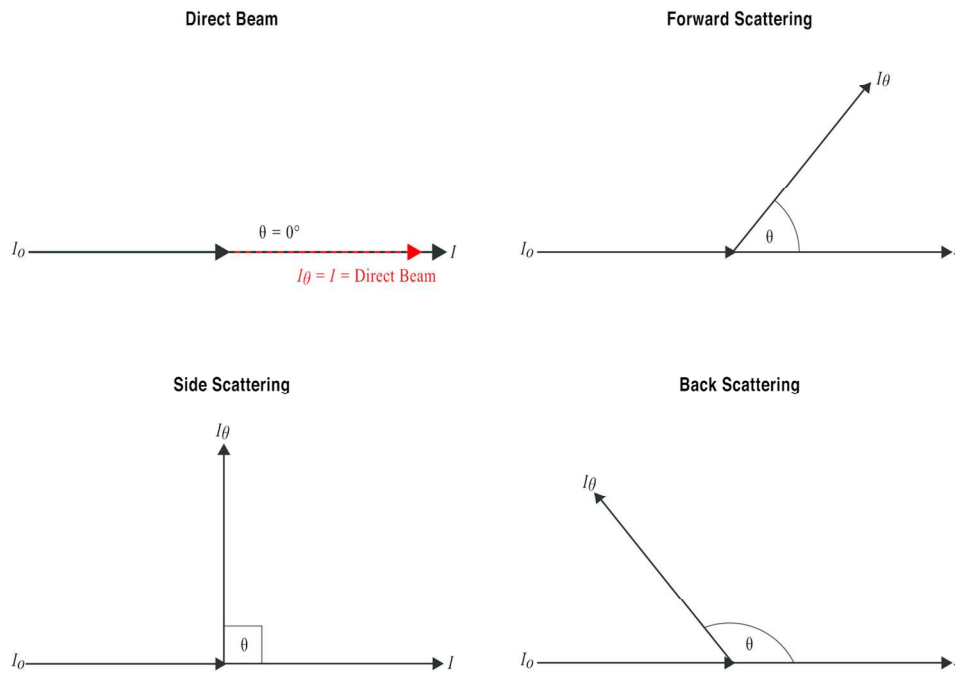


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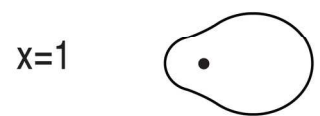
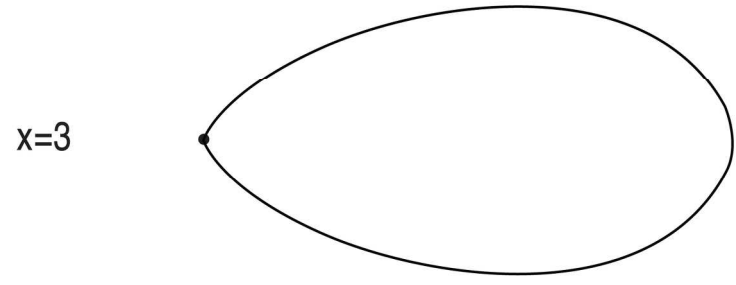
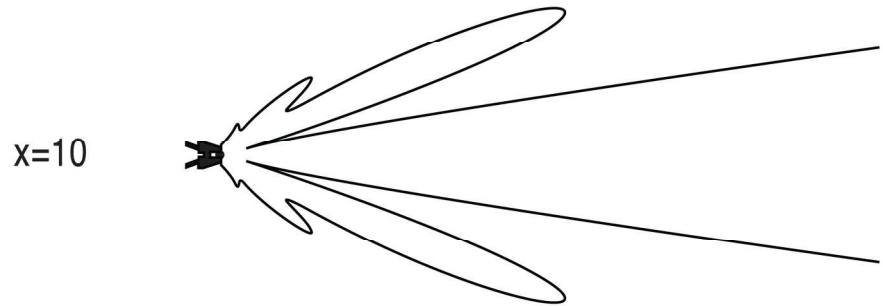
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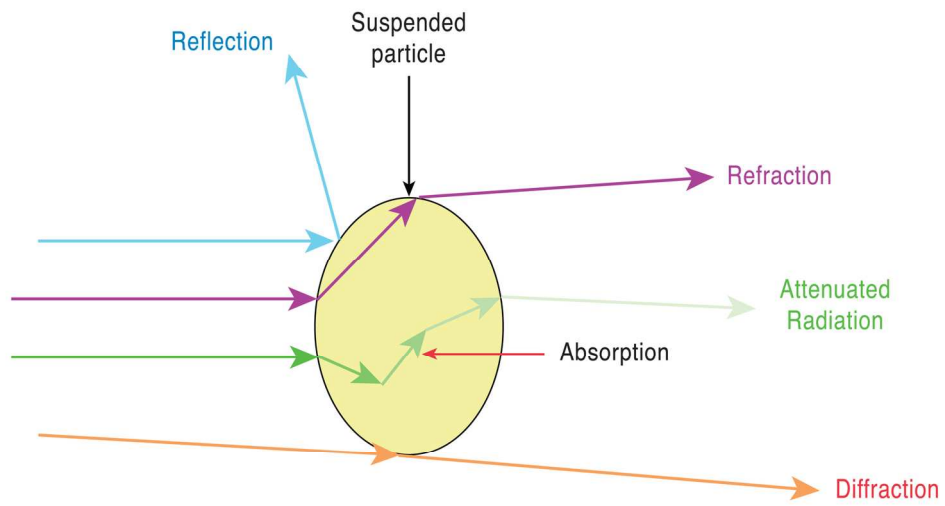
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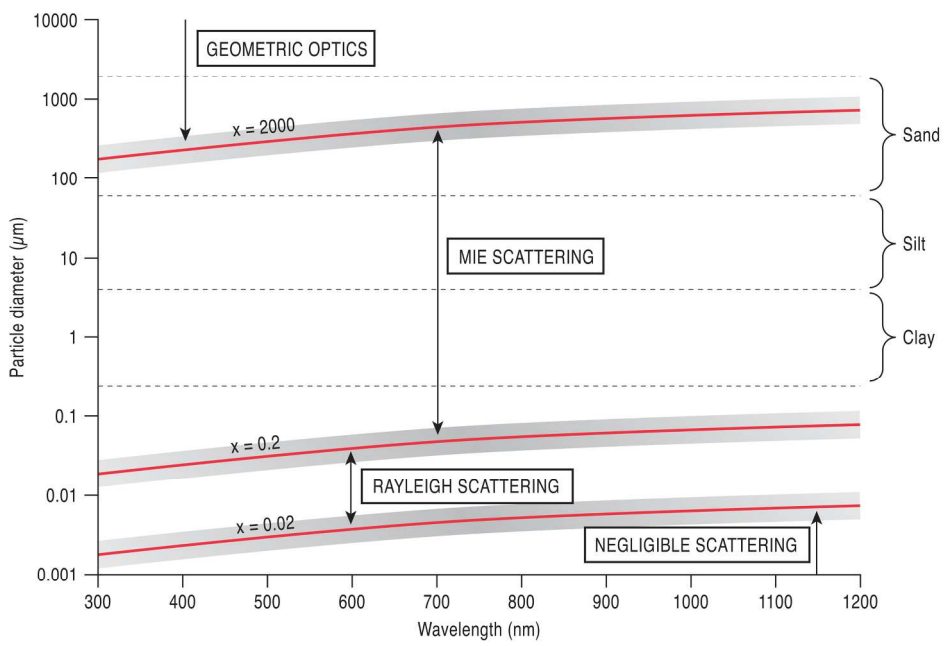
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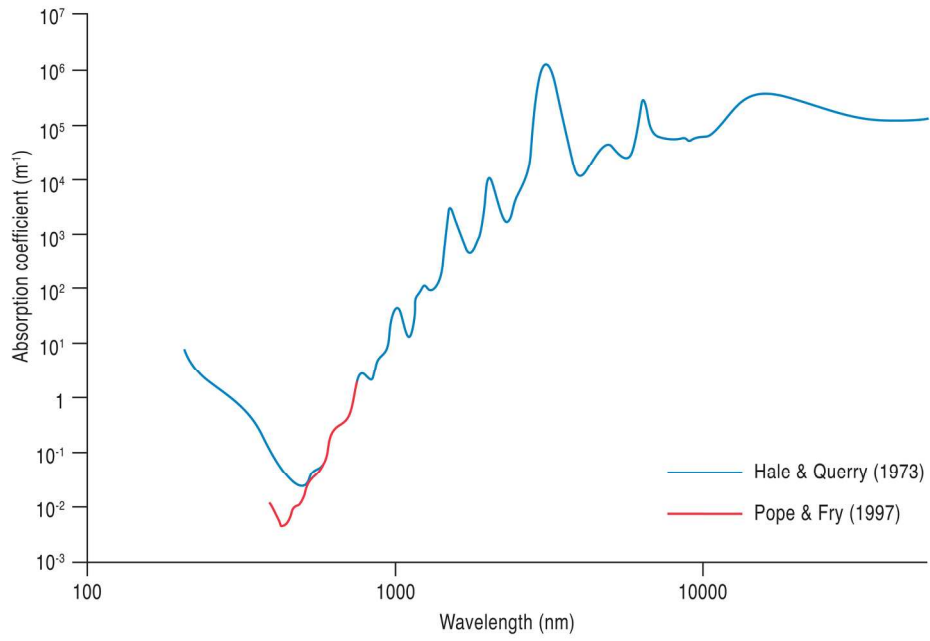
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103x72mm (600 x 600 DPI)

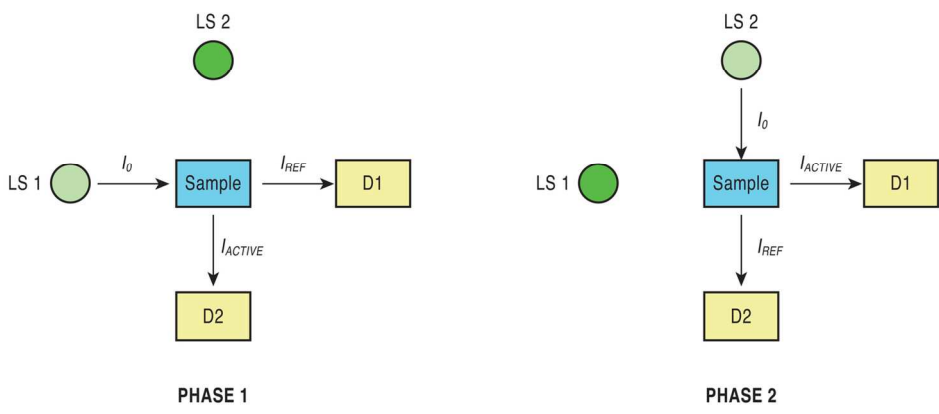
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96x69mm (600 x 600 DPI)

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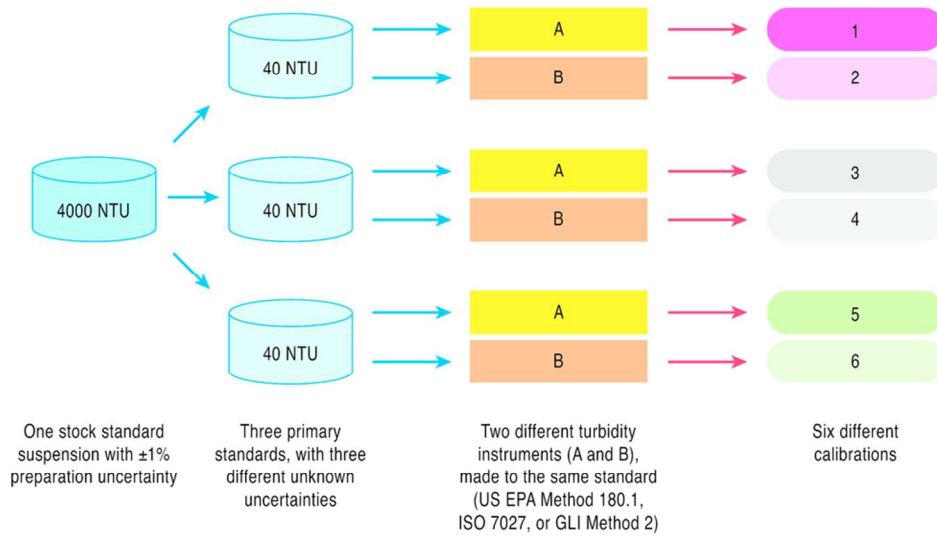


66x31mm (600 x 600 DPI)

Peer Review



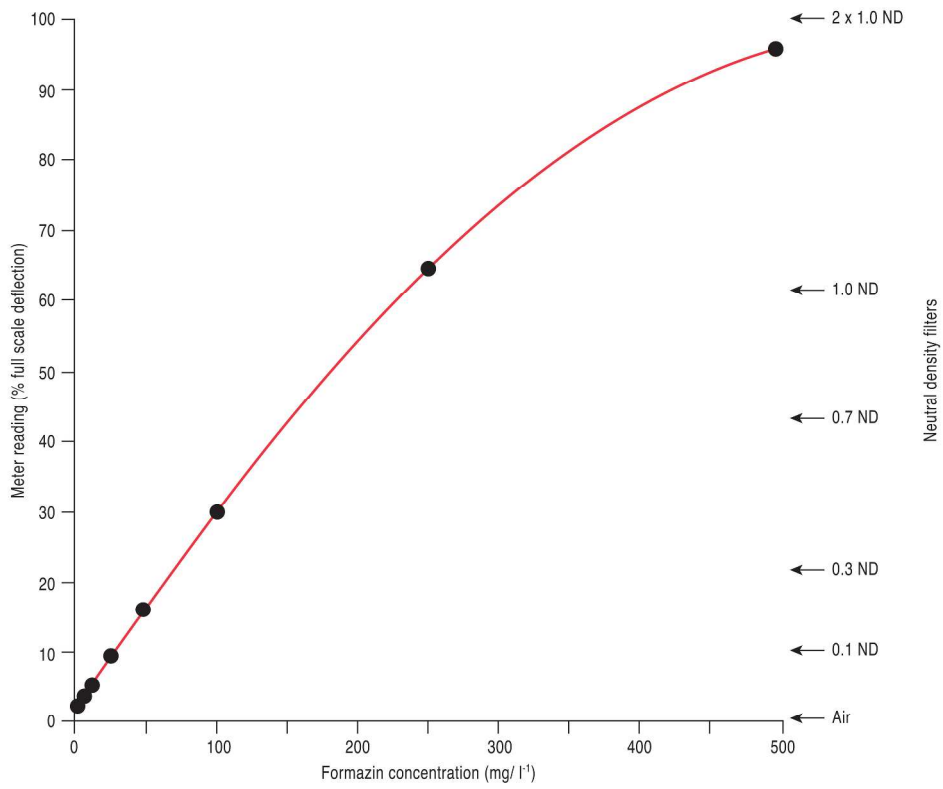
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88x52mm (300 x 300 DPI)

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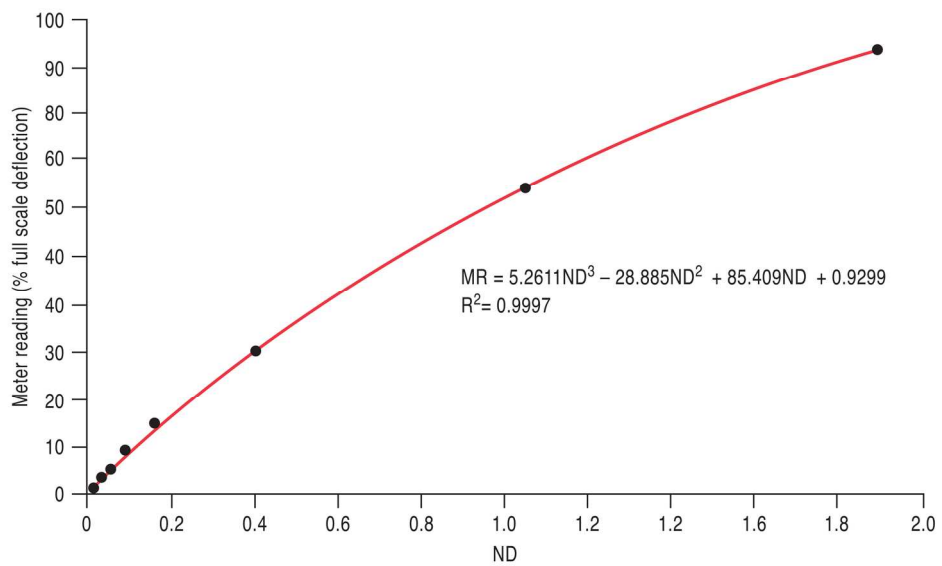
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133x112mm (600 x 600 DPI)

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85x54mm (600 x 600 DPI)

Review

Reference	Scattering-regime and scattering-angle referenced in the text.		
	Transmitted	Back-scattered	Forward-scattered
Agrawal et al. (2008)	<i>Implied 0°</i>		<10°
Bilro et al. (2010)	180°		
Campbell et al. (2005)		180°	
Fugate & Friedrichs (2002)		Angle not defined	
Green & Boon (1993)		>150°	
Guillén et al. (2000)		Angle not defined	
Gumprecht & Sliepcevich (1953)	Angle not defined		
Jansson (1992)	<i>Implied 0°</i>		12°
Morais et al. (2006)	Angle not defined		
Pavanelli & Bigi (2005)		90°	
Sadar (2004, Fig.4, p.8)	180°		
Sadar (2004, Fig.5, p.9)	<i>Implied 0°</i>	140°	
Xu (1997)		Angle not defined	
Yang & Hogg (1979)	Angle not defined		

Units of Measurement for Turbidity Sensors		
	Wavelength of Light Source	
	White or broadband: peak spectral output of 400-680 nm	Infrared, monochromatic: typical output in 780-900 nm range
<b>Single Illumination Beam Light Source</b>		
90° to incident beam; single detector	<b>Nephelometric Turbidity Unit (NTU)<sup>a</sup></b>	<b>Formazin Nephelometric Unit (FNU)<sup>b</sup></b>
90° and other angles; multiple detectors; instrument algorithms use combination of detector readings and ratio techniques	Nephelometric Turbidity Ratio Unit (NTRU)	Formazin Nephelometric Ratio Unit (FNRU)
30°±15° to incident beam (backscatter)	Backscatter Unit (BU)	Formazin Backscatter Unit (FBU)
30°±15° and other angles; multiple detectors; instrument algorithms use combination of detector readings and ratio techniques	Backscatter Ratio Unit (BRU)	Formazin Backscatter Ratio Unit (FBRU)
0° to incident beam (attenuation)	Attenuation Unit (AU)	Formazin Attenuation Unit (FAU)
<b>Multiple Illumination Beam Light Source</b>		
90° and possibly other angles; multiple detectors; instrument algorithms use combination of detector readings	Nephelometric Turbidity Multibeam Unit (NTMU)	Formazin Nephelometric Multibeam Unit (FNMU)
<sup>a</sup> <b>NTU:</b> limited to instruments that comply with EPA Method 180.1. <sup>b</sup> <b>FNU:</b> pertains to instruments that comply with ISO 7027, the European drinking-water protocol. This includes many of the most commonly used submersible turbidimeters.		

Characteristic	USEPA Method 180.1 (non-ratio mode)	ISO Method 7027 (diffuse radiation)	ISO Method 7027 (attenuated radiation)	GLI Method 2
Use of data	Drinking water	Drinking water	Wastewater	Drinking water
Range of method	0-40 NTU (dilution permitted)	0-40 FTU (dilution permitted)	40-4000 FAU	0-40 NTU (dilution permitted)
Light source	Tungsten lamp	Photodiode	Photodiode	Photodiode
Wavelength	400-600 nm	860 nm	860 nm	860 nm
Spectral bandwidth	Not specified	60 nm	60 nm	60 nm
Detector orientation measurement angle	$90^\circ \pm 30^\circ$	$90^\circ \pm 2.5^\circ$	$90^\circ \pm 2.5^\circ$	Two sources, two detectors at $90^\circ \pm 2.5^\circ$
Aperture angle	Not specified	$20^\circ$ - $30^\circ$	$20^\circ$ - $30^\circ$	Unknown
Path length	Less than 0.1 m	Less than 0.1 m	Less than 0.1 m	Less than 0.1 m
Primary standards	Formazin polymer	Formazin polymer	Formazin polymer	Formazin polymer
Secondary standards	Polymer microspheres	Polymer microspheres	Polymer microspheres, cubes, or filaments	Polymer microspheres

Formazin standard concentration	Stability duration
> 400 NTU	1 year
20 – 400 NTU	1 month
2 – 20 NTU	12 – 24 hours
< 2 NTU	<= 1 hour
<= 1 NTU	Difficult to prepare accurately

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3 Figure 1 Illustrations of the light-scattering angle convention: the “direct beam” where  $I_{\vartheta} = I$ ,  
4 forward-scattering, side-scattering and back-scattering. The incident beam is  
5 denoted  $I_0$  and the direct transmitted beam at  $0^\circ$  to the incident beam is denoted  $I$ .  
6 The scattered beams are denoted  $I_{\vartheta}$ , where  $\vartheta$  is the scattering angle with respect to  
7 the incident beam.  
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14 Figure 2 Scattering phase functions derived from Mie theory, with light incident from the left  
15 of the diagrams. Forward scattering becomes more pronounced as  $x$  increases.  
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19 Figure 3 The scattering processes of reflection, refraction and diffraction, and the  
20 attenuation process of absorption of light due to a particle suspended in water.  
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24 Figure 4 Light scattering theory regimes as a function of particle diameter and wavelength of  
25 light. Also shown are sediment particle size bands according to the American  
26 Geophysical Union Sediment Classification System.  
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31 Figure 5 The light absorption spectrum of water. After Hale & Querry (1973) and Pope & Fry  
32 (1997).  
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37 Figure 6 Beam-ratio process as described in GLI Method 2. LS 1 & LS 2 are the light sources;  
38 D1 and D2 are the detectors.  $I_0$  is the light beam incident on the sample;  $I_{\text{ACTIVE}}$  is the  
39  $90^\circ$  scattered light and is considered to be the actual nephelometric measurement;  
40  $I_{\text{REF}}$  is the  $0^\circ$  transmitted light and is used purely as a reference value for use in a  
41 ratio-metric calculation.  
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48 Figure 7 An example of the effect of indeterminate PSD due to identically defined but  
49 potentially physically dissimilar primary turbidity standards on the calibration of  
50 turbidity instruments. Results are further confounded by the variability in response  
51 between different instruments to the same PSD.  
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3 Figure 8 Laboratory calibration of a turbidity meter with Formazin standards. Meter readings  
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5 of the neutral density filters used in the field are shown also (Finlayson 1985).  
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8 Figure 9 A reproduction of the data contained in Figure 8 showing the meter reading vs. the  
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10 ND filter value (after Finlayson 1985). The ND value is equivalent to  $d$ , the optical  
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12 density.  
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For Peer Review

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3 Table 1 A selection of turbidity literature references illustrating the ambiguity associated  
4 with the assignment of scattering-regime nomenclature to the actual scattering-  
5 angle.  
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10 Table 2 Units of Measurement for Turbidity Sensors, after USGS website (USGS 2013).  
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13 Table 3 Summary of turbidity test methods after Ziegler (2003), where NTU are  
14 nephelometric turbidity units, FTU are formazin turbidity units, and FAU are  
15 farmazin attenuation units.  
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20 Table 4 Stability of formazin standards, after Buzoianu (2000).  
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