# **Natural Hazards**

## Exploring representativeness and reliability for late medieval earthquakes in Europe --Manuscript Draft--

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| Abstract:  | Seismic catalogues of past earthquakes have compiled a substantial amount of information about historical seismicity for Europe and the Mediterranean. Using two of the most recent European seismic databases (AHEAD and EMEC), this paper employs GIS spatial analysis (Kernel density estimation) to explore the representativeness and reliability of data captured for late medieval earthquakes. We identify those regions where the occurrence of earthquakes is significantly higher or lower than expected values and investigate possible reasons for these discrepancies. The nature of the seismic events themselves, the methodology employed during catalogue compilation and the availability of medieval written records are all briefly explored.                                      |
| Response to Reviewers:                           | Dear Editor in Chief,   |
|  | We accepted and integrated the minor corrections suggested by reviewer #2.<br>We just added a very minor integration into the conclusions of our work. We integrated<br>archaeoseismology with palaeoseismology as we believe that also this discipline has<br>the tools for improving our knowledge about past earthquakes in those regions that<br>suffer for the lack of written documents.<br>We also added a missing reference about the source we used for identifying the<br>largest medieval European cities (Jotischky and Hull 2005) and updated the map<br>crossing KDE of late medieval earthquakes with these urban centres (editing the<br>legend and the colour of the point shape file).<br>In our opinion, this paper is now in its final version.<br>Yours sincerely,<br>Paolo Forlin |

#### David Petley

| Exploring representativeness and reliability for late medieval earthquakes in Europe   |
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| Abstract: 150 – 250 words  |
| Seismic catalogues of past earthquakes have compiled a substantial amount of<br>information about historical seismicity for Europe and the Mediterranean. Using<br>two of the most recent European seismic databases (AHEAD and EMEC), this paper<br>employs GIS spatial analysis (Kernel density estimation) to explore the<br>representativeness and reliability of data captured for late medieval earthquakes.<br>We identify those regions where the occurrence of earthquakes is significantly<br>higher or lower than expected values and investigate possible reasons for these<br>discrepancies. The nature of the seismic events themselves, the methodology<br>employed during catalogue compilation and the availability of medieval written<br>records are all briefly explored.<br>Keywords: earthquakes, historical seismicity, Late Medieval Europe, GIS, kernel   |
| density estimation, risk   |
| A key initiative in historical seismology in recent years has been the collection of<br>earthquake data at a continental scale, especially for Europe. AHEAD (Archive of<br>Historical Earthquake Data; Locati <i>et al</i> 2014;<br>http://www.emidius.eu/ahead/main/) and SHEEC (SHARE European<br>Earthquake Catalogue 1000-1899; Stucchi <i>et al</i> 2013;<br>http://www.emidius.eu/SHEEC/sheec_1000_1899.html) have developed<br>systematic catalogues of past seismic events between AD 1000 and 1899,<br>generating and publishing a robust archive of macroseismic information. A third<br>project, EMEC (the European-Mediterranean Earthquake Catalogue; Grünthal<br>and Wahlström 2012), consists of a unified catalogue of earthquakes with an M <sub>W</sub><br>higher than 3.5 in Europe, Mediterranean Africa, Turkey and Cyprus up to 2006.<br>In contrast to the other two catalogues, EMEC is mainly based on instrumental<br>recording of recent solution. |
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43 These meta-data archives are primarily intended as inputs into the seismic hazard assessments which remain fundamental to the development of strategies 44 45 for earthquake risk reduction (through, in the case of the SHARE project, Seismic 46 Hazard Harmonization in Europe). The valuable contribution of historical 47 seismology to national and international earthquake catalogues has long been 48 recognised (Ambraseys 1971; Gürpinar 1989; Voght 1991; Caputo and Helly 49 2008) but these new larger databases now open up fresh possibilities for 50 research. Not only do they add a remarkable volume of data which has been 51 standardised according published criteria, but both AHEAD to 52 (http://www.emidius.eu/ahead/main/) EMEC and (http://emec.gfz-53 potsdam.de) also operate on open access online platforms and embed useful 54 tools for geographical and chronological interrogation.

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56 For more than a generation research tools have been available to seismologists 57 to help evaluate the completeness of historical earthquake catalogues as time 58 series data (eg. Stepp 1972; Weichert 1980; Woessner, Wiemer 2005; Hakimhashemi, Grünthal 2012; Alamilla et al. 2014) but it is now possible to 59 60 supply a spatial as well as a chronological assessment of past events. In this paper therefore we explore the use of kernel density estimation (KDE) to 61 62 investigate the representativeness of the historical seismic activity in Europe in 63 the late Middle Ages (here defined as AD 1000-1550) from a geographical 64 perspective. We identify those European regions where our knowledge of 65 medieval seismicity is especially weak and we ask whether medieval seismicity 66 is sometimes overestimated.

67

## 68 The earthquake record over time

69 It is well understood from numerous case studies that the available information 70 for some European regions and periods is better than it is for others (for 71 instance Guidoboni and Comastri 2005 for 11<sup>th</sup> to 15<sup>th</sup> Century earthquakes in 72 the Mediterranean region) and this point is quickly underlined by an analysis of the AHEAD dataset for the last 1000 years (Figure 1). As has been noted 73 74 previously for other datasets (Daniell et al. 2011), the number of recorded 75 earthquakes per year increases through time: there are far more earthquakes 76 known from the modern period of instrumental monitoring than there are from 77 patchy historical records. Thus, the number of recorded earthquakes in the 11<sup>th</sup> 78 century (n=30), for example, represents less than 2% of the earthquakes 79 catalogued for the 19th century (n=2432). In fact, the number of known 80 earthquakes approximately doubles with each passing century.

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The reasons for this disparity are also well rehearsed (Guidoboni and Ebel 2009,
for example). They include the comprehensiveness and reliability of any
individual account of a historical seismic event, the preservation and

85 transmission of that record (which may be one of a number which provide 86 evidence of a single event) and the capacity of modern compilers and analysts to 87 recognise and catalogue the event. When researchers claim that 'libraries may 88 hide hundreds of treasures that are mostly unknown to seismologists' (Vogt 89 1991) they concede the degree to which research intensity varies across 90 European regions. In short, it cannot be assumed that current catalogues, vital 91 though they may be, are homogeneous in their representation of past seismicity. 92 The key question to ask is precisely *where* the strengths and weaknesses of the 93 data might lie.

94

## 95 **KDE analysis**

96 Point density analysis is a technique that permits the visualization and 97 consideration of clusters in a spatial dataset and facilitates comparison of trends 98 (Conolly and Lake 2006). In this case we have used this approach to undertake 99 an evaluation of earthquake distribution across Europe in the late Middle Ages. 100 matched against later seismic activity. By taking earthquake epicenters and 101 magnitudes as the input point layer, a continuous density surface is created. Applying kernel density estimation (KDE), a non-parametric technique (Illian et 102 al 2008; Wand and Jones 1995), the probabilistic density of earthquake 103 104 epicenters is then calculated within a circular area (the KDE 'search radius'). The 105 density value of each output raster cell is obtained by summing the values of all 106 the kernel surfaces calculated for the population of points, the kernel function 107 being based on the quadratic function described in Silverman (1986, p.76, 108 equation 4.5) and available in Esri ArcGis 10.3. The resultant KDE maps (Figures 109 2-5) apply a search radius of 200 km to measure densities in point distribution at 110 a regional scale with an output resolution (pixel dimension) of 5km.

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112 In order to obtain a mean to compare earthquake distributions over time, KDE 113 analysis was undertaken for selected datasets showing medieval earthquakes 114 (AD 1000-1550), post-medieval or early modern historical earthquakes (AD 115 1551-1899), and 20<sup>th</sup> century (AD 1900-1999) earthquakes. Calculations were 116 applied both to the entire number of the recorded earthquakes collected within each dataset and for the earthquakes with  $Mw \ge 5$ , introducing a threshold which 117 excludes events that cause little damage. The KDE maps are then displayed using 118 119 a coloured key which defines density trends. To avoid redundancies during comparison, the density values in each case were homogenised to a range of 120 121 values ranging from 0 to 100. The mean values of density were extracted from 122 the maps using a zonal statistic analysis and assigned to a shape file displaying 123 the provinces (1248 in total) of all European countries. This allows differences in 124 mean value density to be calculated and then displayed.

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Figure 2 is the KDE map of known late medieval earthquakes for the period AD
1000-1550. Some 567 events are shown, displaying a density peak across

128 northern and central Italy. The highest values are concentrated in Umbria, 129 Northern Tuscany and Marche, and the central and the east Po Plain, including 130 the Eastern Alps between Friuli and Slovenia. High values are also identified for 131 the central Apennines and south towards Campania, eastern Sicily (including the 132 area of Reggio Calabria in Calabria) and the Rhine Graben area, which 133 corresponds to the Basel region, Lower Alsace and Baden-Württemberg. Lower 134 values are found in the eastern Pyrenees, Transylvania in central Romania, 135 Central and Southern Dalmatia, around Aachen (Germany) and between Brussels 136 (Belgium) and Cologne (Germany). If a threshold for stronger earthquakes equal to  $Mw \ge 5$  is applied to this dataset (Figure 3), the map shows a marked 137 138 concentration of high values in central and northern Italy.

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140 Figure 4 processes the epicentres of post-medieval earthquakes (AD 1551-1899; n=3840) using the same methodology. Although central Italy and the Strait of 141 Messina are still characterised by high values, higher density values still are 142 143 visible in both Switzerland and Slovenia. For this period, medium values are 144 found across Transylvania (Romania) and along the western border between 145 Slovakia and Hungary. Again, Greece, Albania and Andalusia show lower values, 146 comparable with those for the Pays de la Loire (France) and Aachen region 147 (Germany). If a threshold of  $Mw \ge 5$  is applied to this dataset, the post-medieval 148 map changes dramatically (Figure 5). Peak values shift to central Italy, 149 Transylvania in Romania, and the Adriatic coast of Albania and Greece. Higher 150 values are also obtained for the Strait of Messina and Calabria (Italy), Slovenia 151 and the Belgrade area in Serbia. Average values are registered for Eastern 152 Bulgaria and southern Andalusia (Spain), and lower ones for Switzerland, the Aachen area, the Pyrenees and southern Portugal. 153

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155 Figure 6 presents the KDE for 20<sup>th</sup> century earthquakes extracted from the EMEC database (n=23,438). Here the picture for the instrumental period is very 156 157 different from the catalogue of historical earthquakes. The area with the highest 158 peak of density values now focuses on Greece and Albania, extending southward 159 to the Hellenic Arc and including Crete and Rhodes. Transylvania in Romania 160 represents another peak but, taken together, this pronounced clustering makes 161 the visualisation of areas with lower values appear undifferentiated. 162 Fortunately, this can be overcome by introducing a non-linear binning technique, 163 such that the application of breaks in the data have geometric rather than linear 164 break to emphasise the distribution of lower value events. This makes it possible 165 to identify other regions for which seismic activity is noteworthy. This includes 166 the Balkan Peninsula as a whole, Slovenia and the eastern Alpine arch between Italy and Austria, central Italy and the area of the Strait of Messina, Switzerland, 167 168 and Iceland. Areas characterised by lower seismic activity are the western 169 Pyrenees, the Aachen region, Andalusia and Murcia and around Lisbon.

171 In this case introducing a magnitude threshold of  $Mw \ge 5$  does not change the 172 picture significantly (**Figure 7**). The highest values still centre on Greece, 173 Albania and the Hellenic Arc, with a new peak now identifiable in Iceland. Italy, 174 northern Switzerland, and the Rhine Graben are still characterised by high 175 seismicity, though Andalusia, Murcia and Lisbon area are also included in this 176 class. Lower values are assigned to the western Pyrenees, northern Portugal and 177 Galicia, with two smaller zones to the north of Aachen and in western Belgium.

178

## 179 **Discussion**

180 The maps presented in Figures 2 to 6 confirm that there are significant differences in our understanding of seismicity for different periods of our 181 182 historic past. To draw out these patterns further, **Figure 8** evaluates differences 183 in the spatial distributions in the density of recorded earthquake events from the 20<sup>th</sup> century and the later medieval period. This map was created through a 184 three-step process. First, given that density maps of different periods have 185 186 differing value ranges, these values have been normalised to obtain a common value ranging from 0 to 100. Second, zonal statistical analyses facilitate the 187 188 interrogation of the KDE maps against a vector shape file of the modern day 189 provinces of the European Union and adjoining territories. This generates mean 190 values of epicentre density for each of the 1248 European districts which, in 191 turn, can be exported to produce tables for late medieval, post-medieval and 20<sup>th</sup> 192 century epicentre densities and then joined with the shape files so as to visualise 193 differences in density trends between the KDE raster maps. Third, for ease of 194 identification, differences in density values between late medieval KDE maps and 195 20<sup>th</sup> century KDE maps have been calculated for all the epicentres as well as for 196 epicentres with  $Mw \ge 5$ .

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198 By comparing several datasets, Figure 8 highlights the extent to which the 199 recorded distribution of 20th century seismic events differs from those that 200 occurred in the Middle Ages. Once more, it is possible to identify where higher 201 and lower than anticipated levels of activity are located. Negative values (in blue 202 on **Figure 8**), which indicate lower than anticipated levels of medieval 203 seismicity, are focused on two areas: eastern Europe and the eastern 204 Mediterranean, including the Balkans, Romania, Greece and Crete, and Iceland. 205 The most under-represented areas lie in the south of Albania, around the Gulf of 206 Corinth and Crete. By contrast, those regions showing a higher than expected 207 level of medieval seismicity (in red on Figure 8), when compared to contemporary seismicity, can be found in Western Europe, especially in 208 209 Andalusia, the eastern Pyrenees, Switzerland, the Aachen region, northern and 210 central Italy, the Strait of Messina, Slovenia and Dalmatia. Peaks in positive 211 values centre on northern and central Italy.

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213 The higher than expected spatial density of earthquakes may in part be explained by the nature of the seismic event itself. While it could be argued that, 214 215 over a period of a century, the spatial and temporal distribution of small to 216 medium earthquakes on a continental scale might be approximately constant, 217 the recurrence interval of seismic events scales with time, such that the largest 218 earthquakes occur the least frequently; the recurrence period of a large 219 earthquake on a given active fault might be typically in the order of a century to a 220 few millennia. The largest earthquakes, because they have much longer return 221 periods, introduce greater temporal and spatial variability. Once more, there 222 may be a high occurrence of aftershocks after a very large seismic event and, 223 where a large earthquake has occurred, it would be expected that a number of 224 small to medium earthquakes might also strike in the same region.

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226 It is also the case that two earthquakes of the same magnitude may not have the 227 same consequences for above ground structures because of the nature of local 228 geology and geomorphology. For example, variations in rupture speed may affect 229 the frequency of the shaking experienced at ground level, changing the damage 230 potential of the earthquake. In addition, different continental areas have 231 different attenuation characteristics which affect the distribution of ground 232 shaking. In central Greece, for example, strong earthquakes have been described 233 with a Mw between 6.5 and 7.2 but with only very localised impacts (Ambraseys 234 and Jackson 1990; Stiros and Pytharouli 2014). Another important influence on 235 our mapping is the method by which events have been recorded by catalogue 236 compilers. The observed peaks of post-medieval earthquakes in Switzerland and 237 Slovenia, for example, are probably due to the recording of a large number of low 238 impact aftershocks as independent earthquake events in these regions and also 239 to some extent the comprehensive research which has been undertaken by the 240 Swiss Seismological Service (Fäh et al. 2011; Živčić 2009, as reported in Stucchi 241 et al. 2013: 533).

242 Question marks concerning over- and under-reporting may apply equally to the 243 later medieval period and, to investigate this possibility further, Figure 9 244 displays the KDE map of medieval earthquakes but this time including all late 245 medieval cities with a population above 10,000 inhabitants (Jotischky and Hull 246 2005: 73). What emerges is a positive relationship between the density of 247 recorded seismic events and the distribution of these more significant 248 settlements. Thus, larger numbers of people in medieval urban areas, 249 particularly those in literate institutions such as monasteries and universities, 250 presented not only greater opportunities for damage to occur but also for that 251 damage to be observed and recorded as a seismic episode. Towns and cities 252 which were better connected to national and European trading networks with 253 substantial numbers of visitors, pilgrims and merchants also multiplied many 254 times over the opportunity for comment well outside the affected region. 255 Furthermore, in a risk sensitive society in which earthquakes occurred more 256 frequently and measures of hazard adaptation and mitigation were better 257 understood (e.g. structural assessments, financial relief, reconstruction, etc) 258 there was perhaps a greater propensity to evaluate and record in order to justify 259 a civic or State response (Gerrard and Petley 2013). As an illustration of this, one 260 of the best-documented and most destructive seismic events in the Late 261 Medieval Italy struck the southern Apennines and the Naples region in December 262 1456 with an estimated  $Mw = 7\pm0.30$  and a maximum Intensity Io=11 (total 263 *destruction*)(Meletti et al. 1988). Information about this earthquake which was 264 probably composed of three distinct but coincident seismic events (Teramo et al. 1999) is derived from more than 60 different historical sources, including royal 265 266 privileges given in the aftermath, ambassadors' letters, reports, chronicles, 267 scientific treatises and inscriptions (Figliuolo 1988; Guidoboni and Comastri 268 2005). Not surprisingly, the number of known affected locations is also 269 remarkably high; 199 different places recorded the event in one form or another. 270 On the other hand, the Xylocastro earthquake, which affected the Gulf of Corinth 271 in June 1402, was one of the strongest earthquakes recorded in the late medieval 272 Greece (Guidoboni and Comastri 2005). This tsunamigenic event had an 273 estimated Mw =  $6.6 \pm 0.35$  and Io=10 (University of Thessaloniki 2003) and yet 274 just eight places are recorded as being affected along the shores of the Gulf, and 275 only two historical sources provide any information at all about the event: one is 276 a letter written by a Venetian merchant, the other a chronicle from the city of 277 Ferrara (in Italy).

278 If the earthquake data for Italy and Greece is examined over time rather than 279 spatially, further patterns emerge. Figure 10 shows the 269 known earthquakes in Italy between 1000 and 1550 AD at 50 year intervals, plotted alongside the 48 280 recorded earthquakes for Greece. Whereas the trend for Italy is quite simple if 281 282 non-linear, with more earthquakes in more recent centuries, that for Greece is 283 more variable. With the exception of limited numbers of monastic archives (such as those at Monte Athos and island of Patmos), documents for the Middle and the 284 285 Late Byzantine period are almost completely absent (Tsougarakis and Angelomatis-Tsougarakis 2012). From the 13<sup>th</sup> century, the situation improves 286 287 as commercial contacts improved with the West, for example with Venice, and 288 the presence of new institutions such as the Military Orders on Rhodes, Cyprus 289 and elsewhere (e.g. during the 1493 earthquake of Kos; Figliuolo 2002). Only 290 from the 15<sup>th</sup> century did archives become richer as a consequence of integration 291 into the Ottoman Empire with a subsequent growth of ecclesiastical and monastic archives (Tsougarakis and Angelomatis-Tsougarakis 2012). 292

293

#### 294 Conclusion

295 This paper highlights some of the strengths and weakness of current historic 296 earthquake meta-datasets. While seismologists have long been aware of the 297 incompleteness of their catalogues, we offer this KDE comparison as another tool 298 in the toolbox, one that provides better geographical definition. The results 299 immediately suggests an agenda for further investigation, particularly across 300 eastern Europe and the eastern Mediterranean where our methodology suggests 301 that there were more and more powerful seismic events during the Middle Ages 302 than have hitherto been recorded. For some of these areas archaeoseismological 303 and paleaosesimological projects might shed new light on historical seismic 304 events, otherwise a more detailed assessment is required of the information gap 305 resulting from a scarcity of written documents. Finally, we also highlight here the 306 issue of over-recording, something which may be explained by the nature of the 307 seismic event and the density of human settlement combined with regional 308 cultural and social factors, including the more sophisticated development of risk-309 sensitive tactics.

310

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#### 326 **Captions**

Figure 1. Numbers of recorded earthquake events across Europe by century (source:AHEAD 2014).

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Figure 2. Distribution of late medieval (1000-1550 AD; source: AHEAD 2014)
earthquake epicentres (sx) and the associated KDE (dx).

333Figure 3. Distribution of late medieval (1000-1550 AD; source: AHEAD 2014)334earthquake epicentres with  $Mw \ge 5$  (sx) the associated KDE (dx).335

Figure 4. Distribution of post-medieval (1551-1899 AD; source: AHEAD 2014)
earthquake epicentres (sx) and the associated KDE (dx).

Figure 5. Distribution of post-medieval (1551-1899 AD; source: AHEAD 2014) earthquake epicentres with  $Mw \ge 5$  (sx) and the associated KDE (dx).

341

- Figure 6. Distribution of 20th century earthquake epicentres (sx; source: EMEC 2012)and the associated KDE (dx).
- Figure 7. Distribution of 20th century earthquake epicentres with  $Mw \ge 5$  (sx; source:346EMEC 2012) the associated KDE (dx).
- 347348 Figure 8. Calculated difference between the KDEs for medieval and 20th century349 earthquakes.
- Figure 9. Medieval cities with a population higher than 10k inhabitants (in AD 1300 ca.)
  and the KDE of late medieval earthquakes.
- Figure 10. Recorded seismic events in Greece and Italy between 1000 and 1550 AD, herecalculated for 50 year intervals

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