Evidence for nitrogen accumulation: the total nitrogen budget of the terrestrial
 biosphere of a lowland agricultural catchment.

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11 Abstract

12 Several national-scale studies have shown that reactive N is accumulating in 13 developed countries even when only the terrestrial biosphere is considered. However, 14 none of these studies was able to consider the total N budget and so any discrepancy in budgets could be dismissed as being accounted for by N2 exchange. This study 15 considered a large (9948 km²), mixed agricultural catchment where records of N flux, 16 17 land use, climate and population go back at least to 1883. The N inputs were: 18 biological nitrogen fixation, food and feed transfers, atmospheric deposition and 19 inorganic fertilizers. The N outputs were atmospheric emissions (NH₃, N₂O, NO, N₂), 20 direct waste losses and fluvial losses at the soil source. The results showed that, prior 21 to the large-scale use of inorganic fertilizers, the total N budget of the catchment was 22 at steady state with only a small net loss of total N. After the widespread introduction 23 of inorganic fertilizers, the balance of the catchment shifts in favour of the net 24 accumulation. Even accounting for losses to groundwater, the catchment was found to

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have accumulated 315 ktonnes N (315 tonnes/km²) at a rate of 5.5 tonnes N/km²/yr
(55 kg N/ha/yr) over 35 years since 1973. We propose that the accumulation of N
could be occurring in subsoils of the catchment.

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29 Keywords: Total N; N₂; fluvial nitrogen; nitrate

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31 **1. Introduction**

32 Several N balance studies have been conducted for individual countries and 33 regions, e.g. UK (Lord et al., 2002, Worrall et al., 2009), Finland (Salo et al., 2007), 34 Canada (Janzen et al., 2003) and even for the entire European Union (de Vries et al., 35 2011). However, these studies, even the most recent (e.g. Sutton et al., 2011, Ti et al., 36 2011) have been limited to considering reactive nitrogen species only (Nr). Billen et 37 al. (2012), when considering the N imprint of Paris, do consider N₂ uptake but did not 38 consider gaseous releases of N to the atmosphere of any species of N. Schlesinger and 39 Bernhardt (2013) have outlined the global total N cycle. There are a number of 40 reasons for budgets being restricted to Nr mainly the difficulty of including all 41 possible uptake and release pathways for N₂. For example, Lord et al. (2002) 42 suggested that estimates of denitrification to N2 were too uncertain to include in their 43 analysis of the N balance of all UK agricultural land. Indeed, several N balance 44 studies have gone on to assume no long-term net accumulation or depletion in the 45 terrestrial biosphere or at the country level (Galloway et al., 2004, Ayres et al., 1994) 46 because it was assumed that there is steady-state with regard to total N balance and 47 that any imbalance in Nr would be balanced by N2 fluxes. For example, Kroeze et al. 48 (2003), in their consideration of the N budget of The Netherlands, showed an 49 equivalent Nr sink of 469 ktonnes N/yr in 1995 across the entire country, which they sourced was balanced by aquatic and terrestrial denitrification to N_2 ; however, they did not estimate either pathway. Equally, a meta-analysis of 217 field studies by Gardner and Drinkwater (2009) which suggest that on average 29% of applied inorganic N fertilizer was in the soil after one year, and Sebilo et al. (2013) have shown that applied inorganic fertilsiers in French soils upto 15% of applied N fertilizer was still present after 30 years.

56 Worrall et al. (2009) calculated a Nr budget for the UK from 1974 to 2005 and 57 showed that, not only is the UK a "hotspot" for fluvial Nr flux, with higher exports of 58 dissolved nitrogen than any other region of comparable size in the world, but that 59 increasing fluvial fluxes were occurring at time when inputs were steady or declining, i.e. the UK remained a net sink of Nr but the size of that sink was diminishing, 60 61 However, Worrall et al. (2009) could estimate neither aquatic nor terrestrial 62 denitrification. Although they did consider atmospheric emissions from industry of Nr 63 species, they did not consider the atmospheric emission, or consumption. of N₂ from 64 or by industrial sources and so could not give the total N balance of the UK. Worrall 65 et al. (in review) have endeavoured to estimate the first total N budget for a nation by 66 estimating fluxes of N2 from aquatic and terrestrial denitrification and from industrial 67 emissions of N₂ and show that the UK is a net source of total N, although the size of 68 this source has shrunk significantly since 1990.

Budgets for a nation or region are dominated by the largest input and the largest output; in the case of most European countries these would be the input of inorganic fertilizer and emission of N to the atmosphere, respectively. Of course, the first of these is an input to the terrestrial biosphere whereas the other is an output largely from fossil fuel burning. In other words, unlocking of long-term geological storage of N is influencing the contemporary N budget. If the current national or regional budget is only a net source because of releases from geological sources, then the terrestrial biosphere could be a net sink of total N. Howden et al. (2011) have modelled the export of nitrate from the Thames catchment and so doing have suggested that the terrestrial biosphere of the catchment has been a net sink of N_r for many decades. However, they did not consider a full N_r or total N budget and so could not confirm whether the terrestrial biosphere was indeed accumulating N or not.

81 Therefore, this study aims to be the first to estimate a total N budget for the 82 terrestrial biosphere of a region (defined here as the River Thames drainage basin – 83 see below). Furthermore, by constructing budgets over series of years, the long-term 84 trend in total N balance can be established.

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86 2. Approach & Methodology

87 This study has included all major N input and output pathways for the terrestrial 88 biosphere over a length of time set by the length of the shortest record. The terrestrial 89 biosphere defined for the study area was bounded laterally by the watershed of the 90 catchment and the river network, i.e. once water containing nitrogen enters the stream 91 network, it was considered to have left the terrestrial biosphere. In the vertical profile 92 the terrestrial biosphere was taken as bounded by the atmosphere above and the 93 bottom of the soil profile below. Although the bottom boundary was not fixed, we 94 consider water below the water table and the unsaturated zone of the geology of the 95 catchment to be outside the terrestrial biosphere. The lateral extent of the terrestrial 96 biosphere was taken to stop at the boundary with the fluvial network.

97 The total N budget of the terrestrial biosphere of a catchment was defined as98 the balance between the following inputs and outputs:

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$$N_{total} = N_{dep} + N_{bnf} + N_{food} + N_{fert} - N_{atm} - N_{fluv} - N_{ground} - N_{direct}$$
 (i)

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102 where: N_x = the flux of total N due to x, where x is dep = atmospheric deposition; bnf 103 = biological nitrogen fixation; food = trans-boundary food and feed transfers; fert = 104 inorganic fertilizer; atm = atmospheric emissions; fluv = losses to the fluvial network; 105 ground = losses to groundwater; direct = losses from sewage and waste effluent.

106 The inputs considered are: atmospheric deposition of nitrogen (wet and dry 107 and would include N fixed from lightning); biological fixation; net trans-boundary 108 food and feed transfer across the catchment boundary; and synthetic inorganic 109 fertilizer applications. The vast majority of animal wastes in the UK are returned to 110 the land on the same farm as they are produced (or on nearby farm units), so represent 111 an internal transfer; this study has assumed that there is no net loss or gain of animal 112 wastes across the catchment's watershed. Equally, by taking the boundary of the study 113 system as the watershed of a catchment, there is no need to consider trans-watershed 114 exchanges of surface and ground water. Clearly, this assumption could not be applied 115 in cases where significant inter-basin transfers are evident or where the system 116 boundary was defined politically rather than hydrologically.

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This study considers the following output pathways:

(1) Atmospheric emissions (including: NH₃ volatilisation, N₂O, NO and N₂
emissions) from agricultural land use (including forestry) but not emissions from
industrial sources. It was assumed that all industrial emissions come from
geological sources (geosphere) and there are now fossilfuel sources currently
being extracted in the study basin.

(2) The fluvial flux of N species is made up of several components. Firstly, the fluxof N species from the terrestrial biosphere as it leaves the soil and enters the

fluvial network. Secondly, some of the nitrogen entering the fluvial network will
be lost to the atmosphere; this aquatic denitrification is a loss to the atmosphere.
The N species considered within the fluvial fluxes include: nitrate, nitrite,
ammonium, dissolved organic nitrogen (DON), and particulate organic nitrogen
(PON).

130 (3) Flux to groundwater was labelled a loss from the terrestrial biosphere in Equation (i). Although nitrogen recharging to groundwater may well eventually return to 131 132 the river and be included in estimating the flux from the river of the catchment at 133 its outlet at the tidal limit, the groundwater only represents a significant source or 134 sink if the concentration of the N in the groundwater is changing; otherwise it is 135 just a just a stable reservoir of N. Stuart et al. (2007) have shown a significant 136 increase in nitrate concentrations in UK groundwater over recent decades. 137 Further, because the lag time represented by flow through groundwater in the 138 study catchment has been shown to be of the order of 35 years (Howden et al. 139 2011), large lag times would exist that net changes in groundwater storage of N 140 would not necessarily be reflected in surface water monitoring within the 141 timescale available to this study.

(4) The study includes direct waste effluent from sewage treatments works as loss from the terrestrial biosphere of the study catchment. The direct waste effluent flux was included because the fluvial flux of N species was considered at the point it left the soil profile and entered the fluvial network and so does not include N flux into the fluvial network from sewage outfalls. Further, sewage effluent represents the reprocessing and return of food and feed transfers. The difference between the human consumption of N and the direct waste effluent is

the amount of N lost in sewage treatment to the atmosphere or the amountreturned to land.

A schematic diagram of the flows and fluxes considered by this study is shown inFigure 1.

The management of crop residues could represent a flux of nitrogen within the catchment. The UK Government banned the burning of crop residues in 1993 and most residues are now left in field after harvest. However, the extent to which this will impact the terrestrial nitrogen budget is currently not known and values of what this represents in terms of N transfer to the soil were not available. Smil (1999) recognized that records of crop residues are not maintained by any country and so this cannot be estimated here either.

160 The quality of the individual records varied and so this study attempted to 161 assess the uncertainty in each input or output. Most of the nitrogen fluxes needed for a 162 complete N budget were only reported within Government and other published data 163 sources: hence published values were used, rather than being calculated within this 164 study. Since the original data were not available, it was necessary to accept the error 165 estimation provided by each individual source. In some cases, no error or uncertainty 166 estimate was given or the error estimate was not credible. In other cases, although the 167 reported flux error was given as a range, it was not always clear what this error 168 actually represented (e.g. range, inter-quartile range, confidence intervals). The 169 uncertainty estimation associated with the calculation of each pathway is discussed 170 with each pathway. Where no credible error was available for an individual pathway, 171 a default value of ±20% was used - this value was chosen as it represented a median 172 of the credible uncertainty values. Given the uncertainty associated with each input 173 and output, the total N budget for any year was calculated as the sum of individual inputs and outputs and the uncertainty in that estimate was calculated where 500 values for each input or output pathway taken at random from within the range that can be defined for each input or output. In this way estimates of the annual total N balance were calculated together with an estimate of the level of uncertainty involved.

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179 2.1. Study site

180 The River Thames is the second largest river basin in the UK with a catchment area of 9948 km² at the Kingston gauging station in south west London, close to the 181 tidal limit at Teddington (Figure 2). There are two important aquifers in the basin: 182 183 groundwater supplies from the Cretaceous Chalk provide the majority of London's 184 water supply. Further north, Jurassic limestones form locally important aquifers. In 185 between these two aquifers are clay vales where surface runoff is much more 186 important; large areas were drained post-World War II to enable poorly-drained 187 meadows to be converted to arable land. At present, urban areas comprise 16% of the 188 catchment area and include the major settlements of Swindon, Oxford and Reading; 189 the gauged area studied here lies largely upstream of London. About 8% of the basin 190 is forested. The great advantage of the Thames catchment is that, not only has there 191 been a very long period of water quality monitoring, but there are also extensive 192 records of potential driver variables. The following records were used:

Water quality: monitoring at Kingston has been ongoing since 1867 making it the longest water quality record for anywhere in the world. With respect to this study, Howden et al (2010) have established the nitrate concentration and flux record over this period and the DOC concentration record (for DON) at the catchment outlet has been reconstructed. For PON flux the suspended solids records for the catchment from HMS records (outlined below) were examined.

199 Flow: daily mean flow records for the Thames at Kingston from 1883 were obtained 200 from the UK National Rivers Flow Archive (NRFA: http://www.ceh.ac.uk/data/nrfa/; 201 station number: 39001). Thus, N flux records can be calculated from 1883 onwards. 202 Land use: annual agricultural census returns have been compiled for each English 203 parish since 1868 until 1988. In 1989 the UK Government moved to annual, nationalscale reporting with reporting for supra-parish units in 1990, 1995 and 1999. From the 204 205 year 2000 to present, the UK Government returned to reporting annually but only for 206 supra-parish units. Therefore, in order to provide a consistent and coherent record 207 across the period of water quality monitoring, the land-use records were compiled 208 across the entire period 1868-2007 for the UK and for the Thames catchment for all 209 possible years. The two records were then compared so that land use in the Thames 210 catchment in unreported years could be estimated. The annual agricultural census 211 does not cover woodland areas and so the area of woodland, including all forestry 212 types, both commercial and semi-natural, was taken from statistics held by the 213 Forestry Commission (Forestry Commission, 2007) for the years 1924, 1947, 1965, 214 1980, 1990, 1998 - 2002, and 2008. Linear interpolation was used to derive an 215 annual estimate of national woodland area. In order to estimate the area of woodland 216 in the Thames catchment, national data were rescaled to Thames catchment area. The 217 area of urban land in the catchment was taken to be the area left unaccounted for by 218 agricultural land or forestry. For simplicity and for comparison with other models 219 (Worrall et al., 2012a), land use was summarised as arable land, grassland (including 220 permanent and temporary pasture; as well as rough grazing), woodland and urban 221 areas. Livestock numbers (overwhelmingly sheep and cattle) were derived from the 222 same sources to give an annual time series of livestock numbers in the catchment. Of 223 particular note with regard to land use is that Thames catchment under went a 224 considerable land use change with the onset of the Second World War in 1939. The

225 <u>land use of the catchment, which had previously been dominated by grassland,</u>

226 <u>underwent conversion to arable (Howden et al., 2010 – Figure S1).</u>

227 Population: census returns were available for every English county for every decade 228 from 1841, with additional projected numbers from 2001 2007 to 229 (http://www.ons.gov.uk). The population for the Thames catchment was then 230 estimated as a weighted proportion of the area of each county within the catchment. 231 Linear interpolation was used between census years in order to get an estimate of the 232 Thames catchment population in each year of the study.

Climate: detailed rainfall and temperature records have been maintained at Oxford in
the centre of the Thames basin, since the 18th century (Burt and Shahgedanova, 1998;
Burt and Howden, 2011).

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237 2.2. N inputs

238 Consistent atmospheric deposition records for the UK have only been 239 maintained since 1986 (Fowler et al., 2005). However, the fluxes reported by Fowler 240 et al. (2005) were not complete with respect to wet and dry deposition of either 241 reduced or oxidized forms of nitrogen. Therefore, this study extended the record of 242 dry deposition by linear interpolation of the ratio of wet to dry deposition in those 243 years where both were reported. Fowler et al. (2005) only give records to 2001 but 244 further records were available from the Centre for Ecology and Hydrology (CEH: 245 www.ceh.ac.uk) for 2004 to 2006. In order to get flux estimates for 2002 and 2003, 246 linear interpolation was used. Neither Fowler et al. (2005) nor CEH quote an error for 247 their deposition values, but CEH quotes deposition to 2 decimal places implying the 248 error to be of the order of $\pm 1\%$: this study did not consider this a credible error, 249 therefore we have ascribed an error of ±20%. No national depositional data were 250 available after 2006. None of the atmospheric deposition estimates of Fowler et al. 251 (2005), Simpson et al. (2011) or CEH (www.ceh.ac.uk) included an estimate of the 252 atmospheric DON or PON. There is no national-scale monitoring of DON in either 253 wet or dry deposition; however, there is at least one site in the UK where DON in 254 total deposition has been monitored since 1992, although the site is in the north of 255 England some 300 km north of the Thames catchment. Annual figures of DON 256 deposition for this one site have been calculated for period of 1992 to 2003 by 257 Worrall et al. (2006) and this can be extended through to 2007. Total deposition 258 estimates were then rescaled for the area of the Thames catchment. It is acknowledged 259 that the rates of DON deposition in the north of England could be very different from 260 those further south. In order to understand deposition back over the entire period of 261 the study, it is necessary to assume a background level of deposition and at what point 262 in time this began to be exceeded. Howden et al. (2011), by considering the long-term 263 history of nitrate flux from this catchment, suggested that, given very little trend in 264 stream nitrate concentrations prior to the outbreak of WWII, the catchment was at a 265 steady state with respect to Nr species (though not necessarily with respect to total N) 266 and so this study has assumed that there was no significant change in atmospheric 267 deposition to the catchment before 1936. This study has assumed linear decline in 268 atmospheric deposition from the existing monitoring data (1986-2006) back to the 269 year 1936 and then a constant value back to 1867.

Biological nitrogen fixation can occur in all ecosystems and can represent a significant input of nitrogen to the terrestrial biosphere. For agricultural systems the approach of Smil (1999) was used; although updates to this method have been published by Herridge et al. (2008), their updates are for crops and land use types not 274 found in the UK. The area of nitrogen-fixing crops for the Thames catchment was 275 considered to consist exclusively of legumes (predominantly beans and peas) and 276 clover, as part of crop rotation. The area of each of these was available from the land-277 use records for the catchment. For both clover and legumes, the middle estimate of N 278 fixation as reported by Smil (1999) was used. For biological fixation in natural 279 ecosystems, as opposed to agricultural systems, the approach of Cleveland et al. 280 (1999) was used. It should be noted that Vitousek et al. (2013) has suggested these 281 values are an exaggeration. For the Thames it was assumed that the majority of 282 natural ecosystems fell into the classes of temperate forest or temperate grassland as 283 defined by Cleveland et al. (1999). The area of the study catchment that was not 284 under forestry, or under clover or under peas and beans, was taken as equivalent to 285 temperate grassland as defined by Cleveland et al. (1999). The error in the biological 286 nitrogen fixation was calculated by using the ranges in fixation published by Smil 287 (1999) and Cleveland et al. (1999).

288 Nitrogen is redistributed across boundaries with food and feed transfers as 289 well as plant and seed transfers. Boyer et al. (2002) have estimated the food and feed 290 transfer flux of nitrogen for the eastern USA by considering human and animal 291 demand relative to production within the region. Alternatively, Worrall et al. (2009) 292 used commodity trade data to estimate the nitrogen export or import for the UK. For 293 the Thames catchment, this study used the approach of Lord et al. (2002). The 294 agricultural census data for the Thames for cattle and sheep was used and scaled 295 according to average values of the amount and N content of livestock outputs (meat, 296 wool, milk) and values of feed inputs. Equally, the values of amount and N content 297 for crop off-take of N were taken from Lord et al. (2002) combined with land use data 298 for the Thames catchment: input to crops were considered under fertilizers (see 299 below). Any consideration of food and feed transfer must also consider human 300 consumption and human sewage outputs. In terms of total human consumption of 301 nitrogen, an average of daily N intake (FAO/WHO, 1973) multiplied by the 302 population of the Thames catchment was used. The human sewage is then returned 303 via waste treatment. The N in waste treatment either discharges to the fluvial network; 304 is denitrified and lost as atmospheric emissions; or is returned to land as sewage 305 sludge (the fate of the human consumption as sewage is discussed under outputs 306 below). By considering the input and outputs of each of crop, livestock and humans 307 within the catchment, it was not necessary to consider what proportion of the 308 agricultural output was used within the catchment. In the absence of other means of 309 uncertainty estimation the error in the food and feed transfers was considered as 310 ±20%.

311 Figures for the use of synthetic inorganic fertilizer in the UK were derived for 312 the period 1962 to 2007 from surveys published by the Fertilizer Manufacturers 313 Association and the Environment Agency of England and Wales (British Survey of 314 Fertilizer Practice, 2008). The use of fertilizer in the UK peaked in 1987 and showed 315 a steady, approximately linear, rise to this year from the beginning of the record in 316 1962. To convert the annual total fertilizer use in the UK to inputs of fertilizer per 317 hectare for each land-use type in the study catchment, the recommended values from 318 the UK Fertilizer Best Practice manual (British Survey of Fertilizer Practice, 2008) 319 were used to scale the total annual fertilizer use for any individual year to the average 320 that would be applied for each land-use type for each year. The values of annual 321 fertilizer input are reported with an estimated standard error of ±9%. Before, 1962, N-322 fertilizer inputs were estimated using data from Mittikalli and Richards (1996), who 323 collated average rates of nitrogen fertilizer use on arable and grassland in England and 324 Wales between 1943 and 1989 based on data published in Cooke (1975), ADAS 325 (1979), Church (1979) and MAFF (1983). Data were reported by Mittikalli and 326 Richards (1996) for "arable" and "grassland" in 1943, 1950, 1957 and 1962 and linear 327 interpolation was used to estimate the values for years in between these dates. Prior to 328 1943 values of fertilizer inputs were estimated by linear interpolation decreasing 329 backward through time until they equalled a value of 25 kg N/ha/yr for all fertilised 330 land. The value of 25 kg N/ha/yr is the input expected from organic manures based on 331 evidence from export coefficient models (Worrall and Burt, 2001).

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333 2.3. N outputs

No estimates of N from industrial sources or any fossil fuel burning were estimated in this study as it was only the terrestrial biosphere that was being considered. Therefore, this study only considered denitrification from non-industrial sources to both N_2O and N_2 ; and the emissions of NH_3 from livestock within the catchment. Since net transfers of food and feed across the catchment boundary were included in the budget, it is not necessary to assume that all livestock consume food from within the catchment.

341 Terrestrial denitrification to N₂ was estimated using the review of Barton et al. 342 (1999). They examined 95 studies of N₂ flux from natural systems and were able to 343 establish significant differences between land uses in the annual N2 export and so 344 distinct land use types could be given an estimate and range of N2 flux to the 345 atmosphere. The distinct land uses considered were forestry, rough grazing land, 346 fertilized grassland and cropland. In a similar fashion, it was possible to give 347 estimates and ranges of N₂O flux from distinct land-use types based upon the UK 348 model of Sozanska et al. (2002) and for NO emissions based upon ranges from 349 Davidson and Kingerlee (1997). Misselbrook et al. (2010) have estimated the range of 350 NH_3 flux from a range of UK livestock types. Given the land use history for the study 351 catchment reconstructed above, it was possible to give estimates of N₂, N₂O, NO and 352 NH_3 flux to the atmosphere back to 1867.

353 The fluvial flux of N from the terrestrial biosphere and the aquatic 354 denitrification cannot be directly estimated from available data. The flux of nitrate at 355 the outlet of the study catchment has been calculated by Howden et al. (2010) back to 356 1883. The Thames water quality record at Kingston does not include DON but records 357 of DOC, or its equivalent, for the river for the catchment have been recorded back to 358 1883. This study assumed that flux of DON can be derived from the flux of DOC, 359 given a knowledge of the C/N ratio typical of fluvial DOC (Neal, 2003). Neal (2003) studied sediment from rivers with catchment areas from 373 to 8231 km²; organic 360 361 carbon contents varied from 5 to 17% with 11% as a preferred value. For ammonium 362 it was possible to calculate a flux back to 1906. In 1906 the flux of ammonium was 363 only 4% of the nitrate flux and so this percentage was assumed back to 1883. Given 364 the pH of the River Thames over the course of the record, it was assumed that the flux of nitrite would be 1% of the nitrate flux (Patrick and Mahapatra, 1968). 365

366 After March 1974, the Thames at Teddington was included in the Harmonised 367 Monitoring Scheme (HMS: Bellamy and Wilkinson, 2001) which includes the 368 analysis of suspended solids. Therefore, the flux of suspended solids could be 369 calculated for each year from 1974 to 2007. No significant trend was found for the 370 flux of suspended solids over the period 1974 - 2007. The annual flux estimates of 371 suspended solids were compared to the range of other known fluxes (nitrate, ammonia 372 and DOC) and to annual water yield: a significant relationship was found with annual 373 water yield and used in order to estimate flux of suspended solids back to 1883. Furthermore, the fit of any such relationship was also used to calculate the error on any estimate of the suspended solids flux. The flux of PON can be derived from the suspended sediment flux coupled with a knowledge of the organic carbon content and C/N ratio typical of suspended sediment (Hillier, 2001).

378 For the fluvial fluxes calculated directly from Thames water quality record 379 using an interpolation method (Littlewood et al., 1995), then the error would be due to 380 the sampling frequency: a minimum sampling frequency of monthly within the HMS 381 means a maximum error of 14% (Worrall and Burt, 2007). For the PON flux, the error 382 was both the error in the extrapolation method and, along with the estimation of DON 383 and PON, the variation in the composition. The calculation of the DON and PON 384 fluxes required use of literature estimates of the C/N ratio of the dissolved and 385 suspended matter and the organic carbon content of the suspended sediment. Hillier 386 (2001) studied suspended sediment throughout the River Don catchment in Scotland (area = 1320 km^2); the average C/N ratio was 8.1 with a range of 5.2 (n=13): this 387 388 range was used here.

389 The fluvial flux as calculated above is the loss at the tidal limit and not the 390 flux as the water enters the fluvial network. Therefore, the fluvial flux at the tidal 391 limit will be an underestimate of the losses from the terrestrial biosphere as it does not 392 account for in-stream losses and would include any in-stream gains in nitrogen. The 393 processing of nitrogen species in streams was not considered by Smil (1999) or Boyer et al. (2002, 2005) but was by Marsh (1980). Nitrogen can be lost from rivers through 394 395 immobilisation in the stream biomass or denitrification to the atmosphere. Rivers can 396 themselves be sources of nitrogen as PON and DON and, given the definition of the 397 terrestrial biosphere, i.e. confined to the soil profile and biomass upon it used in this 398 study, then both groundwater influxes and direct sewage inputs represent in-stream 399 sources. However, available methods for calculating in-stream losses of nitrogen 400 species have differing approaches to distinguishing between these sources and some 401 merely estimate loss of N within the stream network from whatever source. Here, the 402 losses of N within the Thames estimated by four different methods. Firstly, Kroeze et 403 al. (2003) reviewed N retention in surface waters, regardless of source, and their 404 figures for rivers, rather than those for lakes, indicate that retention was between 11 405 and 50% of the input. Secondly, Seitzinger et al. (2002) proposed an empirical 406 relationship relating %N removed to depth of water body and residence time: again 407 regardless of source. Thirdly, Worrall et al. (2012a) have assessed the changes in the 408 flux of DON, nitrate and ammonia for up to 169 catchments. By comparing 409 differences in soil and land use between catchments with known dissolved N flux, it 410 was then possible to assess the extent of net loss with increasing catchment size. Once 411 the net loss was estimated, then the loss at the soil source was readily calculated – in 412 this case a net loss in-stream across a catchment of 63%. The approach of was also 413 independent of source as it considered a range of catchments with and without 414 groundwater influence, and included urban land use which may be considered 415 indicative of sewage inputs. Fourthly, Worrall et al. (2012b) have applied an export 416 model to UK land use history since 1925 to estimate the flux of nitrate from UK soils 417 at source. The export model applied across the UK can be applied to the Thames 418 catchment and for the land use records back to 1867. The export model can only be 419 applied to nitrate and not other N species in the fluvial flux – this would give an in-420 stream removal rate of 52%. Similarly, Worrall et al. (2007a) have proposed methods 421 for the correction of DOC fluxes for in-stream losses; their method was used here to 422 correct DON fluxes for in-stream losses. Worrall et al. (2009) have shown that nitrate 423 and DON comprise over 90% of the total fluvial N_r flux. Worrall et al. (2014) have 424 studied losses of POM from across 80 catchments across the UK using methods 425 similar to those of Worrall et al. (2012a) and found an average in-stream loss of POM 426 33.5%. Given the C/N ratio outlined above it was then possible to calculate PON 427 contribution to in-stream losses and PON flux at source. The error in the total fluvial 428 flux at source for DON, nitrate, nitriate and ammonia was then taken as the variation between the local estimates of in-stream losses across a catchment (52 to 63% loss -429 430 an uncertainty of $\pm 6.5\%$) and the error in the flux at the tidal limit (14%) giving a 431 percentage error of 20.5%. Therefore, the method used to assess in-stream losses has 432 implications for whether in-stream sources and losses from these sources need to be 433 included, but given that these are independent of the in-stream sources, therefore this 434 study does take the approach of considering groundwater and sewage fluxes as 435 possible in-stream sources.

436 Groundwater can represent an important store for dissolved nitrogen and thus 437 also a possible source of nitrogen to surface waters. A 1 mg N/l rise in average 438 groundwater nitrate concentration since 1990 has been observed in the UK by Stuart 439 et al. (2007). It was possible that large amounts of N are being stored in the aquifers 440 underlying the catchment. In order to assess the amount of storage over the course of 441 the study period, this study assumed that the input of nitrogen to the aquifer was 442 dominantly in the form of nitrate but other forms of dissolved N could also be 443 transported into the aquifer. For this study it was assumed that the time course of 444 dissolved N species would be as predicted from the reconstruction of the loss at the 445 soil source using the method outlined above. The aquifer was assumed to consist of 446 two parts, saturated and unsaturated. In each case, the storage of nitrate (and other 447 dissolved N species) was considered to be due to diffusion into the matrix of the 448 aquifer from readily mobile transport pathways. The dissolved N diffuses from the 449 fracture or cracks into the matrix and while there, no adsorption of dissolved N to 450 aquifer materials was assumed, but denitrification was allowed. To estimate this sink, 451 the problem was modelled as 1D-transport into water-filled porous media. The diffusion coefficient was taken as between 3.1 and 8.5 x 10^{-8} m²/s (based on nitrate -452 453 Gooddy et al. 2007). The denitrification rate was taken as 0.5 to 3% per year based 454 upon studies of aquifer denitrification by Hiscock et al. (2003). Fracture spacing was 455 assumed to be between 10 and 12 cm (Bloomfield, 1996); this is large in comparison 456 to the diffusion distance over the times of the model. The equation was solved by 457 Crank-Nicholson method with a time step of 1 day for the period since 1883 and 458 spatial step of 0.25 cm with an initial concentration of nitrate in the aquifer assumed 459 to be 1 mg N/1 (Limbrick, 2003). Once the concentration profile for the aquifer 460 material had been calculated, it was possible to calculate the mass of material stored if 461 the following were known: the porosity of the matrix, the percentage of the total 462 porosity that is matrix, the thickness of the saturated zone and the area of the aquifers 463 within the basin. The porosity of matrix was taken from measurements of Chalk as 464 between 3 and 55% (Bloomfield et al., 1995) which also encompasses the range 465 observed for the Jurassic limestones (Neumann et al., 2003). There are no published 466 measurements of the proportion of fracture verses matrix porosity in the aquifers of 467 Thames basin and so this study used values between 95 and 99% for Chalk elsewhere 468 in the UK (Burgess et al. 2005). The thickness of the active aquifers within the basin 469 was taken as up to 30 m. If both unconfined and confined aquifers within the Thames 470 basin were considered, then between 50 and 100% of the catchment is underlain by 471 aquifers that could act as sinks for dissolved N. Given the ranges outlined above, the 472 calculation was performed 100 times drawing randomly from the ranges defined and 473 assuming uniform distribution between the ranges and thus the uncertainty in the estimation was taken from the range of these 100 values. Equally, there would be storage in the unsaturated zone as well as in the saturated zone of aquifers. It was assumed that an unsaturated zone covers between 50% and 100% of the basin area with depth between 0 and 60 m and a moisture content between 5 and 95%. The dissolved N stored in the unsaturated zone can then be calculated as for the saturated zone.

480 The direct flux of sewage and industrial wastes to the streams of the catchment 481 was estimated using an export coefficient approach, i.e. a nitrogen load from sewage 482 per head of population in the catchment was assumed based upon the review of export 483 coefficients by Worrall et al. (2012b). The value of the per capita sewage export was 484 1.2 to 5.6 kg N/yr/ca (Worrall and Burt, 1999; Weber et al., 2006; Johnes, 1996; 485 Johnes et al., 1996), with a preferred value of 4.5 kg N/yr/ca based on data from the 486 smallest catchment. The population history of the catchment could be calculated from 487 census returns; then the direct sewage inputs could be calculated back to 1861 with 488 the error set by the range in the export coefficient. Gaseous emissions from sewage 489 treatment were calculated based upon emissions factors published for the UK and 490 Western Europe (IPCC, 2000). The difference between the amount of N input to the 491 catchment via human consumption and the amounts lost from sewage treatment as 492 either discharge to the river or predicted as emitted to the atmosphere was returned to 493 land within the catchment as sewage sludge solids. It was assumed that there is no net 494 transfer of sewage across the watershed and that all N discharged from sewage 495 treatment within the catchment was discharged into the Thames and its tributaries, 496 returned as sludge to land within the catchment, or lost to the atmosphere. The 497 uncertainty in this pathway was estimated from range in the per capita sewage export 498 coefficients and the range in the published emissions factors.

499

500 **3. Results**

501 <u>There is not space within the manuscript to give the detail and time series of each</u>
502 input and output pathway and where the time series exist they are supplied in

503 supplementary material.

504

505 3.1. N inputs

506 The inorganic N deposition interpolated and extrapolated from values reported by 507 Fowler et al. (2005) suggests a value of 15.6 ktonnes N/yr in 2006; no value was 508 available for 2007, as discussed above. No reasonable error estimate can be provided 509 from the original data and, given the assumption of a steady-state up to 1936, this 510 would be an inorganic N deposition of 6.3 ktonnes N/yr - because this is extrapolated 511 data no detailed time series is provided in the supplementary material. This value does 512 not consider deposition of DON. The deposition of DON at the Moor House site was reported by Worrall et al. (2006) as being between 0.01 and 0.15 tonnes N/km²/yr 513 514 with no significant trend between 1993 and 2005. If the measured import at Moor 515 House were re-scaled to the study catchment, then the input to the catchment would 516 be 0.8 ± 0.7 ktonnes N/yr: therefore total N deposition in 2006 of 17.1 ktonnes N/yr 517 (Table 1).

The biological nitrogen fixation (BNF) varied from 13 ± 3 ktonnes N/yr in 1883 to 10.4 ± 2.5 ktonnes N/yr in 2007 (Table 1 – supplementary material – Figure 20 S2) with a peak year of 1960 when it peaked at 14.6 ± 3 ktonnes N/yr and has been 21 declining ever since. **Formatted:** Font: Not Bold, (Asian) Chinese (PRC), (Other) English (U.S.)

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⁵²² The trans-boundary transfer of food and feed to agriculture was an N input to 523 the study catchment projected to have been 19.4 ± 3.9 ktonnes N/yr in 1867 but only

 9.9 ± 2 ktonnes N/yr in 2007 (Table 1 – supplementary material – Figure S3). It was 524 525 projected to have peaked in 1878 and there has been no significant trend on this 526 transfer since 1994. Human consumption of N has increased in line with population 527 growth in the catchment where in 1861 the population of the catchment was 911000 528 represented an intake of 22.1 ktonnes N/yr but by 2007 this had increased to an intake 529 of 91.4 ktonnes N/yr based on a population of 3.77 million people. Working within 530 the Seine catchment, Billen et al. (2012) found that the food and feed transfers 531 represented a net export of N from the catchment.

The input of synthetic inorganic fertilizer was by far the largest nitrogen input into the basin, varying from an estimated 10.1 ktonnes N/yr in 1867 to a peak input in 1987 at 67.3 ktonnes N/yr, with values declining since then to 44.8 ktonnes N/yr by 2007 at a rate of 1 ktonnes N/yr² (Table 1 – Figure S4). At the national scale, the decline in inorganic fertilizer input has been occurring since 1984, but the particular land uses of the Thames basin means that fertilizer inputs peaked slightly later than the UK as a whole.

539

540 3.2. N outputs

541 The total N₂O emissions from the terrestrial biosphere track the projected 542 inorganic fertilizer inputs projected to peak in 1987 at 2.9 ktonnes N/yr from a value 543 of 0.5 ktonnes N/yr in 1867 and declining to 2.3 ktonnes N/yr by 2007 (Table 1 – 544 Figure S5). The emissions of NH₃ were predicted to be largest at the beginning of 545 study period at 7.5 ktonnes N/yr, declining to a minimum in 1932 at 4.1 ktonnes N/yr, 546 rising to 5.0 ktonnes N/yr in 2007 (Figure S6). Terrestrial denitrification to N2 was 547 estimated as a minimum in 1904 of 3.3 ktonnes N/yr rising to a maximum of 5 548 ktonnes N/yr in the year 2000 (Figure S7). For NO emissions was estimated as a Formatted: (Asian) Chinese (PRC), (Other) English (U.S.), Not Superscript/ Subscript 549 minimum of 0.8 ktonnes N/yr in 1913, and a maximum of 1.2 ktonnes N/yr in 1995 550 (Figure S8). In 2006 the total terrestrial atmospheric emissions were 13.1 ktonnes 551 N/yr with an median $\frac{N_2 O - N}{(N_2 O - N + N_2 - N - N)} = 0.31$ within the ranges reported by 552 Schlesinger (2009) for agricultural soils.

553 The total fluvial N flux at the tidal limit for 1904 (the first year for which 554 complete records were available was 5.2 ktonnes N/yr (with the forms of fluvial N as 555 61:20:15:0.6 for nitrate-N, DON, PON, ammonia-N and nitrite-N respectively). There 556 are two peak values: in 1977 the total fluvial N flux peaked at 27.3 ktonnes N/yr and then in 2001 at 36.6 ktonnes N/yr (equivalent to 3.7.tonnes N/km²/yr, or 37 kg 557 558 N/ha/yr) - Figure S9. In-stream losses of nitrogen follow those of the fluvial losses 559 and are critically dependent upon the model used. However, using the model developed for the UK gives values of in-stream losses of total N from a minimum of 560 5.7 ktonnes N/yr in 1904 to the maximum in 2001 of 86.3 ktonnes N/yr. This means 561 562 that the flux of total N from the soil source to the stream network in 2006 was 16.7 563 ktonnes N/yr.

564 The sink to the groundwater store is shown in Figure 3a and b shows that for 565 most of the period of the record there was a net flux into groundwater storage in the 566 saturated zone. Between 1883 and 1903, there was a net annual sink but this may be 567 the result of not being able to know the concentration in the aquifer matrix in 1883 for 568 which a uniform value of 1 mg N/l was assumed (Limbrick, 2003). After 1903, there 569 was an approximately constant flux, i.e. near steady-state conditions occur once a 570 period of establishment within the model has been achieved. The period of steady-571 state ends in 1939, i.e. at the beginning of the ploughing up of grassland with the 572 onset of WWII. The flux into groundwater storage reached a maximum in 1972 at 66 573 ±33 ktonnes N/yr (Figure 3b); this sink has declined ever since with groundwater 574 predicted to become a net source in 1993. The predicted sink in the unsaturated zone 575 of the aquifers will parallel the time course of flux into the saturated zone where the 576 maximum annual sink was between 1.4 and 5.5 ktonnes N/yr. The accumulated 577 storage in the aquifers of the basin (both in the saturated and unsaturated zones) 578 shows distinct changes (Figure 3a). The step changes can be associated with the step 579 changes observed in the nitrate concentration of the River Thames as observed by 580 Howden et al. (2010). The accumulated dissolved N in the aquifer peaks at a 581 maximum of 1571 ± 608 ktonnes N between 2000 and 2004 and has since begin to 582 decline. The approach here suggests that in 2006 the increase in concentration of N 583 species in groundwater of the catchment represented a net additional store of 15.9 584 ktonnes N/yr.

585 The total flux of N from human sewage and industrial wastes has increased 586 over the period in line with the population of the catchment from a low in 1867 of 3.1 587 ktonnes N/yr to 12.4 ktonnes N/yr in 2007 (Figure S10). Worrall et al. (2009) showed 588 that due the implementation of the Urban Wastewater Directive (European 589 Commission, 1991) the UK-wide direct flux of N declined by 50% between 1990 and 590 2003. The amount emitted to the atmosphere closely followed the population and by 591 2007 the amount released to the atmosphere from sewage treatment at 36 ktonnes 592 N/yr with 42.5 ktonnes N/yr returned to land as sludge.

593

594 **3.3.** Total N budget

Given the uncertainties within each pathway, then the maximum annual source observed was in -59 ± 10 ktonnes N/yr (where the uncertainty is expressed as the inter-quartile range) and the maximum sink was in 2006 at $+110 \pm 26$ ktonnes N/yr (Figure 4a and b) – the maximum sink was equivalent to 111 ± 27 kg N/ha. When 599 preferred values are considered, then the maximum sink was 112 ktonnes N/yr. The 600 total N balance can also be viewed as the cumulative sink or source over the course of 601 the study period (Figure 4b), in which case it can be seen that within the range 602 estimated for each flux pathway, the basin was a net source until somewhere between 603 1959 and 1973 when the cumulative curve reaches a minimum and between 1994 and 604 2004 the basin was a net accumulating sink and that accumulation increases to the 605 present day and by 2007 was 315 ± 379 ktonnes N – which is equivalent to 32 tonnes N/km² or 320 kg N/ha, and accumulating at an average rate of 55 ktonnes N/yr since 606 1973 (the minimum in the accumulation time series - Figure 4b) - equivalent to 5.5 607 tonnes N/km²/yr, or 55 kg N/ha/yr). 608

609

610 4. Discussion

611 The results predict a very large and ongoing storage of total N within the terrestrial 612 biosphere of the Thames catchment. This raises a number of important questions. 613 Firstly, is there evidence from other studies that such a storage could be happening? 614 Several studies have concluded that developed countries are net sinks of reactive 615 nitrogen (Nr - e.g. Sutton et al., 2011) and, indeed, this has been already demonstrated 616 for this catchment (Howden et al., 2011), but those conclusions of these studies were 617 not for the terrestrial biosphere nor for total N. When it comes to specific 618 environments, then it is possible to assess total N budgets although there are a limited 619 number of such studies to refer to. Hemond (1988) was the first to consider a total N 620 budget for a specific environment - a peat bog; however, since the peat bog was a 621 functioning sink of carbon, it is not too surprising that it was also a net sink of total N 622 in line with the C/N ratio of the humified organic matter. Hemond (1988) recorded a net sink of total N of 0.58 tonnes N/km²/yr. At a catchment scale, other studies have 623

624 implied that the change in nitrate flux from the river over time is indicative of 625 nitrogen storage within the catchment. Goolsby et al. (1999) estimated a net annual sink in the Mississippi Basin of 19 tonnes N/km² over a period of 40 years. 626 627 Furthermore, Basu et al. (2010) showed a widespread occurrence of biogeochemical 628 stationarity in large anthropogenically-disturbed catchments from a range sites across 629 the northern hemisphere, but not for small undisturbed catchments (e.g. Hubbard 630 Brook, New Hampshire, USA). This biogeochemical stationarity was ascribed to 631 widespread saturation within anthropogenically-disturbed catchments meaning that, 632 no matter what flow paths were operating, the result was the same. Therefore, 633 although the sink predicted here is larger than those suggested by Goolsby et al. 634 (1999) and Basu et al. (2010), this study did consider total N and not just reactive N.

635 Second, are the results sensitive to the considerable uncertainty in a number of 636 the pathways being considered? The main features of the results is that they are not 637 sensitive to the uncertainty or assumptions of the approach. The two main features of 638 the study results are that there is now a large accumulated sink of total N within the 639 terrestrial biosphere of the catchment and that there was an inflexion point in the 640 behavior of the accumulated total N budget in the 1960s or early 1970s (Fig. 4). Any 641 uncertainty would have to sufficiently large to change the scale of the accumulation or 642 the time series of that accumulation. The largest source of uncertainty was in the 643 terrestrial denitrification estimate but the study has already included that uncertainty 644 and, even if the largest possible value of terrestrial denitrification was used, then all 645 that would happen would be that the total accumulation would be lower by a value of 646 approximately 800 ktonnes N by 2007, but this would only mean a slightly smaller 647 rate of accumulation and the inflexion point would be offset by several years. The 648 period of time for which information available to the study was the least uncertainty 649 was the most recent period, i.e. the period when the largest net sinks and net 650 accumulation were predicted. The most uncertain period was the period at the 651 beginning of the record and so to offset the accumulation observed a larger source 652 would have to be predicted in the second half of the record.

653 The time series that was most uncertain, as opposed to the uptake or release 654 pathways, was the record of atmospheric deposition where it was necessary to assume 655 a rate of increase from a period of steady-state and a year in which the steady state 656 was broken (presently taken as 1936). Both of these assumptions have been based 657 upon observations from the basin and so it would necessary to substantiate a different 658 rate of increase; a different time at which the major increase started and value of the 659 deposition during that period of steady-state. However, it should remembered that 660 atmospheric deposition is an input and so to change the result it would have to be 661 substantially lower, start increasing later than presently assumed and then increase 662 faster the currently observed values.

Finally, it should be considered whether there was a flux pathway missing from the budget? It is always impossible to have a complete budget; however, in order to account for the estimated accumulation in the terrestrial biosphere suggested here, it would have had to have been a sink of the order of 100 ktonnes N/yr – it is just as likely that the study has failed to consider a source of total N. The major part of the terrestrial biosphere which we could not consider were the subsoils of the catchment where accumulation could be occurring.

670 If accumulation is occurring, then where is it accumulating? Conversely, if net 671 loss were occurring then where was the loss coming from? In a catchment which was 672 under intensifying agriculture, urbanization and climate change, it is easy to consider 673 that the disturbance of soils stores means that there is a tendency to lose nitrogen just 674 as there is to lose carbon (e.g. Bell et al., 2011, Barraclough et al., in press). Given the 675 values of carbon loss predicted by Barraclough et al. (in press), this suggests that soils 676 would be losing 1 tonnes N/km²/yr (10 ktonnes N/yr for the Thames catchment). The 677 average N loss from 1883 to 1959 predicted in this study was 20 ktonnes N/yr, i.e. 678 50% of the loss estimated by this study could be predicted by climate change alone, 679 independent of intensification of agriculture or urbanization. Therefore, it is 680 reasonable to assume that when accumulation does occur it is in the soils of the 681 catchment and the so far this study has not considered the subsoils. It is easy to 682 conceive that nitrogen released from topsoils could, in part, be absorbed in subsoils (if 683 it were released in the form of DON) or stimulate biomass if it were released in 684 inorganic forms.

685

686 5. Conclusions

The study has considered the total N budget of the terrestrial biosphere of a large mixed agricultural catchment dominated by mineral soils. The study shows that since the late 1950s the terrestrial biosphere and since 1973 has been accumulating total N at an average rate of 55 ktonnes N/yr (equivalent to 55 kg N/ha/yr), peaking in 2007 at 112 kg N/ha/yr. The accumulation of total N in the catchment was estimated to be 315 ktonnes N by 2007 (315 kg N/ha) even allowing for accumulation in groundwater. We propose that this accumulation is in sub-soils of the catchment.

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905

	Flux in 2006	Export in 2006
	(ktonnes N/yr)	(kg N/ha/yr)
Input		
Atmospheric deposition	17.1	17.2
Biological nitrogen fixation	10.7	10.8
Food & feed transfers	99.2	99.7
Inorganic fertilisers	46.1	46.3
Sub-total	173.1	
Output		
Terrestrial atmospheric emissions	13.1	13.2
Fluvial loss at soil source	16.7	16.8
Groundwater storage	15.9	16.0
Sewage inputs	12.3	12.4
Sub-total	58.0	58.4
Total N budget	115.1	115.8

907 Table 1. Summary of preferred values of N inputs and outputs for 2006.

909 Figure. 1. A schematic diagram of the flows and fluxes considered by this study.

910

- 911 Figure 2. Location of the study catchment; study monitoring point at Teddington and
- 912 location of the long term climate monitoring station at Oxford.
- 913
- 914 Figure 3.

a) The estimated accumulated groundwater store of total dissolved N over the courseof the study period.

- b) The estimated annual flux to groundwater of total dissolved N over the course of
- 918 the study. The values are given as the interquartile range with a negative value being a
- 919 a net discharge from ground to surface water.

- 921 Figure 4.
- 922 a) The estimated annual total N budget of the terrestrial biosphere of the catchment.
- 923 The bar is given as the inter-quartile range based upon the stochastic combination
- 924 within the uncertainties described.
- b) The cumulative total N budget of the terrestrial biosphere of the catchment giving
- 926 the median and inter-quartile range based upon range of annual values shown in
- 927 Figure 4a.