

1 Seismically enhanced solute fluxes in the Yangtze River
2 headwaters following the 2008 Wenchuan earthquake

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13 **ABSTRACT**

14 Large earthquakes alter physical and chemical processes at the Earth's surface,
15 triggering landslides, fracturing rock, changing large-scale permeability, and influencing
16 hydrologic pathways. The resulting effects on global chemical cycles are not fully
17 known. Here we show changes in the dissolved chemistry of the Min Jiang, in the
18 Yangtze River headwaters, following the 2008 M_w 7.9 Wenchuan earthquake. Total
19 solute fluxes transported by the Min Jiang increased after the earthquake, accompanied
20 by a ~4x increase in Na*/Ca ratios (where Na* is Na⁺ corrected for atmospheric and
21 evaporite contributions) and a 0.000644 ± 0.000146 increase in ⁸⁷Sr/⁸⁶Sr isotopic ratios.
22 These changes are consistent with enhanced contribution from silicate sources. We infer

23 that the CO₂ consumption rate via silicate-derived alkalinity increased 4.3 ± 0.4 times. If
24 similar changes are associated with other large earthquakes, enhanced solute export could
25 directly link tectonic activity with weathering and alkalinity fluxes that supply nutrients
26 to ecosystems, influence seawater chemistry evolution, and steer Earth's long-term
27 carbon cycle and climate.

28 **INTRODUCTION**

29 Rivers collect and transport material from across Earth's surface, so their
30 chemistry integrates geological and environmental processes over space and time (e.g.,
31 Berner and Berner, 1996; Gaillardet et al., 1999). Variability in river chemistry over
32 timescales from days to seasons to decades is increasingly well understood (e.g.,
33 Raymond and Cole, 2003; Tipper et al., 2006; Godsey et al., 2009; Torres et al., 2015), as
34 are the global implications (Maher and Chamberlain, 2014). The effects of infrequent
35 events such as large earthquakes on river chemistry are less well known, since these
36 events are not typically represented in observational records even though their cumulative
37 effects may be significant over the centennial or longer timescales of their recurrence.

38 Large earthquakes on continental faults perturb hydrologic pathways (Rojstaczer
39 et al., 1995; Montgomery and Manga, 2003; Claesson et al., 2004; Skelton et al., 2014),
40 and changes in solute chemistry of groundwaters and streams have been proposed as
41 possible earthquake warning indicators (e.g. Rojstaczer et al., 1995; Skelton et al., 2014).
42 Earthquake-driven changes in river solute fluxes could be important for quantifying
43 (bio)geochemical fluxes and for understanding terrestrial geochemical and hydrological
44 processes.

45 The chemical response of the Min Jiang in China to the 2008 M_w 7.9 Wenchuan
46 earthquake provides a rare opportunity, allowing us to directly observe and quantify how
47 a seismic event affects river chemical signatures. The Wenchuan opportunity arises in
48 part because of data on Min Jiang river chemistry prior to 2008 (Qin et al., 2006; Yoon et
49 al., 2008; Huh, 2010), providing the basis for direct comparison with samples collected
50 after the earthquake.

51 The 2008 Wenchuan earthquake occurred along the Longmen Shan mountain
52 range that forms the eastern margin of the Tibetan Plateau (Robert et al., 2010). The
53 Longmen Shan is characterized by steep slopes and high relief, increasing in mean
54 elevation by ~3500 m over less than 100 km distance (Fig. 1a). The geology of the region
55 is dominated by bedrock with alumino-silicate minerals, including metamorphic
56 argillaceous sandstone and flysch, granite and monzonitic granite, and detrital sediments,
57 as well as limestones (Fig. DR1). The Wenchuan earthquake was generated along the
58 Yingxiu–Beichuan and Pengguan faults that run S40–50°NE along the Longmen Shan
59 (Xu et al., 2009). The earthquake and associated aftershocks caused more than 56,000
60 landslides over ~200 km length of the mountain range, whereas the extent of landslides in
61 the region was limited prior to the earthquake (Li et al., 2014). The climate of the region
62 is dominated by the Asian and Indian summer monsoons, with 75% of annual
63 precipitation (600–1100 mm/yr) during May to October. The annual water discharge of
64 the Min Jiang, the major river draining the Longmen Shan, was $1.06 \times 10^{10} \text{ m}^3$ in 2000–
65 2011, and the difference of annual water discharge before and after the earthquake was <
66 20% (Fig. 2a), smaller than inter-annual variability (Wang et al., 2015).

67

68 **MATERIALS AND METHODS**

69 River water samples for this study were collected weekly between December
70 2009 and the end of 2011 from two hydrological stations, at Weizhou and
71 Zhenjianguan, both on the main course of the Min Jiang (Fig. 1). Hydrologic parameters
72 at both stations are monitored regularly by the Chinese Hydrology Bureau (CHB). The
73 Weizhou station in the town of Wenchuan lies in the zone of high peak ground
74 acceleration (PGA) during the 2008 earthquake, and consequently in a region
75 significantly affected by co-seismic landslides and other visible damage (Fig. 1). The
76 Zhenjianguan station is farther upstream, still within the zone of measurable earthquake-
77 associated ground acceleration, but where PGA was much lower than at Weizhou and
78 where there were few earthquake-triggered landslides (Fig. 1). In addition, single samples
79 were collected in 2011 from nine revisited sites distributed throughout the river basin
80 (Fig. 1), matching sites that had been sampled at least once before the earthquake (Qin et
81 al., 2006; Yoon et al., 2008; Huh, 2010). Filtered water samples were analyzed for
82 dissolved major ions and Sr^{2+} concentrations, and $^{87}\text{Sr}/^{86}\text{Sr}$ isotopic ratios (Table DR1,
83 Figs. DR2-4; see Methods in the GSA Data Repository). The chemistry of the samples
84 collected in this study was compared to dissolved chemistry data from prior to the
85 earthquake at equivalent sites. The data set from the Weizhou station enables comparison
86 of pre- and post-earthquake annual time-series (Table DR2). For the other samples,
87 collection season was matched as closely as possible to previous studies (Table DR3).

88 **RESULTS**

89 **Pre- and Post-Earthquake Times-Series from Weizhou Station**

90 When comparing data at Weizhou from 2010–2011 versus data from 2001–2002,
91 prior to the earthquake, a systematic increase is observed in Na*/Ca molar ratios (Fig. 2b,
92 where Na* = Na⁺ derived from silicate weathering; see Methods in the GSA Data
93 Repository) and K/Ca ratios (Fig. DR5). These differences are significantly greater than
94 the variability observed in the long-term record of dissolved chemistry for the Min Jiang
95 acquired by the CHB between the 1970s and 2000 (Qin et al., 2006), suggesting that the
96 higher ratios observed in the 2010–2011 data reflect a distinct post-earthquake change in
97 solute composition. Unlike concentrations, elemental and isotopic ratios are not directly
98 influenced by dilution, so the observed changes cannot be explained by changes in
99 precipitation or discharge amounts. The changes in elemental ratios are also accompanied
100 by an increase in the average dissolved ⁸⁷Sr/⁸⁶Sr isotope ratios from 0.712663 ± 0.000200
101 to 0.713307 ± 0.000207 . Coincident increases in dissolved Na*, K⁺, and ⁸⁷Sr/⁸⁶Sr are
102 consistent with silicate mineral sources. Ratios of Mg/Ca and Sr/Ca do not show similar
103 changes (Fig. DR3), most likely because river waters are at carbonate saturation both
104 before and after the earthquake (Yoon et al., 2008).

105 **Paired Data from Other Sites**

106 Pre-earthquake time series data are not available at Zhenjianguan (cf. Fig. DR6),
107 but we can compare individual paired data from before and after the earthquake at this
108 site and several others. Increases in both Na*/Ca and Sr isotope ratios (Figs. 1b, c and 3a,
109 b) are observed at all of the sites within the zone of significant PGA, and the direction of
110 change is consistent with that observed at Weizhou (Fig. 2). The magnitude of post-
111 earthquake Na*/Ca increase (0.02–0.22, mean 0.07 ± 0.07) at these paired sites is
112 significantly larger than the annual variability at Weizhou (0.02, 1σ pre-earthquake; 0.04,

113 1 σ post-earthquake) and Zhenjianguan (0.03, 1 σ post-earthquake), suggesting that the
114 observed changes are not artifacts of the times of sample collection. Sites 1 and 2,
115 farthest from the Wenchuan earthquake epicenter and activated faults, show relatively
116 little change in Na*/Ca after the earthquake (less than 0.02), reflecting their greater
117 distance from the region of strongest PGA (Fig. 3).

118 The relationship between the magnitude of observed change (Δ , difference
119 between post- and pre-earthquake ratios) in solute chemistry and mean PGA during the
120 Wenchuan earthquake in the catchment area upstream of each sampling site is not
121 straightforward (Fig. 3c, d). Nonetheless, for the main stem, sites showing highest Δ
122 values for both Na*/Ca and $^{87}\text{Sr}/^{86}\text{Sr}$ are ones where mean PGA was relatively high (e.g.,
123 sites 7 and 8, and at Weizhou), while sites with lower PGA are associated with least
124 chemical change (e.g., site 2). For the tributary sites, the relationship is less clear, with
125 sites 3, 4 and 5 all showing large changes in Na*/Ca despite a range in catchment-
126 averaged PGA (Figs. 1b, 3c). The larger magnitude of $\Delta\text{Na}^*/\text{Ca}$ at tributary sites
127 compared to the main stem and the lack of relationship to PGA might be explained by
128 smaller tributary catchment areas, such that smaller perturbations have a greater effect
129 than on the main stem. Overall, the first-order relationship with PGA for the main stem
130 (i.e., significant change within the earthquake-affected zone, but less change farther from
131 the center of this zone at sites 1 and 2) is consistent with a link between the earthquake
132 and observed changes in river chemistry.

133 **DISCUSSION**

134 **Causes of Changing River Chemistry Following the Wenchuan Earthquake**

135 The dissolved Na^*/Ca , K/Ca , and $^{87}\text{Sr}/^{86}\text{Sr}$ of the Min Jiang systematically shifted
136 towards more silicate compositions across multiple sites and over multiple years
137 following the Wenchuan earthquake. Observed changes persisted at least for 3 years
138 (through December of 2011, when the last samples in this study were collected), so any
139 perturbation must have responded rapidly and been sustained for several years. Release
140 of deep basinal brines or greater contribution from human activities are unlikely causes
141 since we do not observe coincident changes in Cl^- or SO_4^{2-} concentrations, as would be
142 expected for such sources (e.g., Gaillardet et al., 1999). Herein we propose that the
143 observed changes are related to the effects of earthquake shaking on fluid pathways and
144 the minerals exposed to fluids by bedrock fracturing and/or by seismic landslides.

145 In more detail, at shallow depths (centimeters to several meters depth),
146 earthquakes trigger extensive co- and post-seismic landslides (e.g., Li et al., 2014), which
147 can act as solute generators by producing reactive fine-grained sediment (Wang et al.,
148 2015) and by focusing flow through this material (e.g., Watanabe et al., 2005; Emberson
149 et al., in press). Leaching of exchangeable cations from finely ground landslide debris
150 could also contribute to the solute load. Co-seismic landslides are prevalent throughout
151 the lower reaches of the Min Jiang study region (Li et al., 2014; Fig. 1), but the spatial
152 extent of observed changes in solute chemistry extends beyond the zone of most
153 concentrated mapped landslide activity, suggesting that other processes play a major role.

154 At greater depths of tens to hundreds of meters, earthquakes can fracture rock
155 (Molnar et al., 2007), alter permeability (Rojstaczer et al., 1995), and perturb hydrologic
156 systems over regional to continental scales, extending beyond regions of visible co-
157 seismic damage (Montgomery and Manga, 2003; Skelton et al., 2014; Shi et al., 2015).

158 Direct measurements of permeability along the Pengguan and Beichuan fault zones
159 showed initial increase followed by healing over exponential decay times of 0.6–2.5
160 years (Xue et al., 2013), but larger spatial scale changes in the hydrologic system
161 following shaking and seismic disturbance may persist for longer (Rojstaczer et al., 1995;
162 Claesson et al., 2004; Skelton et al., 2014). Groundwaters are typically concentrated in
163 silicate-derived cations as a result of prolonged water-rock contact (e.g., Tipper et al.,
164 2006). Post-seismic discharge of such groundwater could shift river chemistry toward
165 higher solute fluxes and more silicate composition. Evacuation of waters close to
166 equilibrium and replacement with more dilute waters could also stimulate higher mineral
167 dissolution rates (Maher and Chamberlain, 2014; Rempe and Dietrich, 2014), as could
168 the exposure of new mineral surfaces (White and Brantley, 2003), for example in bedrock
169 micro-fractures.

170 With the current data, we are not able to distinguish definitively the contribution
171 from seismically altered flowpaths versus enhanced dissolution in landslide debris, nor
172 can we distinguish whether additional solutes have their immediate source from primary
173 minerals, exchangeable sites, or concentrated groundwaters. Greater mechanistic insight
174 might be gained by future studies tracking spatial patterns of hydrochemical change and
175 longer-term evolution, including the return to pre-earthquake river composition.

176 **Implications for Geochemical Fluxes**

177 Independent of mechanism, the results from the Min Jiang offer empirical
178 evidence that a high-magnitude, low-frequency earthquake can have a significant,
179 previously-unrecognized effect on river chemistry. The increased flux of alkalinity
180 following the Wenchuan earthquake enhanced CO₂ drawdown and contributed more

181 radiogenic ^{87}Sr to the oceans. The post-earthquake dissolved $^{87}\text{Sr}/^{86}\text{Sr}$ ratios of Min Jiang
182 main stem river waters increased by a mean of 0.000644 ± 0.000146 relative to those
183 prior to the earthquake (excluding site 2). Sr concentrations did not significantly change
184 at the same time (Figs. DR3 and 4), implying a net increase in the delivery of ^{87}Sr to the
185 oceans, with potential implications for interpreting the geologic record of seawater
186 $^{87}\text{Sr}/^{86}\text{Sr}$ composition (Raymo et al., 1988; Edmond, 1992). The rates of net CO_2
187 consumption associated with silicate alkalinity (CO_2) after the earthquake (2010–2011)
188 are 4.3 ± 0.4 times higher than prior to the earthquake, based on average monthly solute
189 concentrations, drainage area, and water discharge (see Methods in the GSA Data
190 Repository). Assumptions about the composition of the silicate mineral end-member do
191 not affect the relative magnitude of CO_2 before and after the earthquake, as long as the
192 composition of silicate minerals did not change significantly.

193 The Wenchuan case suggests that changes in tectonic activity have the potential to
194 increase riverine alkalinity fluxes to the oceans by changing the frequency of
195 earthquakes. The extent to which such seismic changes in solute flux affect seawater
196 composition and the global carbon cycle (e.g., Raymo et al., 1988; Berner, 2004) remains
197 to be assessed. The magnitude of long-term change will depend on the duration that
198 earthquake-triggered changes persist and how the extent of chemical change varies for
199 different earthquakes. Whether other earthquakes cause similar effects may relate to a
200 number of factors including event magnitude, regional seismicity and earthquake return
201 times, the extent of induced landslides, the nature of fracture development and changes in
202 hydrological pathways, and regional lithology. These questions will only be answered by
203 further investigation to explore how the magnitude and duration of change vary for

204 different earthquakes, and to understand the underlying mechanisms causing observed
205 changes.

206

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214 **REFERENCES CITED**

215 Berner, E., and Berner, R.A., 1996, *The Global Environment: Water, Air, and*
216 *Geochemical Cycles*: Prentice-Hall, Upper Saddle River, New Jersey, 376 p.
217 Berner, R.A., 2004, *The Phanerozoic Carbon Cycle: CO₂ and O₂*: Oxford University
218 Press, 158 p.
219 Claesson, L., Skelton, A., Graham, C., Dietl, C., Mörth, M., Torssander, P., and Kockum,
220 I., 2004, Hydrogeochemical changes before and after a major earthquake: *Geology*,
221 v. 32, p. 641–644, doi:10.1130/G20542.1.
222 Edmond, J.M., 1992, Himalayan tectonics, weathering processes, and the strontium
223 isotope record in marine limestones: *Science*, v. 258, p. 1594–1597,
224 doi:10.1126/science.258.5088.1594.

- 225 Emberson, R., Hovius, N., Galy, A., and Marc, O., in press, Chemical weathering in
226 active mountain belts controlled by stochastic bedrock landsliding: *Nature*
227 *Geoscience*, doi:10.1038/ngeo2600.
- 228 Gaillardet, J., Dupré, B., Louvat, P., and Allègre, C., 1999, Global silicate weathering and
229 CO₂ consumption rates deduced from the chemistry of large rivers: *Chemical*
230 *Geology*, v. 159, p. 3–30, doi:10.1016/S0009-2541(99)00031-5.
- 231 Godsey, S.E., Kirchner, J.W., and Clow, D.W., 2009, Concentration–discharge
232 relationships reflect chemostatic characteristics of US catchments: *Hydrological*
233 *Processes*, v. 23, p. 1844–1864, doi:10.1002/hyp.7315.
- 234 Huh, Y., 2010, Estimation of atmospheric CO₂ uptake by silicate weathering in the
235 Himalayas and the Tibetan Plateau: A review of existing fluvial geochemical data, *in*
236 Clift, P.D., et al., eds., *Monsoon Evolution and Tectonic–Climate Linkage in Asia*:
237 Geological Society of London, Special Publications, v. 342, p. 129–151,
238 doi:10.1144/SP342.10.
- 239 Li, G., West, A.J., Densmore, A.L., Jin, Z., Parker, R.N., and Hilton, R.G., 2014, Seismic
240 mountain building: Landslides associated with the 2008 Wenchuan earthquake in the
241 context of a generalized model for earthquake volume balance: *Geochemistry*
242 *Geophysics Geosystems*, v. 15, p. 833–844, doi:10.1002/2013GC005067.
- 243 Maher, K., and Chamberlain, C.P., 2014, Hydrologic regulation of chemical weathering
244 and the geologic carbon cycle: *Science*, v. 343, p. 1502–1504,
245 doi:10.1126/science.1250770.

- 246 Molnar, P., Anderson, R.S., and Anderson, S.P., 2007, Tectonics, fracturing of rock, and
247 erosion: *Journal of Geophysical Research*, v. 112, F03014,
248 doi:10.1029/2005JF000433.
- 249 Montgomery, D.R., and Manga, M., 2003, Stream flow and water well responses to
250 earthquakes: *Science*, v. 300, p. 2047–2049, doi:10.1126/science.1082980.
- 251 Qin, J., Huh, Y., Edmond, J., Du, G., and Ran, J., 2006, Chemical and physical
252 weathering in the Min Jiang, a headwater tributary of the Yangtze River: *Chemical*
253 *Geology*, v. 227, p. 53–69, doi:10.1016/j.chemgeo.2005.09.011.
- 254 Raymo, M.E., Ruddiman, W.F., and Froelich, P.N., 1988, Influence of late Cenozoic
255 mountain building on ocean geochemical cycles: *Geology*, v. 16, p. 649–653,
256 doi:10.1130/0091-7613(1988)016<0649:IOLCMB>2.3.CO;2.
- 257 Raymond, P.A., and Cole, J.J., 2003, Increase in the export of alkalinity from North
258 America's largest river: *Science*, v. 301, p. 88–91, doi:10.1126/science.1083788.
- 259 Rempe, D.M., and Dietrich, W.E., 2014, A bottom-up control on fresh-bedrock
260 topography under landscapes: *Proceedings of the National Academy of Sciences of*
261 *the USA*, v. 111, p. 6576–6581, doi:10.1073/pnas.1404763111.
- 262 Robert, A., Pubellier, M., de Sigoyer, J., Vergne, J., Lahfid, A., Cattin, R., Findling, N.,
263 and Zhu, J., 2010, Structural and thermal characters of the Longmen Shan (Sichuan,
264 China): *Tectonophysics*, v. 491, p. 165–173, doi:10.1016/j.tecto.2010.03.018.
- 265 Rojstaczer, S., Wolf, S., and Michel, R., 1995, Permeability enhancement in the shallow
266 crust as a cause of earthquake-induced hydrological changes: *Nature*, v. 373, p. 237–
267 239, doi:10.1038/373237a0.

- 268 Shi, Z., Wang, G., Manga, M., and Wang, C.Y., 2015, Continental-scale water-level
269 response to a large earthquake: *Geofluids*, v. 15, p. 310–320, doi:10.1111/gfl.12099.
- 270 Skelton, A., et al., 2014, Changes in groundwater chemistry before two consecutive
271 earthquakes in Iceland: *Nature Geoscience*, v. 7, p. 752–756, doi:10.1038/ngeo2250.
- 272 Tipper, E., T., Bickle, M.J., Galy, A., West, A.J., Pomiès, C., and Chapman, H.J., 2006,
273 The short term climatic sensitivity of carbonate and silicate weathering fluxes:
274 Insight from seasonal variations in river chemistry: *Geochimica Cosmochimica Acta*,
275 v. 70, p. 2737–2754, doi:10.1016/j.gca.2006.03.005.
- 276 Torres, M.A., West, A.J., and Clark, K.E., 2015, Geomorphic regime modulates
277 hydrologic control of chemical weathering in the Andes–Amazon: *Geochimica*
278 *Cosmochimica Acta*, v. 166, p. 105–128, doi:10.1016/j.gca.2015.06.007.
- 279 Watanabe, N., Yonekura, N., Sagara, W., Cheibany, O.E., Marui, H., and Furuya, G.,
280 2005, Chemical Weathering and the Occurrence of Large-Scale Landslides in the
281 Hime River Basin, Central Japan. In: Sassa, K., et al. eds. *Landslides*. Berlin:
282 Springer Heidelberg, pp. 165–171.
- 283 Wang, J., Jin, Z., Hilton, R.G., Zhang, F., Densmore, A.L., Li, G., and West, A.J., 2015,
284 Controls on fluvial evacuation of sediment from earthquake-triggered landslides:
285 *Geology*, v. 43, p. 115–118, doi:10.1130/G36157.1.
- 286 White, A., and Brantley, S., 2003, The effect of time on the weathering of silicate
287 minerals: Why do weathering rates differ in the laboratory and field? *Chemical*
288 *Geology*, v. 202, p. 479–506, doi:10.1016/j.chemgeo.2003.03.001.

- 289 Xu, X., Wen, X., Yu, G., Chen, G., Klinger, Y., Hubbard, J., and Shaw, J., 2009,
290 Coseismic reverse- and oblique-slip surface faulting generated by the 2008 Mw 7.9
291 Wenchuan earthquake, China: *Geology*, v. 37, p. 515–518, doi:10.1130/G25462A.1.
- 292 Xue, L., et al., 2013, Continuous permeability measurements record healing inside the
293 Wenchuan earthquake fault zone: *Science*, v. 340, p. 1555–1559,
294 doi:10.1126/science.1237237.
- 295 Yoon, J., Huh, Y., Lee, I., Moon, S., Noh, H., and Qin, J., 2008, Weathering processes in
296 the Min Jiang: Major elements, $^{87}\text{Sr}/^{86}\text{Sr}$; $\delta^{34}\text{S}_{\text{SO}_4}$; and $\delta^{18}\text{O}_{\text{SO}_4}$: *Aquatic*
297 *Geochemistry*, v. 14, p. 147–170, doi:10.1007/s10498-008-9030-7.

298

299 **FIGURE CAPTIONS**

300

301 Figure 1. (a) Map of the 2008 M_w 7.9 Wenchuan earthquake region, showing sampling
302 sites, epicenter, co-seismic landslides (yellow polygons, from Li et al., 2014), and
303 contours for peak ground accelerations (PGA, from USGS earthquake hazard program,
304 <http://earthquake.usgs.gov/earthquakes>). Dots with stars inside are river water sample
305 sites at Weizhou and Zhenjianguan hydrological stations; dots (main stem) and
306 diamonds (tributaries) are revisited sites along the Min Jiang (1, Source of the Min Jiang;
307 2, Songpan; 3, Riwu Qu; 4, Heishui; 5, Zagunao; 6, Wenchuan; 7, Yingxiu; 8, Guan
308 Xian; 9, Dujiangyan). The magnitudes of observed change (Δ) in (b) Na^*/Ca and (c)
309 $^{87}\text{Sr}/^{86}\text{Sr}$ are shown by color shading.

310

311 Figure 2. (a) Seasonal patterns of Min Jiang river water discharge before and after the
312 2008 Wenchuan earthquake. Water discharge totaled $11.5 \times 10^9 \text{ m}^3$ (2005) and 11.9×10^9
313 m^3 (2010) at the Weizhou station. (b) Annual time series of Na^*/Ca ratios, with means
314 and standard deviations (1σ) on the right of the panel; pre-earthquake (EQ) data are from
315 samples collected near the Weizhou station in 2001 and 2002 (Qin et al., 2006).

316

317 Figure 3. Comparison of Min Jiang river dissolved (a) Na^*/Ca and (b) $^{87}\text{Sr}/^{86}\text{Sr}$ before
318 and after the 2008 Wenchuan earthquake. Data are for water samples collected from the
319 same sites and during the same seasons. Pre-EQ data are from Qin et al. (2006), Yoon et
320 al. (2008) and Huh (2010); post-EQ data are from the samples collected in 2011 in this
321 study. Pre-EQ $^{87}\text{Sr}/^{86}\text{Sr}$ data are not available for tributary sites. Analytical uncertainties
322 are smaller than symbol sizes. Gray shading in (a) shows the pre-EQ Na^*/Ca variability
323 of 0.02 (1σ). The site numbers correspond to Figure 1; ZJG and WZ are Zhenjiangguan
324 and Weizhou stations, respectively. Magnitude of changes (Δ) between pre-EQ and post-
325 EQ can be compared to the mean PGA in each catchment (c and d). Bars for WZ in (c)
326 show standard deviations (1σ) of the data in Figure 2b.

327

328 ¹GSA Data Repository item 2015xxx, xxxxxxxx, is available online at
329 www.geosociety.org/pubs/ft2015.htm, or on request from editing@geosociety.org or
330 Documents Secretary, GSA, P.O. Box 9140, Boulder, CO 80301, USA.