Aridland constructed treatment wetlands II: Plant mediation of surface hydrology
enhances nitrogen removal.
Authors: Christopher A. Sanchez ¹ , Daniel L. Childers ^{1,4} , Laura Turnbull ² , Robert F. Upham ³ ,
and Nicholas Weller ¹
¹ School of Sustainability, Arizona State University, Tempe AZ USA
² Department of Geography, Durham University, Durham, UK
³ City of Phoenix Water Services Department, Phoenix AZ USA
⁴ Corresponding author

ABSTRACT

15 Constructed treatment wetlands have been well established as effective and sustainable 16 solutions to the problem of urban water treatment and reuse. However, treatment wetlands 17 located in aridland cities may behave differently relative to their more mesic and humid 18 counterparts, and this could potentially impact their ability to deliver the ecosystem services that 19 are expected of them. Specifically, in hot, dry climates large water losses via evaporation and 20 plant transpiration may comprise a major component of whole-system water budgets. Our 21 primary goal was to develop a rigorous and informed model of how well these "working 22 wetlands" function in hot, arid climates by developing and comparing robust water and nutrient 23 budgets, as our process-based understanding of how mesic constructed wetlands function may 24 not be readily transferred to arid climates where constructed wetlands are becoming increasingly 25 widespread. At the Tres Rios constructed treatment wetland in Phoenix AZ USA, we quantified 26 water losses via plant transpiration and open water evaporation as well as inorganic N loads into 27 and from the whole wetland system and into the vegetated marsh. We found that water losses 28 due to transpiration and evaporation were remarkably high when compared to most mesic constructed wetlands. Total water losses via evaporation and transpiration peaked at 300,000 m³ 29 mo^{-1} (714 L H₂O m⁻² mo⁻¹) in the hot, dry summer months and averaged more than 70% of the 30 31 whole-system water losses over a 27 month time period. At the same time, the vegetated marsh 32 removed nearly all of the inorganic N that was supplied to it. Large transpirative water losses 33 moved large volumes of replacement water into the marsh via a "biological tide" that provided more opportunities for vegetation and soil microbes to process N and other target solutes. This 34 35 enhanced the N treatment efficacy of the Tres Rios constructed treatment wetland relative to 36 humid, mesic systems. To our knowledge, this is the first time that biotically-mediated surface 37 hydrology has been demonstrated in any wetland.

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40 Keywords: Constructed treatment wetlands, water budgets, nutrient budgets, nitrogen removal,
41 transpiration, urban wetlands

42 **1. INTRODUCTION**

43 The last two centuries have been characterized by a dramatic shift in human 44 demographics and land use practices, with more people choosing cities and urban areas as their 45 primary habitat. Specifically, since 1900 the human population living in urban areas has 46 increased from 10% to over 50% globally, with that proportion projected to grow to 80% by 47 2050 (Grimm et al. 2008). As urban centers grow in size and number, so do their demands for 48 resources and services. In particular, the last century has seen the transformation of many cities 49 into "sanitary cities", with highly centralized, capitalized, and expensive infrastructure designed 50 to keep inhabitants healthy (Melosi 2000; per Grove 2009). This development of infrastructure 51 has imparted considerable inertias on cities that make it more difficult for decision-makers to 52 envision, let alone implement, novel or transformative new solutions to growing problems 53 (Childers et al. 2014). Still, there are a number of ways in which urban infrastructure may make 54 cities more resilient and sustainable by designing to optimize key ecosystem services (Pickett et 55 al. 2013). One example of using infrastructure to optimize key ecosystem services is the 56 increasing use of constructed treatment wetlands as part of wastewater treatment plants in place 57 of expensive and energy-intensive treatment technologies.

58 Constructed treatment wetlands are a relatively low cost and low maintenance solution to 59 the age-old problem of wastewater treatment and, more recently, water reclamation. They are 60 especially attractive due to the fact that variable wastewater pathways and often relatively remote 61 placement of wastewater treatment systems demand simple technologies that require little active 62 management (Wallace and Knight 2006; Nivala et al. 2013). Most treatment wetlands are 63 characterized by the presence of macrophytic vegetation, waterlogged or saturated soils, and 64 municipal effluent that must be treated (Fonder and Headley 2013). While these wetlands share 65 similar ecosystem service expectations, or "jobs", of removing nutrients or other solutes (Tanner 66 et al. 2012), they may be further differentiated with regards to the level of "hard" engineering 67 components designed into them. These range from totally passive (horizontal flow) systems to 68 intensely managed systems featuring pulse loading or active aeration (Fonder and Headley 2010). 69 Design choices for treatment wetlands are often highly dependent on local or regional variables, 70 including imposed statutory wastewater treatment regulations or site-specific conditions. 71 Importantly, basic differences, such as climate, may limit general conclusions that might be 72 drawn from comparative studies of treatment wetland efficacy (Nivala et al. 2013). While

treatment wetlands are relatively similar in design and expectations, particular attention must be
paid to the way these systems function in their respective climatic settings.

75 In arid environments, which occupy more than 30% of the earth's land surface, one of the 76 biggest challenges facing cities is the scarcity of water. As a result, many of these cities are 77 increasingly turning to the reuse of treated municipal effluent as a means of reclaiming and 78 recycling water for various urban uses (Greenway 2005). Notably, in the aridland city of Phoenix 79 AZ USA, virtually all municipal effluent is reused within the city (Metson et al. 2012), and the 80 only significant export of water from the city is via evaporative losses to the atmosphere. The 81 challenge facing water reuse is that using reclaimed water in densely populated areas requires 82 that this water be clean. Increasingly, water service providers are turning to treatment wetlands 83 that, in addition to cleaning effluent, can also provide auxiliary and serendipitous ecosystem 84 services such as habitat for birds and other fauna and recreational spaces for urban dwellers. 85 However, constructed treatment wetlands in hot, arid cities may face unique challenges relative 86 to similar systems in more mesic and humid settings. In particular, treatment wetlands in hot, dry 87 climates should lose large amounts of water to evaporation and plant transpiration. How aridland 88 treatment wetlands perform under these conditions is not well quantified nor well understood, 89 and our primary goal for this research was to better understand these dynamics. To address this 90 goal, we quantified water and nutrient budgets for a constructed treatment wetland in Phoenix.

91 In Phoenix, annual precipitation averages 180 mm while potential wetland 92 evapotranspiration is over 2100 mm (Kadlec 2008). Thus, the objectives driving our research in 93 this wetland were: 1) to use a nutrient budget approach to determine net inorganic nitrogen (N) 94 removal by the vegetated marsh and the whole system; 2) to measure transpirational water losses 95 from the emergent macrophytes; 3) to calculate a dynamic water budget for the system based on 96 surface water inflows and outflows, precipitation inputs, and estimates for open-water 97 evaporation and plant transpiration, and; 4) to determine how well this aridland wetland removed 98 N from municipal wastewater effluent.

Historically, phase-change atmospheric losses of water from wetlands have been
presented as a single integrated process: evapotranspiration (E_t). Efforts to quantify and
understand E_t in wetlands have been confounded by large differences in estimates across
wetlands as well as divergences in the many ways that E_t may be calculated (Goulden et al.
2007). For the purposes of our study, we conceptually and empirically separated plant-mediated

104 transpirational water losses from open-water evaporation. Transpirational water losses have been 105 reported as both larger and smaller than water losses due to open-water evaporation, but it seems 106 probable that transpiration dominates water loss in wetlands found in hot, arid climates 107 (Pauliukonis and Schneider 2001). In fact, this has been documented for a wetland near our study 108 system (Kadlec 2006). Open-water evaporation should also be high in these systems, and there 109 are a variety of approaches to this calculation (Kumar et al. 2012). Many approaches to 110 estimating E_t in wetlands have calculated it as a residual term in whole-system water and energy 111 budgets that are based on meteorological data. However, a fundamental limitation with this 112 approach for our study was that this "subtraction" approach did not articulate spatial 113 heterogeneity and site-specific nuances that affect E_t , such as water level, plant community 114 composition, and plant canopy structure (Lott and Hunt 2001).

115 The objectives we addressed in this study are clearly not new to the constructed treatment 116 wetland literature. However, researchers have seldom rigorously studied wetlands that are 117 specifically designed to perform particular ecosystem services but are located in hot, arid 118 climates. Our research sought to tease apart the key components of the wetland water budget and 119 combine them with a nutrient budget in order to understand how well the constructed treatment 120 wetland we studied was performing. Our ultimate goal was to arrive at a more informed model of 121 aridland wetland ecosystem service performance that can facilitate the best possible adaptive 122 management of these systems.

123

124 **2. METHODS**

125 2.1. Site Description

126 Our study was conducted at the constructed treatment wetland associated with the 91st 127 Avenue Wastewater Treatment Plant, which is the largest such facility in Phoenix. Construction 128 of the 3-cell Tres Rios wetland system was completed in 2009-2010 and we began our work 129 there in Summer 2011. We have focused our work on the largest of three wetland cells, which 130 was planted first and came online in Summer 2010 (Figure 1). Each wetland cell is bounded by 131 levee roads, and the cell we studied is 42 ha in size, half of which is fringing vegetated marsh 132 and half of which is mostly open water with several small upland "islands". Water depths in the 133 marsh were consistently about 25 cm while open water depths were 1.5 to 2 m; because of the 134 way water was managed in the system, these depths did not vary. Our study cell received from

95,000 to over 270,000 m³ d⁻¹ of effluent, depending on the time of year. The marsh areas were
planted with seven emergent wetland species that are native to Arizona: *Typha latifolia, Typha domingensis, Schoenoplectus acutus, Schoenoplectus americanus, Schoenoplectus californicus, Schoenoplectus maritimus,* and *Schoenoplectus tabernaemontani. Schoenoplectus maritimus* was
a minor component of the plant community and was out-competed within the first year of our
research (Weller et al. this issue), so we focused our transpiration measurements only on the
other six species.

142

INSERT FIGURE 1 HERE

143 2.2. Transpiration Measurements and Evaporation Calculations

144 We utilized a dual-gradient experimental design to study the two major hydrologic 145 pathways of the wetland system: A whole-system gradient from the inflow to the outflow, and a 146 within-marsh gradient that included 10 transects from the shore to the open water that were 147 distributed evenly along the whole system gradient (Figure 1; for details, see Weller et al. this 148 issue). All of our sampling followed a bimonthly schedule (January, March, May, July, 149 September, and November), and we present data from July 2011 through September 2013 in this 150 paper. We measured leaf-specific gas flux along marsh transects that included all species groups 151 using a LICOR LI-6400 handheld infrared gas analyzer (IRGA). Measurements were made on 152 individual T. latifolia, T. domingensis, S. acutus, S. americanus, S. californicus, and S. tabernaemontani plants at 50-cm intervals from near the water surface to the top of the canopy. 153 154 We collected gas flux data continuously from early morning until as late in the day as was 155 feasible. For logistical reasons, we sampled one transect at a time, with individual plants chosen 156 haphazardly until all species present in that transect had been sampled a number of times. Due to 157 site access restrictions, we were unable to collect nighttime transpiration data and thus assumed 158 that nighttime transpirative water loss was negligible. Notably, this assumption is one reason that 159 our transpiration estimates are conservative.

160 Several of the *Schoenoplectus* species have thick stems which required us to modify the 161 stock LI-6400 IRGA sampling chamber in order to get a gas-tight seal without crushing the plant 162 stems. To remedy this, we used custom-made foam pads that allowed us to seal the leaf chamber 163 around these stems while minimizing damage to the plant material. Specifically, we were 164 interested in IRGA measurements of leaf-specific transpiration rate (Tr; mmol H₂O m⁻² sec⁻¹), 165 photosynthetically active radiation (PAR, µmol photons), ambient air temperature (°C), and

relative humidity (%). All of these measurements were taken using the IRGA's default internal sensors. Daily transpiration data were expressed in units of leaf surface area, and we converted surface area to dry weight biomass by drying and weighing 8-10 samples of different surface areas and generating relationships between surface area and dry weight biomass for each species ($r^2>0.90$ and p<0.01 in all cases).

Half of the 42 ha wetland cell we sampled was open water, as determined by digitizing
and measuring aerial imagery in ArcGIS (ver 10.0, ESRI, Redlands, CA). We calculated openwater evaporation (E, in mm hr⁻¹) using the Shuttleworth (1993) variation of the Penman (1948)
equation:

175 Open Water Evaporation = ((S*PAR) + C*6.43*(1+0.536*V)*VPD)/(2.39*(S*C))176 using hourly data for air temperature (T), relative humidity (RH), irradiance (PAR), barometric 177 pressure (P), vapor pressure deficit (VPD), slope of the saturation vapor pressure curve (S), 178 psychrometric constant (C), and wind speed (V) logged at an on-site meteorological station 179 operated by the City of Phoenix (hereafter referred to as the "meteorological station"). If the 180 hourly meteorological station data were missing or not reliable for any necessary variable (<5% 181 of all hourly readings), we did not calculate transpiration or open-water evaporation. Because of 182 these missing data, we are confident that our evaporation water loss values are underestimates 183 and are further confident that our transpiration estimates are conservative. Finally, we calculated 184 internal water loss using a dynamic monthly water budget for the whole system by summing the 185 whole-system hourly water fluxes via open water evaporation (E) and plant transpiration (Tr) to 186 monthly totals:

187 Internal Water Loss = E + Tr

and compared these data with the monthly whole-system water deficit, which we calculated by
summing daily inflows (SW_i), rainfall (R), and outflows (SW_o) for the month and subtracting the
latter from the former:

191 Water Budget Deficit = $SW_i + R - SW_o$

SW_i, SW_o, and R measurements were taken from whole-system data sets provided by the City of
Phoenix (see Section 2.4 for details). After consultation with on-site engineers, we determined
that vertical losses via percolation to groundwater were negligible so this parameter was not
included in our water budget deficit calculations.

196 2.3. Water Quality Sampling

197 On the same bimonthly schedule described above, we collected water quality samples 198 along both gradients. Triplicate water samples were collected by hand in acid-washed Nalgene 199 bottles at the inflow and outflow points of the study cell, and at the shore and open-water ends of 200 three marsh transects that represented the inflow-outflow whole system gradient (Figure 1). 201 Nitrate (NO_3) , nitrite (NO_2) , and ammonium (NH_4) samples were centrifuged to remove 202 particulates and analyzed on a Lachat Quick Chem 8000 Flow Injection Analyzer (detection 203 limit 0.85 µg NO₃-N/L and 3.01 µg NH₄-N/L). For simplicity, we will restrict our water quality 204 analysis to dissolved inorganic nitrogen (NH_4^+, NO_3^-) , and $NO_2^-)$. Differences in water quality 205 data were determined using a paired t-test for whole-system inflow versus outflow data, while an 206 unequal variance t-test was used for shore versus open water marsh transect data.

207

208 2.4. Whole System Data

209 Our research at the Tres Rios constructed treatment wetland is in partnership with the 210 City of Phoenix Water Services Department, and they provided us with key whole system 211 datasets for our analyses. We used their daily inflow (SW_i) and outflow (SW_o) data (in mgd, converted to m³ d⁻¹), measured using standard acoustic Doppler flow meters, to calculate total 212 213 monthly water flux into and out of our study cell from January 2012 through September 2013. 214 For the months when we sampled the system, we combined total monthly water flux with 215 inorganic N concentrations to calculate whole system monthly N loads. We used the previously 216 discussed hourly meteorological station data to scale our transpiration estimates in time and to 217 calculate evaporative water losses from the open-water component of our study cell (see Sections 218 2.2 and 2.5 for details). Because of inconsistencies in the rainfall data collected by the on-site 219 meteorological station, we summed daily rainfall data from the nearby Goodyear airport to 220 calculate monthly rainfall contributions (R) to the whole system water budget.

221

222 2.5. Data-Scaling and Water Budget Calculations

An important step in our whole-system water budget estimates was the scaling of leafspecific measurements of transpiration water losses from individual plants to 21 ha of marsh, and from transpiration measurements made over the course of individual days to a 2+ year timeseries. To scale transpiration water flux across space (Daily Tr_{System} ; m³ H₂O d⁻¹ whole system⁻¹),

- 227 we combined our IRGA measurements (Instantaneous Tr_{IRGA} ; mmol H₂O m⁻² sec⁻¹), corrected for
- dry-weight plant biomass (Area/Bio; $cm^2 gdw^{-1}$), with the whole system live macrophyte
- 229 biomass (Bio_{System}; Mg whole system⁻¹) calculated from the bimonthly data collected along our
- 230 10 marsh transects (see Weller et al. this issue):
- 231 Dail

Daily $Tr_{System} = Instantaneous Tr_{IRGA} * Area/Bio * Bio_{System} * 74.06$

where 74.06 was a dimensional analysis conversion for mmol H_2O to $m^3 H_2O$, seconds to days, 232 233 and gdw biomass to Mg biomass for the 21 ha marsh. To estimate whole system biomass 234 (Bio_{System}) we developed phenometric models that allowed us to non-destructively estimate live 235 biomass for all plant species by making simple allometric measurements in the field (per Daoust 236 and Childers 1998; Childers et al. 2006). Every two months, we measured all of the plants in five 0.25 m² quadrats that were randomly located along each of the 10 marsh transects shown in 237 Figure 1, for a total of 50 0.25 m^2 quadrats sampled (for details Weller et al. this issue). We used 238 239 simple linear interpolation to extrapolate plant biomass between bimonthly samplings, producing 240 daily estimates of live macrophyte biomass from July 2011 through September 2013.

241 To scale our plant-specific transpirational water losses in time, we used the micrometeorological data generated by the IRGA at the time of sampling and the corresponding 242 243 hourly data from the on-site meteorological station. Because plant transpiration flux is driven 244 largely by T, PAR, and RH (Sánchez-Carrillo et al. 2001), we first regressed hourly 245 meteorological station data for these three variables against the mean of all IRGA measurements 246 for the same variables for that same hour for each transpiration sampling day. We then generated 247 multiple regression models that related transpiration flux for each macrophyte species (per Mg 248 dry weight of plant biomass) to PAR, T, and RH as measured by the IRGA. These models 249 combined with the IRGA-meteorological station regressions allowed us to use continuous data 250 from the meteorological station to interpolate our species-specific transpiration measurements through time, resulting in whole-system daily transpirative losses (in m³ H₂O day⁻¹ for the 21 ha 251 of vegetated marsh) from July 2011 through September 2013. See Table 1 for a summary of all 252 253 parameters, applications, methods, and sampling frequencies used to for these calculations. 254 **INSERT TABLE 1 HERE** 255

256 3. RESULTS AND DISCUSSION

257 *3.1. Transpiration*

258 The phenometric biomass models were the same for T. latifolia and T. domingensis, and 259 for S. acutus and S. tabernaemontani (see Weller et al. this issue for details). For this reason we 260 combined our leaf-specific transpiration measurements for both species of each genus in both 261 cases. Transpiration rates for S. californicus and S. americanus were treated separately and 262 discretely. On a typical hot, dry day in July, these rates were as high as 10, 14, 13, and 9 mmol $H_2O m^{-2}$ leaf area s⁻¹ for *Typha* spp., *S. acutus* + *S. tabernaemontani*, *S. californicus*, and *S.* 263 americanus, respectively. In January, when air temperatures were low and plant biomass was at 264 its annual minimum, the highest transpiration rates were only 2, 4, 4, and 6 mmol $H_2O \text{ m}^{-2}$ leaf 265 area s^{-1} for Typha spp., S. acutus + S. tabernaemontani, S. californicus, and S. americanus, 266 267 respectively. These rates equated to July 2011 averages of 30 ± 0.8 , 16 ± 0.4 , 0.6 ± 0.3 , and $9.6 \pm$ 0.02 mm day^{-1} and January 2014 averages of 1.7 ± 0.2 , 1 ± 0.03 , 0.4 ± 0.03 , and $0.4 \pm 0.03 \text{ mm}$ 268 day⁻¹ for the entire standing stock of *Typha* spp., *S. acutus* + *S. tabernaemontani*, *S. californicus*, 269 270 and S. americanus, respectively. By comparison, Pedescoll et al. (2013) found that a T. 271 angustifolia-dominated constructed wetland in a similarly arid climate transpired a maximum of 23 mm day⁻¹, which is similar to the maximum rates from our treatment wetland. In contrast, 272 273 individual *T. domingensis* plants in the humid, mesic Florida Everglades transpired up to 11 mmol m^{-2} sec⁻¹ (Koch and Rawlik 1993)—well below the maximum plant-specific rates we 274 275 measured for this species. Reported transpiration rates for Typha spp. are even lower in cooler climates: 5.8 ± 0.9 mm day⁻¹ in marshes along Oneida Lake NY USA (Pauliukonis and 276 Schneider 2001) and 6.5 mm day⁻¹ in marshes in Ohio, USA (Martin et al. 2003). These findings 277 278 illustrate the dramatic differences in transpiration water loss between hot, arid and cooler, mesic wetlands. 279

We scaled these leaf-specific IRGA gas flux measurements in time to generate estimates of daily water loss by plant transpiration for the entire 21 ha vegetated marsh using hourly data from the on-site meteorological station. Regression relationships between meteorological station data and measurements within the canopy allowed us to extrapolate transpiration measurements between our bimonthly samplings. Multivariate regressions relating transpiration rate to T, RH, and PAR found that only T and PAR were significant predictors of transpiration by all four plant species groups, with relatively little interspecies variation in the strength of the relationships ($r^2 =$

287 0.57 to 0.67; all p-values < 0.01). We scaled leaf-specific transpiration measurements in space 288 using the live macrophyte biomass estimates for each species group from Weller et al. (this issue) after converting transpiration rates from mmol $H_2O \text{ m}^{-2}$ leaf area s⁻¹ to ml H_2O gdw 289 biomass⁻¹ day⁻¹ and ultimately to $m^3 H_2O day^{-1}$ for the entire 21 ha marsh. We found 290 291 considerable variation in transpirative water losses among the four species groups, with Typha spp. having the largest average transpiration rate $(2614 \pm 58 \text{ m}^3 \text{ H}_2\text{O} \text{ day}^{-1} \text{ or } 12.45 \pm 0.28 \text{ L} \text{ H}_2\text{O}$ 292 $m^{-2} day^{-1}$) and S. californicus the lowest average rate (221 ± 5 m³ H₂O day⁻¹ or 1.05 ± 0.02 L 293 $H_2O \text{ m}^{-2} \text{ day}^{-1}$). Most, but not all, of this variation was driven by differences in the amount of 294 295 live biomass (Weller et al. this issue). In fact, although Typha spp. has been steadily taking over 296 the macrophyte community in our study wetland (Weller et al. this issue), S. acutus and S. 297 tabernaemontani continued to make a contribution to overall daily transpirative water losses that 298 was disproportionate to their contributions to total biomass (Figure 2). This disproportionate 299 contribution to daily transpirative water loss was likely due to the fact that these large-stemmed 300 bulrushes transpire considerably more water per unit biomass than do cattails.

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INSERT FIGURE 2 HERE

302 As expected, all species groups showed strong seasonality, with the highest 303 transpirational water losses in the hot, dry summer months and the lowest losses in the winter 304 months (Figures 2 and 3). While close relationships between transpiration and seasonal variation 305 in climate and biomass have been well documented in the literature (e.g. Moro et al. 2004; 306 Pedescoll et al. 2013), the rates of water transpired from our Tres Rios system were considerably 307 higher than even those from other arid wetlands. For example, Drexler et al. (2008) found that an arid wetland in California, USA transpired on average 6 mm day⁻¹, and Goulden et al. (2007) 308 reported peak annual transpiration of less than 5 mm day⁻¹ for another marsh in California. A 309 310 study conducted by Bialowiec et al. (2014) in a cooler, more mesic wetland in Poland reported a maximum system-wide transpiration rate of 4.6 mm d⁻¹, and Abtew (1996) found that *Typha*-311 312 dominated marsh systems in the warm, humid Florida Everglades had maximum transpirational losses of only 3.6 mm day⁻¹. Overall, Tres Rios had transpirative water losses that were as much 313 314 as an order of magnitude greater than has been reported for other wastewater treatment wetlands. 315 Whole-system macrophyte biomass exhibited regular seasonality, with peak biomass in 316 July of approximately 400 to 600 Mg for the 21 ha marsh (Figure 3). As expected, whole-system 317 transpiration losses closely mirrored plant biomass (Figure 3). The highest total monthly

transpiration losses of approximately 200,000 to 300,000 m³ for the 21 ha marsh were in June and July, when ambient air temperatures regularly exceeded 45°C and relative humidity was typically below 10%, and often below 5%. As we noted above, during this study, the plant community was tending towards a bi-specific stand of *T. domingensis* and *T. latifolia*, but the contribution of several *Schoenoplectus* species to transpiration flux remained strong (Figure 2).

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INSERT FIGURE 3 HERE

324 3.2. Whole System Water Budget

325 Daily inflow and outflow data for our wetland study cell were used to calculate whole-326 system monthly water deficits. The inflow sensor was not fully operational until January 2012, 327 and therefore whole-system water deficit was not calculated until January 2012 (Figure 4). The 328 whole-system water budget was dominated by transpirational water losses and open water 329 evaporation, the latter of which represented a lower fraction of our total water budget—on the 330 order of a third or less—during the hotter months when plant biomass was higher and the plants 331 were most productive (Figure 4). Notably, precipitation was only a small fraction of the water budget. During our study, Tres Rios received 105 mm (105 L H₂O m⁻² yr⁻¹ or 44,100 m³ yr⁻¹), 332 122 mm (122 L H₂O m⁻² yr⁻¹ or 51,240 m³ yr⁻¹), and 167 mm (167 L H₂O m⁻² yr⁻¹ or 70,140 m³ 333 yr⁻¹) of rainfall annually in 2011, 2012, and 2013 respectively. These inputs were very small 334 when compared with transpiration and evaporation losses of 2.5 - 3.0 million m³ y⁻¹. In 2012 and 335 336 2013, precipitation comprised on average 0.8% (with a maximum of 4.1% in January 2013) of 337 the whole system monthly water deficit, while from 2011 to 2013 precipitation comprised on 338 average only 1.9% (with a maximum of 16.2% in December 2012) of the total water losses via 339 transpiration and open water evaporation.

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INSERT FIGURE 4 HERE

341 In many months our transpiration + evaporation-based water budget was remarkably 342 consistent with the inflow-outflow deficits, particularly in 2013 (Figure 4). We found that 343 transpiration + evaporation represented an average of 70% of the whole-system water deficit 344 throughout the study period (67% if March-May 2012 are included; see Figure 4 legend for 345 explanation). This consistency lends confidence in our scaled-up transpiration measurements and 346 calculations of open water evaporation. Notably, these contributions of evaporation and 347 transpiration are much higher than those reported for constructed wetlands in mesic, humid 348 climates. For example, Meuleman et al. (2003) reported that transpiration and evaporation

comprised only 13% of the water budget for a constructed wetland in the Netherlands, where the
climate is considerably cooler, more humid, and wetter than in Phoenix. Similarly, Kadlec et al.
(2010) reported that transpiration and evaporation represented only 3% of total water outputs for
a treatment wetland in Columbia, Missouri, USA and Favero et al. (2007) reported that these two
water losses comprised only about 10% of the water budget for a constructed wetland near
Venice, Italy.

355

356 *3.3. Water Quality and Nitrogen Budget*

357 We focused our analysis of nutrient sequestration on nitrogen (N), specifically dissolved 358 inorganic nitrogen (NO₃⁻ and NH₄⁺), because this was the focus of the City's Clean Water Act 359 permit for the treatment wetland. At the whole-system scale, inflow loads of both NO_3^- and NH_4^+ were nearly always higher than outflow loads (Figure 6a and 6b), with inflow loads averaging 360 532 ± 84 (SE) kg d⁻¹ and 169 ± 19 (SE) kg N d⁻¹ and outflow loads averaging 417 ± 67 (SE) kg N 361 d^{-1} and 87 ± 17 (SE) kg d^{-1} for NO₃⁻ and NH₄⁺, respectively. The monthly whole-system N 362 uptake rates averaged 3.51 ± 0.98 (SE) (16.71 ± 4.66 g N m⁻² mo⁻¹) and 2.48 ± 0.61 (SE) (11.81 363 ± 2.90 g N m⁻² mo⁻¹) Mg N mo⁻¹ for NO₃⁻ and NH₄⁺, respectively, from July 2011 through 364 365 September 2013. On average, 22% of nitrate inputs and 48% of ammonium inputs were taken up 366 between the whole-system inflow and outflow in spite of the short water residence time of the 367 system and the likelihood that much of this water did not come into physical contact with the 368 vegetated marsh that makes up 50% of the system by area.

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INSERT FIGURE 5 HERE

370 Within the vegetated marsh, we were able to compare N concentrations along the three 371 marsh transects but did not have water flux data to calculate actual nutrient removal rates. Still, 372 we consistently found much lower concentrations of dissolved inorganic N in the near-shore 373 water overlying the marsh, compared with concentrations at the marsh-open water interface. This 374 pattern was most dramatic for NO_3^{-} , with near-shore concentrations that were often at or below the detection limit (Figure 6a). Marsh removal of NH_4^+ was also consistent through time and 375 376 across seasons, but was less dramatic than with NO₃⁻ (Figure 6b). Weller et al. (this issue) 377 estimated N uptake rates by the plants themselves based on productivity and tissue stoichiometry 378 and found that 19% of the whole-system N removal could be directly accounted for by plant

- uptake. Even without actual nutrient flux estimates, it seems clear that the vegetated marsh wasactive and efficient in N removal, regardless of time of year.
- 381

INSERT FIGURE 6 HERE

382 3.4. A Plant-Mediated "Biological Tide"

383 The N concentration data from along our marsh transects (Figure 6) suggested that 384 whole-system efficiency might easily be improved if more N-rich water could directly interact 385 with the vegetated marsh. One possible mechanism for this is the lateral movement of water into 386 the marsh to replace losses via plant transpiration, particularly in the summer when the plants are 387 most productive, soil temperatures are highest, and transpiration is maximal. This lateral 388 movement of water into the vegetated marsh is what we are calling a plant-mediated "biological 389 tide." We refer to it as a "tide" because it is an active horizontal advection from open water areas 390 into the vegetated marsh that is analogous to astronomical tides seen in coastal wetlands. To our 391 knowledge, this is the first time that such control of surface hydrology by plants has been 392 documented in a wetland.

393 We investigated this "biological tide" by calculating the volume of water overlying the 394 marsh that was lost daily to transpiration. Water depth measurements at the 50 transect points 395 where aboveground plant biomass was measured showed that water depth was consistent across 396 space and time, with a mean of 24 cm. We subtracted the portion of this water volume that was 397 occupied by plant stems by scaling up stem culm diameter measurements, and used this 398 corrected volume to estimate transpirational losses. During the summer months, as much as 15 -20% (approximately 8,000 to 11,000 m³ H₂O day⁻¹) of the total volume of water overlying the 399 400 marsh was transpired daily and must have been replaced by a lateral flow from the open-water 401 areas to maintain the observed constant water depths (Figure 7). These values are conservative 402 estimates because we were not able to account for the volume of water occupied by extensive 403 thatched dead vegetation (i.e., wrack) on the marsh surface, thus suggesting that summertime 404 water residence times in the marsh were likely less than 4-5 days.

405

INSERT FIGURE 7 HERE

These transpirative losses from the vegetated marsh, and lateral flow to replace them, were thus substantial. We argue that this plant-driven control of wetland hydrology, or "biological tide", may be an important mechanism in arid systems to move additional nutrients into the marsh for uptake and processing (Figure 8). This "biological tide" is pulling water and

410 nutrients both laterally into the marsh and vertically down into the soils, creating more 411 opportunities for plants to take up N and for soil microbes to process N and other bioactive 412 solutes as Martin et al. (2003) and others have suggested. In fact, we argue that this "biological 413 tide" phenomenon makes the Tres Rios treatment wetland more efficient at nutrient removal, and 414 thus more effective at providing desired ecosystem services, than its counterparts in more mesic, 415 humid climates. This phenomenon may actually be at work in many other treatment wetlands 416 located in hot, dry climates. However, the high volume of water loss that drives the "biological 417 tide" means that there is a trade-off between the provision of water-cleansing ecosystem services 418 and the volume of water with improved water quality that can be exported from these systems for 419 local reuse.

420

INSERT FIGURE 8 HERE

421 *3.5. Summary*

422 We found that water losses due to transpiration and evaporation were remarkably high 423 compared with constructed wetlands in more humid, mesic climates. Total water losses via both processes peaked at over 300,000 m³ month⁻¹ in the hot, dry summer months and averaged more 424 425 than 70% of the whole-system water losses from June 2011 through September 2013. We found 426 that large transpirative water losses actually appeared to enhance N removal efficacy relative to 427 humid, mesic systems by drawing large volumes of replacement water and solutes into the marsh 428 via a "biological tide". This plant-mediated "tide" appears to be replacing more than 20% of the 429 total volume of water overlying the marsh during hot, dry summer months. Transpiration-driven 430 movement of shallow subsurface water has been documented in a number of wetlands, including 431 into tree islands in the Okavango Delta in Botswana (Bauer-Gottwein et al. 2007, Ramberg and 432 Wolski 2008) and in the Florida Everglades (Bazante et al. 2006, Troxler-Gann and Childers 433 2006, Sullivan et al. 2012). But, to our knowledge, this is the first time that biotically-mediated 434 surface hydrology has been demonstrated in any wetland.

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585	

586	6 FIGURE LEGENDS				
587	Figure 1: The Tres Rios constructed treatment wetland cell used in this study. Blue arrows show				
588	the effluent inflow and outflow points and white lines are the approximate locations of				
589	the 10 marsh transects (each $50 - 60$ m long). Water quality samples were collected at the				
590	inflow and outflow points and at 1) the northernmost transect closest to the inflow, 2) th				
591	southernmost transect nearest to the outflow, and the 3) eastern transect in between them.				
592	Total cell area = 42 ha; 21 ha of vegetated marsh and 21 ha of open water.				
593					
594	Figure 2: Total daily water loss via plant transpiration from the 21 ha of vegetated marsh.				
595	TYPHA=Typha domingensis and T.latifolia; SAC/STAB=Schoenoplectus acutus and S.				
596	tabernaemontani; SAM=S. americanus, and; SCAL=S. californicus.				
597					
598	Figure 3: Total live aboveground macrophyte biomass and total monthly transpiration water				
599	losses from the 21 ha marsh. Note that MT is equivalent to Mg.				
600					
601	Figure 4: Monthly total water losses due to open water evaporation and plant transpiration				
602	(stacked bars) compared with the inflow – outflow whole system water deficit. Note that				
603	the inflow sensor was not working properly prior to January 2012 or in March through				
604	May 2012.				
605					
606	Figure 5: Whole system nitrate (A) and ammonium (B) flux into the study cell (circles) and out				
607	of the study cell (squares).				
608					
609	Figure 6: Vegetated marsh concentrations of nitrate (A) and ammonium (B) flux near the				
610	shoreline (circles) and at the marsh-open water interface (squares) of the three water				
611	quality marsh transects.				
612					
613	Figure 7: Estimates of the volume of water overlying the 21 ha of vegetated marsh that was				
614	transpired daily by the plants and had to be replaced by the "biological tide".				
615					

Figure 8: Conceptual schematic of our hypothesized "biological tide" in which high rates of plant
transpiration, particularly during the hot, dry summer months, drive a lateral movement
of water and solutes into the marsh and vertically down into marsh soils from the openwater areas. This biotically-mediated water flow may be replacing the entire volume of
water overlying the marsh every 4 – 5 days at peak transpiration rates.

FIGURES

Figure 1:

625









644 Figure 5:









671	TABLE LEGENDS
672	Table 1: Summary of the parameters, applications, methods, and sampling frequencies for all
673	variables used to estimate whole-system plant transpiration, open water evaporation, and
674	water budgets.
675	

TABLES

678 Table 1:

Parameter and Application	Methods	Data Frequency
Transpiration, leaf area, leaf-level	Infrared Gas Analyzer (IRGA)	Bi-monthly
RH, leaf-level PAR, leaf-level T;		
used to estimate whole system		
plant transpiration		
Ambient RH, ambient PAR,	On-site meteorological station	Hourly
ambient T, wind speed, vapor	(City of Phoenix)	
pressure deficit, R; used to		
calculate open water evaporation		
and to scale plant transpiration in		
time		
Surface water inflow, surface water	Acoustic Doppler flow gauges	Daily
outflow; used to calculate water	(City of Phoenix)	
budget deficit		
Aboveground plant biomass for 4	Non-destructive phenometric	Bi-monthly
species groups; used to scale plant	technique (Weller et al. this	
transpiration in space	issue)	