- 1 APATITE FISSION TRACK AND RE-OS
- 2 GEOCHRONOLOGY OF THE XUEFENG UPLIFT,
- **3** CHINA: TEMPORAL IMPLICATIONS FOR DRY GAS
- 4 ASSOCIATED HYDROCARBON SYSTEMS
- 5 Xiang Ge^{1,2}, Chuanbo Shen^{1*}, David Selby², Dafei Deng^{1,3}, Lianfu Mei¹
- ⁶ ¹Key Laboratory of Tectonics and Petroleum Resources (China University of Geosciences),
- 7 Ministry of Education, Wuhan, 430074, China
- 8 ²Department of Earth Sciences, Durham University, Durham, DH1 3LE, UK
- 9 ³Institute of China National Offshore Oil Company Ltd, Shenzhen, Guangzhou, 510240,
- 10 *China*

11 ABSTRACT

12 Hydrocarbon evolution is extremely challenging to determine, both temporally and spatially, 13 in complex tectonic settings. Here we investigate the western margin of the Xuefeng Uplift, 14 China, which records multiple and protracted tectonic and hydrocarbon generation events. This timing of initial oil generation is recorded by low maturity bitumen (Type A), which 15 yields a Re-Os bitumen date of ~430 Ma, consistent with basin models and a ~405 Ma 16 bitumen Rb-Sr date. In contrast, Apatite Fission Track (AFT) data yield considerably younger 17 18 dates that reflect the timing and tectonic evolution of the Yanshan Orogeny from the northwest (~150 Ma) to southeast (~70 Ma). The youngest AFT date coincides with the 19 western margin of Xuefeng Uplift, where high maturity bitumen (Type B) occurs, which 20 21 yields a ~70 Ma Re-Os date. The AFT and Re-Os dates imply that both the last stage of the 22 Yanshan Orogeny and, by inference, the cessation of dry gas generation, occurred at ~70 Ma. 23 The Re-Os data of this study imply that the Re-Os chronometer can aid in constraining the 24 timing of oil generation, and of secondary and/or more mature hydrocarbon processes (e.g., thermal cracking/gas generation) in hydrocarbon systems worldwide. 25

26 INTRODUCTION

Being able to accurately constrain the timing of hydrocarbon generation is vital for 27 28 understanding the evolution of a petroleum system and for hydrocarbon exploration. Although the timing of oil generation can be estimated by numerous methods, establishing 29 the timing of gas formation during hydrocarbon evolution is currently less definitive 30 (Schaefer, 2005). To date, basin modelling, as well as hydrocarbon fluid inclusion analysis, 31 has been widely used to evaluate the timing of gas generation from oil cracking (Roberts et 32 al., 2004). However, imperfect kinetic models and poorly constrained parameters, such as the 33 34 palaeo-geothermal gradient, pressure conditions, and the physical properties of the strata, 35 hamper the accuracy of any derived age model (Braun and Burnham, 1992). Although studies are limited (Qiu et al., 2011), quartz fluid inclusion Ar-Ar dating has shown potential to yield 36 37 the timing of gas emplacement; however this method is hampered by requiring abundant gasliquid inclusions and the challenging analytical protocol (Liu et al., 2011; Qiu et al., 2011). 38 39 Apatite fission track (AFT) dating has also proven useful, as it can directly track the thermal history of a sedimentary basin (Donelick et al., 2005). Hitherto, the rhenium and osmium 40 (Re-Os) isotope system has provided temporal constraints for oil generation and migration 41 (Lillis and Selby, 2013; Selby and Creaser, 2005), but without any reference to the timing of 42 gas formation. In this study, we apply both AFT and Re-Os geochronology to the Majiang 43 and Wanshan reservoirs of the Xuefeng Uplift, China, to quantitatively constrain the timing 44 of tight gas generation, which is currently highly debated (Schenk et al., 1997). 45

46 GEOLOGICAL SETTING

The Majiang and Wanshan hydrocarbon reservoirs are the two largest in Southern China, containing a reserve exceeding 10 BBL of oil and 200 BCM of gas (Deng et al., 2014; Wu, 1989) (Fig.1). The gas reserve is close to that of the second largest gas field in China, Puguang (Ma et al., 2007), and the Ledovoe gas field, Russia (Pavlenko and Glukhareva,

51	2010), which each contain ~300 BCM of gas. The Majiang and Wanshan reservoirs are
52	located in the foreland basin belt along the western margin of the Xuefeng Uplift of the mid-
53	Yangtze block (Deng et al., 2014) (Fig. 1). The foreland basin is characterized by multiple
54	tectonic events (the Paleozoic Caledonian, Triassic Indosinian, and the Cretaceous-present
55	Yanshan-Himalaya orogenies) (Mei et al., 2012), which has resulted in numerous tectonic
56	models and timing constraints for hydrocarbon generation (Bai et al., 2013; Liu et al., 2011;
57	Tang and Cui, 2011). Oil and bitumen generation from Palaeozoic strata is proposed to have
58	occurred during tectonic subsidence associated with the Late Silurian Caledonian and
59	Triassic Indosinian Orogenies (Bai et al., 2013; Liu, 2011). However, the timing of gas
60	generation is less certain, with estimates tied to both the Triassic Indosinian and Cretaceous
61	Yanshan Orogenies (Liu, 2011; Xiang et al., 2008). For example, a ~228 Ma date derived
62	from Ar-Ar analysis of hydrocarbon-bearing fluid inclusions in the Majiang reservoir have
63	been used to suggest that hydrocarbon generation of both oil and wet gas was
64	contemporaneous with the Indosinian Orogeny (Liu, 2011). In contrast, basin modelling- of
65	the burial history of the foreland basin belt, (gas generated ~100Ma) coupled with fluid
66	inclusion analysis (Hhomogenization temperatures being > 150 °C)-, suggests that dry gas
67	formation occurred during the Cretaceous (Xiang et al., 2008), which coincides with the
68	Yanshan Orogeny as constrained by AFT dates (~92 Ma) from ~100 km to the north of the
69	study area, and quartz electron spin resonance (ESR) dates (75-62 Ma) along the Jiangnan
70	fault Zone in the Xuefeng Uplift, ~500 km to the east of the study area (Mei et al., 2010; Zhu
71	et al., 2011) (Fig.1).

Bitumen is commonly hosted within the pore spaces of Cambrian to Ordovician limestone
units, but is also present along fractures/cleavage planes of Silurian sandstone units (Zhou,
2006). Organic geochemistry defines two types of bitumen in the Majiang and Wanshan
reservoirs. In the Kaili area only, which represents a structural high, of the Majiang reservoir

(Fig.1), bitumen (Type A) has low maturity (Ro< 1.0), low T_{max} (~450°C), yellow 76 77 fluorescence, and high H/C ratio (>0.8) (Han et al., 1982; Zhou, 2006). In contrast, bitumen 78 (Type B) throughout the rest of the Majiang and Wanshan reservoirs (Fig. 1) has high 79 maturity (Ro>2.0), high T_{max} (~550°C), low H/C ratio (<0.6), no fluorescence, and contains high adamantine (3-4+ diamondoids) concentration (~50-300 ppm), which classifies the 80 bitumen as pyrobitumen (Shi et al., 2015; Wang et al., 2013). Though several processes can 81 form pyrobitumen (Lewan, 1997), dry gas and pyrobitumen can form contemporaneously by 82 thermal cracking of low maturity bitumen (e.g., Type A of this study) and crude oil (Huc et 83 al., 2000). To evaluate the timing of both bitumen types of the Majiang and Wanshan 84 reservoirs, and by inference the timing of oil and gas generation, we conduct AFT analyses 85 on the Paleozoic and Jurassic sandstones, and Re-Os analyses on both Type A and B bitumen. 86

87 SAMPLES AND METHODOLOGY

Samples for AFT analysis were collected from the Ordovician Dawan Formation (sample 2-88 89 33; Majiang reservoir), Jurassic Ziliujing Formation (sample 2-3; Majiang reservoir), Silurian Shiniulan Formation (sample 3-67; Wanshan reservoir), Triassic Guanling Formation (sample 90 3-97; Wanshan reservoir), and the Triassic Xiaojiangkou Formation (sample 3-18; Xuefeng 91 Uplift) (Fig.1A, Table DR1). Three samples (samples 3-97, 3-67 and 3-18) were broadly 92 collected along a NW-SE profile in the north of the study area, parallel to the propagation 93 direction of the Yanshan Orogeny. Samples 2-33 and 2-3 were collected ~200 km to SW of 94 95 samples 3-97, 3-67 and 3-18 (Fig. 1A).

Type A and B bitumen were sampled from outcrops across the Majiang and Wanshan
reservoirs (Fig.1A; Table DR2)._Five-Type A bitumen issamples, only known in outcrop in
the Majiang reservoir_; in-In_a ~20 m long exposure in the Ordovician Honghuayuan
Limestone Formation (O₁h) which is part of a structural high, -~20 km west from the city of
Kaili, <u>5 samples</u> were collected (Fig.1A). In contrast to Type A bitumen, Type B bitumen is

101 widely found, and predominantly occurs in a northeast-southwest trending belt in Early 102 Paleozoic limestone in both the Majiang and Wanshan reservoirs west of the Xuefeng Uplift 103 (Fig.1A). Nine samples were collected from the Ordovician Honghuayuan (O_1h) Formation 104 in the Majiang reservoir. Samples MJ-S1-B, MJ-S2-B, and MJ-S4-B were collected with an \sim 10 m interval spacing from an \sim 30 m section \sim 2 km south of the city of Majiang. Sample 105 XR-S1-B is from ~3 km west of the town of Xingren. Samples, HBZ-S1-B, HBZ-S4-B, 106 HBZ-S6-B and HBZ-S7-B are from a 20 m section, collected ~4 m apart, ~1 km northeast of 107 108 the town of Huoba. Sample MJ1-2 was collected from ~1 km northeast of the town of Pojiao. In addition, three samples were collected from the Cambrian Aoxi Formation (\in_{2a}) in the 109 Wanshan reservoir. Samples WS-S3-B, WS-S4-B and WS-S4-B6 were collected with an 110 ~5m interval spacing from an ~30 m section ~3 km from the city of Wanshan (Fig.1A). 111

112 The detailed analytical protocol and data for both the AFT and Re-Os analysis are presented113 in the DR.

114 **RESULTS**

115 The AFT data for the samples collected along a profile parallel to the propagation direction of the Yanshan Orogeny yield dates of 155 ± 13 Ma (sample 3-97), 97 ± 7 Ma (sample 3-67), 116 and 71 ± 6 Ma (sample 3-18) from NW to SE, respectively (Fig.1A; Table DR1). The two 117 118 samples in the southwest of the study area yield AFT dates of 150 ± 7 Ma (sample 2-3), and 119 123 ± 7 Ma (sample 2-33) (Fig.1A; Table DR1). The mean track lengths of all samples are 12.66-13.71 μ m (1 SD = 1.43 - 2.21 μ m; Table DR1). Thermal history modelling of the AFT 120 data shows continuous cooling from ~ 160 to 70 Ma from NW to SE (Fig.1B). 121 Type A bitumen samples possess 1.5-4.3 ppb Re, and 76.4-206.4 ppt Os, with ¹⁸⁷Re/¹⁸⁸Os and 122

123 $^{187}\text{Os}/^{188}\text{Os}$ compositions of 113-121 and 1.60-1.66, respectively (Table DR2). The Type A 124 bitumen Re-Os data yield a Model 1 (assumes that the assigned uncertainties are the only 125 reason for any scatter in the fit of the data) date of 429 ± 140 Ma, with an initial $^{187}\text{Os}/^{188}\text{Os}$ 126 (Os_i) value of 0.79 ± 0.27 (MSWD = 0.41) (Fig.2A). Type B bitumen possess typically higher 127 Re (2.5-15.2 ppb) and Os (40.0-498.1 ppt) abundances, and have a greater variability in their 128 ¹⁸⁷Re/¹⁸⁸Os and ¹⁸⁷Os/¹⁸⁸Os compositions (87-497 and 1.52-1.97, respectively; Table DR2). 129 The Type B bitumen Re-Os data yield a Model 3 (assumes that the scatter in the degree of fit 130 of the data is a combination of the assigned uncertainties, plus a normally distributed 131 variation in the ¹⁸⁷Os/¹⁸⁸Os values) date of 69 ± 24 Ma (Os_i = 1.45 \pm 0.09, MSWD = 9.6) 132 (Fig.2B).

133 DISCUSSION AND IMPLCATIONS

134 The Xuefeng Uplift and adjacent districts to the west are part of a piggyback thrust system 135 that has a NW transport direction (and therefore youngs towards the SE), which developed from the late Mesozoic, as a result of the Pacific-Yangtze plate collision (Yangshan Orogeny) 136 (Yan et al., 2003) (Fig.1A). Restoration of the section across the Majiang reservoir shows 137 ~20 % (~12 km) shortening occurred during the Cretaceous (Zhang, 2010). The AFT dates 138 (~150-70 Ma) obtained in this study coincide with the estimated timing of the Yanshan 139 Orogeny and support the SE younging direction of the thrust system (Mei et al., 2010; Zhu et 140 al., 2011). For example, samples 2-3 and 3-97 from 150 km to the northwest of the Xuefeng 141 Uplift belt possess similar AFT dates of ~150 Ma. Significantly younger AFT dates (~120 142 and ~100 Ma) were obtained for samples 2-33 and 3-67 from ~90 km NW of the Xuefeng 143 144 Uplift. An AFT date of 92 Ma was also obtained from ~50 km NW of Xuefeng Uplift (Mei et 145 al., 2010). The ~70 Ma AFT date (sample 3-18) recorded within the Xuefeng Uplift (Fig.1; Table DR1) is similar to the ESR dates (75-62 Ma) for the northern part of Xuefeng Uplift 146 (Zhu et al., 2011). The AFT dates indicates that the Yangshan Orogeny was a prolonged (~80-147 90 Ma) tectonic event, with the last phase occurring at \sim 70 Ma. 148

A thermal history derived from the AFT lengths using the HeFTy (v.1.8.1) software and employing a fanning curvilinear annealing model (Ketcham et al., 2009), with the parameters 151 of; 1) the annealing temperature (60-120 °C) of the apatite fission tracks, and 2) a present day 152 surface temperature of 20 °C, yields the same outcome as the determined by the AFT dates – 153 e.g., uplift events at \sim 150, \sim 100 and \sim 70 Ma from the NW to the SE across the western 154 margin of the Xuefeng Uplift (Fig.1; Table. DR1). The AFT dates and thermal history model also yield, by inference, a timeframe for the cessation of hydrocarbon generation. Given that 155 the thermal cracking of oil/bitumen requires temperature of $\geq 150^{\circ}$ C (Huc et al., 2000), the 156 spatial association of Type B bitumen and gas within/near the Xuefeng Uplift suggests that 157 158 hydrocarbon generation had ceased by ~70 Ma during the final stage of the Yanshan 159 Orogeny.

For the Re-Os analysis, <u>A</u> a Late Silurian-Early Devonian <u>Re-Os date timing</u> is recorded for 160 161 Type A bitumen (429 ± 140 Ma) (Fig. 2a). Although possessing a large uncertainty likely because of the low Re and Os abundance and limited spread in ¹⁸⁷Re/¹⁸⁸Os and ¹⁸⁷Os/¹⁸⁸Os 162 163 values, the nominal date calculated by the Re-Os data for Type A bitumen broadly coincides 164 with the timing for the records the initial timing of oil generation_in the reservoirs along the western margin of the Xuefeng Uplift in the Majiang reservoir based on and coincided well 165 with previous results. For example, burial basin models of the basin, ling oil-bearingand 166 hydrocarbon fluid inclusions with homogenization temperatures of ~100°C analysis in 167 Majiang area show Cambrian shales got matured and began to generate oil during late 168 Ordovician to early Devonian (Bai et al., 2013), and besides, Rb-Sr bitumen geochronology 169 that yields an age of $(405 \pm 20 \text{ Ma})$ on the bitumen from the Majiang reservoir suggests 170 hydrocarbon generated at late Silurian (Tang and Cui, 2011). 171

In agreement with the ~70 Ma AFT date is the Re-Os date $(69 \pm 24 \text{ Ma}, \text{MSWD} = 9.6)$ for Type B bitumen (Fig. 2b). As noted above Type B bitumen possess different geochemical characteristics from that of Type A bitumen, e.g., Type B bitumen possesses no fluorescence and contains adamantane, which indicate that the bitumen is pyrobitumen. Pyrobitumen could

176	form by the thermal cracking of bitumen and/or oil at temperatures >150 °C (Huc et al.,
177	2000). Along the western margin of the Xuefeng Uplift, continuous subsidence since the Late
178	Triassic lead to the burial of Paleozoic strata to a depth of \sim 5000 m, with temperatures >150
179	°C being attained during the Late Jurassic to Early Cretaceous (Bai et al., 2013; Han et al.,
180	1982). At this temperature, coincident with the formation of pyrobitumen, gas dominated by
181	methane, which is found in the Majiang and Wanshan reservoirs, also formed (Zhou, 2006).
182	<u>AlEven-</u> though spatially close to the Type B bitumen (~40 km), Type A bitumen in the Kaili
183	area did not suffer thermal cracking because it occurs structurally higher on the footwall of
184	the Shanban Fault (Fig. 1), and as a result, only experienced temperatures associated with oil
185	generation during the late stages of the Caledonian Orogeny (Latest Silurian) (Zhang, 2010).
186	Our AFT dates indicate that the Yanshan Orogeny caused the uplift of both the Majiang and
187	Wanshan reservoirs to a level where the ambient temperature was between 120 and 60 °C by
188	\sim 70 Ma. The AFT dates, coupled with basin <u>burial</u> models, fluid inclusion analysis
189	homogenization temperatures, and hydrocarbon composition-numerical modelling (evolution
190	of C_{14}^+ to C_1 compounds with time), suggest that thermal cracking of hydrocarbons (e.g.,
191	Type A bitumen and oil) may have occurred over a duration of 75 myrs (Late Jurassic to Late
192	Cretaceous) (Huc et al., 2000; Xiang et al., 2008). Although the Re-Os date for Type B
193	bitumen agrees with the AFT dates, the statistical fit of the Re-Os data yields a large MSWD
194	of 9.6 and a large date uncertainty (34%), which suggests that the Re-Os data have not fully
195	met the criteria to obtain a precise isochron-age. For example the requirements to yield a
196	precise isochron are;, the sample setrepresent contemporaneous formation, possess identical
197	initial isotope compositions (IOs), and have experienced no disturbance to the that the Re-Os
198	systematics post formationhave not been disturbed. Nevertheless, the agreement of the AFT
199	and Re-Os chronometers suggests that the process of thermal cracking has reset the Re-Os
200	systematics in the thermal cracked Type A bitumen. As such, the Re-Os Type B bitumen date

Formatted: Font color: Text 1, English (U.S.)

is recording the end of pyrobitumen formation and, by inference, the cessation of dry gas 201 202 generation. The agreement of the AFT and Re-Os Type B bitumen dates may also indicate 203 that the Re-Os systematics in hydrocarbons have a similar closure temperature range to that 204 of AFT (e.g., 120-60 °C; (Kohn and Green, 2002)). This range also coincides with the temperature condition (~100-140 °C) of thermochemical sulfate reduction (TSR) (Machel, 205 2001), a process which has also been proposed to disturb/even reset the Re-Os isotope 206 systematics in oil (Lillis and Selby, 2013). Assuming no significant disturbance to the Re-Os 207 208 systematics of the Type B bitumen since the uplift related to the Yanshan Orogeny, the scatter in the data and the large date uncertainty may be controlled by the IOs of the Type B bitumen 209 210 sample set.

211 Calculating IOs values for the samples at 70 Ma results in a range from 1.32 to 1.51 (Table. DR2), which is broadly defined by two groups (Group 1 = 1.32 to 1.39, n = 3; Group 2 = 1.42212 213 and 1.53, n = 9; Fig. 2). Treated as two groups, the Re-Os data yield identical, but more 214 precise (± 16 %) Model 1 Re-Os dates (Group 1 = 80 \pm 13 Ma, IOs = 1.30 \pm 0.05, MSWD = 1.3; Group $2 = 78 \pm 13$ Ma, $IOs = 1.45 \pm 0.04$, MSWD = 1.7) (Fig. 2b). Although more 215 precise, the Re-Os dates are still in agreement with the ~70 Ma AFT date and AFT thermal 216 217 modelling in the Xuefeng Uplift. The coupled AFT and pyrobitumen Re-Os dates both record 218 the last stage of tectonic uplift of the Xuefeng Uplift and yield the best estimate for the cessation of dry gas formation. The long-lived Yanshan Orogeny controlled the final 219 evolution of the hydrocarbon system and led to the near surface exposure of pyrobitumen. 220 221 Uplift, coupled with erosion, may indicate that dry gas could also have accumulated in 222 relatively deep regions, for example, the ramp or foredeep of the foreland basin. Traps under 223 the faults could also be potential gas reservoirs.

As illustrated by this study in the Xuefeng Uplift, the evolution of hydrocarbon systems can be extremely challenging to understand when affected by multiple tectonic events. This is 226 particularly the situation in the Yangtze Plate block, Southern China (Mei et al., 2012; Yan et al., 2003). Although many hydrocarbon shows (e.g., bitumen, oil and gas) are known (e.g., 227 228 Weng'an, Nanshanpin and Pingtang reservoirs) (Deng et al., 2014), only the Sichuan Basin 229 currently has producing fields (e.g., Puguang gas field (Ma et al., 2007)). Further, pyrobitumen occurs widely in basins all over the world, for example, the Alberta basin, 230 Canada, Dahoney basin, Nigeria and Basque-Cantabrian Basin, Spain. As such, Re-Os 231 232 bitumen and pyrobitumen geochronology (coupled with AFT dating) show potential to yield 233 quantitative timing of oil and gas generation, which may aid in the understanding of both the 234 temporal and spatial evolution of hydrocarbon systems.

235 ACKNOWLEDGMENTS

This work was supported by the NNSF of China (# 41372140), the Fundamental Research
Fund for the Central Universities, CUG (# 201536), Wuhan Science and Technology Project,
111 project and a CSC post-graduate award to XG. <u>DS acknowledges the support of the</u>
<u>TOTAL endowment fund.</u> We also thank Paul Lillis and Rachael Bullock for comments. <u>We</u>
thank 3 GEOLOGY reviewers for their comments, which helped improve the content of this
<u>paper.</u>

 1 GSA Data Repository item 201Xxxx.

243 REFERENCES CITED

- Bai, S., Peng, J., Liu, G., and Wang, Y., 2013, Hydrocarbon accumulation features and
 exploration potentials in Anshun Sag, Southern Guizhou Depression: Petroleum
 Geology & Experiment, v. 1, p. 24-28.
- Braun, R. L., and Burnham, A. K., 1992, PMOD: a flexible model of oil and gas generation,
 cracking, and expulsion: Organic Geochemistry, v. 19, no. 1, p. 161-172.
- Deng, D., Mei, L., Shen, C., Liu, Z., Tang, J., and Fan, Y., 2014, Characteristics and
 distributions of marine paleo-reservoirs in the northern margin of the Jiangnanxuefeng uplift, Southern China: Oil Shale, v. 31, no. 3, p. 225-237.
- Donelick, R. A., O'Sullivan, P. B., and Ketcham, R. A., 2005, Apatite fission-track analysis:
 Reviews in Mineralogy and Geochemistry, v. 58, no. 1, p. 49-94.
- Han, S., Wang, S., and Hu, w., 1982, The Discovery of a Paleopool in Majiang and its
 Geological Significance: Oil & Gas Geology, v. 3, no. 4, p. 316-327.
- Huc, A. Y., Nederlof, P., Debarre, R., Carpentier, B., Boussafir, M., Laggoun-Défarge, F.,
 Lenail-Chouteau, A., and Bordas-Le Floch, N., 2000, Pyrobitumen occurrence and
 formation in a Cambro–Ordovician sandstone reservoir, Fahud Salt Basin, North
 Oman: Chemical Geology, v. 168, no. 1, p. 99-112.
- Ketcham, R. A., Donelick, R. A., Balestrieri, M. L., and Zattin, M., 2009, Reproducibility of
 apatite fission-track length data and thermal history reconstruction: Earth and
 Planetary Science Letters, v. 284, no. 3, p. 504-515.
- Kohn, B. P., and Green, P. F., 2002, Low temperature thermochronology: from tectonics to
 landscape evolution: Tectonophysics, v. 349, no. 1–4, p. 1-4.
- Lewan, M., 1997, Experiments on the role of water in petroleum formation: Geochimica et
 Cosmochimica Acta, v. 61, no. 17, p. 3691-3723.

- 267 Lillis, P. G., and Selby, D., 2013, Evaluation of the rhenium-osmium geochronometer in the
- Phosphoria petroleum system, Bighorn Basin of Wyoming and Montana, USA:
 Geochimica et Cosmochimica Acta, v. 118, p. 312-330.
- Liu, Z., 2011, Hydrocarbon Accumulation in Marine Strata at Critical Tectonic Moment in
 Southern Middle and Upper Yangtze Intra-Continental Structural Belt, South China
 [Doctor: China University of Geoscience (Wuhan), 157 p.
- Liu, Z., Mei, L., Qiu, H., Shen, C., Tang, J., and Yun, J., 2011, 40Ar/39Ar geochronology
 constraints on hydrocarbon accumulation and destruction periods in the Bankeng
 paleo-reservoir in the southern margin of the middle Yangtze block: Chinese Science
 Bulletin, v. 56, no. 26, p. 2803-2812.
- Ma, Y., Guo, X., Guo, T., Huang, R., Cai, X., and Li, G., 2007, The Puguang gas field: New giant discovery in the mature Sichuan Basin, southwest China: AAPG bulletin, v. 91, no. 5, p. 627-643.
- Machel, H., 2001, Bacterial and thermochemical sulfate reduction in diagenetic settings—old
 and new insights: Sedimentary Geology, v. 140, no. 1, p. 143-175.
- Mei, L., Deng, D., Shen, C., and Liu, Z., 2012, Tectonic Dynamics and Marine Hydrocarbon
 Accumulation of Jiangnan-Xuefeng Uplift Geological Science and Technology
 Information, v. 31, no. 5, p. 85-93.
- Mei, L., Liu, Z., Tang, J., Shen, C., and Fan, Y., 2010, Mesozoic Intra-Continental Progress
 Deformation in Western Hunan-Hubei-Eastern Sichuan Province of China : Evidence
 from Apatite Fission Track and Balanced Cross-Section Earth Science- Journal of
 China University of Geoscience no. 2, p. 161-174.
- Pavlenko, V., and Glukhareva, E., Development of Oil and Gas Production and
 Transportation Infrastructure of Russian West Arctic Offshore Regions, *in*

- 291 Proceedings The Ninth ISOPE Pacific/Asia Offshore Mechanics Symposium2010, 292 International Society of Offshore and Polar Engineers.
- 293 Qiu, H.-N., Wu, H.-Y., Yun, J.-B., Feng, Z.-H., Xu, Y.-G., Mei, L.-F., and Wijbrans, J., 2011, 294 High-precision 40Ar/39Ar age of the gas emplacement into the Songliao Basin: Geology, v. 39, no. 5, p. 451-454. 295
- 296 Roberts, L. N., Lewan, M. D., and Finn, T. M., 2004, Timing of oil and gas generation of petroleum systems in the Southwestern Wyoming Province: The Mountain Geologist. 297
- 298 Schaefer, B. F., 2005, When do rocks become oil?: Science, v. 308, no. 5726, p. 1267-1268.
- Schenk, H., Di Primio, R., and Horsfield, B., 1997, The conversion of oil into gas in 299 300 petroleum reservoirs. Part 1: Comparative kinetic investigation of gas generation from 301 crude oils of lacustrine, marine and fluviodeltaic origin by programmed-temperature closed-system pyrolysis: Organic Geochemistry, v. 26, no. 7, p. 467-481. 302
- 303 Selby, D., and Creaser, R. A., 2005, Direct radiometric dating of hydrocarbon deposits using rhenium-osmium isotopes: Science, v. 308, no. 5726, p. 1293-1295. 304
- Shi, C., Cao, J., Bao, J., Zhu, C., Jiang, X., and Wu, M., 2015, Source characterization of 305 306 highly mature pyrobitumens using trace and rare earth element geochemistry: Sinian-307 Paleozoic paleo-oil reservoirs in South China: Organic Geochemistry, v. 83-84, p. 77-308 93.
- Tang, L., and Cui, M., 2011, Multiphase tectonic movements, cap formations and evolution 309 of the Majiang paleo-reservoir: Petroleum Science, v. 8, no. 2, p. 127-133. 310
- Wang, G., Li, N., Gao, B., Li, X., Shi, S., and Wang, T., 2013, Thermochemical sulfate 311 312 reduction in fossil Ordovician deposits of the Majiang area: Evidence from a 313 molecular-marker investigation: Chinese Science Bulletin, v. 58, no. 28-29, p. 3588-3594. 314

Wu, W., 1989, The Formation and Destruction of Palaeo-Oil-Reservoirs in the East of
Guizhou Province Geology of Guizhou, v. 1, p. 9-23.

317	Xiang, C., Tang, L., Li, R., and Pang, X., 2008, Episodic fluid movements in superimposed
318	basin: Combined evidence from outcrop and fluid inclusions of the Majiang ancient
319	oil reservoir, Guizhou Province: Science in China Series D: Earth Sciences, v. 38, no.

320 1, p. 70-77.

- Yan, D.-P., Zhou, M.-F., Song, H.-L., Wang, X.-W., and Malpas, J., 2003, Origin and
 tectonic significance of a Mesozoic multi-layer over-thrust system within the Yangtze
 Block (South China): Tectonophysics, v. 361, no. 3, p. 239-254.
- Zhang, J., 2010, The Research of Tectonic Evolution in Southern Guizhou Depression
 [Master: China University of Petroleum 97 p.
- Zhou, F., 2006, Hydrocarbon Accumulation and Petroleum Formation in the North Margin of
 Jiangnan Uplift [Master: China University of Geoscience (Wuhan), 88 p.
- 328 Zhu, Q., Yang, K., and Cheng, W., 2011, Structure evolution of Northern Jiangnan Uplift:
- Evidence from ESR Dating Geoccience v. 25, no. 1, p. 31-38.

330

331 FIGURE CAPTIONS

Figure 1. A. Simplified geological map of western margin of Xuefeng Uplift showing the location of the Type A and Type B bitumen in the Majiang and Wanshan reservoirs, and sandstone samples for Re-Os and AFT analysis, respectively. B. Thermal history inversions from AFT data for sandstone samples using HeFTy software (Table DR1). See text for discussion.

- 337 Figure 2. A. ¹⁸⁷Re-¹⁸⁷Os isochron for five Type A bitutmen samples from Kaili area, Majiang
- reservoir. B: ¹⁸⁷Re-¹⁸⁷Os isochron diagram for twelve Type B bitumen samples from Majiang
- and Wanshan reservoir. All data include 2 sigma level uncertainties and the error correlation
- 340 function, rho (Table DR 2). See text for discussion.