1	Establishment of modern circulation pattern at <i>c.</i> 6000 cal a BP in Disko Bugt,	
2	central West Greenland: Opening of the Vaigat Strait	
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19 20	Keywords: Disko Buat, Hold	pcene. Titanium, Magnetic susceptibility, Iceberg, West Greenland
21		
22	Abstract	
23	Variations in the Holocene circulation of the West Greenland Current (WGC) ir	
24	the Disko Bugt region have been reconstructed from a suite of sediment cores.	
25	Palaeoceanographic proxies include magnetic susceptibility (MS) and X-ray	
26	fluorescence Titanium counts, which document a major shift in circulation at c. 6000	
27	cal a BP. Before this date, sediments in southern Disko Bugt were characterised by	
28	high terrigeneous and basaltic input, suggesting widespread influence of meltwater	
29	plumes. Our data show that the WGC re-circulated in the southern Disko Bugt area	
30	because a potential northern pathway, the narrow Vaigat Strait, was blocked by	
31	icebergs that calved from marine outlet glaciers in eastern Disko Bugt. Sediments in	
32	southern Disko Bugt deposited after c. 6000 cal a. BP have significantly lower	
33	terrestrial and basaltic sediment input, which coincides with minimum Holocene ice	
34	sheet extent. The reduced meltwater and iceberg discharge to the embayment	
35	caused the Vaigat Strait to become free of blocking icebergs and terrigeneous input	
36	was partly diverted to the outer shelf. Thus, the modern circulation pattern of the	
37	WGC was established in the Disko Bugt region through the opening of the Vaigat	
38	Strait c. 6000 cal a	a BP.
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40 Introduction

On its pathway into northern Baffin Bay the relatively warm and saline West 41 Greenland Current (WGC) determines the modern oceanographic and climatic 42 conditions along the West Greenland shelf and its coastal areas (e.g. Ribergaard et 43 al., 2006). As it passes over the West Greenland shelf a branch of this water mass 44 penetrates into Disko Bugt (Figure 1). In Disko Bugt, Jakobshavn Isbræ one of 45 Greenland's largest outlet glaciers (draining today about 7 % of the Greenland Ice 46 Sheet (GIS)), produces through calving more than 10 % of icebergs from the ice 47 sheet (Bindschadler, 1984; Rignot and Kanagaratnam, 2006; Weidick and Bennike, 48 2007). Today, the majority of these icebergs are carried by the WGC northwards in 49 Disko Bugt, through the Vaigat Strait into the Baffin Bay. 50

Temperature and salinity variability of the WGC exerts an important control 51 52 (ocean forcing) on ice sheet dynamics; such as calving rates, meltwater-/iceberg production and glacier flow (e.g. Holland et al., 2008; Rignot et al., 2010) and 53 54 modulates the palaeoceanographic development of the region on a Holocene time scale (e.g. Moros et al., 2006; Seidenkrantz et al., 2008; Krawczyk et al., 2010; Lloyd 55 et al., 2007; 2011; Andresen et al., 2011; Perner et al., 2011, 2013). During the last 56 glacial (MIS 2) central West Greenland, and Disko Bugt itself, was occupied by a 57 large marine terminating ice stream, presumably an enlarged Jakobshavn Isbræ, 58 which extended approximately 30 to 50 km westwards across the continental shelf of 59 Disko Bugt (Kelly, 1985; Funder, 1989; Ingólfsson et al., 1990; Bennike et al., 1994). 60 This ice stream became unstable sometime before 10500 cal a BP, and started to 61 retreat into the eastern part of the embayment (Long et al., 2006; O'Cofaigh et al., 62 2012). Between 10000 and 9000 cal a BP re-appearance of the WGC in the Disko 63 Bugt area is reported (Funder and Weidick, 1991; Lloyd et al., 2005; Long et al., 64 2006; Weidick and Bennike, 2007; Funder et al., 2011). Initial WGC warming at this 65 time has also been reported from the far north of Baffin Bay (Knudsen et al., 2008; 66 67 Jennings et al., 2011). Recent benthic foraminifera based studies from the Disko Bugt area (Lloyd et al., 2011; Perner et al., 2011, 2013) highlight the prominent 68 influence of ocean temperature variability on ice margin stability/position on a range 69 70 of time scales during the Holocene. However, yet it is unclear when in the Holocene 71 the large-scale modern circulation pattern in Disko Bugt was established and which 72 role regional ocean-ice sheet interactions played during this process. A better

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understanding of the interaction between the ocean and cryosphere is fundamentalto predict future ice sheet behaviour and the sea level change involved.

In this study we present a reconstruction of large scale regional changes in ocean circulation in the Disko Bugt region, which uses magnetic susceptibility (MS) and X-ray fluorescence Titanium (Ti) data. Based on distinct variability in the bedrock geology in the area (Figure 2) the provenance of sediments is investigated from a suite of sediment cores along a transect, following the modern flow path of the WGC through the Disko Bugt area during the Holocene.

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82 Study area

Disko Bugt is a large marine embayment of about 40,000 km² located within 83 the relatively wide and shallow shelf area in central west Greenland (Figure 1). The 84 85 embayment is characterised by a rugged terrestrial and submarine topography, with two troughs across the central bay that join to form the deepwater trough, 86 87 Egedesminde Dyb, at the south-western entrance of the bay. This deep water trough, up to 900 m water depth (Brett and Zarudzki, 1979), is thought to be of glacial origin 88 (Long and Roberts, 2003; Roberts and Long, 2005) and extends south-westward 89 across the inner shelf adjacent to Disko Bugt. Outside the deep trough water depths 90 typically vary between 200 and 400 m (Kuijpers et al., 2001; Long and Roberts, 91 2002). Disko Island lies to the north of Disko Bugt itself and is separated from the 92 Nuussuag Peninsula and the mainland by the Vaigat Strait. This narrow strait, not 93 more than c. 25 km wide, is relatively shallow (average 300 m deep), is separated 94 from central Disko Bugt by a wide and shallow threshold (245 m deep; Andersen, 95 1981) and connects the north-eastern part of Disko Bugt with the open shelf 96 northwest of Disko, where the Uummannag Trough is found (Figure 1). 97

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99 Geological setting

100 The bedrock geology of eastern and south-eastern Disko Bugt is characterized 101 by Precambrian crystalline rocks (Figure 2). Common rocks are metasedimentary 102 and metavolcanic rocks, orthogneisses and granites (Escher *et al.*, 1976, Larsen and 103 Pulvertraft, 2000). To the east of a divide running approximately north-south, close to 104 the present day coast line of inner Disko Bugt the geology consists largely of Late 105 Archaean orthogneisses (Grade and Steenfeldt, 1999). Disko Island and the north-106 western part of the Nuussuag Peninsula are largely composed of Tertiary flood

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107 basalts with more restricted areas of Upper Cretaceous-Tertiary clastic sediments,

particularly along the eastern coastline of Disko Island (Henricksen *et al.*, 2000).

109 Tholeiitic picrites and olivine-rich basalts also underlie the outer continental shelf of

Disko Bugt (Figure 2). The seafloor of inner Disko Bugt and the Vaigat Strait are

111 mainly composed of Cretaceous-Palaeocene sediments (Figure 2). Precambrian

basement forms the seafloor in south-eastern Disko Bugt (Bonow, 2005).

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114 Modern oceanographic setting

The present day oceanographic conditions along the coast of West Greenland 115 are dominated by the WGC. The WGC is a combined water mass of: i) relatively 116 117 warm and saline Atlantic-sourced water from the Irminger Current (IC), a branch of the North Atlantic Current (NAC); ii) Arctic-sourced cold, low-salinity water from the 118 119 East Greenland Current (EGC; Buch, 1981; Tang et al., 2004); and iii) local meltwater discharge along the SW Greenland coast (Figure 1). A branch of the WGC enters 120 121 Disko Bugt from the southwest and flows northwards exiting primarily through the Vaigat Strait into Baffin Bay (Andersen, 1981; Bâcle et al., 2002; Ribergaard et al., 122 123 2006). Along this flow path icebergs and meltwater from a variety of outlet glaciers such as Jakobshavn Isbræ, Semerg Avangnardleg, Sermeg Kujatdleg and 124 Kangersuneg, which have calving fronts at the eastern margin of Disko Bugt into the 125 embayment, are carried into Baffin Bay (Figure 1). Fjords of these outlet glacier are 126 separated by shallow thresholds (Jakobshavn Isbræ: 150 m deep (Hogan et al. 127 2011); Semerg Avangnardleg: 200 m deep (Rignot et al., 2010)) from Disko Bugt. As 128 129 keel depths of calved icebergs can reach more than 200 m, they can get stuck at the fjords threshold. 130

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133 Material and Methods

During the cruise of the German research vessel *Maria S. Merian* in Greenland waters (Harff *et al.*, 2007), sediment sequences were recovered using multi (MUC) and gravity (GC) cores from the shelf southwest of Disko Bugt (sites 343340 and 343300), the Egedesminde Dyb (site 343330), the Vaigat Strait (site 343390) and the Uummannaq Trough (site 343520). Location and core information is given in Figure 1 and Table 1. Age-depth relationships for all sediment sequences are provided by accelerator mass spectrometry AMS ¹⁴C dates based on benthic foraminifera and molluscs (Table 2), calibrated with the Marine09 (Reimer *et al.*, 2009) calibration curve in Calib6.02 (Stuiver and Reimer, 1993). Following previous results from Lloyd *et al.* (2011), we applied a reservoir age correction ΔR of 140±35 years, which represents the modern ΔR value for the Disko Bugt area. West Greenland.

X-ray fluorescence scanning (XRF) provides semi-quantitative geochemical 146 sediment composition, and was performed on the Avaatech XRF core scanner from 147 the Royal Netherlands Institute for Sea Research (NIOZ). XRF measurements were 148 carried out directly on the split sediment core surface of the gravity cores during the 149 150 cruise. The sediment surface was carefully flattened and smoothed, and covered with a thin (4 µm) Ultralene film. All measurements were carried out at 10 kV at a 1 cm 151 152 step size. The analysed elements comprise AI, Si, S, CI, K, Ca, Ti, Mn, and Fe (Richter et al., 2006). 153

Magnetic susceptibility (MS) measurements were carried out onboard on all split sediment core sections, covered with a plastic film, using a Bartington MS2E1 sensor coupled to a TAMISCAN-TS1 automatic logging conveyor (Sandgren and Snowball, 2002). Down core resolution was set at 0.5 cm intervals for MUCs and GCs. Additional SIRM/X (saturation isothermal remanent magnetization) measurements from all core sites (I. Snowball, unpublished data) indicate that there have been no major grain sizes changes that can affect the MS measurements.

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163 **Results and interpretation**

In this section sediment provenance changes, inferred from Ti and MS data, 164 are presented for all gravity and multi cores that were taken on a transect following 165 the present day flow path of the WGC through the Disko Bugt area (Figure 1 and 2); 166 167 i) the shelf southwest of Disko Bugt (Figure 3); ii) the inner Egedesminde Dyb (Figure 4); iii) the Vaigat Strait and Uummannag Trough (Figure 5). Sediment deposition at 168 169 sites in the southern Disko Bugt region is influenced by the WGC and presently not subjected to direct meltwater discharge from outlet glaciers in eastern Disko Bugt 170 (e.g. Andersen, 1981; Buch, 1981; Buch et al., 2004). Icebergs and meltwater 171 discharged by outlet glaciers such as Jakobshavn Isbræ in the eastern Disko Bugt 172 173 region are routed along the flow path of the WGC and exit the embayment via the

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Vaigat Strait. The geological setting of the Disko Bugt area, see Figure 2, allows to use the MS and Ti data as proxies for terrestrial sediment input, but also to identify provenance of the terrigeneous material. Tertiary basalts on Disko Island and the Nuussuaq Peninsula (Figure 2) are a source of easily magnetized minerals such as magnetite and ilmenite to the sediments, producing a relatively high MS signal, whereas Ti provides an index of terrigeneous sediment input originating from Precambrian basement rocks (orthogneisses and granites) in eastern Disko Bugt.

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Modern distribution pattern of magnetic susceptibility data (MS) from the Disko Bugt area

The surface sections of the series of multi cores shows the modern regional 184 variation in the sediment MS signals and hence in the content of magnetic minerals 185 186 (Figure 2). On the south-western Disko Bugt shelf (343340-M, 343300-M) and in Egedesminde Dyb (343310-M, 343320-M, and 343330-M) a relatively low MS signal 187 is found, ranging between 40 and c. 80 (10^{-6} SI). In contrast to this, multi cores from 188 the Vaigat Strait (343380-M, 343390-M) and from the Uummannag Trough (343520-189 190 M) exhibit significantly higher MS signals, with the highest signal in the Vaigat Strait, averaging 400 (10⁻⁶ SI) and a slightly lower of 200 (10⁻⁶ SI) in the Uummannag 191 Trough (Figure 2). This contrasting spatial distribution in the modern sediment 192 magnetic properties suggests that sites north of Disko Island (343380-M, 343390-M, 193 and 343520-M) receive proportionally more eroded material from the basaltic terrains 194 of Disko Island and the Nuussuag Peninsula compared to sites in the south. The 195 distinct lower MS signals found at sites south of Disko Island suggests a generally 196 197 reduced influence of terrestrial sediment input from basaltic terrain. Thus, a distinct southeast to northwest gradient in the magnetic mineral content of sediments, linked 198 to the provenance of terrestrial sediments, is apparent along the modern flow path of 199 the WGC in the Disko Bugt region (Figure 2). 200

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- 203 Sediment provenance changes during the Holocene in the Disko Bugt area

204 Shelf southwest of Disko Bugt. Core 343340-GC is located mid-way across the 205 shelf southwest of Disko Bugt (Figure 1) and AMS ¹⁴C dates show that sediments 206 cover approximately the last 12000 years (Table 2). Sediments deposited between *c*. 207 12000 to 8000 cal a BP are composed of light to medium olive grey to brown silty

clay with darker brown mottling, occasional dropstones, shell fragments and 208 polycheate tubes. This sediment facies exhibits relatively high MS values, for most of 209 the section varying from 125 to 250 (10⁻⁶ SI), indicating abundant occurrence of 210 magnetic minerals in the sediments (Figure 3). Characterised by high Ti counts 211 ranging between 10000 to 8000 cps this site received large quantities of terrestrial 212 sourced material during this time interval. The age-depth profile indicates a rapid 213 sedimentation from c. 12000 to 8000 cal a. BP at a rate of c. 220 cm ka⁻¹. Initial 214 decrease of the MS signal (average 40 (10^{-6} SI)), starting from *c.* 9000 cal. a BP, is 215 accompanied by a notable decrease in the Ti counts (average 5000 cps) towards the 216 top of the core from c. 8000 cal a BP (Figure 3). This indicates an initial reduction of 217 218 high magnetic minerals, followed by an overall reduction in the contribution of terrigeneous clastic material to the site. Thus, a distinct shift in the sediment 219 provenance occurred, which is accompanied by a decrease of the sedimentation rate 220 to *c*. 25 cm ka⁻¹. 221

Core 343300-GC was collected from the inner south-western Disko Buat shelf 222 (Figure 1), which covers the last c. 10000 cal a BP (Table 2). From c. 10000 to 9000 223 224 cal a BP sediments are composed of light olive grey to olive grey clay, and include shell fragments and dropstones. In this sediment unit, similarly high MS (100 to 220 225 (10⁻⁶ SI)) and Ti (average 8000 cps) values are found as seen at the site on the outer 226 shelf southwest of Disko Bugt (Figure 3). The age-depth profile shows rapid 227 sedimentation at an average rate of c. 260 cm ka⁻¹ between c. 10000 to 6000 cal a 228 BP. The overlying sediments are composed of olive grey and moderate olive brown 229 rich clay with occasional occurrence of shell fragments. From c. 9000 to 6000 cal a 230 BP a prominent decline of the MS signal to a level of c. 40 (10^{-6} SI) is found, implying 231 a decreased input of sediment rich in higher magnetic minerals. During this interval Ti 232 values gradually decline to about 6000 cps, interrupted by a negative excursion at c. 233 8000 cal a BP to a level of 4000 cps. The overall Ti trend indicates a gradual 234 reduction in the supply of terrigeneous clastic sediments to the site. The MS signal 235 remains constantly low from c. 6000 cal a BP onwards, averaging c. 20 (10^{-6} SI), 236 accompanied by low Ti values (average 5000 cps). While the MS and Ti values as 237 well as the sedimentation rate (c. 60 cm ka⁻¹) are decreasing, the TOC content 238 increases, reaching peak values from c. 3000 cal a BP onwards (Figure 3). This can 239 be interpreted as representing a change in sediment provenance changed from 240

terrestrial to marine sources, with evidence of hydrographic conditions in the lateHolocene that favour marine productivity (Figure 3).

Outer Disko Bugt (Egedesminde Dyb). The core 343330-GC was collected 243 from Egedesminde Dyb, in the outer Disko Bugt (Figure 1). AMS ¹⁴C dates from this 244 core indicate that sediments below c. 4.30 m core depth were deposited between 245 8000 to 9000 cal a BP, and record a relatively high sedimentation rate (> 200 cm ka 246 ¹) at the site (see Table 2). From the base of the core up to 4.30 m the sediments are 247 composed of light to medium olive grey clay. Pebbles and dropstones are common 248 and occasionally shell fragments occur. This core section is characterised by 249 relatively high MS values (average c. 70 (10^{-6} SI)) and Ti counts (7500 cps), 250 accompanied by a low TOC (< 1.5 %) content (Figure 4). The MS signal at this site is 251 252 notably weaker than seen at sites 343340-GC and 343300-GC on the shelf 253 southwest of Disko Bugt, while Ti values are at a comparable level. This implies that in Egedesminde Dyb sedimentation is also subjected to terrestrial sediment input, 254 though containing less magnetic minerals. Overlying this unit between 4.30 to 4.20 m 255 core depth a distinct layer of moderate olive brown clay with an intercalated band of 256 light olive grey clay and well sorted medium sand is found. An AMS ¹⁴C date. 257 obtained just above this layer, is dated to c. 6000 cal a BP. This layer is 258 characterised by a marked decrease in the MS signal (drop to 5 (10⁻⁶ SI)). Ti counts 259 (drop to 3000 cps) and TOC content (drop below 0.5%). In the following from c. 6000 260 cal a BP onwards the TOC content increases to an average of c. 2 % and Ti counts 261 gradually decline to an average of c. 6000 cps, while the MS signal remains at a low 262 level (Figure 4). This distinct shift in the sediment properties indicates a change in the 263 sediment provenance, from terrestrial supply to a mainly marine source. 264

The Vaigat Strait. Core 343390-GC was obtained from the central Vaigat Strait north of Disko Island (Figure 1). Three available AMS ¹⁴C dates indicate that this core covers approximately the last 2000 cal a BP (Table 2). Sediments from this site are composed of light olive grey clay with occasional occurrence of dropestones.

²⁶⁹ Throughout the core, little variation is seen in MS values (average 380 (10⁻⁶ SI)) and

these indicate a high magnetic mineral concentration (Figure 5). The MS signal from

this core is about two times higher compared to that of sediments recovered from the

shelf southwest of Disko Bugt and in Egedesminde Dyb (Figures 3 and 4). In

addition, Ti counts varying between 12000 and 9000 cps are also significantly higher

than found at sites south of Disko Island (Figure 5). A stable MS signal and

consistently high Ti counts imply no changes in sediment properties or provenances
at the site through the last *c*. 2000 cal a BP. This is interpreted as reflecting a
consistent input of terrigeneous material rich in basaltic material from Disko Island
and the Nuussuag Peninsula.

Uummannag Trough. Core 343520-GC was recovered to the northwest of 279 Disko Island from the Uummannag Trough system (Figure 1). According to AMS ¹⁴C 280 dates shown in Table 2, sediments from this site cover approximately the last 10000 281 cal a BP. Sediments are composed of olive grey clay with minor silt and occasional 282 shell fragments. The MS signal ranges from 160 to 250 (10⁻⁶ SI) between 10000 to 283 8000 cal a BP and decreases to an average level of c. 180 (10^{-6} SI) from 8000 to 284 6000 cal a BP (Figure 5). During the period from 10000 to c. 6000 cal a BP the Ti 285 counts average c. 10000 cps, being significantly higher than found on the shelf 286 287 southwest of Disko Bugt during this time period. The sedimentation rate averages c. 100 cm ka⁻¹ during this interval. In the period from c. 6000 to c. 4000 cal a BP an 288 increase in the MS (up to 280 (10⁻⁶ SI)) and Ti (10700 cps) values is found, indicating 289 that more terrestrial sourced sediments, containing more magnetic minerals, were 290 291 deposited at the core site. This is accompanied by a decrease in the sedimentation rate to *c*. 66 cm ka⁻¹. From *c*. 4000 cal a BP towards the top of the core, the gradual 292 decrease of the Ti counts suggests reduction in the terrigeneous sediment supply to 293 the core site. Also a decreasing MS signal indicates reduced contribution of magnetic 294 minerals until c. 1000 cal a BP. During this interval the sedimentation rate increases 295 slightly to c. 78 cm ka⁻¹. It should be noted, however, that sediment provenance and 296 accumulation at this site is controlled by sediment supply from both Vaigat Strait 297 sources and from the Uummannag fjord system to the north. 298

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301 Discussion

Distinct regional variations have been observed in the MS signal of modern sediments along the core transect, which are linked to the source of sediments deposited at respective sites, with high MS values relating to proportionally greater input of basaltic material (Figure 1 and 2). Modern sediments on the shelf southwest of Disko Bugt and in Egedesminde Dyb exhibit distinctively lower MS signals than those found in the Vaigat Strait and in the Uummannaq Trough. These distinct contrasts follow the modern WGC flow path and reflect differences in sediment

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provenance related to geological variations in the Disko Bugt area (Figures 1 and 2). 309

The higher magnetic mineral content of the sediments in the Vaigat Strait and Uummannag Trough points to a significant contribution of basaltic material from 311

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Tertiary Basalts on Disko Island and the Nuussuag Peninsula. This basaltic material 312

is presumably washed into the Vaigat by meltwater, transport from glacier-fed rivers 313

as well as from meltout of material from icebergs, and then under the influence of the 314

WGC via the Vaigat further dispersed towards the shelf northwest of Disko Island. 315

However, this modern spatial distribution pattern of the MS signal found in 316 surface sediments was different in the past, as demonstrated by the gravity core 317 records represented here. These cores display a marked change in sediment 318 319 provenance and sedimentation rate, which occurs at c. 6000 cal a BP (Figures 3 and 4). High sedimentation rates (> 200 cm ka⁻¹), MS signals and Ti counts (Figures 3 320 321 and 4) prior to c. 6000 cal a BP indicate rapid sediment deposition from large meltwater plumes emerging from melting glaciers in eastern Disko Bugt, which also 322 323 included glacial (basaltic) sources on Disko Island. Previous terrestrial and marine studies show that following the LGM the ice sheet retreated from the shelf west of 324 325 Disko Bugt eastwards into the embayment and reached the eastern coast by c. 10000 to 9000 cal a BP (Long and Roberts, 2002; Lloyd et al., 2005; Weidick and 326 Bennike, 2007; Hogan et al., 2011). This retreat is thought to be in response to a 327 pronounced temperature rise over the GIS (Dahl-Jensen et al., 1998; Vinther et al., 328 2009). It is likely that a relatively warm WGC also supported regional deglaciation 329 during this period in the Disko Bugt area (Lloyd et al., 2005). A similar early initiation 330 of the WGC system has been reported from northern Baffin Bay (Knudsen et al., 331 2008). We postulate that the high content of terrigeneous sediments prior to c. 6000 332 cal a BP reflects discharge of large volumes of meltwater from the retreating ice 333 sheet in the eastern Disko Bugt area that primarily exited the embayment via the 334 (outer) shelf to the southwest rather than through the Vaigat Strait (Figure 6A). 335 336 Relatively high atmospheric temperatures over Baffin Bay (Kerwin et al., 2004) and a relatively warm WGC (Lloyd et al., 2005; Perner et al., 2013) supports significant ice 337 338 sheet melting within the area, which led to the production of large volumes of meltwater and enhanced iceberg calving. Such a strong early Holocene meltwater 339 340 influence on sediment deposition in inner Disko Bugt has also been reported by Seidenkrantz et al. (2012). The gradual decrease in deposition of basaltic material 341 342 (decreasing MS signals) between 8000 to 6000 cal a BP, on the south-western shelf

(Figure 3), implies an initial decline in the meltwater discharge from the ice sheet, 343 which might be linked to an initial opening of the Vaigat Strait or to an overall 344 reduction in meltwater production and iceberg calving. The strong meltwater imprint 345 on the sediment composition south of Disko Island prior to c. 6000 cal a BP provides 346 evidence of markedly reduced WGC flow into central and eastern Disko Bugt, which 347 suggests that the WGC was unable to exit Disko Bugt through the Vaigat Strait. We 348 assume that the narrow and shallow Vaigat Strait mainly acted as a meltwater 349 conduit and was largely blocked by icebergs discharged from the outlet glaciers. 350 351 Thus, the WGC was forced to re-circulate in Disko Bugt, and to continue its northwards flow via the shelf southwest of Disko Island (Figure 6A). 352

353 The marked shift in sediment provenance at c. 6000 cal a BP; i.e. reduced deposition of terrigeneous clastics rich in basaltic material, a lower sedimentation rate 354 (< 100 cm ka⁻¹) and increase of the TOC content, implies that the influence of 355 meltwater discharge from eastern Disko Bugt was significantly reduced on the south-356 357 western shelf and in Egedesminde Dyb (Figure 3 and 4). From c. 6000 cal a BP onwards the WGC is the dominant water mass controlling environmental conditions 358 359 and sediment deposition in this area. Benthic foraminiferal assemblage data from Perner et al. (2013) provide further support for this change to dominant WGC 360 influenced conditions after c. 6000 cal a BP in the southern Disko Bugt region. This 361 shift in the palaeoceanographic circulation pattern of the WGC occurs simultaneously 362 with the reported Holocene minimum ice sheet extent in eastern Disko Bugt from c. 363 6000 to 4000 cal a BP (Weidick and Bennike, 2007; Briner et al., 2010, Young et al., 364 2011) and maximum sea surface temperatures in the north-western part of Baffin Bay 365 (Ledu et al., 2010). As the ice sheet and outlet glaciers became progressively land-366 based, the relatively warm WGC could no longer cause basal tidewater glacier 367 melting on a large scale. Consequently, calving activity of outlet glaciers (e.g. 368 Jakobshavn Isbræ, Sermeg Avangnardleg, and Sermeg Kujadtleg; Figure 1) and 369 370 meltwater supply to the embayment was markedly reduced. This allowed the WGC to circulate via the Vaigat Strait, which led to further enhanced melting of icebergs and 371 372 in turn, increased WGC flow-through (Figure 6B).

Once the WGC was able to exit Disko Bugt through the Vaigat Strait, establishing the modern circulation pattern, iceberg and meltwater discharge from outlet glaciers was routed via the Vaigat Strait into Baffin Bay (Figure 6B). The available sediment record from the Vaigat only extends back until *c*. 2000 cal a BP

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and another record from the Vaigat Strait (Andresen *et al.*, 2011) covers the last *c*. 377 5200 cal a BP. These cores are too short to cover the change at c. 6000 cal a BP 378 shown here from the cores from the southern Disko Bugt region. The Uummannaq 379 Trough site (343520-GC), which covers the last c. 10000 cal a BP, exhibits similar 380 high MS and Ti values as found on the shelf southwest of Disko Bugt and in 381 Egedesminde Dyb prior to c. 6000 cal a BP. Following the opening of the Vaigat 382 383 Strait, the sedimentation rate decreases and a higher proportion of basaltic material from Vaigat Strait origin is deposited until c. 4000 cal a BP at site 343520-GC (Figure 384 5). This site in the Uummannaq Trough receives not only sediments from the Vaigat 385 Strait, but also from the Uummannaq fjord system to the north. Therefore, changes in 386 387 sediment provenance recorded in this core may also be influenced by changes in input from glaciers flowing into the Uummannaq Trough system. 388

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391 Conclusions

Investigations of sediment provenance from a suite of sediment cores from the 392 393 Disko Bugt region reveal that significant meltwater production during Deglaciation from the Greenland Ice sheet influenced local environmental conditions within the 394 area, lasting until c. 6000 cal a BP. Our data show a marked shift in sediment 395 provenance at *c*. 6000 cal a BP recording the development of the modern circulation 396 pattern of the West Greenland Current (WGC) in the Disko Bugt area. Before this 397 time, meltwater and icebergs discharged from outlet glaciers in eastern Disko Bugt 398 (e.g. Jakobshavn Isbræ) filled the Vaigat Strait, the modern exit of the WGC from 399 Disko Bugt and prevented the WGC reaching the innermost part of the embayment. 400 We find that establishment of the modern circulation pattern by the opening of the 401 Vaigat Strait at c. 6000 cal a BP coincided with minimum ice sheet extent in eastern 402 Disko Bugt. It is likely that, before the opening of the Vaigat Strait a warm WGC 403 404 supported basal melting of outlet glaciers. Thus extensive meltwater plumes and iceberg calving forced the WGC to re-circulate in the southern Disko Bugt area and to 405 406 continue north over the shelf southwest of Disko Island. Once the ice sheet had 407 retreated to a predominantly land-based position decreased meltwater and iceberg 408 discharge permitted the WGC to enter and flow through the Vaigat Strait, similar to the present day situation. Our results show that the interaction between ice sheet and 409

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- 410 ocean circulation changes (WGC) modulated the local environmental conditions
- during the Holocene within the Disko Bugt region.
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414 Acknowledgements

- The authors thank captain and crew of the RV Maria S. Merian for assistance and
- 416 help during core retrieval (cruise MSM05/03). This work is part of the 'Disco Climate'
- 417 project (MO1422/2-1), which was funded by the Deutsche Forschungsgemeinschaft418 (DFG).
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Figure 1: Map of study area with schematic sea floor topography and coring sites
from MSM05/03 on the shelf southwest of Disko Bugt (sites 343340 and 343300), in
Egedesminde Dyb (site 343330), the Vaigat Strait (site 343390) and the Uummannaq
Trough (site 343520), along the modern flow path (green arrows) of the West
Greenland Current (WGC) in the Disko Bugt area. The inset shows the WGC's
source currents, the East Greenland Current (EGC) and the Irminger Current (IC).



Figure 2: Geological setting of the Disko Bugt area adapted from Weidick and

Bennike (2007). Full colours display the onshore, whereas pastel shades display the

offshore geology. Additionally, the range of magnetic susceptibility (MS) signals is

626 plotted for multi cores from the area, as listed in table 1.

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Figure 3: MS (solid line), Ti (circles), and TOC (crosses) data plotted against core

depth (cm) from site 343340-GC on the outer south-western Disko Bugt shelf and

from site 343300-GC on the inner shelf. Additionally AMS ¹⁴C radiocarbon ages are
 shown from the respective core sites. Red shaded bar marks depth in cores of

sediment provenance changes, displaying the time of the opening of the Vaigat

634 Strait. Average sedimentation rates are presented for the time period before and after

635 *c.* 6000 cal a BP.

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Outer Disko Bugt (Egedesminde Dyb)

Figure 4: MS (solid line), Ti (circles), and TOC (crosses) data plotted against core
depth (cm) from site 343330-GC from Egedesminde Dyb. Additionally, AMS ¹⁴C
radiocarbon ages are shown. Red shaded bar marks depth in cores of sediment
provenance changes, displaying the time of the opening of the Vaigat Strait. Average
sedimentation rates are presented for the time period before and after *c*. 6000 cal a
BP.

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Figure 5: MS (solid line) and Ti (circles) data plotted against core depth (cm) and AMS ¹⁴C radiocarbon ages, plotted against core depth, from site 343520-GC in the Uummannaq Trough and site 343390-GC from the Vaigat Strait. Red shaded bar marks depth in cores of sediment provenance changes, displaying the time of the opening of the Vaigat Strait. Average sedimentation rates for core 343520-GC are

presented for the time period before and after *c*. 6000 cal a BP.



Before *c.* 6000 cal a BP: WGC re-circulates in southern Disko Bugt

At c. 6000 cal a BP: Opening of the Vaigat Strait as exit for the WGC from Disko Bugt

- Figure 6: Schematic illustration of the evolution of the West Greenland Current 656 (WGC) in Disko Bugt during the Holocene. A) Conditions before c. 6000 cal a BP. 657 The Vaigat Strait is blocked by ice bergs. The flow path of the WGC is shown by 658 green arrows entering the western region of Disko Bugt. The WGC re-circulates and 659 660 exits back out of Disko Bugt and continues flowing northwards. B) Conditions from c. 6000 cal a BP onwards. Modern circulation system in Disko Bugt develops as the 661 Vaigat Strait becomes unblocked. The flow path of the WGC is shown in green 662 arrows flowing through Disko Bugt and the Vaigat Strait into the Uummannag system 663 to the northwest. C) Magnetic susceptibility transect of gravity cores (343340-GC, 664 343300-GC, and 343330-GC) south of Disko Island. Cores mark clearly shift in 665 sediment provenance from high to reduced terrigeneous and basaltic sediment input 666 at c. 6000 cal a BP. 667
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