Re-Os geochronology and oil-source correlation of the Duvernay petroleum system, Western Canada sedimentary basin: Implications for the application of the Re-Os geochronometer to petroleum systems

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Abstract

The rhenium-osmium (Re-Os) geochronometer has been applied to many petroleum systems worldwide. However, it is debated if the Re-Os systematics in petroleum actually records the timing of oil generation. Here, we investigate the Re-Os isotope systematics of the Duvernay petroleum system in the Western Canada sedimentary basin, which has been shown to be a relatively simple petroleum system that is associated with oil generated during the Late Cretaceous - Early Eocene Laramide orogeny from a single source.

The organic geochemistry of the Duvernay oils (pristane/phytane ratios of ~1.5, smooth homohopane profile, C_{29}>C_{27}>C_{28} regular sterane distribution, and predominance of diasteranes over regular steranes) strongly suggest the oil source to be that of Type I-II marine organic matter of the Late Devonian Duvernay Formation.

The asphaltene fraction Re-Os isotope data of the Duvernay oil yield an age of 66 ± 31 Ma, which is in excellent agreement with the timing of the main-stage hydrocarbon generation of the Duvernay Formation based on basin models. Further, the similarity between the $^{187}Os/^{188}Os$ compositions of the Duvernay Formation source rock (0.46 to 1.48) and the oil (0.55 to 1.06) at the time of oil generation supports the hypothesis that the $^{187}Os/^{188}Os$ composition of an oil is inherited from the source unit at the time of oil generation, and therefore shows no, or limited influence from interaction with basin fluids. This study supports that the Re-Os isotope systematics of an oil can yield the timing of oil generation, and be used to trace its source.

1. Introduction

Establishing an accurate and precise age of hydrocarbon generation is crucial for understanding the evolution of a petroleum system. Timing constraints for hydrocarbon generation can be achieved through many techniques, for example, basin modelling with paleotemperature and petroleum generation kinetics; or the timing of oil generation can be inferred by constraining the earliest migration of oil using radiometric dating ($^{40}Ar/^{39}Ar$) of fluid inclusions and diagenetic minerals, e.g., authigenic illite and K-feldspar (Eadington et

al., 1991; Kelly, 2002; Mark et al., 2005; Mark et al., 2010; Welte et al., 2012). Despite the contribution of these methods to the understanding of petroleum system evolution, they are limited by poorly constrained parameters (e.g., paleogeothermal gradient) or difficulties associated with sampling and analyzing (e.g., grain size versus purity of authigenic minerals for radiometric dating). During the past decade, the rhenium-osmium (Re-Os) chronometer has been demonstrated to be a promising radiometric tool to directly date petroleum samples and for oil-source correlation (Selby and Creaser, 2005b; Selby and Creaser, 2005a; Selby et al., 2007; Finlay et al., 2011; Finlay et al., 2012; Rooney et al., 2012; Lillis and Selby, 2013; Cumming et al., 2014; Ge et al., 2017).

Rhenium-Os analysis demonstrating chronologic information in petroleum systems was first shown by Selby et al. (2005), who derived a Re-Os date of 374.2 ± 8.6 Ma (MSWD = 11.7, Model 3) for bitumen formation in the Polaris Mississippi Valley Type deposit, Nunavut, Canada, which is in agreement with contemporaneous Pb-Zn mineralization dated at 366 ± 15 Ma (Rb-Sr sphalerite methodology). In addition, Re-Os data derived from highly biodegraded oil from the Oil Sand deposits of Alberta, Canada, give an age of 111.6 ± 5.3 Ma (MSWD = 2.2, Model 3) (Selby and Creaser, 2005a), which is in agreement with the basin model of Riediger et al. (2001). Consequently, the Re-Os date is taken to record the timing of oil generation and migration in the world's largest oil sand system. However, the Re-Os age for the oil sands is inconsistent with some other basin models proposed for this complex oil system, e.g. Higley et al. (2009), Berbesi et al. (2012), and a recent study on trap restoration

and the bitumen present in kimberlites (Tozer et al., 2014) which could relate to secondary migration of oil in the Alberta basin.

Oil Re-Os dates (68 ± 13 Ma) from the UK Atlantic Margin (UKAM) petroleum system (Finlay et al., 2011) also correspond to the age of petroleum generation determined through basin modelling (Lamers and Carmichael, 1999) and K-feldspar cement $^{40}Ar/^{39}Ar$ dating (Mark et al., 2005). A Re-Os study of oils from the Phosphoria petroleum system, Bighorn basin, USA produced an age of 211 ± 21 Ma, corresponding to the start of petroleum generation according to basin modelling rather than the re-migration of oil (during the Laramide orogeny which starts from 80 - 70 Ma and ended until 55 - 35 Ma, although the exact dates are in dispute; English and Stephen, 2004) or the deposition of the source rock (during the Permian, ~270 Ma) (Lillis and Selby, 2013). Although the UKAM and Phosphoria oils are variably biodegraded, the process has not appreciably disturbed the Re-Os systematics from recording the timing of generation of oil. This is supported by the fact that Re and Os in oil are hosted predominantly by the asphaltene fraction (Selby et al., 2007), which is the most resistant phase to biodegradation (Wenger et al., 2001; Peters et al., 2005). More recently, Re-Os bitumen data has provided insights into the temporal hydrocarbon evolution in the Longmen Shan Thrust Belt, southwest China, i.e. the multiple oil generation episodes (486 Ma and 172 - 162 Ma) and sources, and its relationship to tectonism, and additionally discussed implications for future exploration of the adjacent petroliferous Sichuan Basin (Ge et al., 2017). Furthermore, a study of the Green River petroleum system in the Uinta basin, USA, demonstrated the capacity of the Re-Os chronometer to yield the

timing of oil generation (19 ± 14 Ma) in lacustrine petroleum systems in addition to marine petroleum systems (Cumming et al., 2014).

Unlike previous studies using multiple oil and/or bitumen samples to establish a Re-Os date, a study focusing on two oil families of the Gela oil field, southern Sicily, Italy, which are sourced from the Noto Formation (Upper Triassic, Rhaetian) and the Streppenosa Formation (Lower Jurassic, Hettangian), respectively, promotes the establishment of Re-Os dates using fractions of one single oil sample, in order to negate possible disadvantages of using multiple oil samples from a wide geographic region (Georgiev et al., 2016). The authors analysed the asphaltene and maltene fractions, which were separated using a series of solvents (*n*-pentane, *n*-hexane, *n*-heptane and *n*-decane). The best-fit lines of the Re-Os data for asphaltenes from one oil family and for maltenes from the other oil family yield reasonable oil generation ages for each family (27.5 ± 4.6 Ma and 200.0 ± 5.2 Ma, respectively). However, this approach of obtaining a Re-Os oil generation age from one single oil sample is potentially confined to oils with high asphaltene contents and Re and Os abundances. It also depends on the difference of asphaltenes and maltenes separated by different *n*-alkanes from the same oil to obtain a range in Re-Os isotopic compositions sufficient to obtain an isochron. As such, oils with low Re and Os abundances will be too analytically challenging to obtain precise data to distinguish these differences.

Hydrous pyrolysis experiments on organic-rich sedimentary rocks (Rooney et al., 2012; Cumming et al., 2014) have shown that during thermal maturation, Re and Os will be

transferred to the generated oil, although the majority of the Re and Os remains in the source rock. The $^{187}Re/^{188}Os$ values show no consistent relationship between the source rock and the generated oil, but the $^{187}Os/^{188}Os$ composition of a source rock is inherited by the generated oil (Selby et al., 2005; Finlay et al., 2011; Rooney et al., 2012; Cumming et al., 2014). The relationship between the $^{187}Os/^{188}Os$ composition of a source rock and the $^{187}Os/^{188}Os$ composition of the generated oil (at the time of generation) is employed in the aforementioned studies for oil-source correlation. For example, the attempt of identifying the source(s) for the Polaris bitumen through the Os isotope composition of a few Phanerozoic and Proterozoic potential source rocks led to the conclusion of the mix of several source rocks or a main source unit that was not sampled in the study for the studied bitumen. In addition, the coupling of Pt/Pd values with $^{187}Os/^{188}Os$ compositions has been employed as a viable inorganic oil-source correlation tool in recognizing the Lower Jurassic Gordondale Member as the predominant source of the Canadian oil sands (Finlay et al., 2012), which is supported by basin and migration modelling studies (Higley et al., 2009; Berbesi et al., 2012).

Although present studies show that biodegradation has little influence on the Re-Os systematics of asphaltene (Selby and Creaser, 2005a; Selby et al., 2005; Finlay et al., 2012; 2014), they can be vulnerable to other secondary alteration processes of petroleum. For example, disturbance of the Re-Os systematics has been shown for oils spatially associated with the main basin-bounding faults of the Viking Graben and East Shetland basin of the North Sea, United Kingdom (Finlay et al., 2010). Here, oils are suggested to have been contaminated by fault-charged mantle-fluids that contain Os with a non-radiogenic

$^{187}Os/^{188}Os$ composition. As a result, the Re-Os systematics of the oil fail to yield a generation age, but instead trace the crustal-scale fluid dynamics and migration. Such processes have also been demonstrated in the laboratory, where experiments have shown that Re and Os in aqueous phase can transfer into oil during oil-water contact under various conditions (Mahdaoui et al., 2015). Thermal cracking may also reset the Re–Os isotope systematics in petroleum and result in only the timing of thermal cracking and generation of dry gas and pyrobitumen being recorded (Lillis and Selby, 2013; Ge et al., 2016). Similarly, thermochemical sulphate reduction (TSR) of oil may also disturb or even reset the Re-Os systematics in oil. For example, the Re-Os data from oils affected by TSR in the Big Horn basin yield a Re-Os date of 9.24 ± 0.39 Ma, which is in agreement with the proposed end of TSR as a result of reservoir cooling due to the major uplift and erosion at ~10 Ma (Lillis and Selby, 2013).

The previous attempts and achievements of using Re-Os as petroleum generation geochronometer and oil-source tracer are often compromised by the lack of understanding, or intrinsic complex geological characteristics, of the petroleum systems. The source units of some petroleum systems are often debated or represent mixed sources. For example: 1) the Phosphoria system records a mixed oil generated from two separate members of the Permian Phosphoria Formation, i.e. Meade Peak Member (ca. 270 Ma) and Retort Member (ca. 265 Ma), which have an age difference of about 5 million years; 2) the Eocene Green River Formation (ca. 52 to 44 Ma) is over 3000 m (10000 ft) thick, within which many members are considered as possible oil sources (Anders et al., 1992; Ruble et al., 2001); 3) the

Canadian Oil Sands may have a mixed source, although the main contributor is highly debated (Riediger et al., 2001; Higley et al., 2009; Berbesi et al., 2012; Finlay et al., 2012; Adams et al., 2013). Inconsistent thermal maturation timings, multiple stages of oil generation and remigration are also often proposed by geological models. Both multiple sources and multiple stages, or prolonged oil generation, are likely to yield a large range of initial $^{187}Os/^{188}Os$ compositions of generated oil. Some oil samples are also altered (e.g., biodegradation, thermal cracking, and thermochemical sulphate reduction), for which the possibility and degree of disturbance to the oil Re-Os systematics has been discussed, but not yet fully understood. Further, the application of the $^{187}Os/^{188}Os$ composition as an oil-source correlation tool suffers from the lack of source rock Re-Os data in many cases. For example, the Green River Formation is over 3000 m (10000 ft) thick, but the shale samples for Re-Os dating were collected from three intervals with most of the samples spanning only 1 to 3 m (3 to 10 ft) (Cumming et al., 2014).

Here, as an examination of the usefulness of the application of the Re-Os chronometer in petroleum systems, we present the study on the simple and well-characterized Duvernay petroleum system of the Western Canada sedimentary basin (WCSB). The oil in this petroleum system is sourced from the Late Devonian Duvernay carbonaceous shale (Stoakes and Creaney, 1984; Li et al., 1998; Fowler et al., 2001). Oil generation from the Late Devonian Duvernay carbonaceous shale occurred during the Late Cretaceous to Eocene, associated with the Laramide orogeny (Deroo et al., 1977; Creaney and Allan, 1990). The generated oil migrated into the adjacent Leduc reef build-ups in the Rimbey-Meadowbrook

reef chain and Bashaw reef complex area, and then into the overlying Winterburn Group Nisku Formation (Creaney et al., 1994; Rostron, 1997). Biomarker analyses indicate that the Duvernay Formation is the only source of this petroleum system and that crude oil has experienced limited secondary alteration (Li et al., 1998; Fowler et al., 2001). We present Re-Os data from oil and source units, coupled with organic geochemistry, and discuss the geochronology and oil-source correlation of the Duvernay petroleum system and the implications for the application of the Re-Os geochronometer to petroleum systems.

2. Geological setting

The Duvernay petroleum system is located in south-central Alberta within the Western Canada sedimentary basin (Fig. 1). The Late Devonian Woodbend and Winterburn strata in the basin represent a thick and lithologically complex accumulation of shales and carbonates (Fig. 2). The two salient features of the Woodbend and Winterburn strata are the high quality source rocks and reservoirs they encompass, for example, the proliferous Duvernay Formation and the porous carbonates of the Leduc and Nisku Formations.

The Late Devonian Duvernay Formation of the Woodbend Group was deposited as marine basin-filling laminated carbonates and shales (Stoakes and Creaney, 1984; Switzer et al., 1994; Chow et al., 1995; Fowler et al., 2001; Fothergill, 2014). The Duvernay Formation correlates with the Muskwa Formation to the north, and the Horn River and Canol Formations in the Northwest Territories. The formation also overlies the platform carbonates of the Cooking Lake Formation and is in turn overlain by the basin-filling calcareous shales

of the Ireton Formation. The thickness varies depending on proximity to Devonian reef build-ups and is predominantly 20 - 60 m (60 to 200 ft) (Fig. 17, Switzer et al., 1994; Fothergill, 2014), but is up to 90 m (300 ft) within some carbonate embayments adjacent to the Killam Barrier reef-edge (Switzer et al., 1994; Fowler et al., 2001). The organic matter is concentrated in thin (typically a few mm) alternating laminae within the Duvernay Formation, which are separated by an organic-poor mudstone. Petrographically, the laminates of the Duvernay Formation contain abundant alganite macerals and some bituminite, often within an amorphous organic matrix, indicating that they represent periods of significant accumulation and preservation of oil-prone Type II marine organic matter (Fowler et al., 2001). This organic-poor mudstone becomes thinner (30 m (100 ft)) in the center of the basin (Switzer et al., 1994) to the west and reduces the separation between the two organic-rich intervals (Chow et al., 1995). The average TOC content ranges from 4 wt.% to 6 wt.%, although occasionally it is as high as 17 wt. % (Stoakes and Creaney, 1984; Chow et al., 1995; Higley et al., 2009; Dunn et al., 2012).

The numerous isolated Leduc reef build-ups (e.g., Rimbey-Leduc-Meadowbrook chain, Redwater, Bashaw) sit stratigraphically above the Cooking Lake and Beaverhill Lake Formations and partially interfinger with the Duvernay Formation (Fig. 1 - 2). The Leduc reefs are up to 275 m (900 ft) thick (Switzer et al., 1994; Chow et al., 1995; Rostron, 1997) and are highly porous and permeable (Amthor et al., 1994). A broad Leduc carbonate shelf also developed on the Cooking Lake Formation to the southeast of the basin. The Winterburn Nisku Formation is dominated by an extensive carbonate shelf in the East Shale basin to the

east of the Rimbey-Meadowbrook Leduc reef trend. Its thickness decreases from 60 m (200 ft) near the underlying Leduc reef trend to about 20 m (60 ft) in the east, with its lithology changing stratigraphically from interbedded shale and limestone to relatively pure carbonates (Switzer et al., 1994).

The hydrocarbon maturation of the majority of the Phanerozoic strata of the WCSB did not occur until the Laramide orogeny (Late Cretaceous-Eocene) and was terminated by subsequent uplift and erosion (Deroo et al., 1977; Creaney and Allan, 1990; Bustin, 1991). As a result of the eastward propagation of crustal down-warping and foreland basin development, deep burial and thermal maturation progressed from the foreland belt northeastward across the plain. Present day isomaturity maps of Jurassic and younger source rocks (Bustin, 1991) display a decrease in thermal maturity from the foredeep to the shallower part of the basin in the northeast, where significant hydrocarbon generation occurred until the Eocene. Also estimated by apatite fission track studies (Issler et al., 1999), the peak paleo-temperature and inferred maximum burial depth is estimated from apatite fission track dating to have occurred at ~ 60 Ma.

It is generally accepted that the Duvernay source rocks have generated large amounts of hydrocarbons, with volume estimations ranging up to 100 million barrels (MMbbl) per square mile (39 MMbbl per square kilometre; Jenden and Monnier, 1997). The generated oil migrated extensively both laterally and stratigraphically within the basin, including in the older Beaverhill Lake and Elk Point reservoirs. But migration of oils was largely confined to

the Cooking Lake-Leduc and Nisku Formations by the low permeability shales, dolomitic carbonates and anhydrites of the Devonian Waterways Formation, Ireton Formation and Wabamun Group (Fig. 3; Stoakes and Creaney, 1985; Creaney and Allan, 1990; Rostron, 1997) over most of central and northern Alberta. The generated oil migrated locally into the Cooking Lake platform margins and Leduc reefs (Rimbey-Meadowbrook trend and Bashaw Reefs), and the stratigraphically older Beaverhill Lake platform present in an updip position due to the regional tilt of the strata (Fig. 31, Creaney and Allan, 1992; Creaney et al., 1994; Fig. 5, Rostron, 1997. The Cooking Lake Formation also provided a migration pathway for the isolated Leduc reef as a "leaky pipeline" dependent on the local presence or absence of permeability barriers (Stoakes and Creaney, 1984). The Ireton shale serves as the seal to the Duvernay petroleum system. However, due to the absence or the thin nature of the Ireton shale in some places, oil also leaked into the overlying Nisku carbonates in the Bashaw area (Rostron, 1997; Li et al., 1998). Geochemical studies (Creaney et al., 1994; Li et al., 1998; Fowler et al., 2001) documented that the organic geochemical composition of the Duvernay Formation extracts is quite similar to most of the crude oils found within the Leduc-Nisku intervals along the Rimbey-Meadowbrook and Bashaw reef trends. In fact, geochemical oil-source correlations have demonstrated that the Leduc and Nisku oils are almost all sourced from the Duvernay Formation with probably only a very small contribution of the Nisku source to the Nisku oils (Rostron, 1997; Li et al., 1998; Harris et al., 2003).

Gas chromatography and biomarker data show that there is no or limited alteration of the Leduc and Nisku oils by biodegradation or water washing. Thermal cracking and

thermochemical sulphate reduction of crude oil occurred only in the foredeep (Evans et al., 1971; Rogers et al., 1974; Creaney and Allan, 1992; Marquez, 1994; Stasiuk, 1997; Fowler et al., 2001). The generated gas migrated up-dip and may have resulted in deasphalting, which is considered responsible for the high H_2S content of some of the stratigraphically-higher crude oil (Simpson, 1999; Fowler et al., 2001).

3. Sample preparation and analysis methodology

3.1 Oil samples

All oil samples analysed in this study were taken from the oil library of the Geological Survey of Canada - Calgary office (GSC-C). The selected oils are from both the Leduc and Nisku reservoirs in the area of the Rimbey-Meadowbrook Leduc reef trend and the Bashaw reef complex between Edmonton and Calgary (Fig. 1; Table 1). Prior to sampling, the archived oil was thoroughly mixed to remove any density separation of oil fractions that had occurred during storage.

The selected Duvernay oil samples typically have low asphaltene contents. Given that Re and Os are both predominantly present in the asphaltene fraction, and that the Re–Os isotopic composition of an asphaltene can be used as an approximation for that of the bulk oil (Selby et al., 2007), the asphaltene fractions were separated from the oils to permit more precise determinations of the Re and Os isotope compositions. The separation of asphaltene from the bulk oil follows the principles and steps outlined by Speight (2004) and Selby et al. (2007). For each gram of crude oil, 40 ml of *n*-heptane (or *n*-pentane) was added, and the mixture

was agitated on a rocker overnight (~ 20 h) at room temperature. Asphaltene is the insoluble fraction of oil in *n*-heptane (or *n*-pentane) and was isolated from the maltene-bearing *n*-heptane (or *n*-pentane) fraction by centrifuging at 3500 rpm for 15 minutes. Following the removal of the *n*-heptane (or *n*-pentane), the asphaltene was extracted using chloroform and dried at 60°C.

3.2 Source rock

In addition to the Duvernay oil asphaltene, the organic-rich Duvernay Formation was also analysed for its Re and Os abundances and Re-Os isotopic composition. The shale and carbonaceous shale samples of the Duvernay Formation were collected from the cores of wells 16-18-52-5W5, 14-29-48-6W5, 2-6-47-4W5 and 10-27-57-21W4 (Fig. 1; Table 2). The Re-Os data from these samples were compiled with previously published Re-Os data from the Duvernay Formation from wells 2-12-50-26W4, 8-35-31-25W4 and 1-28-36-3W5 (Selby et al., 2007; Finlay et al., 2012). The entire sample set covers a large geographical area of 300 × 150 km^2 (1866 × 93 sq mi). As a result, the sample set, with respect to hydrocarbon maturation, ranges from immature to highly mature with an average R_o from 0.5 % of 10-27-57-21W4 to 1.25 % of 2-6-47-4W5 and 1.45 % of 1-28-36-3W5 (Stasiuk and Fowler, 2002), and possesses highly variable TOC contents (from less than 1 wt.% to over 10 wt.%).

Depending on the thickness of the Duvernay Formation interval of each well, most samples were taken equally from each well at 0.5 m (1.6 ft; 16-18-52-5W5, ~1.5 m (4.9 ft)) and 2 m (7 ft; 14-29-48-6W5, ~8 m (26 ft); 2-6-47-4W5, ~20 m (60 ft)) or as equal as possible

depending on availability (10-27-57-21W4) in order to obtain a Re-Os data set that was representative of the entire Duvernay Formation organic-rich interval (Table 2). The collected core samples were cut and polished to remove any drilling marks and materials, and dried at 60˚C. Approximately 30 g of rock for each sample was crushed to a fine homogeneous powder using a Zr dish.

3.3 Organic geochemical analyses

All geochemical analyses were performed at the organic geochemistry laboratory of the Geological Survey of Canada - Calgary office. Pyrolysis of pulverized rock samples was carried out using Delsi Rock-Eval II and Rock-Eval 6 instruments (for measurement details see Espitalie et al., 1985; Behar et al., 2001). Biomarker analyses were performed on low- to high-maturity Duvernay sourced oil samples and source rock solvent extracts. Soluble organic matter was extracted from powdered rock samples of the Duvernay Formation for 24 hours using a chloroform:methanol (87:13) mixture. The extracts and crude oils were treated with approximately 40 volumes of *n*-pentane to precipitate the asphaltenes. The deasphalted extracts and oils were fractionated using open column chromatography. Saturates were recovered by eluting with pentane, aromatics by eluting with a 50:50 mixture of pentane-dichloromethane, and resins were recovered with methanol.

The gasoline range fractions (i-C_5-n-C_8) of the crude oils were analysed on a HP5890 Gas Chromatograph (50m HP-1 column, siloxane gum used as a fixed phase) connected to an OI Analytical 4460 Sample Concentrator. A small amount of the whole crude oil was mixed with

deactivated alumina and transferred to the Sample Concentrator. The temperature was initially held at 30 °C for 10 minutes, then increased at a rate of 1 °C/min to 45 °C and held for 25 minutes. The eluting hydrocarbons were detected using a flame ionization detector.

Saturate fractions were analysed using a Varian 3700 FID gas chromatograph equipped with a 30m DB-1 column. The temperature was programmed from 60 °C to 300 °C at a rate of 6 °C/min and then held for 30 min at 300 °C. Gas chromatography-mass spectrometry (GCMS) was performed on a VG 7070 mass spectrometer with a gas chromatograph attached directly to the ion source (30 m (100 ft) DB-5 fused silica column used for GC separation). The temperature was initially held at 100 °C for 2 min and then programmed at 40 °C/min to 180 °C and at 4 °C/min to 320 °C. After reaching 320 °C, the temperature was held for 15 min. The mass spectrometer was operated with a 70 eV ionization voltage, 100 mA filament emission current, and interface temperature of 280 °C.

3.4 Re-Os analyses

The measurement of asphaltene and shale Re and Os abundances, and Re and Os isotopic compositions were undertaken at Durham University in the Laboratory for Source Rock and Sulphide Geochronology and Geochemistry (a member of the Durham Geochemistry Centre) by isotope dilution - negative thermal ionization mass spectrometry (ID-NTIMS). The chemical processing of samples to isolate and purify Re and Os and the measurement by NTIMS follow previously published protocols (Selby and Creaser, 2001; Selby and Creaser, 2003; Selby, 2007; Rooney et al., 2011; Cumming et al., 2014). In brief, a known amount of

shale (~0.1 to 1 g) and asphaltene (~0.15 g), together with a known amount of mixed spike containing ^{185}Re and ^{190}Os, was digested by a Cr^{VI}-H_2SO_4 solution for shale and inverse *aqua-regia* (2:1 16 N HNO_3 and 12 N HCl, 9 ml) for asphaltene in carius tubes at 220˚C for 48 hrs. For some of the samples, the amount of asphaltene digested is limited by the amount of asphaltene obtained from the oil. The Os was extracted from the digested samples using chloroform, and then back-extracted from the chloroform using 9N HBr, before being purified by micro-distillation. Following the Os extraction, for shale samples, a 1 ml aliquot of the Cr^{VI}-H_2SO_4 solution was evaporated to dryness, and the Re fraction was isolated using sodium hydroxide-acetone solvent extraction, before being further purified by anion chromatography. For the asphaltene samples, the *aqua-regia* was evaporated to dryness and the Re fraction purified by anion chromatography. The purified Re and Os were loaded onto Ni and Pt filaments, respectively (Selby et al., 2007). Isotopic measurements were conducted using negative thermal ionization mass spectrometry (Creaser et al., 1991; Völkening et al., 1991) on a Thermo Scientific TRITON mass spectrometer via ion-counting, using a secondary electron multiplier in peak-hopping mode for Os, and static Faraday collection for Re. In-house solution standards are 0.1611 ± 0.0004 (1SD, n = 126) for DROsS (Durham Romil Osmium Solution) and 0.5989 ± 0.0019 (1SD, n = 116) for Re Standard solution, which are in agreement with those previously reported (Luguet et al., 2008; Nowell et al., 2008; Cumming et al., 2012).

The measured difference in $^{185}Re/^{187}Re$ values for the Re standard solution and the accepted $^{185}Re/^{187}Re$ value (0.5974; Gramlich et al., 1973) is used for mass fractionation correction of

the Re sample data. All Re and Os data are oxide and blank corrected. The total procedural blanks of inverse *aqua-regia* are ~ 1.6 pg Re and 0.04 ~ 0.1 pg Os, with an average $^{187}Os/^{188}Os$ of 0.24 ± 0.06 (n = 3). The total procedural blanks of Cr^{VI}-H_2SO_4 digested samples are ~12 pg Re and 0.02 ~ 0.05 pg Os, with average $^{187}Os/^{188}Os$ of 0.22 ± 0.06 (n = 4). All uncertainties are determined by error propagation of uncertainties in Re and Os mass spectrometer measurements, weighing, blank abundances and isotopic compositions, spike calibrations, and the reproducibility of standard Re and Os isotopic values. The Re-Os data are regressed using *Isoplot* (V. 4.15; Ludwig, 2012) with the ^{187}Re decay constant value of 1.666×10^{-11} yr^{-1} (Smoliar et al., 1996). This decay constant is also used for the calculation of the $^{187}Os/^{188}Os$ values of the shales at the time of deposition and oil generation, and the $^{187}Os/^{188}Os$ values of the asphaltene at oil generation.

4. Results

4.1 Asphaltene Re-Os data

The asphaltene contents and Re-Os data of asphaltene fractions of this study, and one analysis reported by Selby et al. (2007), are shown in Table 3. The asphaltene abundance of the sampled Duvernay oils is between 0.23 and 14.10 %. Coupled with low asphaltene contents are low Re and Os abundances of the asphaltene fractions, which range from 0.01 to 3.78 ppb Re, 0.6 to 41.2 ppt Os, and 0.2 to 14.9 ppt ^{192}Os, respectively. Oil samples possessing asphaltene abundances up to ~4 % show a positive correlation between asphaltene abundance and both Re and ^{192}Os (representing unradiogenic common Os) abundances (Fig. 4).

However, this relationship does not hold for the remaining oil samples that possess higher asphaltene content (6 – 8% and ~15%) (Fig. 4). Though, in general, oil samples possessing higher asphaltene contents have higher Re and ^{192}Os abundances (Fig. 4). The low Re and Os abundances, coupled with the limited asphaltene obtained for Re-Os analyses, make precise measurements very challenging. The majority of the samples possess blank corrections between 0.3 – 3.3 % for Re and 1 - 10 % for Os (L02221 to L02155), although a few samples (e.g., L01810) have significant Re and Os corrections (40% Re and 50% Os). The samples with such large blank corrections also possess large uncertainties. Samples with uncertainty greater than 55% for ^{187}Re/^{188}Os and 30% for ^{187}Os/^{188}Os are excluded when determining a Re-Os age and are not further discussed. The ^{187}Re/^{188}Os and ^{187}Os/^{188}Os values of the asphaltene used for isochron construction range between ~67 and 642, and 0.78 and 1.62, respectively (Table 3). Together with the sample of Selby et al. (2007), the Re-Os data yield a Model 3 date (a model assuming variable initial ^{187}Os/^{188}Os values for samples) of 66 ± 31 Ma (n = 14, MSWD = 6.7, initial ^{187}Os/^{188}Os = 0.77 ± 0.20; Fig. 5-6).

4.2 Re-Os data for organic-rich intervals of the Duvernay Formation

The Re and Os abundances and isotope data for the Duvernay source rock analysed from four wells of this study, plus those of previous studies (Selby et al., 2007; Finlay et al., 2012) are reported in Table 2. The Re and Os abundances for the Duvernay source range between 0.3 and 68 ppb Re, and 29 and 1965 ppt Os. The more Re- and Os- (^{192}Os) enriched sectors of the Duvernay source rock broadly correlate with intervals possessing higher TOC abundances

(Fig. 7). The $^{187}Re/^{188}Os$ and $^{187}Os/^{188}Os$ ratios range from ~26 to 229 and 0.49 to 1.74, respectively.

We have calculated the $^{187}Os/^{188}Os$ composition of the Duvernay Formation at the time of deposition (initial $^{187}Os/^{188}Os$ (Os_i)), a value that is equal to the $^{187}Os/^{188}Os$ composition of the contemporaneous seawater (Cohen et al., 1999). The biostratigraphy of the Duvernay Formation shows mainly the occurrence of *Palmatolepis punctata* and *Palmatolepis hassi* (Ziegler and Sandberg, 1990; McLean and Klapper, 1998; Whalen et al., 2000), which constrains the deposition age to be between 376.5 and 380.5 Ma based on the Geologic Timescale 2016 (Ogg et al., 2016). The Os_i values for the shale samples are calculated at 378.5 Ma. The Os_i values range between 0.28 and 0.38, with a Tukey's Biweight mean (which weighs the points according to their scatter from the iteratively-determined mean) of 0.33 ± 0.01 (n = 33; Table 2; Fig. 7).

4.3 Organic geochemistry of crude oil and rock samples

The Leduc-Nisku oils, although typically not altered by biodegradation or water-washing, show a wide range of thermal maturities, often resulting in variable geochemical characteristics, and have been classified into three main subgroups that include low-, normal- and high-maturity oils (Fig. 8). These hydrocarbons, however, represent a continuum that can be easily documented by examples of crude oils having maturities intermediate between the three subgroups. The following overview pertains mostly to the saturated hydrocarbon fractions of the three oil sub-types resulting from differences in thermal maturity (although

similar compositional variations have also been observed in the aromatic fraction by Li et al., 1998). Main compositional characteristics of the organic extracts and a brief oil-source correlation summary are also included.

The least mature Duvernay sourced oils are found in the Leduc and Nisku reservoirs in central eastern Alberta, to the south of the Bashaw Reef Complex. These oils are characterized by low saturate/aromatic ratio, relatively high amounts of NSO compounds, and relatively high sulphur content (1-1.5%). Gasoline range chromatograms show high *n*-heptane and methylcyclohexane peaks (n-C_7>MCH), with toluene often in relatively high abundance compared to most other oils sourced from Devonian units. Saturate fraction gas chromatograms display smooth distribution of n-alkanes with slight odd carbon number preference of n-C_{15} – n-C_{19} members, high abundance of acyclic isoprenoids relative to n-alkanes, low pristane/phytane ratio (~0.80), and evident biomarkers.

Biomarker distributions show a smooth homohopane (C_{31}–C_{35}) profile, gammacerane present in moderate abundance, low amounts of rearranged hopanes (i.e. low Ts/Tm (abbreviations are annotated in Table 4)) and tricyclic and tetracyclic terpanes relative to 17α(H)-hopane, C_{24} tetracyclic terpane in similar abundance to the two C_{26} tricyclic terpane isomers, relatively low abundances of diasteranes and short-chain steranes relative to regular steranes, and a C_{27}-C_{29} sterane abundance of C_{29}>C_{27}>>C_{28}.

The majority of the oil pools along the Rimbey-Meadowbrook trend contain normal-maturity Duvernay Formation sourced oils. Compared to low maturity oils, these oils have higher

saturate/aromatic ratios, lower amounts of NSO compounds and lower sulphur content.

Gasoline range and saturate fraction gas chromatograms are generally similar in both subgroups, with notably lower abundance of acyclic isoprenoids relative to *n*-alkanes and some biomarker peaks in more mature oils.

Furthermore, normal maturity oils contain higher amounts of rearranged hopanes (hence higher Ts/Tm), tricyclic and tetracyclic terpanes compared to the hopanes, C_{24} tetracyclic terpane slightly greater than C_{26} tricyclic terpanes, very low abundance of gammacerane, higher abundances of diasteranes and short-chain steranes, and similar carbon number distributions for C_{27}-C_{29} regular steranes.

The high maturity Duvernay sourced oils and condensates tend to be those from the deeper, westerly pools with higher API gravities, typically occurring in the southern half of the province. Compared with other Duvernay sourced oils, they have higher saturate/aromatic ratios and lower amounts of NSO compounds. Although these oil can still be correlated to oils from the other two subgroups by gasoline range and saturate fraction gas chromatography data (though maturation parameters indicate higher maturity and *n*-alkane profiles shows more bias towards lower molecular weight homologues), they are quite easily distinguished based on the lack of or very low abundance of biomarkers such as hopanes and steranes. The visible peaks in the m/z 191 and m/z 217 chromatograms are typically tricyclic terpanes, Ts and C_{29} Ts, low molecular weight steranes, high diasteranes and secosteranes with an extremely low abundance of regular steranes.

The analysis of the available Duvernay Formation Rock-Eval and TOC data reveals this unit contains on average 4% of organic carbon with the maximum of up to 18%, enough to be considered as petroleum source rock (Fig. 9). In thermally unaltered and low maturity samples hydrogen index values are up to 700 indicating presence of Type I-II organic matter (Fig. 10). Although the extract yields are commonly above 100 mg total extract/g TOC, and greater than 200 mg in about 1/3 of the samples, the hydrocarbon yields typically range from 20 to 80, with few above 100 mg HC/g TOC (Fig. 11), which overall is indicative of good source rock potential. The extract data show no evidence of staining, as the proportions of hydrocarbons in extracts are normally less than 50% (Fig. 11)

The gas chromatograms of the extracts typically show *n*-alkane distributions dominated by shorter-chain homologues (n-C_{15} – n-C_{20}) (Fig. 12). Amounts of acyclic isoprenoids relative to *n*-alkanes are high in low maturity samples and decrease with increasing level of thermal maturity. Pristane/phytane ratios generally fall within 1-2 range. Saturate biomarker signatures of mature extracts show relatively high amounts of rearranged hopanes and steranes with no homohopane prominence, low tricyclic terpanes, Ts/Tm ratios greater than 1.0 and high diasteranes. The geochemical character of the Duvernay Formation extracts is different compared with other Devonian units in Alberta that show petroleum source potential (Creaney et al., 1994).

Correlation of biomarker data indicates that organic geochemical composition of the thermally mature Duvernay Formation solvent extracts is very similar to that of the normal

maturity crude oils found within the Leduc-Nisku intervals along the Rimbey-Meadowbrook and Bashaw reef trends (Fig. 13). Both extracts and oils have higher concentrations of n-C_{17} and n-C_{18} relative to pristane and phytane, and average Pr/Ph ratios typical of suboxic depositional conditions (average of 1.66 and 1.38 for extracts and oils, respectively), show predominace of Ts over Tm, a smooth homohopane distributions without C_{34} or C_{35} prominence, relatively high amounts of rearranged steranes and common C_{29}>C_{27}>C_{28} regular sterane profile (Fig. 14). Their overall similarity and genetic relationship are further supported by regional geology and stratigraphic data. As noted previously (Stoakes and Creaney, 1984 and 1985; Li et al., 1998), the Duvernay strata is the only prolific source rock interval that occurs in close stratigraphic proximity to the Leduc reefs to account for the large volume of oil found in the Leduc reservoirs and, therefore, is considered as the main source of these oils.

5 Discussion

5.1 Petroleum Re-Os geochronology

The burial history of the Albertan part of the Western Canada sedimentary basin has been constructed with isopach maps of the stratigraphic formations (Deroo et al., 1977; Higley et al., 2005) and basin modelling based on oil generation kinetic parameters (Higley et al., 2009; Berbesi et al., 2012), which indicate that the Duvernay Formation started to generate oil at the onset of the Laramide orogeny, i.e., around 80 Ma. According to the most recent basin models (Higley et al., 2009; Berbesi et al., 2012), the majority of the oil generation occurred between

~70 and ~58 Ma, with a significant kerogen transformation ratio (>50%) of the Duvernay Formation being reached by ~65 Ma, and cessation of oil generation by ~50 Ma due to basin uplift.

The Re-Os oil date (66 ± 31 Ma, n = 14, MSWD = 6.7; Fig. 5) obtained from the asphaltene fraction Re-Os data, although possessing a large uncertainty, is in good agreement with the timing of peak oil generation indicated by geological modelling. It is clear that the Re-Os oil date does not in any way reflect the age of the Duvernay Formation which is the source of Duvernay oil (this study; Deroo et al., 1977; Creaney et al., 1994; Li et al., 1998), i.e. early Late Devonian (Frasnian; this study; Switzer et al., 1994).

The Re-Os date is derived from a Model 3 *Isoplot* solution (Ludwig, 2012), which means that the scatter about the best-fit line (the isochron), as represented by the MSWD (mean square weighted deviation) value, is due to a combination of the calculated uncertainties for the $^{187}Re/^{188}Os$ and $^{187}Os/^{188}Os$ values and the error correlation, rho, plus an unknown but normally distributed variation in the initial $^{187}Os/^{188}Os$ (Os_i) values. The uncertainty in the Re-Os date spans the entire duration of oil generation; however, it is not to say that the uncertainty of the Re-Os oil age will necessarily be indicative of the duration of oil generation. The MSWD value (6.7) being >1 for the best-fit line suggests that the scatter about the best-fit line is controlled by geological factors rather than purely analytical uncertainties (Ludwig, 2012). The ~25% uncertainty in the calculated initial Os_i composition (0.77 ± 0.20, Fig. 5; 0.55 to 1.06, Fig. 6-7, Table 3) suggests that the primary factor causing the degree of scatter about the best-fit of the Re-Os data is the variation in the Os_i

composition of the oil sample set. The geological cause of the variation in the Os_i composition of the Duvernay oil is discussed below (Section 5.2). Although the Re-Os date for the generation of Duvernay sourced oil is not as precise or detailed as the basin modeling, Re-Os oil geochronology works directly on oil samples and can provide comparison and validation of the oil generation dating by traditional geological tools in a fast and economic way.

5.2 $^{187}Os/^{188}Os$ composition oil-source fingerprinting

Previous research has proposed that at the time of oil generation, an oil inherits the $^{187}Os/^{188}Os$ composition of the source rock (Selby et al., 2005; Selby and Creaser, 2005; Finlay et al., 2011; Finlay et al., 2012; Rooney et al., 2012; Cumming et al., 2014). Compared to the previous attempts to apply the $^{187}Os/^{188}Os$ composition of oil and source rock as an absolute fingerprinting tool, the Duvernay petroleum system benefits from its single source and simple thermal maturation history. Here we have also obtained source rock Re-Os data for the entire stratigraphic interval of the Duvernay Formation over a large geographical area, to reflect the source rock Re-Os systematics as completely as possible.

Although not sampled for shale geochronology, for which an intensive sampling strategy should be adopted to minimize the time span of samples and limit any possible variability in the initial $^{187}Os/^{188}Os$ value, the Duvernay Formation source rock $^{187}Re/^{188}Os$ and $^{187}Os/^{188}Os$ data do display a positive correlation and yield Late Devonian ages (well 10-27-57-21W4, 10 samples over 55 m (180 ft), Re-Os data yield 371 ± 22 Ma, initial $^{187}Os/^{188}Os$ = 0.33 ± 0.03,

MWSD = 66; well 16-18-52-5W5, 4 samples over 1.5 m (5 ft), Re-Os data yield 372 ± 13 Ma, initial $^{187}Os/^{188}Os$ = 0.33 ± 0.02, MWSD = 0.06; well 14-29-48-6W5, 5 samples over 6 m (20 ft), Re-Os data yield 377 ± 62 Ma, initial $^{187}Os/^{188}Os$ = 0.33 ± 0.14, MWSD = 84; well 2-6-47-4W5, 11 samples over 20 m (60 ft), Re-Os data yield 360 ± 15 Ma, initial $^{187}Os/^{188}Os$ = 0.38 ± 0.03, MWSD = 33; Fig. 15). Further, given that the samples analysed are considered broadly the same geological age (Fig. 15A-D), we also present a calculated Re-Os date from all 33 samples. The Re-Os data for the entire sample set yield 373.9 ± 9.4 Ma, with an initial $^{187}Os/^{188}Os$ = 0.34 ± 0.02 (MWSD = 82; Fig. 15E). The uncertainties of the ages are due to the variation in the Os_i compositions (Table 2, Fig. 7), which range from 0.28 to 0.38, and the limited spread in the $^{187}Re/^{188}Os$ values (≤160 $^{187}Re/^{188}Os$ units; Table 2 and Fig. 15). Although the Re-Os dates are imprecise, they are nominally in agreement with the Late Devonian biostratigraphic age constraints for the Duvernay Formation.

To evaluate the relationship between the oil Os_i (Fig. 5-6) composition and that of the source rock at the time of oil generation, the Duvernay source rock $^{187}Os/^{188}Os$ compositions at 66 Ma (Os_g) are calculated from the present day Re-Os data of shale samples. The Duvernay Formation Os_g values range between 0.46 and 1.48 (n = 33 from 7 wells), and yield a Tukey's Biweight mean of 0.83 ± 0.01 (Tables 2; Fig. 7). The oil Os_i defined by the Re-Os isochron is 0.77 ± 0.20 (Fig. 5). The range of the individual oil Os_i values is between 0.55 and 1.06 (Fig. 6). Seventy-nine percent (26 of the 33 samples) of the shale Os_g values are indistinguishable from the individual oil Os_i values (Fig. 7). The evidence from organic geochemistry of the Duvernay petroleum system indicates that the Duvernay Formation is the principal source of

the Duvernay oils. As such, the similarity of the oil Os_i and Duvernay Formation Os_g compositions supports the research hypothesis that the oil Os_i is inherited from the source at the time of oil generation, and therefore oil Os_i and source Os_g values can be used as an oil-source correlation tool.

An oil-source correlation based on Os isotopic compositions is supported by the comparison of Duvernay petroleum system oil Os_i and shale Os_g values with the Os_g values of other oil-prone strata of the WCSB. As most of the oil-prone Phanerozoic strata of the WCSB underwent hydrocarbon maturation during the Laramide orogeny (Bustin, 1991; Higley et al., 2009), the Os_g values of these strata with available Re-Os data are also calculated at 66 Ma (Table 5; Fig. 16). However, in contrast to the Duvernay petroleum system oil Os_i and shale Os_g values, the Os_g values of the other oil-prone strata of the WCSB are distinctly more radiogenic (e.g., the Gordondale Member of the Early Jurassic Fernie Formation (Selby et al., 2007; Finlay et al., 2012) has Os_g values are between 0.99 and 4.06, with an arithmetic average and median of 2.04 and 1.51; the Devonian-Mississippian Exshaw Formation (Creaser et al., 2002; Selby and Creaser, 2003; Finlay et al., 2012) has Os_g values between 1.35 and 4.96 with an arithmetic average and median of 2.22 and 1.93; the Middle Devonian Keg River Formation (Miller, 2004) laminates of La Crete basin have Os_g values between 1.34 and 4.25, with an arithmetic average and median of 1.97 and 1.64). The exceptions to the latter are the Os_g values of the kukersites of the Ordovician Yeoman Formation, Williston basin (Miller, 2004), which are between 0.46 and 5.36, with an arithmetic average and median of 1.21 and 0.84. The Ordovician kukersites Os_g values are only slightly more

radiogenic than the Duvernay petroleum system oil Os_i and shale Os_g values, however, it can be spatially (eastern Montana - western North Dakota and southern Saskatchewan versus south-central Alberta) excluded as the source of the Duvernay oils (Creaney et al., 1994). The Os_i and Os_g data, coupled with the organic geochemistry, support the Duvernay Formation as the key source unit of the Duvernay oils, and eliminate the other oil-prone Phanerozoic strata of the WCSB as sources of the Duvernay petroleum system.

Although we show a strong similarity between the Os_i and Os_g values for the Duvernay oils and Formation, we observe that the Os_g values in the cores studied here and the oil Os_i values show some heterogeneity (Fig. 5-7). If we consider each core separately, the average Os_g values are nominally skewed either positively or negatively to the overall average of the Os_g values (0.83 ± 0.01; Fig. 15; Table 2). For example, the Tukey's Biweight mean of Os_g values for core 10-27-57-21W4 is 0.72 ± 0.12, for core 12-6-47-4W5 it is 0.90 ± 0.17, for core 14-29-48-6W5 it is 0.96 ± 0.07, and for core 16-18-52-5W5 the arithmetic average is 0.66 ± 0.22 (2 SD; weighted by assigned errors only as Tukey's Biweight mean is not suitable for its small sample set).

The change in the $^{187}Os/^{188}Os$ composition from the time of source rock deposition to the time of oil generation is dependent on the duration between deposition and hydrocarbon maturation, and the Re abundance in the source rock. Variability in the Re abundance stratigraphically, plus the variations in local seawater $^{187}Os/^{188}Os$ values inherited by the organic-rich rock, such as that shown by the Duvernay Formation (Table 2), results in heterogeneous $^{187}Os/^{188}Os$ values with ^{187}Re decay (Table 2; Fig. 7 and 15).

As such, if the heterogeneous Os_g is inherited by the generated oils then our data propose that the oil generated from different geographical sections of the same source interval could lead to reservoired oils possessing variable Os_i. The primary migration and expulsion of oil within and from a source rock, which is often broadly contemporaneous with or occurs shortly after oil generation, could result in the homogenization of the $^{187}Os/^{188}Os$ values of oil. Although, the extent of homogenization should depend on many factors (e.g. permeability of source rock and carrier beds). In the case of the Duvernay Formation, the oils generated could possess a range in Os_i from 0.46 to 1.48, which is similar to what is observed in this study with the Duvernay oil Os_i values, ranging from ~0.55 to 1.06 (Table 2 - 3; Fig. 5-7). The variability in the Os_i values makes it analytically challenging to obtain a precise isochron (66 ± 31 Ma). The determined Os_i by the isochron (0.77 ± 0.20) demonstrates the variability in the Os_i values. As such, if we select the samples with the most similar Os_i (e.g., L00864, L00873, L02196, L02220 and L02221), a Re-Os date of 56 ± 18 Ma is obtained, with an Os_i of 0.74 ± 0.10 and MSWD of 0.33. The Re-Os date still coincides with the timing of hydrocarbon maturation associated with the Laramide orogeny, but is almost 50% more precise.

5.3 Oil Re-Os systematics: effect of oil-water contact

Although we propose that the Re and Os in oil are inherited from a source rock, based on the similarity of the Os_i values of the asphaltene fractions and the Os_g values of the Duvernay Formation shales, an experimental study has suggested that Re and Os in oil can also be derived from contact with formation water (Mahdaoui et al., 2015). The outcome of this

experimental work has recently been questioned (Wu et al., 2016), particularly based on the several orders of magnitude higher Re and Os concentrations applied in the experiment (0.001 to 100 μg/g Re and 1 or 10 ng/g Os) compared with the known low Re and Os concentrations of basin fluids (typically 4 pg/g Re and 70 fg/g Os; Mahdaoui et al. (2015)) in most oil-bearing sedimentary basins. Nevertheless, we consider here if the oil Re-Os systematics may have been affected by or originated from the basin fluids in the Leduc reef trend and Nisku Formation. There are very limited Re and Os data for basin fluids available at present and if we also use the typical formation water Re and Os abundances of 4 pg/g and 70 fg/g, respectively (Mahdaoui et al., 2015), approximately 2 to 40 g of such water contains equivalent Re and Os in 1 gram of Duvernay oil (Table 3). Such oil-water contact ratios are highly probable in a petroleum system (Magoon and Dow, 1991). As such, based on the experiments of Mahdaoui et al. (2015), the Re-Os systematics of the low Re- and Os-bearing Duvernay oils could have been easily influenced by oil-basin fluid interaction and reflect the Os isotopic composition of the basin fluid. However, comparing the $^{87}Sr/^{86}Sr$ values of Late Devonian seawater of 0.7080 - 0.7084 (Burke et al., 1982) with the $^{87}Sr/^{86}Sr$ values of calcites and dolomite cements (0.7088 - 0.7108) of the Leduc Rimbey-Meadowbrook reef trend (Stasiuk, 1997), and that of the current brines of Nisku Formation (0.70998-0.71206) and Leduc Formation (0.70872 - 0.70981) (Connolly et al., 1990), it is suggested that the basin fluids could have interacted with the radiogenic crystalline basement and/or Proterozoic - Lower Cambrian clastics (Mountjoy et al., 1999). As such, the $^{187}Os/^{188}Os$ compositions of the basin fluids should be radiogenic, considering the high $^{187}Os/^{188}Os$ initial values of these

strata (Kendall et al., 2009) and the long duration between the depositional age and the timing of oil generation. For example, the Neoproterozoic Old Fort Point Formation has an initial $^{187}Os/^{188}Os$ value of ~ 0.62 at deposition of 608 Ma (Kendall et al., 2004), which will have radiogenically evolved to $^{187}Os/^{188}Os$ values >3 at 66 Ma (Table 5). Further, considering a lower end member $^{187}Re/^{188}Os$ value of 100 for the crystalline basement and/or Proterozoic - Lower Cambrian strata and an initial $^{187}Os/^{188}Os$ composition of 0.5, the decay of ^{187}Re over ≥450 million years will yield a $^{187}Os/^{188}Os$ value > 1.25 at the time that the Duvernay Formation underwent thermal maturation. The deep-originated basinal fluids interacted with old strata and therefore should have very radiogenic $^{187}Os/^{188}Os$ compositions. The minimum estimate of the $^{187}Os/^{188}Os$ composition of the basinal fluid is more radiogenic than the present day (0.78 – 1.62) and initial (0.77 ± 0.20) values of the Duvernay oil . As a result, we propose that the predominant Re and Os budget and the isotope systematics of the Duvernay oil represent that inherited from the Duvernay Formation source rocks during thermal maturation, and not Re and Os sequestered from basin fluids. And therefore the Os_i values are inherited at the time of oil generation, and are ultimately linked to the source rock rather than formation fluids.

Conclusions

Despite the uncertainty of Re-Os asphaltene isotopic data of the Duvernay oils, which is a direct result of the low Re and Os abundance, the limited range in the $^{187}Re/^{188}Os$ values, and the variation in the Os_i values of the asphaltene, the Re-Os age derived from asphaltene

fractions agrees well with the widely accepted burial and hydrocarbon maturation models that indicate that hydrocarbon generation occurred long after source rock deposition during the Laramide orogeny. Further, the similarity between the oil Os_i values and the Duvernay Formation Os_g values, and their difference to other oil-prone strata of the WCSB, coupled with the evidence from organic geochemistry, confirms the Duvernay Formation as the source and the ability of the $^{187}Os/^{188}Os$ composition as an oil-source correlation tool. Moreover, the Duvernay-sourced oils Re-Os systematics are not considered to have been significantly affected by basin fluids, indicating that the majority of Re-Os systematics in the oil (asphaltene fraction) is inherited from its source.

We suggest that one of the major factors controlling the precision of a Re-Os date for oil relates to the variation in Os_i values inherited by an oil from its source at the time of oil generation, which is independent of the radiogenic evolution of the $^{187}Os/^{188}Os$ composition in the source rock. For example, the relative homogeneous $^{187}Os/^{188}Os$ of Duvernay Formation at the time of deposition (Os_i = ~0.33) evolved with time to a heterogeneous $^{187}Os/^{188}Os$ composition due to the variable Re abundance in the source unit. The heterogeneous $^{187}Os/^{188}Os$ composition of the Duvernay Formation was inherited by the generated oil. Our data suggests that basin wide thorough homogenization of the Os isotopic composition did not occur during migration.

The agreement of the Re-Os asphaltene date, coupled with the similarity of the Os_i and Os_g values of the asphaltene and Duvernay Formation, respectively, of the relatively simple

Duvernay petroleum system of Western Canada sedimentary basin with basin models and organic geochemistry confirms the potential ability of the Re-Os oil (asphaltene fraction) systematics to record the timing of hydrocarbon generation, and to serve as an effective oil-source correlation tool.

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Figure Captions

Figure 1. Map of the study area showing the location of the oil and Duvernay Formation shale samples used in this study (modified from Fowler et al., 2001 with the permission of Canadian Society of Petroleum Geologists). The oils are from the Devonian Leduc and Nisku formations. The majority of the shale samples are located in the mature zone (16-18-52-5W5, 14-29-48-6W5, 2-12-50-26W4 and 2-6-47-4W5) with some from the immature zone (10-27-57-21W4) and the highly mature zone (1-28-36-3W5 and 8-35-31-25W4). Solvent extracts of well 10-27-57-21W4 (low maturity) and 13-14-35-25W4 (high maturity) are used for organic geochemical analysis.

Figure 2. Devonian stratigraphy of the study area, south-central Alberta (information is collectively from the Table of Formations, Alberta by Energy Resources Conservation Board based in Calgary, Canada, Aitken and Norris, 2014, Li et al., 1998 and Fowler et al., 2001). Late Devonian Woodbend Group Duvernay Formation interfingers with Leduc reefs. Also shown here are the organic-rich Devonian-Mississippian Exshaw Formation and the equivalent of Middle Devonian Keg River Formation, the Winnipegosis Formation.

Figure 3. Migration of the Duvernay Formation generated oil (Modified from Switzer et al., 1994 and Rostron, 1997). Following oil generation the Duvernay oil migrated to adjacent Leduc reefs with the underlying Cooking Lake carbonates working as a “pipe line”. In the Bashaw reef area, the differential compaction of the Ireton aquitard due to the underlying Leduc reefs leads to the thinning or absence of the aquitard. This facilitates the breach of Duvernay sourced oil from Leduc reefs to the overlying Nisku formation.

Figure 4. Asphaltene content versus asphaltene Re (ppb) and unradiogenic Os (ppt, represented by ^{192}Os) abundance of the Duvernay oils.

Figure 5. ^{187}Re/^{188}Os vs ^{187}Os/^{188}Os plot for the asphaltene fractions of the Duvernay oil. See text for discussion.

Figure 6. The individually calculated Os_i values of asphaltene samples of the Duvernay oil. determined at 66 Ma.

Figure 7. Duvernay Formation total organic carbon (TOC) (Stasiuk and Fowler, 2002), Re and unradiogenic Os (^{192}Os as representative) abundances, and calculated ^{187}Os/^{188}Os compositions at the time of deposition (Os_i, 378.5 Ma) and oil generation (Os_g, 66 Ma). The depth of TOC does not correlate exactly with the Re and Os data (Fowler et al., 2003). Broadly the Re and Os abundances positively co-vary with the TOC content. The Os_i values show a limited (0.28 to 0.38) range, and yield a weighted average of ~0.33, reflecting a relatively nonradiogenic ^{187}Os/^{188}Os composition for contemporaneous seawater. The Os_g values of the Duvernay Formation calculated at 66 Ma are between 0.46 and 1.48, and yield a Tukey's Biweight mean of 0.833 ± 0.009.

Figure 8a. Representative gas chromatograms showing distributions of: 1) gasoline range hydrocarbons and 2) normal alkanes and isoprenoids in low (sample L02221), medium (sample L02155) and high (sample L02177) maturity Duvernay-sourced crude oils. The high maturity oils have very low asphaltene content and hence were not selected for Re-Os analysis. For peak annotations see Table 4.

Figure 8b. Representative gas chromatograms showing distributions of: 3) terpanes and 4) steranes in low (sample L02221), medium (sample L02155) and high (sample L02177) maturity Duvernay-sourced crude oils.

Figure 9. Histogram showing distribution of total organic carbon (TOC) content in the Duvernay Formation samples.

Figure 10. A cross-plot of Rock-Eval hydrogen (HI) and oxygen (OI) indices for the 4 wells from which Re-Os data were collected during present study. The data indicate presence of Type I-II organic matter in most of the samples; notice a systematic decrease in HI values corresponding to an increase in thermal maturity levels from 10-27-57-21W4 and 16-18-52-5W5 through 14-29-48-6W5 to 2-6-47-4W5.

Figure 11. Extract yield data showing relative maturity and source rock potential of the Duvernay Formation samples (criteria by Powell, 1978).

Figure 12. Saturate fraction gas chromatograms showing distributions of 1) normal alkanes and isoprenoids, 2) terpanes and 3) steranes in low and high maturity Duvernay Formation organic extracts. For peak annotations see Table 4.

Figure 13. Oil-source correlation of a Nisku crude oil from Wood River field (within the Bashaw reef complex, 10-28-42-23W4, sample L02203) and Duvernay Formation organic

extract from 13-14-35-25W4 showing strong similarities in the distributions of terpane and sterane biomarkers. For peak annotations see Table 4.

Figure 14. Ternary diagram showing relative normalized abundance of C_{27}:C_{28}:C_{29} regular steranes based on αββ isomers in the analyzed oil samples. Data for 29 Duvernay extracts (from GSC-C database) are shown for comparison.

Figure 15. $^{187}Re/^{188}Os$ vs $^{187}Os/^{188}Os$ plots for the Duvernay formation. A) core 10-27-57-21W4; B) core 16-18-52-5W5; C) core 14-29-48-6W5; D) core 2-6-47-4W5; and E) all data including samples from references (Selby et al., 2007; Finlay et al., 2012). See text for discussion.

Figure 16. Comparison of the currently available $^{187}Os/^{188}Os$ compositions at 66 Ma for the WCSB Phanerozoic organic-rich intervals shown with the Duvernay shale Os_g and oil Os_i values.

Table 1. Well locations and reservoir formations of the oil samples of this study

Table 2. Re-Os data synopsis for the Duvernay Formation shale samples from multiple locations (wells) and their Os_i (378.5 Ma) and Os_g (66 Ma) values

Table 3. Re-Os data synopsis for asphaltene fractions of the oil samples from the Duvernay petroleum system and their Os_i (66 Ma) values.

Table 4. The compounds represented by the peak annotations in the geochemistry figures

(Fig. 8, 12 and 13)

Table 5. Re-Os data and calculated $^{187}Os/^{188}Os$ composition at 66 Ma for other oil prone strata of the WCSB.

Table 1 Well locations and reservoir formations of the oil samples of this study

Sample ID	Well ID	Latitude	Longitude	Depth (m)	Depth (ft)	Formation
L00864	14-35-48-27W4	53.19	-113.85	1920.0-1923.9	6299.2-6312.0	Leduc
L00873	10-9-48-27W4	53.13	-113.89	1781.9-1783.4	5846.1-5851.0	Nisku
L01558	01-3-37-20W4	52.15	-112.77	1604.8-1622.1	5265.1-5321.9	Nisku
L01810	12-11-49-12W5	53.21	-115.66	2997-3069	9833-10069	Nisku
L01827	11-14-42-2W5	52.62	-114.18	2373-2387	7785-7831	Leduc
L01831	8-35-48-12W5	53.18	-115.64	3059-3086	10036-10125	Nisku
L02037	16-16-41-2W5	52.54	-114.22	2431-2442	7976-8012	Nisku
L02083	12-13-41-23W4	52.53	-113.19	1742-1743	5715-5719	Nisku
L02155	10-31-39-23W4	52.4	-113.3	1781.3	5844.2	Nisku
L02196	9-11-33-26W4	51.81	-113.59	2497.8	8194.9	Leduc
L02197	9-14-33-26W4	51.83	-113.58	2202	7224	Nisku
L02203	10-28-42-23W4	52.65	-113.25	1720	5643	Nisku
L02220	11-35-36-20W4	52.14	-112.76	1606.3	5270.0	Nisku
L02221	7-23-36-20W4	52.1	-112.75	1569.7-1617.9	5149.9-5308.1	Leduc/Nisku
L02225	15-22-38-20W4	52.29	-112.78	1563.6-1645.9	5129.9-5399.9	Leduc
L02226	12-10-38-20W4	52.25	-112.79	1569.7-1624.6	5149.9-5330.1	Leduc
L01822	2-27-37-20W4	52.2	-112.78	1599-1603	5246-5259	Leduc

Table 2 Re-Os data synopsis for the Duvernay Formation shale samples from multiple well locations

Sample	Depth (m)	Depth (ft)	Re (ppb)	±	Os (ppt)	±	^{192}Os (ppt)	±	^{187}Re/^{188}Os	±	^{187}Os/^{188}Os	±	rho [a]	Os_i [b]	±	Os_g [c]	±
10-27-57-21W4			*53.96°N*		*113.03°W*												
DS18-12	1106	3629	5.84	0.02	631.2	2.2	244.5	1.0	47.5	0.2	0.633	0.004	0.590	0.33	0.01	0.58	0.01
DS20-12	1125	3691	14.37	0.04	1245.3	4.8	478.7	2.2	59.7	0.3	0.693	0.005	0.604	0.32	0.01	0.63	0.01
DS22-12	1128	3701	13.23	0.04	1299.2	5.1	500.7	2.4	52.5	0.3	0.673	0.005	0.600	0.34	0.01	0.62	0.01
DS24-12	1140	3740	12.36	0.04	1121.1	4.4	431.3	2.1	57.0	0.3	0.687	0.005	0.600	0.33	0.01	0.62	0.01
DS26-12	1144	3753	16.86	0.05	947.8	3.8	355.5	1.5	94.4	0.5	0.899	0.005	0.598	0.30	0.01	0.80	0.01
DS28-12	1146	3760	14.68	0.04	1392.9	5.2	537.7	2.4	54.3	0.3	0.659	0.004	0.598	0.32	0.01	0.60	0.01
DS30-12	1151	3776	26.33	0.07	1965.2	6.9	751.1	2.9	69.8	0.3	0.742	0.004	0.584	0.30	0.01	0.67	0.01
DS32-12	1154	3786	24.15	0.06	1537.6	5.6	579.5	2.2	82.9	0.4	0.859	0.005	0.582	0.34	0.01	0.77	0.01
DS34-12	1160	3806	10.91	0.03	397.1	2.0	142.7	0.7	152.0	0.9	1.267	0.009	0.608	0.31	0.02	1.10	0.01
DS36-12	1161	3809	25.43	0.07	1188.0	4.7	436.5	1.7	115.9	0.5	1.074	0.006	0.582	0.34	0.01	0.95	0.01
16-18-52-5W5			*53.5°N*		*114.72°W*												
DS1-12	2343.0	7687.0	0.31	0.01	58.8	2.1	23.2	1.9	26.3	2.2	0.488	0.055	0.687	0.32	0.07	0.46	0.06
DS3-12	2343.5	7689.6	1.25	0.01	242.5	2.5	95.6	1.9	26.0	0.5	0.491	0.014	0.683	0.33	0.02	0.46	0.02
DS5-12	2344.0	7690.3	2.03	0.01	178.1	1.3	68.5	0.8	58.9	0.7	0.693	0.012	0.671	0.32	0.02	0.63	0.01
DS7-12	2344.5	7691.9	4.03	0.01	248.0	1.3	93.4	0.7	85.8	0.7	0.862	0.009	0.657	0.32	0.01	0.77	0.01
14-29-48-6W5			*53.18°N*		*114.85°W*												
DS9-12	2646.6	8683.1	48.46	0.14	1370.7	8.2	474.1	2.9	203.3	1.4	1.611	0.014	0.646	0.32	0.02	1.39	0.02
DS11-12	2648.6	8689.6	8.75	0.03	362.4	1.9	132.7	0.8	131.3	0.9	1.107	0.010	0.622	0.28	0.02	0.96	0.01
DS13-12	2650.6	8696.2	10.82	0.03	482.0	2.3	176.0	0.9	122.3	0.7	1.129	0.008	0.609	0.36	0.01	0.99	0.01
DS15-12	2652.6	8702.8	21.94	0.06	1827.3	6.6	700.6	2.8	62.3	0.3	0.716	0.004	0.576	0.32	0.01	0.65	0.01
DS17-12	2654.6	8709.3	36.35	0.09	1726.5	8.7	636.2	3.4	113.7	0.7	1.050	0.009	0.589	0.33	0.01	0.93	0.01

2-6-47-4W5			*53.02°N*		*114.57°W*												
DS37-12	2629	8625	0.53	0.01	29.0	1.0	10.8	0.9	98.2	8.0	0.942	0.107	0.701	0.32	0.16	0.83	0.12
DS38-12	2631	8632	0.40	0.01	30.7	0.7	11.7	0.6	68.2	3.6	0.792	0.056	0.677	0.36	0.08	0.72	0.06
DS39-12	2633	8638	1.36	0.01	62.2	0.9	22.8	0.6	118.5	3.2	1.110	0.042	0.694	0.36	0.06	0.98	0.05
DS40-12	2635	8645	41.52	0.14	1218.3	4.2	423.1	1.0	195.2	0.8	1.575	0.005	0.403	0.34	0.01	1.36	0.01
DS41-12	2637	8652	67.56	0.22	1721.5	7.0	587.5	1.7	228.8	1.0	1.736	0.007	0.476	0.29	0.01	1.48	0.01
DS42-12	2639	8658	13.07	0.04	900.4	2.9	340.3	1.1	76.4	0.3	0.836	0.004	0.486	0.35	0.01	0.75	0.01
DS43-12	2641	8665	11.27	0.04	746.9	3.7	282.3	1.8	79.4	0.6	0.835	0.007	0.620	0.33	0.01	0.75	0.01
DS44-12	2643	8671	5.99	0.02	430.7	1.8	163.0	0.8	73.1	0.4	0.823	0.006	0.576	0.36	0.01	0.74	0.01
DS45-12	2645	8678	8.74	0.03	598.6	3.5	226.1	1.8	76.9	0.7	0.841	0.010	0.648	0.35	0.01	0.76	0.01
DS46-12	2647	8684	28.75	0.09	1707.7	5.7	639.0	2.0	89.5	0.4	0.919	0.004	0.482	0.35	0.01	0.82	0.01
DS47-12	2649	8691	41.79	0.14	1774.0	8.1	649.0	3.1	128.1	0.7	1.112	0.008	0.581	0.30	0.01	0.97	0.01
2-12-50-26W4 [d]			*53.30°N*		*113.67°W*												
DS44-03	1751.6	5746.7	7.41	0.03	383	2	140	1	104.4	0.6	1.037	0.006	0.800	0.38	0.01	0.92	0.01
8-35-31-25W4 [d]			*51.70°N*		*113.43°W*												
DS69-03	2340.9	7680.1	11.55	0.04	569	2	208	1	109.5	0.5	1.036	0.004	0.800	0.34	0.01	0.92	0.01
1-28-36-3W5 [e]			*52.20°N*		*114.60°W*												
DS45-03-1-4	3013.1	9885.5	8.10	0.03	219.6	1.2	75.3	0.4	214.1	1.2	1.689	0.012	0.542	0.34	0.02	1.45	0.01

[a] rho: error correlation value between $^{187}Re/^{188}Os$ and $^{187}Os/^{188}Os$;

[b] Os_i = initial $^{187}Os/^{188}Os$ ratio of shales during deposition which is calculated at 378.5 Ma (see text for discussion);

[c] Os_g = $^{187}Os/^{188}Os$ ratio of shales at oil generation calculated at 66 Ma (see text for discussion);

[d] Data from Finlay et al. (2012);

[e] Data from Selby et al. (2007).

Table 3 Re-Os data synopsis for asphaltene fractions of the oil samples from the Duvernay petroleum system

Sample ID	Asphaltene (%)	Re (ppb)	±	Os (ppt)	±	^{192}Os (ppt)	±	^{187}Re/^{188}Os	±	^{187}Os/^{188}Os	±	rho [a]	Os_i [b]	±
L00864	1.61%	0.12	0.07	9.4	0.5	3.6	0.5	66.8	37.1	0.775	0.111	0.211	0.70	0.15
L00864rpt	1.16%	0.13	0.05	8.2	0.4	3.0	0.3	84.9	31.1	0.980	0.118	0.270	0.89	0.15
L00864rpt2	1.81%	0.12	0.04	7.4	0.4	2.7	0.4	84.1	32.0	1.036	0.164	0.291	0.94	0.20
L00873	5.65%	0.42	0.03	8.8	0.3	3.3	0.2	256.1	26.1	1.006	0.086	0.621	0.72	0.12
L01558	3.60%	1.42	0.03	20.4	0.5	7.6	0.4	371.1	20.0	0.970	0.063	0.698	0.56	0.09
L02083	2.46%	1.26	0.04	11.2	0.4	3.9	0.3	641.6	51.8	1.521	0.131	0.814	0.82	0.19
L02155	1.45%	0.32	0.04	3.6	0.3	1.2	0.3	511.4	131.1	1.623	0.373	0.844	1.06	0.52
L02196	0.25%	0.04	0.00	0.6	0.1	0.2	0.1	351.4	81.0	1.105	0.275	0.918	0.72	0.36
L02203	1.39%	0.26	0.00	4.6	0.5	1.7	0.5	298.8	85.4	0.878	0.273	0.919	0.55	0.37
L02220	3.00%	1.41	0.03	19.1	0.5	7.0	0.4	400.4	22.9	1.105	0.073	0.712	0.66	0.10
L02221 [c]	14.10%	3.78	0.04	41.2	1.6	14.9	1.2	504.7	42.3	1.206	0.139	0.712	0.65	0.19
L02225 [c]	6.75%	0.51	0.04	6.4	0.4	2.3	0.3	442.6	74.5	1.285	0.219	0.760	0.80	0.30
L02226 [c]	5.84%	1.11	0.05	12.0	0.6	4.2	0.5	524.6	68.6	1.475	0.216	0.766	0.90	0.29
L01822 [d]	14.50%	2.03	0.07	25.3	1.6	8.8	0.6	457.1	72.3	1.488	0.236	0.500	0.99	0.32
Data not applied on isochron														
L01810	0.35%	0.04	0.11	1.0	0.8			218.2	705.6	0.63	1.33	0.54		
L01831	0.52%	0.01	0.11	23.3	1.0			2.0	25.1	1.11	0.12	0.01		
L02196rpt	0.78%	0.02	0.05	0.6	0.4			210.4	562.6	0.52	0.92	0.51		
L02037	0.28%	0.05	0.17	-	-			-	-	-	-	-		
L02197	0.27%	0.07	0.17	-	-			-	-	-	-	-		

[a] rho: error correlation value between ^{187}Re/^{188}Os and ^{187}Os/^{188}Os;

[b] Osi = initial ^{187}Os/^{188}Os ratios of oil at the timing of generation which are calculated at 66 Ma (see text for discussion);

[c] these samples are asphaltene fractions precipitated by *n*-pentane at the GSC;

[d] data from Selby et al. (2007).

Table 4 The compounds represented by the peak annotations in the geochemistry figures (Fig. 8-10)

Peak	**Compound**	**Peak**	**Compound**
n-C_5	C_5 *n*-alkane	C_{26}TT	C_{26} tricyclic terpane
n-C_6	C_6 *n*-alkane	Ts	18α(H),22,29,30-trisnorhopane (Ts)
MCYC5	methylcyclopentane	Tm	17α(H),22,29,30-trisnorhopane (Tm)
n-C_7	C_7 *n*-alkane	C_{29}	17α(H),21β(H)-30-norhopane
MCYC6	methylcyclohexane	C_{29}Ts	18α(H),21β(H)-norneohopane
n-C_8	C_8 *n*-alkane	C_{30}H	17α(H),21β(H)-hopane
Pr	pristane	G	gammacerane
Ph	phytane	C_{34}	17α(H),21β(H)-tetrakishomohopanes
n-C_{15}	C_{15} *n*-alkane	C_{35}	17α(H),21β(H)-pentakishomohopanes
n-C_{20}	C_{20} *n*-alkane	C_{21}P	pregnane
n-C_{25}	C_{25} *n*-alkane	dC_{27}S20S	diacholestane 20S
C_{23}TT	C_{23} tricyclic terpane	dC_{29}S20S	diastigmastane 20S
C_{24}TeT	C_{24} tetracyclic terpane	C_{29}S20R	5α(H),14β(H),17β(H)-stigmastane 20R

Table 5 Re-Os data and calculated $^{187}Os/^{188}Os$ composition at 66 Ma (Os_g) for other oil prone strata of the WCSB

Lower Jurassic Gordondale Member

Finlay et al. (2012)

wells	depth (m)	depth (ft)	Re (ppb)	Os (ppt)	$^{187}Re/^{188}Os$	$^{187}Os/^{188}Os$	Os_g [a]
07-31-79-10W6	1557.0	5108.3	574.9	6.34	550.4	2.116	1.51
14-24-80-7W6	1183.5	3882.9	369.2	1.86	1674	5.900	4.06
14-24-80-7W6	1190.4	3905.5	229.9	1.83	885.8	3.655	2.68
8-26-69-7W6	2271.4	7452.1	195.3	6.13	178.6	1.372	1.18
10-17-84-22W5	2361.0	7746.1	196.2	1.12	1352	4.748	3.26
6-29-85-11W5	1148.7	3768.7	402.6	3.05	933.2	3.716	2.69

Selby et al. (2007) [b] *well 6-32-78-5W6 1218.0-1218.6 m (3996.1-3998.0 ft)*

sample	$^{187}Re/^{188}Os$	$^{187}Os/^{188}Os$	Os_g [a]
TS22b.DD.GB	765.25	2.925	2.08
TS23c.DD.GB	432.40	1.914	1.44
TS24b.DD.GB	291.71	1.405	1.08
TS25b.DD.GB	424.18	1.926	1.46
TS26a.DD.GB	182.21	1.186	0.99

Devonian-Mississippian Exshaw Formation

Selby and Creaser (2005) 51°05'29''N 115°09'29''W

within ~4 m (13 ft) lateral interval and 10 cm (0.4 in) of vertical stratigraphy

sample	Re (ppb)	Os (ppt)	$^{187}Re/^{188}Os$	$^{187}Os/^{188}Os$	Os_g [a]
DS53	15.6	367.8	253.5	1.948	1.67
DS54	17.5	375.9	283.8	2.141	1.83
DS55A	15.5	341.4	273.7	2.073	1.77
DS55B	16.4	363.5	277.3	2.100	1.80
DS55C	21.2	426.3	306.1	2.265	1.93
DS56	16.4	493.3	190.4	1.570	1.36
DS57	36.0	595.8	391.8	2.787	2.36
DS58A	15.2	458.7	189.3	1.563	1.35
DS58B	16.6	490.6	194.5	1.597	1.38

Selby et al. (2003); Creaser et al. (2002); Finlay et al. (2012)

wells	depth (m)	depth (ft)	Re (ppb)	Os (ppb)	$^{187}Re/^{188}Os$	$^{187}Os/^{188}Os$	Os_g [a]
3-19-80-23W5	1752.0	5748.0	70.0	0.69	487.1	3.517	2.98
3-19-80-23W5	1753.0	5751.3	88.8	1.62	350.7	2.656	2.27
3-19-80-23W5	1754.0	5754.6	22.8	0.50	279.8	2.195	1.89
3-19-80-23W5	1756.5	5762.8	31.7	0.48	441.1	3.220	2.73
13-18-80-23W5	1748.8	5737.5	47.9	0.64	518.3	3.546	2.98
13-18-80-23W5	1750.9	5744.4	68.9	1.11	408.9	2.987	2.54
8-29-78-01W6	2099.1	6886.8	128.6	1.70	538.4	3.806	3.21

8-29-78-01W6	2099.2	6887.1	20.0	0.69	166.1	1.535	1.35
14-22-80-02W6	2058.3	6753.0	31.8	0.67	287.5	2.095	1.78
6-19-78-25W5	2006.2	6582.0	52.0	0.84	405.1	2.885	2.44
4-23-72-10W6	3570.4	11713.9	31.3	0.29	916.8	5.968	4.96
4-23-72-10W6	3567.7	11705.1	42.3	0.80	327.8	2.389	2.03

Middle Devonian Keg River Formation, La Crete Basin

Miller (2004)

sample ID	Re (ppb)	Os (ppb)	$^{187}Re/^{188}Os$	$^{187}Os/^{188}Os$	Os_g [a]
CM-KR-1	50.8	1.24	238	1.775	1.51
CM-KR-2	34.6	0.94	208	1.568	1.34
CM-KR-3	32.8	0.92	204	1.575	1.35
CM-KR-4	20.7	0.47	259	1.920	1.64
CM-KR-5	57.2	0.95	386	2.733	2.31
CM-KR-6	224.2	2.32	757	5.087	4.25
CM-KR-7	26.7	0.67	232	1.754	1.50
CM-KR-I-D	42.4	0.78	336	2.392	2.02
CM-KR-I-E	37.9	0.78	294	2.124	1.80

Ordovician Kukersites, Williston Basin

Miller (2004)

sample ID	Re (ppb)	Os (ppb)	$^{187}Re/^{188}Os$	$^{187}Os/^{188}Os$	Os_g [a]
CM-13	2.60	0.085	146	1.608	1.45
CM-14	0.94	0.033	136	1.390	1.24
CM-15	1.01	0.031	155	1.742	1.57
CM-16 (rpt.)	0.60	0.069	44.7	0.776	0.73
CM-17 (rpt.)	0.27	0.074	18.5	0.476	0.46
CM-19 (rpt.)	0.40	0.063	31.4	0.514	0.48
CM-22 (rpt.)	0.40	0.067	30.2	0.580	0.55
CM-IV-A-iii.iv.v (rpt.)	0.22	0.019	54.7	0.899	0.84
CM-IV-A-v,vi,vii (rpt.)	0.24	0.023	50	0.894	0.84
CM-IV-A-xiii,xiv,xv	0.23	0.015	75.6	0.943	0.86
CM-IV-A-xvi	0.14	0.030	22.4	0.593	0.57
CM-24	6.58	0.070	752	6.190	5.36
CM-28	0.50	0.048	49.5	0.845	0.79

Neoproterozoic Old Fort Point Formation, Jasper, Alberta

Kendall et al. (2004) *52.96°N* *118.48°W*

samples	Re (ppb)	Os (ppt)	$^{187}Re/^{188}Os$	$^{187}Os/^{188}Os$	Os_g [a]	Os_i [c]
BK-01-014B	15.43	249.4	524.7	5.954	5.38	0.61
BK-01-015A	6.36	164.7	262.9	3.299	3.01	0.62
BK-01-015B	12.53	218.9	463.7	5.352	4.84	0.63
BK-01-015C	8.54	192.0	318.7	3.859	3.51	0.61
BK-01-015D	7.28	144.8	379.6	4.475	4.06	0.61

[a] Os_g = $^{187}Os/^{188}Os$ ratios of strata at the approximate timing of thermal maturation of most of the strata (adopting the Re-Os age of 66 Ma);
[b] No Re and Os abundance data available;
[c] Os_i = initial $^{187}Os/^{188}Os$ ratios of Old Fort Point Formation calculated at its Re-Os depositional age of 608 Ma.

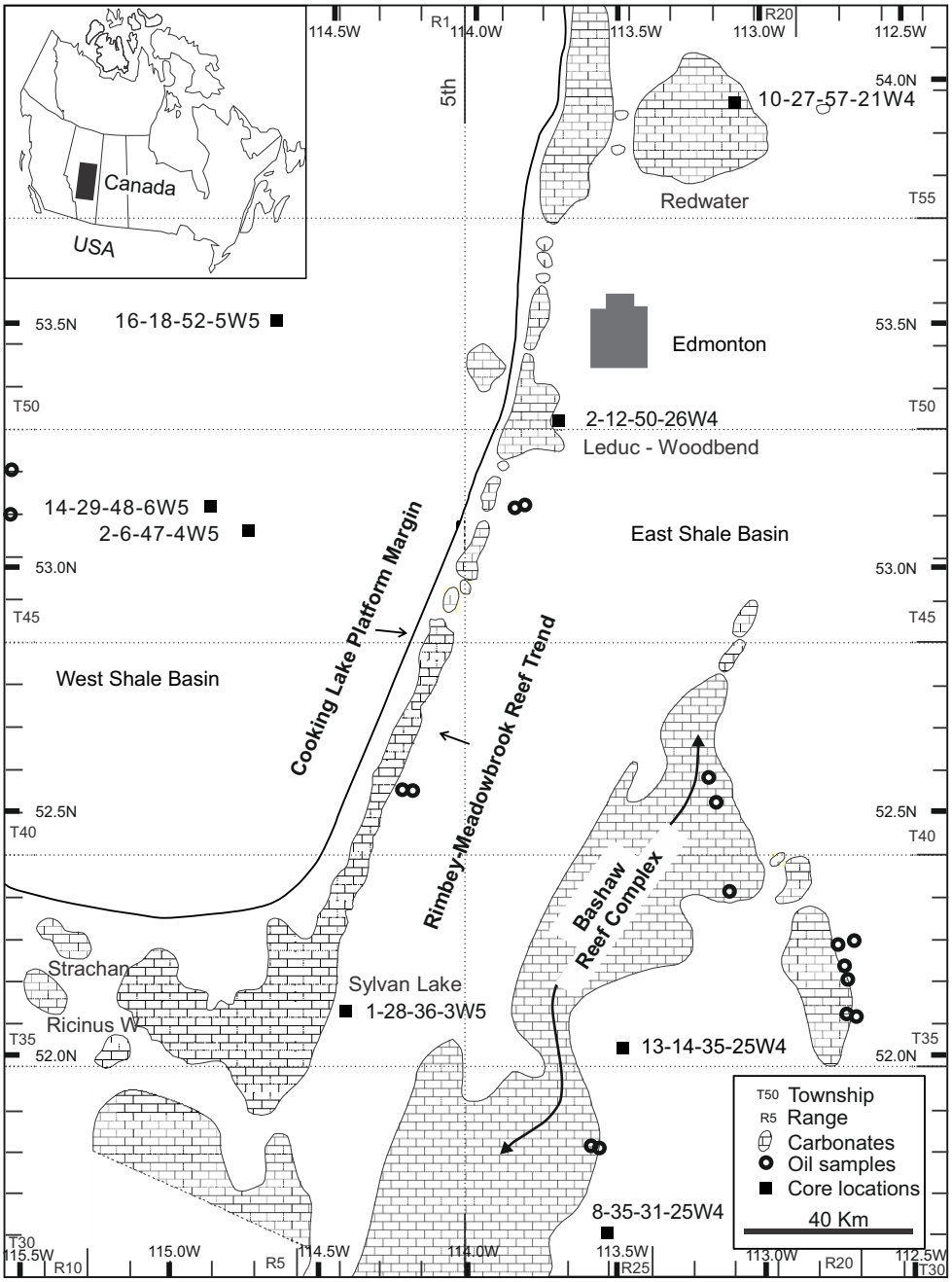
Canada
USA
114.5W
R1
114.0W
113.5W
R20
113.0W
112.5W
5th
54.0N
10-27-57-21W4
Redwater
T55
53.5N
16-18-52-5W5
Edmonton
T50
2-12-50-26W4
Leduc - Woodbend
14-29-48-6W5
2-6-47-4W5
East Shale Basin
53.0N
T45
Cooking Lake Platform Margin
Rimbey-Meadowbrook Reef Trend
West Shale Basin
52.5N
T40
Bashaw Reef Complex
Strachan
Sylvan Lake
1-28-36-3W5
Ricinus W
T35
52.0N
13-14-35-25W4
T50 Township
R5 Range
Carbonates
Oil samples
Core locations
40 Km
8-35-31-25W4
T30
115.5W
R10
115.0W
R5
114.5W
114.0W
113.5W
R25
113.0W
R20
112.5W
T30

Figure 1

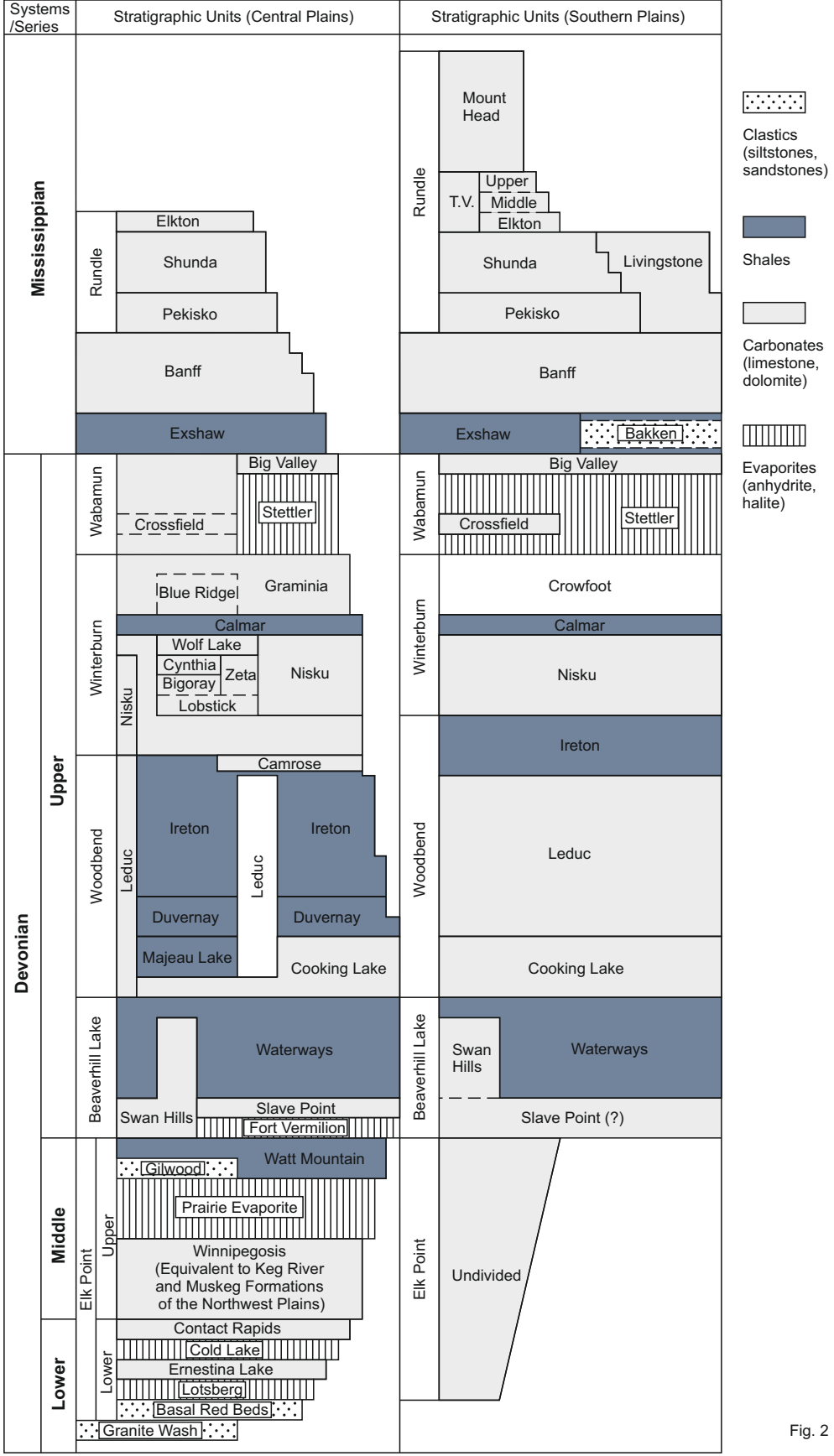
Systems /Series
Stratigraphic Units (Central Plains)
Stratigraphic Units (Southern Plains)
Mississippian
Rundle
Elkton
Shunda
Pekisko
Banff
Exshaw
Mount Head
T.V.
Upper
Middle
Elkton
Shunda
Livingstone
Pekisko
Banff
Exshaw
Bakken
Devonian
Upper
Wabamun
Big Valley
Stettler
Crossfield
Winterburn
Blue Ridge
Graminia
Calmar
Wolf Lake
Cynthia
Bigoray
Zeta
Nisku
Lobstick
Woodbend
Camrose
Ireton
Leduc
Duvernay
Majeau Lake
Cooking Lake
Beaverhill Lake
Waterways
Swan Hills
Slave Point
Fort Vermilion
Crowfoot
Slave Point (?)
Middle
Elk Point
Watt Mountain
Gilwood
Prairie Evaporite
Winnipegosis
(Equivalent to Keg River
and Muskeg Formations
of the Northwest Plains)
Undivided
Lower
Contact Rapids
Cold Lake
Ernestina Lake
Lotsberg
Basal Red Beds
Granite Wash
Clastics
(siltstones,
sandstones)
Shales
Carbonates
(limestone,
dolomite)
Evaporites
(anhydrite,
halite)

Fig. 2

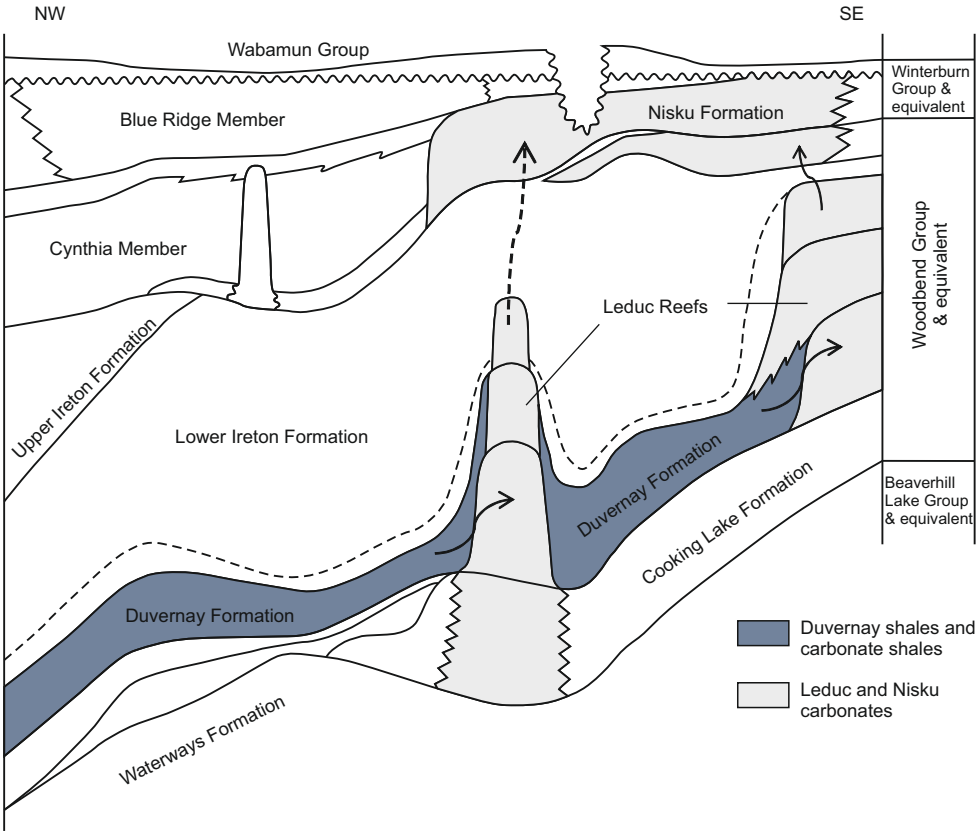
NW
SE
Wabamun Group
Blue Ridge Member
Nisku Formation
Winterburn Group & equivalent
Cynthia Member
Leduc Reefs
Woodbend Group & equivalent
Upper Ireton Formation
Lower Ireton Formation
Duvernay Formation
Cooking Lake Formation
Beaverhill Lake Group & equivalent
Duvernay Formation
Waterways Formation
Duvernay shales and carbonate shales
Leduc and Nisku carbonates

Figure 3

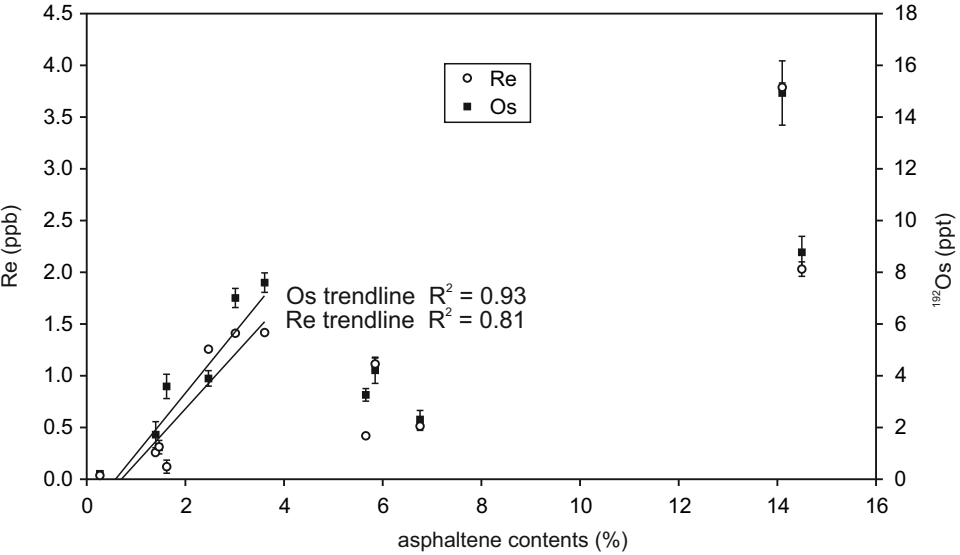
Re
Os
Os trendline R^2 = 0.93
Re trendline R^2 = 0.81
Re (ppb)
^{192}Os (ppt)
asphaltene contents (%)
0.0
0.5
1.0
1.5
2.0
2.5
3.0
3.5
4.0
4.5
0
2
4
6
8
10
12
14
16
18

Figure 4

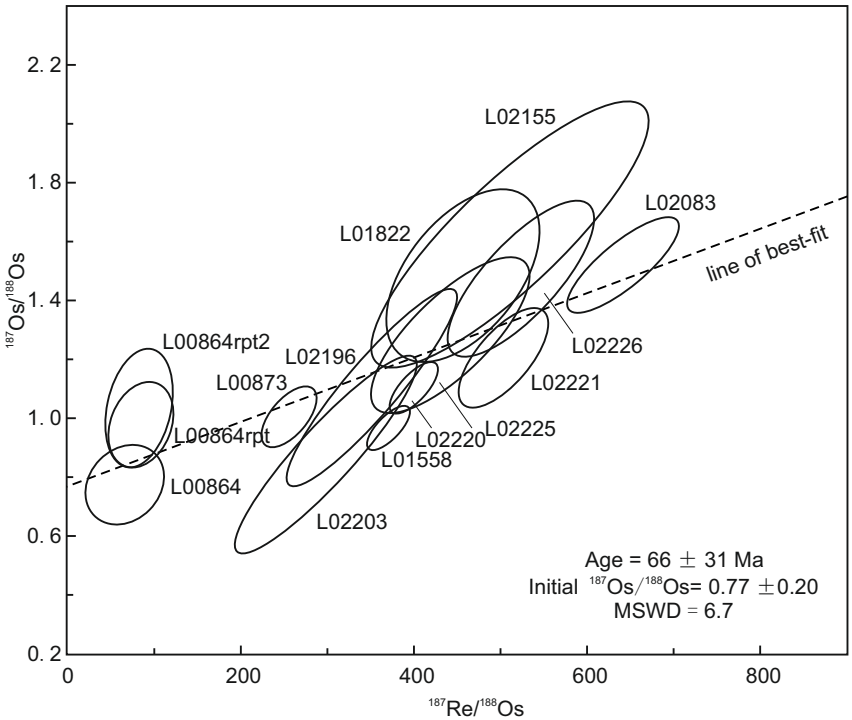
2.2
1.8
1.4
1.0
0.6
0.2
0
200
400
600
800
$^{187}Os/^{188}Os$
$^{187}Re/^{188}Os$
L02155
L02083
L01822
line of best-fit
L00864rpt2
L02196
L02226
L00873
L02221
L00864rpt
L02220
L02225
L01558
L00864
L02203
Age = 66 ± 31 Ma
Initial $^{187}Os/^{188}Os$= 0.77 ±0.20
MSWD = 6.7

Figure 5

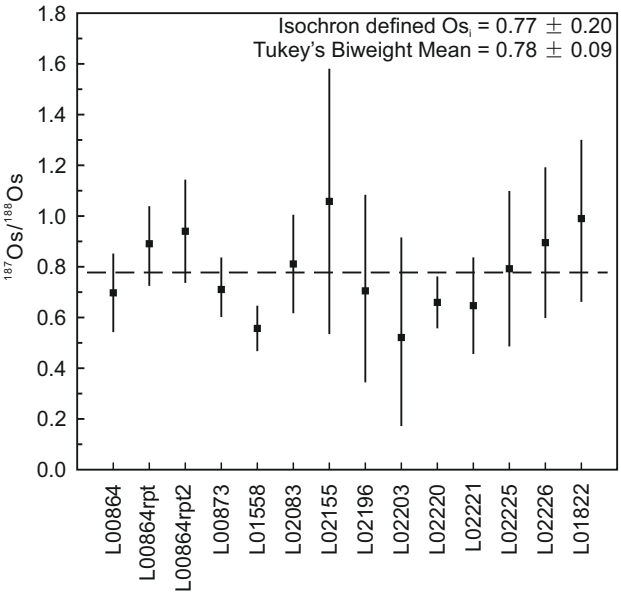
Isochron defined Os$_i$ = 0.77 ± 0.20
Tukey's Biweight Mean = 0.78 ± 0.09
^{187}Os/^{188}Os
1.8
1.6
1.4
1.2
1.0
0.8
0.6
0.4
0.2
0.0
L00864
L00864rpt
L00864rpt2
L00873
L01558
L02083
L02155
L02196
L02203
L02220
L02221
L02225
L02226
L01822

Figure 6

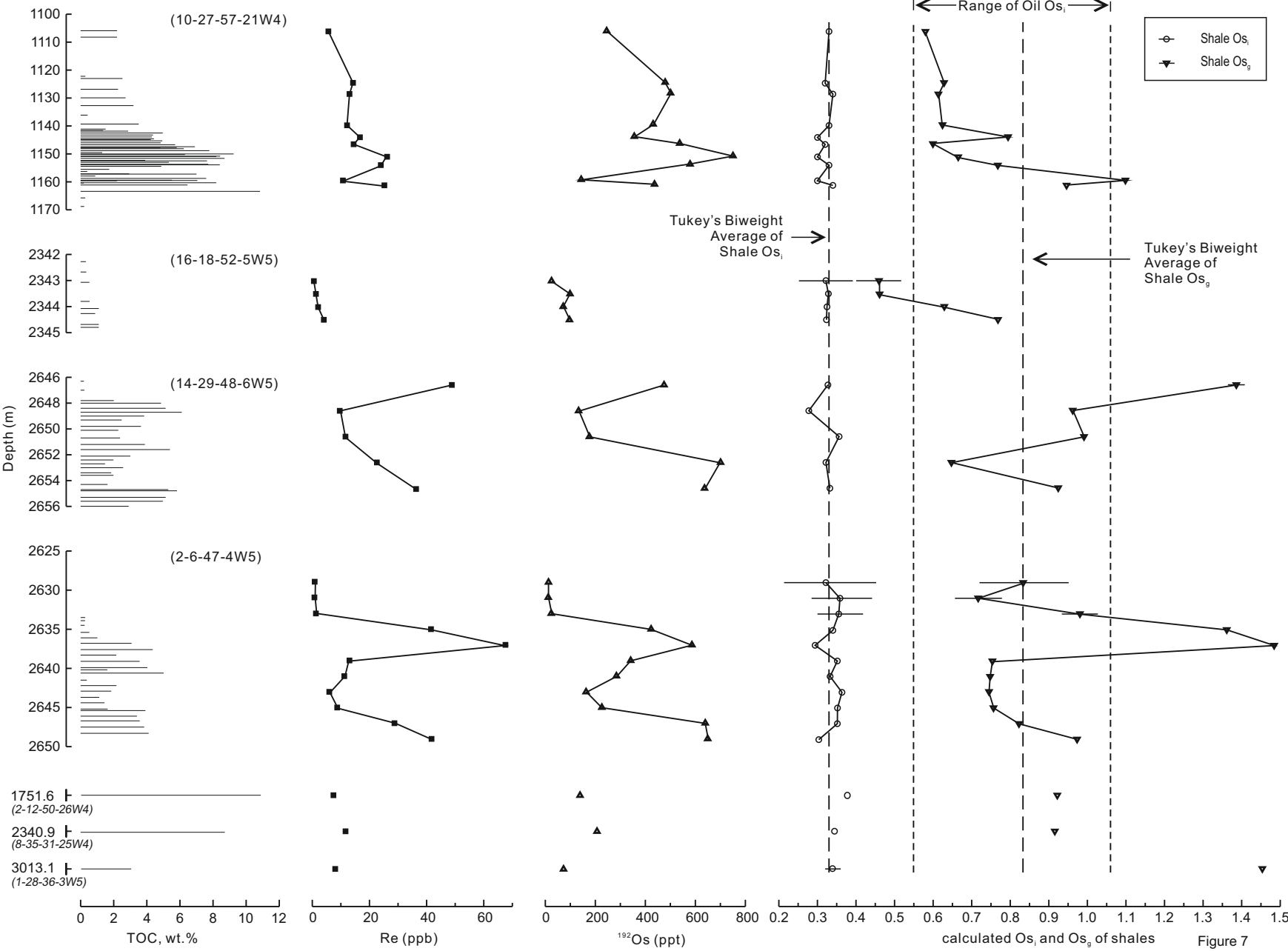
(10-27-57-21W4)
(16-18-52-5W5)
(14-29-48-6W5)
(2-6-47-4W5)
Depth (m)
1751.6
(2-12-50-26W4)
2340.9
(8-35-31-25W4)
3013.1
(1-28-36-3W5)
Range of Oil Os$_i$
Shale Os$_i$
Shale Os$_g$
Tukey's Biweight Average of Shale Os$_i$
Tukey's Biweight Average of Shale Os$_g$
TOC, wt.%
Re (ppb)
^{192}Os (ppt)
calculated Os$_i$ and Os$_g$ of shales

Figure 7

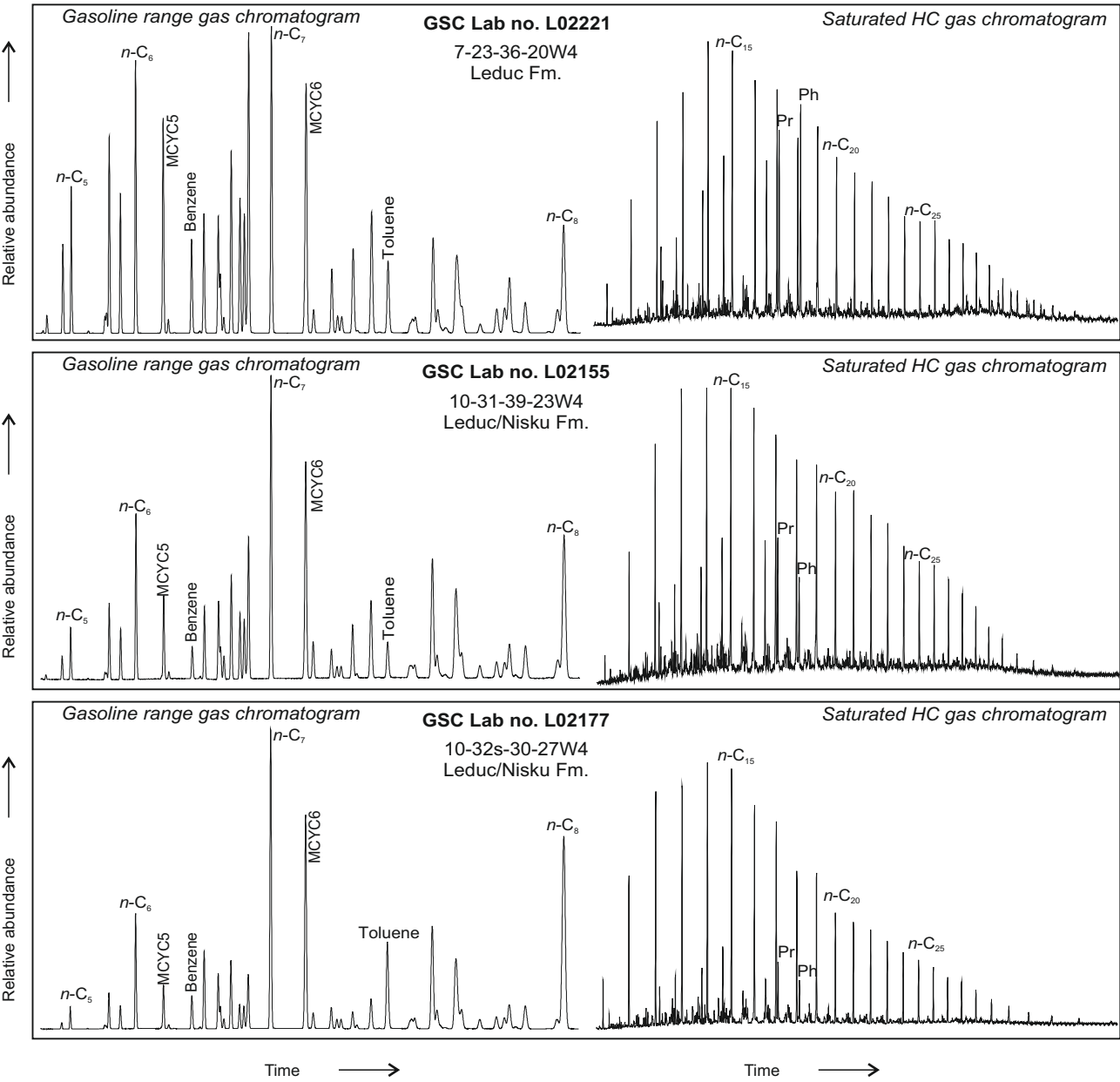
Gasoline range gas chromatogram
GSC Lab no. L02221
7-23-36-20W4
Leduc Fm.
Saturated HC gas chromatogram
n-C5
n-C6
MCYC5
Benzene
n-C7
MCYC6
Toluene
n-C8
n-C15
Pr
Ph
n-C20
n-C25
Gasoline range gas chromatogram
GSC Lab no. L02155
10-31-39-23W4
Leduc/Nisku Fm.
Saturated HC gas chromatogram
Gasoline range gas chromatogram
GSC Lab no. L02177
10-32s-30-27W4
Leduc/Nisku Fm.
Saturated HC gas chromatogram
Relative abundance
Time

Figure 8a

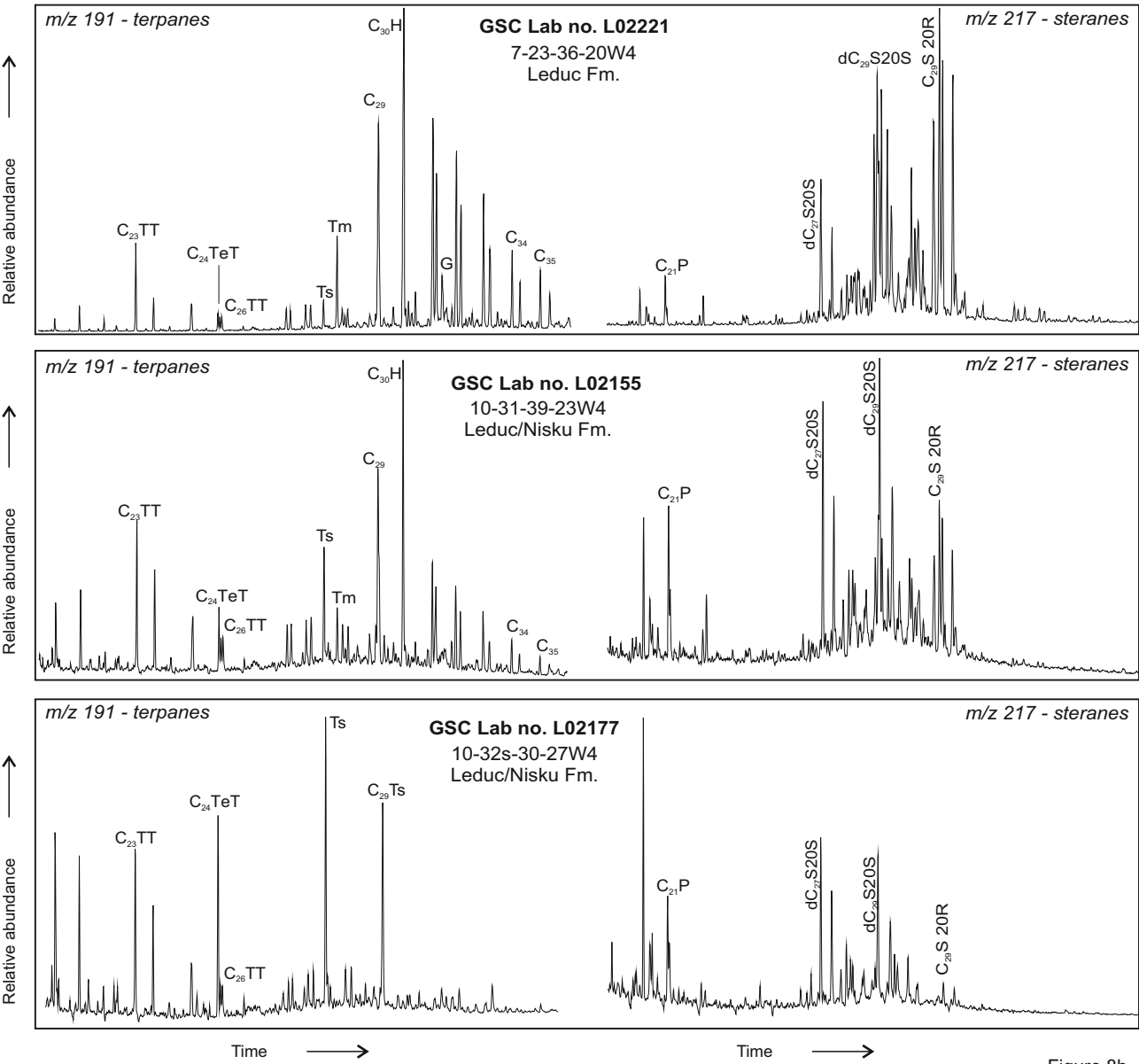
m/z 191 - terpanes
GSC Lab no. L02221
7-23-36-20W4
Leduc Fm.
m/z 217 - steranes
C30H
C29
C23TT
C24TeT
C26TT
Tm
Ts
G
C34
C35
C21P
dC27S20S
dC29S20S
C29 S 20R
m/z 191 - terpanes
GSC Lab no. L02155
10-31-39-23W4
Leduc/Nisku Fm.
m/z 217 - steranes
C30H
C29
C23TT
Ts
C24TeT
Tm
C26TT
C34
C35
C21P
dC27S20S
dC29S20S
C29 S 20R
m/z 191 - terpanes
GSC Lab no. L02177
10-32s-30-27W4
Leduc/Nisku Fm.
m/z 217 - steranes
Ts
C29Ts
C24TeT
C23TT
C26TT
C21P
dC27S20S
dC29S20S
C29 S 20R
Relative abundance
Time
Time

Figure 8b

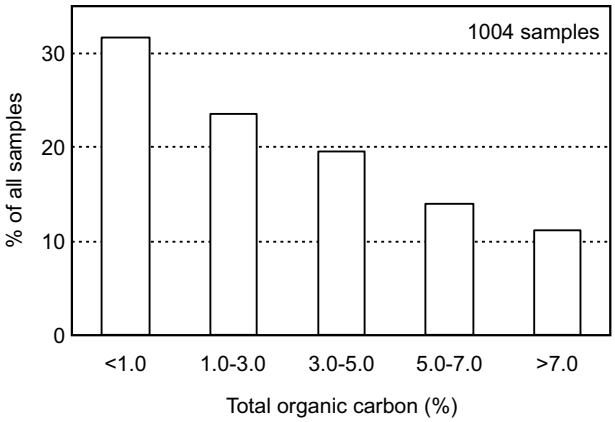
1004 samples
% of all samples
30
20
10
0
<1.0
1.0-3.0
3.0-5.0
5.0-7.0
>7.0
Total organic carbon (%)

Figure 9

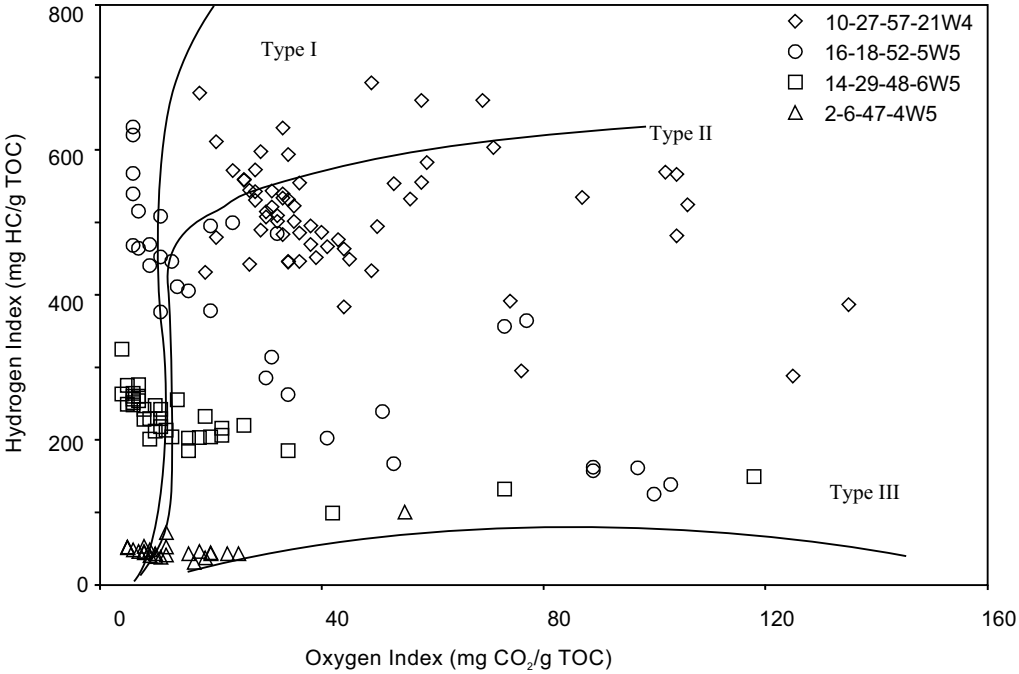
800
600
400
200
0
Hydrogen Index (mg HC/g TOC)
0
40
80
120
160
Oxygen Index (mg CO2/g TOC)
Type I
Type II
Type III
10-27-57-21W4
16-18-52-5W5
14-29-48-6W5
2-6-47-4W5

Figure 10

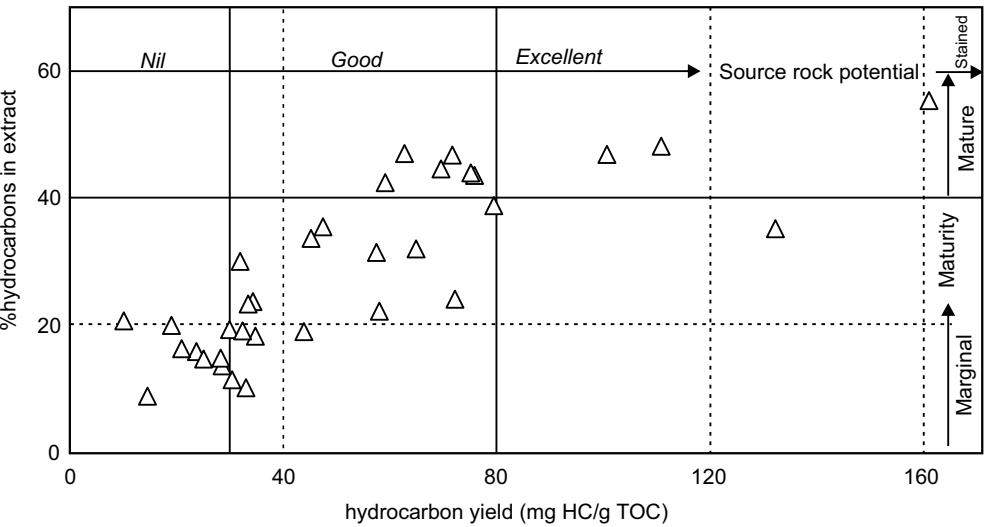
Nil
Good
Excellent
Source rock potential
Stained
Mature
Maturity
Marginal
%hydrocarbons in extract
hydrocarbon yield (mg HC/g TOC)
0
20
40
60
0
40
80
120
160

Figure 11

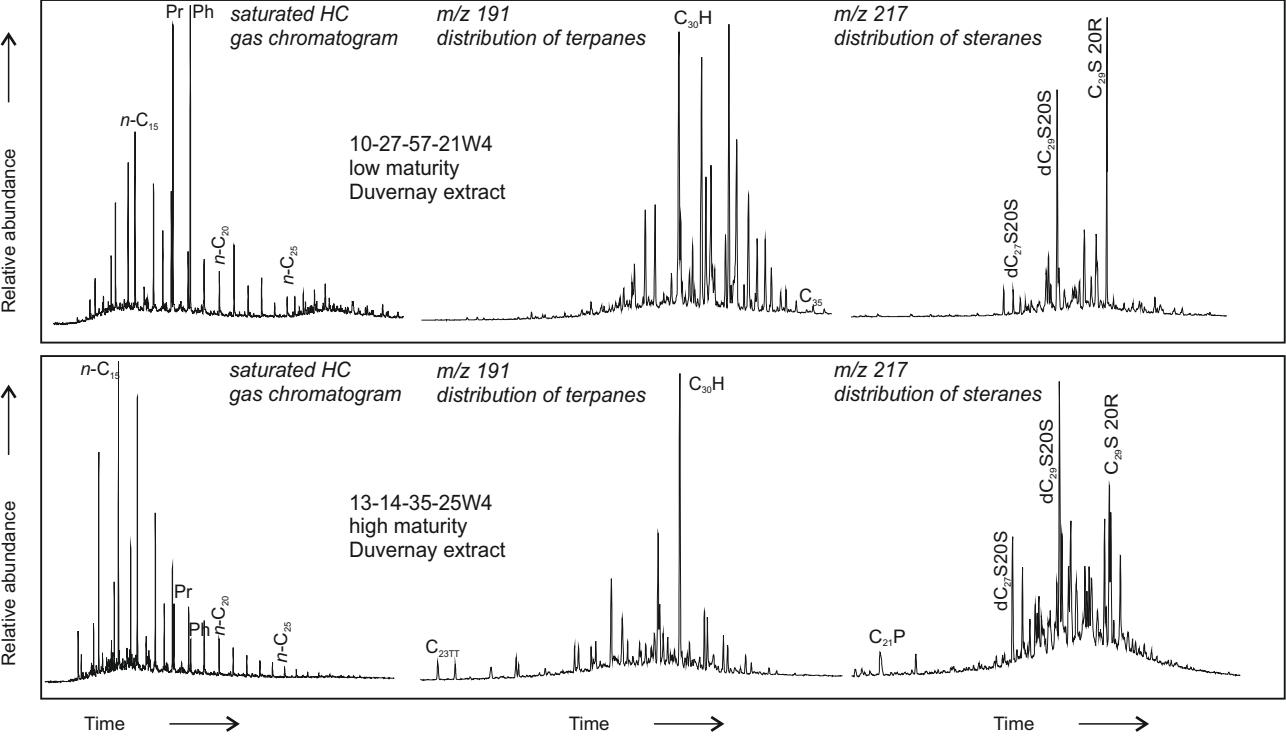
saturated HC
gas chromatogram
m/z 191
distribution of terpanes
m/z 217
distribution of steranes
Pr
Ph
n-C15
n-C20
n-C25
C30H
C35
dC27S20S
dC29S20S
C29S 20R
10-27-57-21W4
low maturity
Duvernay extract
13-14-35-25W4
high maturity
Duvernay extract
C23TT
C21P
Relative abundance
Time

Figure 12

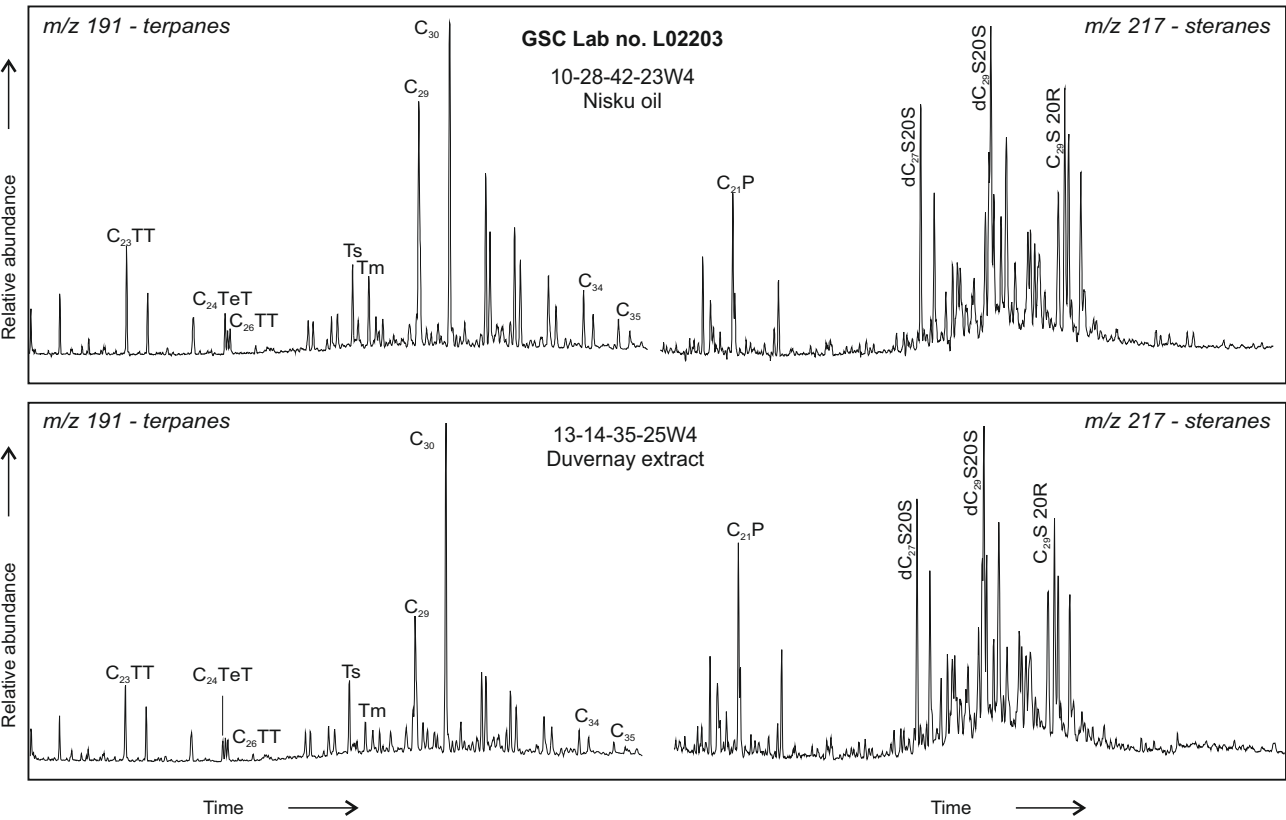
m/z 191 - terpanes
GSC Lab no. L02203
10-28-42-23W4
Nisku oil
m/z 217 - steranes
C30
C29
C23TT
C24TeT
C26TT
Ts
Tm
C34
C35
C21P
dC27S20S
dC29S20S
C29S 20R
Relative abundance
m/z 191 - terpanes
13-14-35-25W4
Duvernay extract
m/z 217 - steranes
C30
C29
C23TT
C24TeT
C26TT
Ts
Tm
C34
C35
C21P
dC27S20S
dC29S20S
C29S 20R
Relative abundance
Time
Time

Figure 13

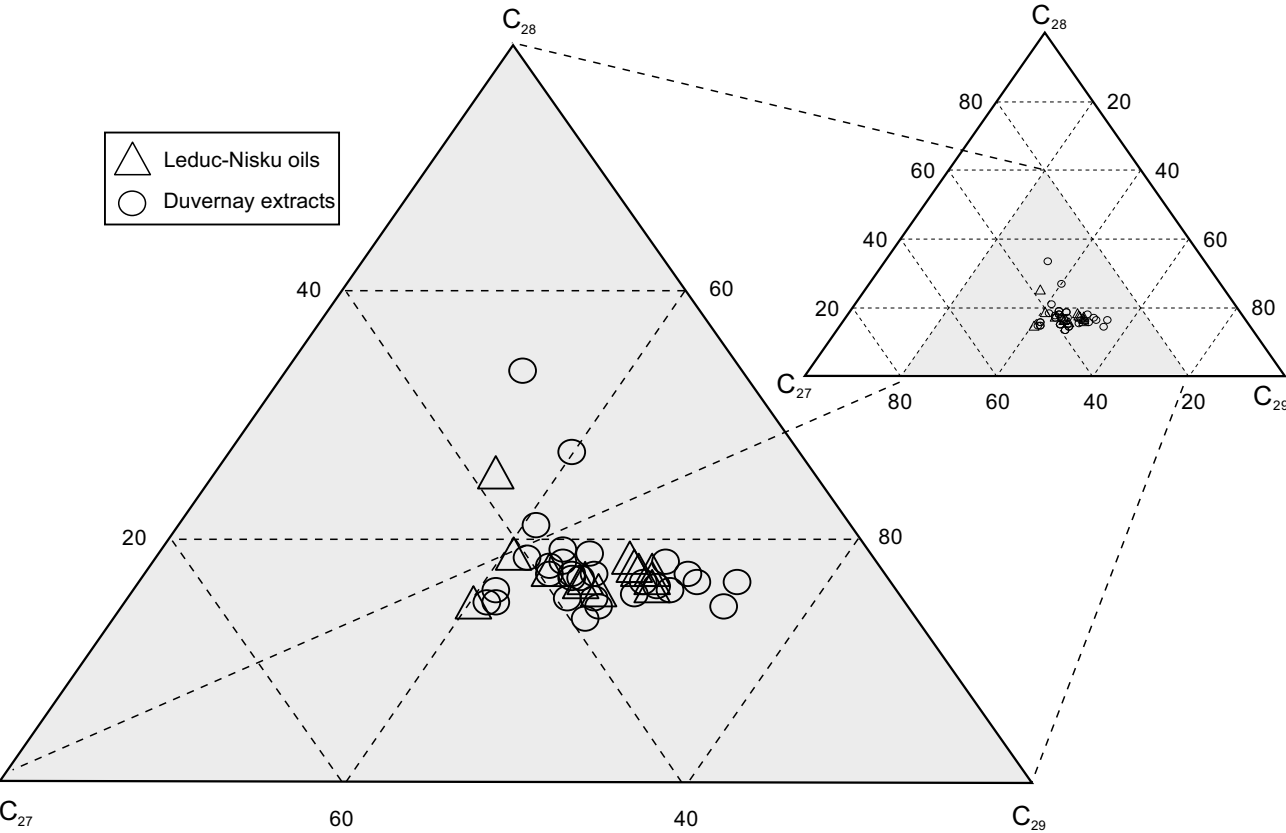
C28
Leduc-Nisku oils
Duvernay extracts
40
60
20
80
C27
60
40
C29
C28
80
20
60
40
40
60
20
80
C27
80
60
40
20
C29

Figure 14

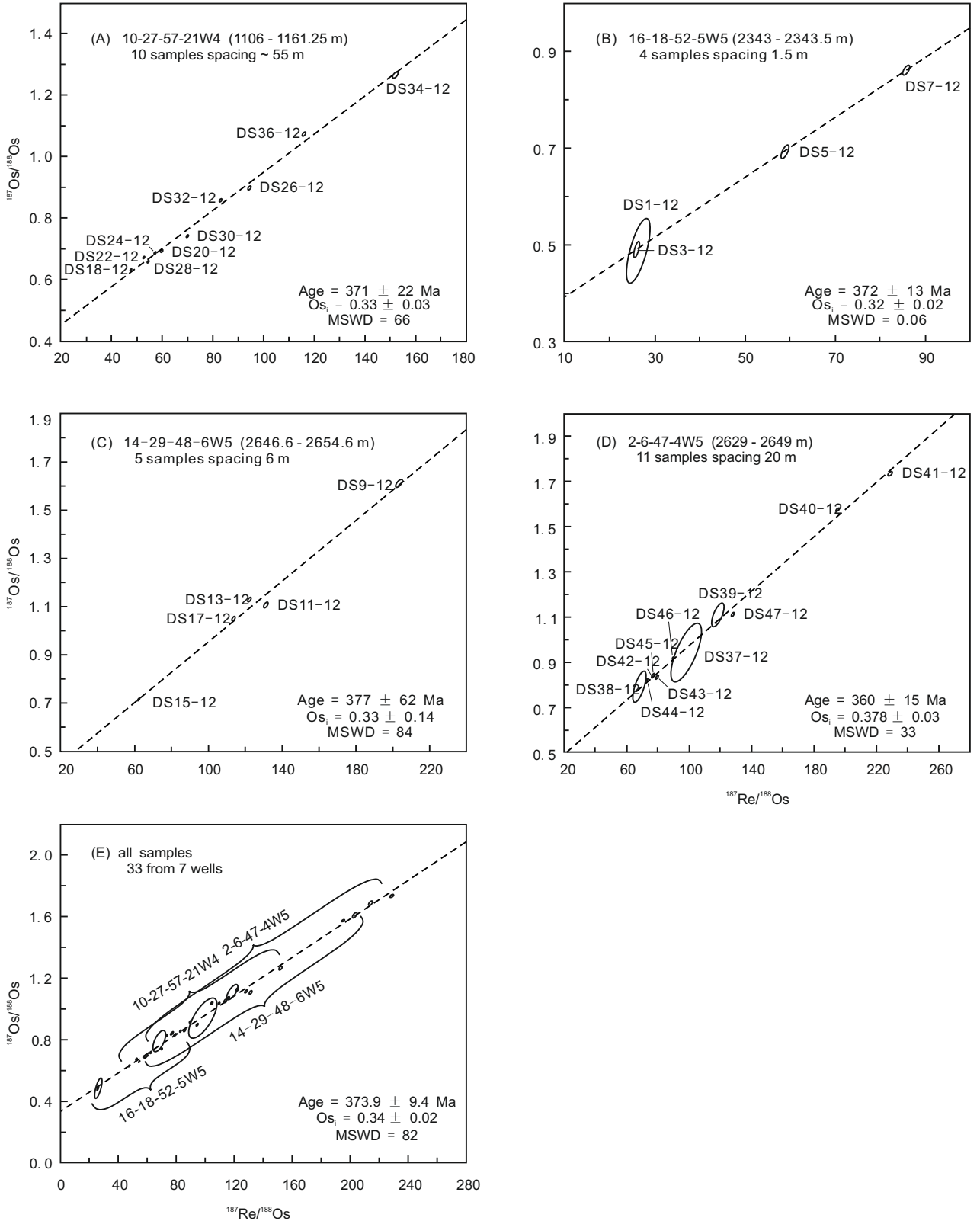
(A) 10-27-57-21W4 (1106 - 1161.25 m)
10 samples spacing ~ 55 m
DS34-12
DS36-12
DS26-12
DS32-12
DS30-12
DS24-12
DS22-12
DS20-12
DS18-12
DS28-12
Age = 371 ± 22 Ma
Os$_i$ = 0.33 ± 0.03
MSWD = 66
(B) 16-18-52-5W5 (2343 - 2343.5 m)
4 samples spacing 1.5 m
DS7-12
DS5-12
DS1-12
DS3-12
Age = 372 ± 13 Ma
Os$_i$ = 0.32 ± 0.02
MSWD = 0.06
(C) 14-29-48-6W5 (2646.6 - 2654.6 m)
5 samples spacing 6 m
DS9-12
DS13-12
DS11-12
DS17-12
DS15-12
Age = 377 ± 62 Ma
Os$_i$ = 0.33 ± 0.14
MSWD = 84
(D) 2-6-47-4W5 (2629 - 2649 m)
11 samples spacing 20 m
DS41-12
DS40-12
DS39-12
DS46-12
DS47-12
DS45-12
DS42-12
DS37-12
DS38-12
DS43-12
DS44-12
Age = 360 ± 15 Ma
Os$_i$ = 0.378 ± 0.03
MSWD = 33
(E) all samples
33 from 7 wells
10-27-57-21W4
2-6-47-4W5
14-29-48-6W5
16-18-52-5W5
Age = 373.9 ± 9.4 Ma
Os$_i$ = 0.34 ± 0.02
MSWD = 82
^{187}Os/^{188}Os
^{187}Re/^{188}Os

Figure 15

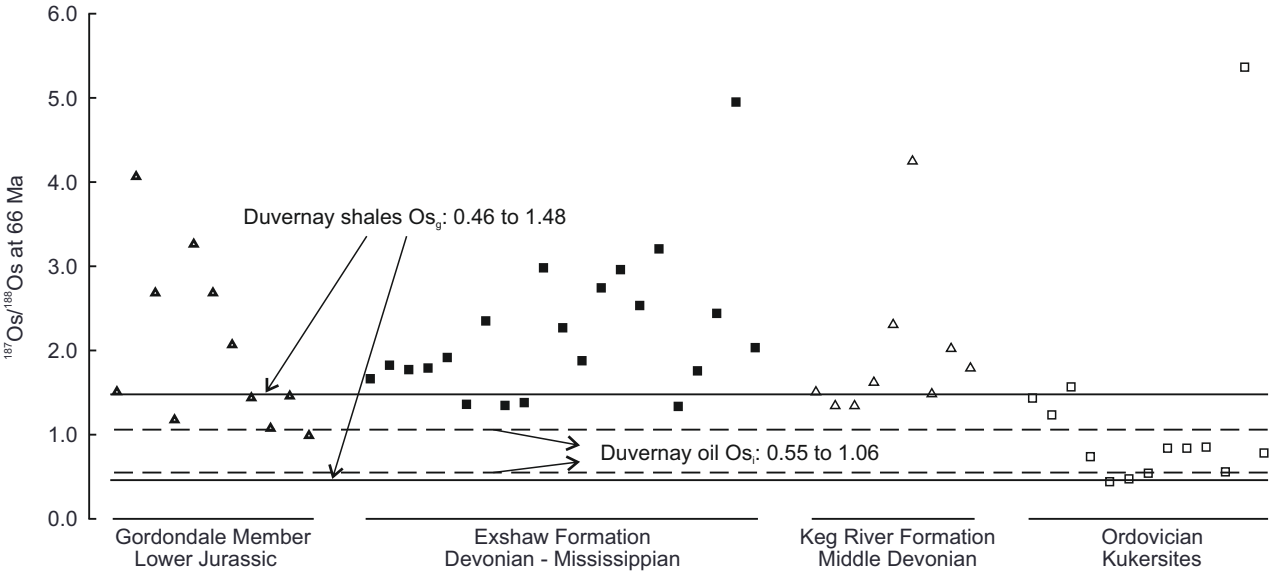
6.0
5.0
4.0
3.0
2.0
1.0
0.0
$^{187}Os/^{188}Os$ at 66 Ma
Duvernay shales Os_g: 0.46 to 1.48
Duvernay oil Os_i: 0.55 to 1.06
Gordondale Member
Lower Jurassic
Exshaw Formation
Devonian - Mississippian
Keg River Formation
Middle Devonian
Ordovician
Kukersites

Figure 16