

1           **Effects of debris-flow magnitude-frequency distribution on**  
2                                   **avulsions and fan development**

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6 **Abstract**

7 Shifts in the active channel on a debris-flow fan, termed avulsions, pose a large threat because  
8 new channels can bypass mitigation measures and cause damage to settlements and infrastruc-  
9 ture. Recent, but limited, field evidence suggests that avulsion processes and tendency may  
10 depend on the flow-size distribution, which is difficult to constrain in the field. Here, we in-  
11 vestigate how the flow magnitude-frequency distribution and the associated flow-magnitude  
12 sequences affect avulsion on debris-flow fans. We created three experimental fans with con-  
13 trasting flow-size distributions: (1) a uniform distribution, (2) a steep double-Pareto distribu-  
14 tion with many flows around the mean and a limited number of large flows, and (3) a shal-  
15 low double-Pareto distribution with fewer flows around the mean and more abundant large flows.  
16 The fan formed by uniform flows developed through regular sequences of stepwise channel-  
17 ization, backstepping of deposition toward the fan apex, and avulsion over multiple flows. In  
18 contrast, the wide range of sizes in the double-Pareto distributions led to distinct avulsion mech-  
19 anisms and fan evolution. Here, large flows could overtop channels, creating levee breaches  
20 that could initiate avulsion immediately or in subsequent events. Moreover, sequences of small-  
21 to moderately-sized flows could deposit channel plugs, triggering avulsion in the next large  
22 flow. This mechanism was most common on the fan formed by a steep double-Pareto distri-  
23 bution but was rare on the fan formed by a shallow double-Pareto distribution, where large flows  
24 were more frequent. We infer that some flow-size distributions are more likely to cause avul-  
25 sions - especially those that produce abundant sequences of small flows followed by a large  
26 flow. Critically, avulsions in our experiments could occur by either large single events or over  
27 multiple flows. This observation has important implications for hazard assessment on debris-  
28 flow fans, suggesting that attention should be paid to flow history as well as flow size.

## 1 Introduction

Debris-flow fans are ubiquitous landforms in high-relief areas around the world [e.g., *Beatty, 1963; Okuda et al., 1981; Whipple and Dunne, 1992; Blair and McPherson, 1994, 2009; De Haas et al., 2014, 2015a; D’Arcy et al., 2015; Schürch et al., 2016*]. They form by deposition in repeated debris flows, and are thus an archive of past flow magnitude, timing, composition and depositional pattern [*Schumm et al., 1987; Harvey, 2011; Dühnforth et al., 2007*]. Extracting such information requires understanding of the spatio-temporal patterns of debris-flow fan evolution, which largely depend on changes in the active-channel position, termed avulsions, that distribute sediment across the fan surface [*Schürch et al., 2016*].

Ongoing expansion of human populations into mountainous regions has led to increasing exposure to debris-flow hazards [*Pederson et al., 2015*]. Avulsions pose an especially severe threat to settlements and other infrastructure on fans, particularly as flow mitigation measures such as check dams and retention basins are typically applied to active channels and cannot prevent damage from flows that establish a new channel pathway. The mechanisms by which debris flows avulse to occupy new flow paths on fans, however, and the controls on avulsion frequency and timing, are poorly understood [e.g., *Pederson et al., 2015; De Haas et al., 2016, 2018*]. One outstanding issue is that the spatio-temporal patterns of deposition on debris-flow fans have been monitored [*Suwa and Okuda, 1983; Wasklewicz and Scheinert, 2016; Imaizumi et al., 2016*] or reconstructed [e.g., *Helsen et al., 2002; Dühnforth et al., 2008; Stoffel et al., 2008; Bollschweiler et al., 2008; Schürch et al., 2016; Zaginaev et al., 2016*] on only a few natural debris-flow fans. Moreover, there have been few attempts to simulate debris-flow fan evolution with physical scale experiments [*Hooke, 1967; Schumm et al., 1987; De Haas et al., 2016*] or numerical models [*Schürch, 2011*]. *De Haas et al.* [2018] summarized and compared the patterns of spatio-temporal debris-flow deposition on natural fans, and identified two important controls on avulsion that operate over separate time scales: (1) during individual flows or flow surges, deposition of sediment locally blocks or plugs channels, forcing avulsion in subsequent flows, and (2) over time scales of tens of flows, the average locus of debris-flow deposition gradually shifts towards topographically lower sectors of a fan. Many, but not all, debris-flow avulsions follow a pattern of channel plugging, backstepping of deposition toward the fan apex, avulsion and establishment of a new active channel. In this conceptual model, sequences of small- to medium-sized flows can progressively deposit sediment within the active channel toward the fan apex, thereby plugging the channel, until a flow occurs that is of sufficient magnitude to leave the main channel upstream of the sediment plug and form a new channel.

62 Plug deposition is a stochastic process that depends on the sequence of flow magnitudes, the  
63 geometry of the channel, and the composition and bulk rheology of the flows. Furthermore,  
64 *De Haas et al.* [2018] showed that large flows can have contrasting impacts on avulsion, de-  
65 pending on whether or not they follow smaller flows that have deposited channel plugs.

66 These observations suggest that avulsions and associated patterns of debris-flow fan for-  
67 mation may depend on the relative numbers of small and large flows - and thus on the magnitude-  
68 frequency distribution - and on the sequence of flows that feed a fan. Each debris-flow fan is  
69 built by a unique, but generally unknown, magnitude-frequency distribution, which could con-  
70 ceivably lead to contrasting spatio-temporal avulsion patterns on different fans. While *De Haas*  
71 *et al.* [2018] speculated on this link, they lacked robust data on flow volumes for most of their  
72 field examples, and they had no information on the underpinning distributions. For hazard mit-  
73 igation, and to effectively decipher the debris-flow fan archive, it is therefore of key impor-  
74 tance to understand how different magnitude-frequency distributions, and the associated se-  
75 quences of flow sizes, can affect avulsions and the spatio-temporal patterns of debris-flow fan  
76 evolution. This is especially relevant in regions where magnitude-frequency distributions have  
77 changed, or may change, as a result of global climate change [e.g., *Rebetz et al.*, 1997; *Stof-*  
78 *fel*, 2010; *Clague et al.*, 2012] or regional factors such as earthquakes [e.g., *Shieh et al.*, 2009;  
79 *Huang and Fan*, 2013; *Ma et al.*, 2017], landslides [e.g., *Imaizumi et al.*, 2016], or wildfires  
80 [e.g., *Cannon et al.*, 2008, 2011].

81 Here, we investigate how flow magnitude-frequency distribution and associated flow se-  
82 quences affect the spatio-temporal patterns of debris-flow-fan development. To do so, we study  
83 and compare the evolution, avulsion mechanisms and compensational tendency of three experimentally-  
84 created debris-flow fans formed by different flow-magnitude distributions. We follow *De Haas*  
85 *et al.* [2016], who investigated avulsions and debris-flow fan evolution on an experimental fan  
86 formed by flows of uniform size and composition. They found that avulsions on this fan fol-  
87 lowed a predictable pattern of gradual backstepping, avulsion and channelization. Phases of  
88 backstepping and channelization were approximately equal in length and developed over mul-  
89 tiple flows. They speculated about the potential effects of varying flow size on avulsion oc-  
90 currence and mechanism, but could not test these ideas. We thus build on this work by cre-  
91 ating two additional debris-flow fans formed by contrasting heavy-tailed magnitude-frequency  
92 distributions using the same experimental setup, and comparing the avulsion mechanisms and  
93 spatio-temporal patterns of activity between these fans.

94 The structure of this paper is as follows. We first describe the methodology, experimen-  
95 tal setup and procedure, and data reduction and analysis methods. Then we describe the spatio-  
96 temporal patterns of development on the three experimental debris-flow fans, and determine  
97 their compensational tendencies as quantified by the compensation index [cf. *Straub and Pyles,*  
98 2012]. Finally, we discuss the potential relationships between debris-flow magnitude-frequency  
99 distribution, flow sequence, and avulsion on debris-flow fans, based on the experimental re-  
100 sults.

## 101 **2 Materials and methods**

102 The fan described by *De Haas et al.* [2016] and the two new experimental fans described  
103 here were generated with the same experimental setup and procedure. The large-scale flow pat-  
104 terns of the experimental debris flows mimic those of natural debris flows [*De Haas et al., 2015b*]:  
105 all experimental debris flows presented here were frictional flows, with coarse particles selec-  
106 tively transported to the flow front and subsequently shouldered aside to form coarse-grained  
107 lateral levees. Each flow produced a distinct depositional lobe wherein coarse particles were  
108 predominantly concentrated at the frontal flow margins [*De Haas et al., 2015b; De Haas and*  
109 *Van Woerkom, 2016*]. Moreover, channel width to depth ratios, and runout length and area rel-  
110 ative to debris-flow volume, are similar to those in natural debris flows [*De Haas et al., 2015b*].  
111 The morphological similarity between the experimental and natural debris flows allows for rep-  
112 resentative interactions between debris flows and evolving fan morphology, which enables us  
113 to study avulsion mechanisms and tendencies, and allows for broad comparisons of the results  
114 to natural debris-flow fans [cf. *Hooke and Rohrer, 1979; Paola et al., 2009*]. We emphasize  
115 that the experiments are not intended as scaled analogues of natural flows or fans, and that our  
116 aim is to examine the morphodynamic behavior of the system in the face of different flow-  
117 magnitude distributions. Thus, it are the differences between experiments, rather than the de-  
118 tailed results of a single experiment, that are of primary interest.

### 119 **2.1 Experimental setup and procedure**

120 The experimental setup was described in detail in *De Haas et al.* [2016], and consisted  
121 of a mixing tank connected to a 30° inclined chute channel, 2 m long and 0.12 m wide, which  
122 at its downstream end was linked to an outflow plain with an inclination of 10° (Fig. 1). The  
123 channel bed and sidewalls of the chute channel were covered with sandpaper to simulate nat-  
124 ural bed roughness (grade 80; average particle diameter 0.19 mm), and the outflow plain was

125 covered by a 0.5 cm deep layer of unconsolidated sand (median particle diameter 0.4 mm).  
126 Sediment and water were added to the mixing tank and then agitated until a coherent mixture  
127 was formed, after which a gate was opened electromagnetically to release the debris-flow mix-  
128 ture into the channel. A hatch in the channel bed, located 0.76 m above the transition from  
129 the channel to the outflow plain, was opened electromagnetically 1.5 s after the release of de-  
130 bris from the mixing tank to cut off the sediment-poor debris-flow tail, which would other-  
131 wise incise unrealistically deep into the fan deposits [De Haas *et al.*, 2016].

132 The experimental fans were created by stacking of consecutive debris-flow deposits on  
133 the outflow plain, leaving base level fixed. The fans were allowed to grow in size until a max-  
134 imum extent was reached, at which point subsequent debris flows were not able to reach the  
135 fan as they were blocked by accumulated debris in the feeder channel. This occurred after 55  
136 to 85 flows, depending on the experiment. All debris flows had a similar composition consist-  
137 ing of clay (kaolinite; 5.8% of total sediment volume), sand (77.2% of the total sediment vol-  
138 ume) and basaltic gravel (17% of the total sediment volume) (Fig. 2). All flows contained 44  
139 vol% of water.

140 Above the flume, a digital camera (Canon PowerShot A640) was set up to image fan to-  
141 pography after each flow. Videos of flow movement and deposition on the fan were captured  
142 with a Canon Powershot A650 IS on a tripod directed obliquely at the channel and fan. De-  
143 posit morphology was measured at sub-millimeter resolution and accuracy after every debris-  
144 flow event using a Vialux z-Snapper 3D scanner that captures a three-dimensional point cloud  
145 from a fringe pattern projector [Hoefling, 2004]. Point clouds from the scanner were converted  
146 into a gridded digital elevation model (DEM) with 1 mm spatial resolution using natural neigh-  
147 bor interpolation (Fig. 1).

## 148 **2.2 Magnitude-frequency distribution**

149 The three experimental fans were formed by selecting flow sizes from three different magnitude-  
150 frequency distributions (Fig. 3). Although the distribution of flow magnitudes from natural debris-  
151 flow catchments is generally not known [e.g., Stoffel, 2010], there is a considerable body of  
152 evidence that landslide magnitudes follow a heavy-tailed distribution [e.g., Hovius *et al.*, 1997;  
153 Hungr *et al.*, 1999; Dai and Lee, 2001; Stark and Hovius, 2001; Malamud *et al.*, 2004; Ben-  
154 nett *et al.*, 2012]. Given the genetic link between shallow landsliding and debris-flow initia-  
155 tion [e.g., Iverson, 1997; Iverson *et al.*, 2011; Bennett *et al.*, 2013], it seems logical to assume

156 that debris-flow magnitudes may also show heavy-tailed behavior. Indeed, *Bardou and Jaboyed-*  
 157 *off* [2008] demonstrated a heavy-tailed magnitude distribution for a compilation of historical  
 158 debris flows from the Swiss Alps, while *Bennett et al.* [2014] compiled observations from the  
 159 Illgraben (Switzerland) that also show heavy-tailed behavior. *Bennett et al.* [2014] also cau-  
 160 tioned, however, that their modeling showed important differences between the magnitude dis-  
 161 tributions of landslides and debris flows in the Illgraben catchment.

162 With these considerations in mind, we developed and compared three distinct flow-magnitude  
 163 distributions: a uniform distribution previously described by *De Haas et al.* [2016] (fan 01),  
 164 a heavy-tailed distribution with a large power-law exponent (corresponding to a rapid decrease  
 165 in exceedance probability with increasing flow magnitude, fan 02), and a heavy-tailed distri-  
 166 bution with a small exponent (corresponding to a more gradual decrease in exceedance prob-  
 167 ability with increasing magnitude, fan 03). For convenience, we extracted flow mass from each  
 168 distribution, using a constant flow composition and thus bulk density. To simulate the two heavy-  
 169 tailed distributions, we followed *Stark and Hovius* [2001] and *Guthrie and Evans* [2004] in adopt-  
 170 ing a double-Pareto formulation. This distribution exhibits power-law behavior in the upper  
 171 and lower tails, and allows for inclusion of a rollover, with extremely large and extremely small  
 172 flows both being less likely [e.g., *Reed*, 2001; *Reed and Jorgensen*, 2004]. The probability den-  
 173 sity function can be written as:

$$f(M) = \begin{cases} \frac{\alpha\beta}{\alpha+\beta} \left(\frac{M}{M_c}\right)^{\beta-1}, & M \leq M_c \\ \frac{\alpha\beta}{\alpha+\beta} \left(\frac{M}{M_c}\right)^{-\alpha-1}, & M \geq M_c \end{cases} \quad (1)$$

174 where  $M$  is flow mass (kg),  $M_c$  is a rollover parameter (kg), and  $\alpha$  and  $\beta$  are empiri-  
 175 cal constants that describe the slope of the density function at small and large magnitudes,  
 176 respectively. For fan 02, we set  $M_c = 4.25$  kg,  $\alpha = 10.05$ , and  $\beta = 30.5$ , and we refer to this  
 177 below for convenience as the ‘steep’ distribution. For fan 03, we set  $M_c = 3.0$  kg,  $\alpha = 3.05$ ,  
 178 and  $\beta = 10.5$ , and we refer to this as the ‘shallow’ distribution. These distributions are not in-  
 179 tended to mimic known field examples, but were rather designed as plausible and contrasting  
 180 end-members.

181 The mean flow mass in all three experiments was fixed at 6.5 kg. For fan 01, the flow  
 182 mass was kept uniform. For fans 02 and 03, the mass of each flow in the sequence was de-  
 183 termined by extracting a random deviate from the distribution described by eq. 1 with the ap-

184 appropriate parameter values. The maximum flow mass in the latter experiments was fixed at 13  
 185 kg due to operational constraints.

## 186 **2.3 Data reduction**

### 187 **2.3.1 Spatio-temporal patterns of activity**

188 The patterns of deposition in each flow were summarized by the flow angle and runout  
 189 distance for each debris-flow snout, the maximum runout among all snouts, the deposit width,  
 190 the deposit width/depth ratio, and the channel depth at the fan apex (Fig. 4). The flow angle  
 191 was defined as the angle between the fan midline and a straight line connecting the fan apex  
 192 and the debris-flow snout. The runout distance per snout was defined as the length of a straight  
 193 line from apex to snout. Deposit width was defined as the maximum width of the deposit, ex-  
 194 cluding individual snouts substantially outside of the main flow direction. Apex channel depth  
 195 was measured 10 cm downstream of the fan apex.

### 196 **2.3.2 Compensation index**

197 The compensation index ( $\kappa_{cv}$ ) describes the tendency of a sedimentary system to occupy  
 198 and fill topographic lows by avulsion [Straub *et al.*, 2009; Wang *et al.*, 2011; Straub and Pyles,  
 199 2012]. This index ranges from 0 to 1, representing a continuum from persistent channel po-  
 200 sitions and vertical (anti-compensational) stacking of deposits ( $\kappa_{CV} = 0$ ), through random chan-  
 201 nel positions ( $\kappa_{CV} = 0.5$ ), to frequent avulsions and perfect topographic compensation ( $\kappa_{CV}$   
 202  $= 1$ ). In other words, in an anti-compensational system previous deposits act as attractors for  
 203 the active channel, while in a compensational system previous deposits act as deflectors. As  
 204 such, the compensation index is a valuable measure for understanding avulsion frequency and  
 205 future flow path prediction. For the experimental debris-flow fans we calculated the compen-  
 206 sation index at 0.05 m increments of distance from fan apex to fan toe following the method  
 207 of Straub and Pyles [2012], which is a revised version of the earlier approach of Straub *et al.*  
 208 [2009] that ignores basin subsidence rates. This index has been previously used to calculate  
 209 the compensational tendency of natural debris-flow fans in Colorado, USA, by Pederson *et al.*  
 210 [2015]. The compensation index depends upon the coefficient of variation of the ratio of lo-  
 211 cal to mean sediment thickness between every pairwise combination of bed boundaries inte-  
 212 grated across the horizontal length ( $L$ ) of the basin:

$$CV = \left( \int_L \left[ \frac{\Delta\eta(x)_{A,B}}{\Delta\bar{\eta}_{A,B}} - 1 \right]^2 dL \right)^{0.5} \quad (2)$$

213 where  $\Delta\eta(x)_{A,B}$  is the local sediment thickness between surfaces A and B, and  $\Delta\bar{\eta}_{A,B}$   
 214 is the mean deposit thickness between surfaces A and B. The compensation index ( $\kappa_{cv}$ ) is the  
 215 exponent in the power law decay of  $CV$  with increasing mean sediment thickness:

$$CV = a\Delta\bar{\eta}_{A,B}^{-\kappa_{cv}} \quad (3)$$

216 where  $a$  is an empirical constant.

## 217 **3 Results**

### 218 **3.1 Spatio-temporal patterns of fan development**

219 In this section we summarize the spatio-temporal evolution of the three experimental debris-  
 220 flow fans and their dominant avulsion mechanisms - more extensive descriptions of the evo-  
 221 lution of the fans on a flow by flow basis can be found in the supplementary materials. We  
 222 end the results by presenting the compensation index calculations. Flow sizes and spatio-temporal  
 223 patterns of debris-flow activity on the fans are summarized in Figure 5. Flows that moved to-  
 224 wards the left-hand side of the fan, when looking downstream from the fan apex, are denoted  
 225 by negative flow angles and flows towards the right are denoted by positive flow angles.

#### 226 **3.1.1 Fan 01: Uniform distribution**

227 Fan 01 evolved in a predictable manner (Figs. 5a-f, 6; supplementary movie S1): fill-  
 228 ing of accommodation induced backstepping sequences which resulted in a searching phase,  
 229 followed by avulsion and re-channelization (Fig. 6). Maximum runout was observed to decrease  
 230 in gradual and near-uniform steps over multiple flows during the backstepping phases, and sim-  
 231 ilarly increased during phases of channel establishment and progradation (Fig. 5b). These progra-  
 232 dation and backstepping phases required an approximately similar number of flows, but the  
 233 total length of these phases increased as the fan apex grew in elevation and more accommo-  
 234 dation had to be filled before a backstepping sequence could be initiated (compare the sequence  
 235 during flows 10-25 with 25-52 in Figs. 5a and supplementary movie S1). During searching  
 236 phases, when multiple channels were active, deposit width was relatively large while the apex  
 237 channel was shallow or even absent (Fig. 5c-e) [see *De Haas et al.*, 2016, for further details].

### 238 **3.1.2 Fan 02: Steep double-Pareto distribution**

239 Runout distances on fan 02 were relatively long during channelized phases and restricted  
240 during unchannelized phases (Figs. 5g-l, 7; supplementary movie S2). Compared to fan 01,  
241 however, this relationship was less well-developed as a result of the varying debris-flow sizes.  
242 Periods of backstepping and the formation of persistent channel plugs that induced avulsion  
243 on fan 02 were generally initiated by (1) sequences of small- to moderately-sized flows (e.g.,  
244 flows 16-18, 26-29 and 35-36; Fig. 7a-h; supplementary movie S2), and by (2) complete fill-  
245 ing of the regional accommodation (e.g., flows 53-58). Very large events (e.g., flows 15 and  
246 59; Fig. 7j) had sufficient magnitude to overflow the main channel, often upstream of chan-  
247 nel plugs, and form a new channel. Additionally, large flows were often observed to overtop  
248 the main channel in multiple locations, creating levee breaches that could be exploited as avul-  
249 sion sites during subsequent flows and develop into new main channels.

### 250 **3.1.3 Fan 03: Shallow double-Pareto distribution**

251 Runout on fan 03 was generally greatest when a well-defined apex channel was present  
252 (Figs. 5m-r, 8; supplementary movie S3). During searching phases when channel depth was  
253 small or an apex channel was absent, runout was restricted and deposits were wide. This trend  
254 is only weak and hard to recognize, however, due to the large variation in runout and deposit  
255 width caused by the broad distribution of flow magnitudes. Channels were more persistent on  
256 fan 03 compared to fans 01 and 02, and small channel plugs were sometimes removed by large  
257 flows (Fig. 8a-d). Clear backstepping sequences over multiple flows in the main channel, which  
258 were frequently observed on fans 01 and 02, were much less frequent on fan 03. Backstep-  
259 ping and plug deposition were, however, generally responsible for the closure of secondary  
260 channels. Periods during which the main locus of deposition migrated towards the fan apex  
261 did occur and were observed to induce avulsion (e.g., flows 56-62; Fig. 8e-f), but these back-  
262 stepping sequences predominantly occurred as a result of filled regional accommodation. More-  
263 over, these sequences were irregular, often showing alternating progradation and retrograda-  
264 tion on an event scale, because of the strongly varying flow sizes. New topographically-favorable  
265 channels were often formed during solitary large events (e.g., flows 63-64 and 69; Fig. 8h),  
266 triggering main channel avulsion in subsequent flows.

## 267 **3.2 Compensational tendency**

268 In this section we calculate the compensation index (eq. 3) for the experimental debris-  
269 flow fans. We first describe the stratigraphy that developed in each experiment at representa-  
270 tive proximal and distal transects located 0.2 and 0.8 m downstream of the apex, respectively.  
271 Next, we determine the compensation index and examine how it varied with distance from the  
272 apex in each experiment.

### 273 **3.2.1 Fan stratigraphy**

274 On fan 01, stratigraphy on both the proximal and distal parts of the fan shows that de-  
275 position was generally persistent for periods of at least  $\sim 20$  flows, after which activity avulsed  
276 to a topographically lower area (Fig. 9a-b). Fan 02 showed roughly similar behavior, but the  
277 variability in debris-flow magnitudes resulted in less clearly-pronounced lateral shifts and a  
278 larger number of overflow events in the fan stratigraphy (Fig. 9c-f). We define an overflow event  
279 as a flow that was able to extensively overtop the main channel levees and deposit substan-  
280 tial amounts of sediment adjacent to the main channel. Overflow events were even more pro-  
281 nounced in the stratigraphy of fan 03 compared to fans 01 and 02, as would be expected from  
282 the shallow flow distribution. This difference resulted in even more persistent deposition and  
283 less pronounced lateral shifts in both the proximal and distal parts of the fan. Deposition was  
284 observed to be persistent on one side of the fan for  $\sim 20$  flows on fans 01 and 02, while it was  
285 typically persistent for  $>30$  flows on fan 03 (Figs 5, 9).

### 286 **3.2.2 Compensation index**

287 The compensation index was roughly similar on the three experimental debris-flow fans  
288 (Fig. 10). In general, the index had a value of  $\sim 0.25$  near the fan apex, and increased roughly  
289 linearly to  $\sim 0.35$  near the fan toe. The compensation index was well-defined and similar for  
290 stratigraphic intervals of between 1 and 20 flows, implying that the compensational behavior  
291 was similar over this range of depositional scales.

## 292 **4 Discussion**

293 In this section we discuss the effects of the flow magnitude-frequency distribution and  
294 associated flow-magnitude sequence on avulsion and experimental fan evolution, and compare  
295 these findings to observations from natural debris-flow fans. Next, we consider the compen-

296 sational tendencies of debris-flow fans, and implications for flow routing. Finally, we detail  
297 the potential implications of our experimental results for mitigation of avulsion hazards on debris-  
298 flow fans.

#### 299 **4.1 Effects of magnitude-frequency distribution on avulsion mechanisms and fan evo-** 300 **lution**

301 Broadly speaking, the three experimental fans followed similar overall patterns of de-  
302 velopment. After a spin-up phase during which flow deposits were largely stacked on top of  
303 each other along the fan midline, each fan showed an alternation of (1) channelized phases,  
304 during which debris flows occupied a well-defined channel and deposited on the medial to dis-  
305 tal parts of the fan, and (2) unchannelized or searching phases, during which flows were spread  
306 widely over the proximal parts of the fan and formed multiple snouts. The searching phases  
307 continued until debris-flow activity avulsed towards a topographically-favorable sector of the  
308 fan. All experimental fans occupied an increasing number of channels and formed more snouts  
309 as the fan surface topography grew more complex over time. Most of these channels were closed  
310 off and abandoned by sequences of backstepping sedimentation over the course of multiple  
311 debris flows.

312 The three different flow distributions also caused marked differences, however, in avul-  
313 sion mechanisms and patterns of fan evolution. Fan 01, with uniform flow magnitudes, devel-  
314 oped through regular sequences of stepwise channelization, backstepping, and avulsion over  
315 multiple flows *De Haas et al.* [2016]. On this fan, backstepping of deposition proceeded from  
316 fan toe to fan apex before substantial shifts in main channel location could occur. This sequence  
317 and the filling of regional accommodation is analogous to the backfilling process that causes  
318 avulsion on fluvial-flow-dominated fans [e.g., *Van Dijk et al.*, 2009, 2012; *Clarke et al.*, 2010;  
319 *Powell et al.*, 2012; *Hamilton et al.*, 2013]. In contrast, large debris flows on fans 02 and 03  
320 were able to overtop channel levees and either occupy new flow paths, or trigger avulsion in  
321 subsequent flows. As a result, avulsions could occur during a single flow of sufficient mag-  
322 nitude, whereas avulsions on fan 01 always occurred over multiple flows. In addition, large  
323 events could cause widespread overbank surges, thereby also contributing to fan construction  
324 and channel embankment. This was particularly evident on fan 03, which had a shallow power-  
325 law decay and thus had relatively more abundant large flows than the other two experiments.

326 Avulsions on fans 02 and 03 were also triggered by specific sequences of flow sizes that  
327 promoted channel plugging and backstepping, enabling rapid channel closure and avulsion with-  
328 out the need for complete backstepping sequences from fan toe to apex. These sequences gen-  
329 erally consisted of a series of flows with similar or progressively-decreasing sizes (e.g., back-  
330 stepping sequences during flows 15-19 and 25-29, and channel plug sequence during flows 35-  
331 36, on fan 02). Interestingly, these sequences were more common on fan 02 compared to fan  
332 03, which we attribute to both (1) the predominance of small to moderate flows on fan 02 and  
333 (2) the relative abundance of large flows on fan 03, which kept the channel swept clean.

334 These experimental observations suggest that new channels on debris-flow fans can form  
335 instantaneously during large flows, but can also form progressively during sequences of small-  
336 to medium-sized flows. New channels generally form after channel blocking by a backstep-  
337 ping sequence, which is either initiated by (1) the deposition of a channel plug or (2) filling  
338 of regional accommodation. In the latter mechanism, sedimentation occurs over the near-full  
339 length of the channel, whereas large parts of the abandoned channel, downstream of the plug,  
340 are preserved by the former mechanism; this distinction may be important for routing of fu-  
341 ture flows down the abandoned channel network [e.g., *Aslan et al.*, 2005; *Reitz et al.*, 2010;  
342 *De Haas et al.*, 2018]. The relative importance of these two mechanisms for channel abandon-  
343 ment appears to depend largely on the debris-flow magnitude-frequency distribution: on fan  
344 02 most backstepping sequences were initiated by channel plugs formed during favorable se-  
345 quences of small- to medium-sized flows, while the absence of variable flow magnitudes on  
346 fan 01 and the relative abundance of large flows on fan 03 both (partly) inhibited the forma-  
347 tion of channel plugs.

348 The observed spatio-temporal patterns of development on the experimental debris-flow  
349 fans are similar to patterns observed on natural debris-flow fans. Channel plugs locally block  
350 channels and force subsequent flows to avulse on the experimental fans (Fig. 7a-c), and this  
351 behavior is also frequently observed in natural fan systems [*Okuda et al.*, 1981; *Suwa and Ya-*  
352 *makoshi*, 1999; *De Haas et al.*, 2018] (Fig. 11b). Over time scales of multiple events, typically  
353 at least 5 to >20 flows on both a range of natural debris-flow fans [*Okuda et al.*, 1981; *Suwa*  
354 *et al.*, 2009; *Imaizumi et al.*, 2016; *De Haas et al.*, 2018] and in the experiments, the average  
355 locus of debris-flow deposition shifts towards the topographically lower parts of a fan (Fig. 12).  
356 Moreover, in both our experiments and on natural debris-flow fans [*De Haas et al.*, 2018], avul-  
357 sion and new channel formation predominantly occur as a result of large flows, especially when  
358 these follow channel-plug formation in previous flows or when they occupy multiple channels

359 that can then become established as a new favorable pathway [Stoffel *et al.*, 2008; Imaizumi  
360 *et al.*, 2016]. By way of comparison, avulsions on fluvial fans are often triggered during high  
361 flood discharges [e.g., Brizga and Finlayson, 1990; Slingerland and Smith, 1998; Edmonds *et al.*,  
362 2009; Reitz *et al.*, 2010; Ganti *et al.*, 2016], and the avulsion frequency on deltas has been shown  
363 to increase with increasing discharge variability in the delta branches [Ganti *et al.*, 2014]. These  
364 common observations emphasize the link between the magnitude–frequency distribution of for-  
365 mative flows and avulsion frequency across a wide range of distributary systems.

366 It is important to temper our interpretations with several caveats. Our experiments show  
367 that avulsion behavior is highly sensitive to the sequence of flow sizes in a series of succes-  
368 sive events. A single magnitude-frequency distribution can, of course, yield contrasting sequences  
369 of flow sizes. We thus anticipate that repeating our experiments, randomly extracting flow mag-  
370 nitudes from a fixed distribution, may result in contrasting fan development. What is impor-  
371 tant here, however, is the way in which avulsions develop over short sequences of flows; we  
372 are interested in the underlying mechanisms by which avulsions occur, not in the final mor-  
373 phology of the experimental fans. Similar sequences should give rise to similar mechanisms,  
374 irrespective of the overall stage of fan development.

375 In addition, our analyses have so far focused only on the effects of varying debris-flow  
376 magnitudes on avulsion and fan evolution, while in reality the debris-flow composition (and  
377 thus bulk or macro-scale rheology) may also vary between flows [e.g., Suwa *et al.*, 2009; Okano  
378 *et al.*, 2012; De Haas *et al.*, 2015b]. For example, flows with a composition that renders them  
379 more immobile, such as low water content or high cobble to boulder fraction [e.g., De Haas  
380 *et al.*, 2015b], may be more likely to cause channel plugging and induce avulsion. Such be-  
381 havior has been documented on the Kamikamihori fan, where debris flows with a bouldery,  
382 matrix-poor, flow front were observed to be relatively immobile and therefore prone to deposit  
383 near the fan apex, jamming the channel and triggering avulsion in subsequent flows [Okuda  
384 *et al.*, 1981; Suwa and Yamakoshi, 1999; Suwa *et al.*, 2009]. The impact of debris-flow com-  
385 position on avulsion mechanisms remains an important avenue towards a better understand-  
386 ing of fan evolution.

#### 387 **4.2 Compensational tendency and implication for flow routing**

388 The compensation index increased from  $\sim 0.25$  at the fan apex to  $\sim 0.35$  near the fan toe  
389 on the three experimental debris-flow fans over temporal intervals of 1 to 20 flows (Fig. 10).

390 This suggests that debris-flow deposition at the scale of one to a few flows fell between anti-  
391 compensational and random tendencies, and that over such short time scales flows were likely  
392 to follow the existing topographic pathway. Depositional behavior became somewhat more com-  
393 pensational with increasing distance from the fan apex. Beyond a time scale of about 10-20  
394 events, however, avulsion (whether initiated by a channel plug or a backstepping sequence,  
395 as described above) often tended to redirect flow towards a topographically low area on the  
396 fan. *Straub et al.* [2009] showed that fan systems dominated by gradual lateral shifting of the  
397 depocentre, punctuated by infrequent large-scale avulsion to the absolute topographic low, tend  
398 towards a compensation index of  $\sim 0.3$ , similar to the values estimated here.

399 *De Haas et al.* [2018] observed that compensational behavior on natural debris-flow fans  
400 typically occurs only across sequences of flows, rather than between successive flows, again  
401 in agreement with our experimental observations. Debris flows on the Kamikamihori fan in  
402 Japan, for example, have been observed to occupy a channel for periods of a few flows un-  
403 til deposition shifts towards a topographically lower part on the fan [e.g., *Okuda et al.*, 1981;  
404 *Suwa and Okuda*, 1983; *Suwa and Yamakoshi*, 1999; *De Haas et al.*, 2018]. Similarly, such flow  
405 routing patterns are expressed in the surface topography of many modern debris-flow fans. We  
406 illustrate this behavior with a debris-flow fan from Saline Valley in the southwestern USA. The  
407 two most recent channel pathways on this fan have formed by persistent deposition over mul-  
408 tiple flows, as inferred from the surface morphology (Fig. 11a, b). This persistent depositional  
409 activity has led to channel superelevation of  $\sim 2-8$  m (see topographic profiles A-A' to D-D'  
410 in Figure 11), which substantially exceeds the  $\sim 1-3$  m channel depth. For comparison, *Mohrig*  
411 *et al.* [2000] found that fluvial channel levees rarely aggrade more than 0.6 times the chan-  
412 nel depth before avulsion. Thus, we tentatively hypothesize that debris-flow fans may be char-  
413 acterized by more persistent, anti-compensational depositional behavior compared to their flu-  
414 vial counterparts. On the other hand, the stochastic nature and formation of channel plugs can  
415 allow for rapid and unexpected channel shifting on debris-flow fans, without the need for widespread  
416 backfilling of accommodation, and enhancing the degree of compensation. On three debris-  
417 flow fans in Colorado, USA, *Pederson et al.* [2015] found intermediate to fully compensational  
418 stacking patterns (compensation indices  $\sim 0.6-1$ ). The compensational tendency of these fans  
419 appeared to increase with flow thickness, flow width, abundance of coarse clasts, percentage  
420 of clay and distance from the fan apex, all of which were inferred to increase the likelihood  
421 of avulsion out of the active channel. Our experiments are at least partly consistent with these  
422 observations, as avulsions were promoted by wider, thicker flows, and the experimental com-

423      pensation index increased toward the fan toes. The dependencies of avulsion behavior on both  
424      flow magnitude and composition make it difficult to attribute one mode of compensational be-  
425      havior to all debris-flow fans, and highlight the need for more systematic understanding of the  
426      link between flow characteristics and compensation.

#### 427           **4.3 Implications for debris-flow hazard mitigation**

428           Debris-flow avulsions are generally difficult to predict, and may therefore have substan-  
429      tial hazardous effects. The experimental results, however, provide some guidelines for iden-  
430      tifying and mitigating potential avulsions. Small to moderately-sized debris flows are unlikely  
431      to leave the main channel and cause avulsion, but they can be important avulsion precursors  
432      or triggers. Avulsion is highly likely to occur when the snouts of one or a series of small to  
433      moderately-sized flows have jammed the proximal channel. Avulsion can then be expected to  
434      occur just upstream of the channel plug in the next large debris flow. In addition, large flows  
435      have been observed to create levee breaches and potential new channel locations during over-  
436      bank surges, thus creating the template for an imminent avulsion in a large subsequent flow.  
437      In terms of hazard mitigation, it is thus important to check for potential channel plugs after  
438      small to moderate-sized flows, and for levee breaches and the onset of potential new channels  
439      after large events. The tendency of avulsions to re-occupy older channels on the flow surface,  
440      and thus for channels to act as flow attractors, should also be considered when assessing po-  
441      tential debris-flow hazard.

442           Our experimental observation that plugging and backstepping are favored during small-  
443      to medium-sized flows, while avulsion subsequently occurs during the next large flow, sup-  
444      ports the hypothesis of *De Haas et al.* [2018] that there may be an optimal magnitude-frequency  
445      distribution for which avulsion frequency is maximized. For hazard mitigation it is important  
446      to understand which types of magnitude-frequency distribution result in a high avulsion fre-  
447      quency. Although still far from a definitive answer, our experimental results suggest that such  
448      a favorable distribution likely includes sufficient small- to moderately-sized debris flows to cause  
449      channel plugs and induce backstepping sequences, but also sufficient large flows to enable the  
450      formation of new pathways. Systems in which large flows are relatively abundant may be less  
451      prone to avulsion because of the paucity of smaller, plug-forming flows and the tendency of  
452      large flows to entrain material and thus enlarge the main channel as they transit the fan [*Schürch*  
453      *et al.*, 2011]. On the other hand, a proportional deficiency of large flows may also result in  
454      a lower avulsion frequency, because there are fewer flows with sufficient size to leave the main

455 channel and form a new channel. At present, we lack the data on flow magnitude-frequency  
456 distributions from natural debris-flow fans with which to test these ideas, but such distribu-  
457 tions would be an important research target in the near future.

## 458 **5 Conclusions**

459 This paper investigates how patterns of debris-flow fan avulsion and evolution are af-  
460 fected by the magnitude-frequency distribution of the flows. We compared the topographic evo-  
461 lution, avulsion mechanisms, and compensational tendencies, of three experimental fans formed  
462 by contrasting flow magnitude-frequency distributions: (1) a uniform distribution, (2) a steep  
463 double-Pareto distribution with many flows around the mean and a limited number of large  
464 flows, and (3) a shallow double-Pareto distribution with fewer flows around the mean and more  
465 abundant large flows.

466 The three experimental fans followed similar overall patterns of development, evolving  
467 through alternating channelized and unchannelized phases that were governed by sequences  
468 of backstepping deposition and avulsion. In detail, however, the differences in magnitude-frequency  
469 distribution also caused marked differences in the avulsion mechanisms, and thus surface evo-  
470 lution, of the three fans. The fan formed by uniform flows developed through regular sequences  
471 of channelization, backstepping from fan toe to fan apex, and avulsion over multiple flows.  
472 In contrast, large debris flows on the fans formed by a double-Pareto distribution were observed  
473 to overtop the active channel and carve new flow paths through the channel levees, at times  
474 initiating avulsion within a single event. On these fans, avulsions were also triggered by se-  
475 ries of similarly-sized or progressively smaller flows which plugged the active channel, lead-  
476 ing to avulsion in the next large flow. This latter mechanism was far more common on the fan  
477 formed by a steep double-Pareto distribution, which we attribute to both (1) the predominance  
478 of moderate flows on this fan and (2) the relative abundance of large flows on the fan formed  
479 by the shallow double-Pareto distribution that kept the channel clear. On all experimental fans,  
480 backstepping sequences were either initiated after filling of regional accommodation or plug  
481 formation, and the relative importance of these processes largely depended on the debris-flow  
482 magnitude-frequency distribution.

483 In short, channel plugs were more likely to be formed by small- to moderately-sized flows,  
484 whereas large flows were more prone to leave the main channel and initiate or exploit a new  
485 pathway down the fan. We infer that there is likely to be an optimal magnitude-frequency dis-

486 tribution that maximizes the avulsion frequency, reflecting a balance between small- to medium-  
487 sized flows that can plug the channel and induce backstepping, and large flows that subsequently  
488 avulse out of the main channel.

489 Our results provide some guidelines for avulsion hazard mitigation; sequences of small-  
490 to moderately-sized flows, especially those that deposit material within the active channel, may  
491 serve as precursors to avulsion on natural fans. Similarly, large flows that cause levee breaches  
492 and incipient development of new channel pathways should also be treated as avulsion pre-  
493 cursors.

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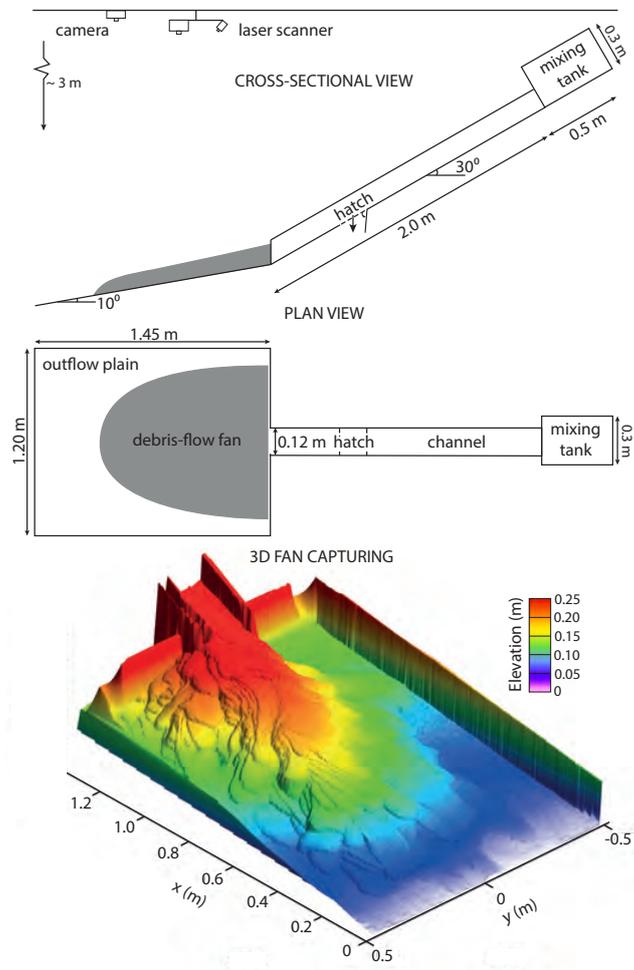


Figure 1: Experimental flume setup. The 3D fan image shows the final morphology of fan 01. The flume setup is similar to that used in *De Haas et al.* [2015b] and *De Haas et al.* [2016]. Figure modified from *De Haas et al.* [2016].

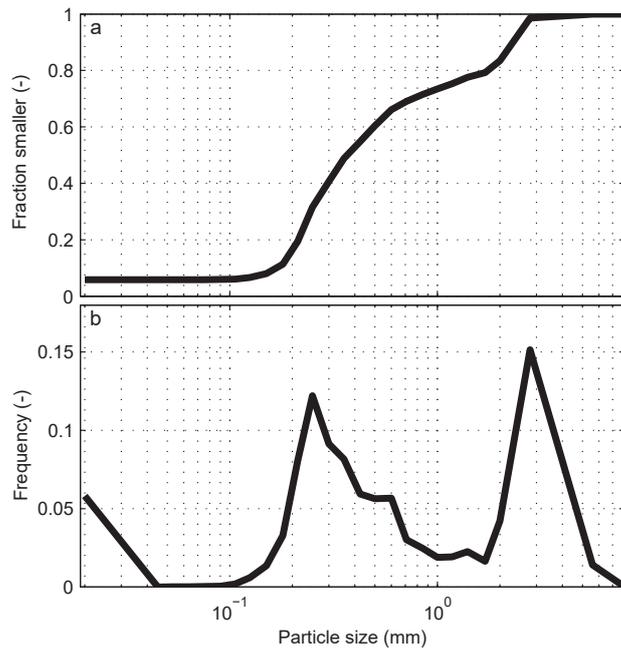


Figure 2: Particle-size distribution of the debris flows forming the experimental fans. (a) Cumulative distribution. (b) Frequency distribution.

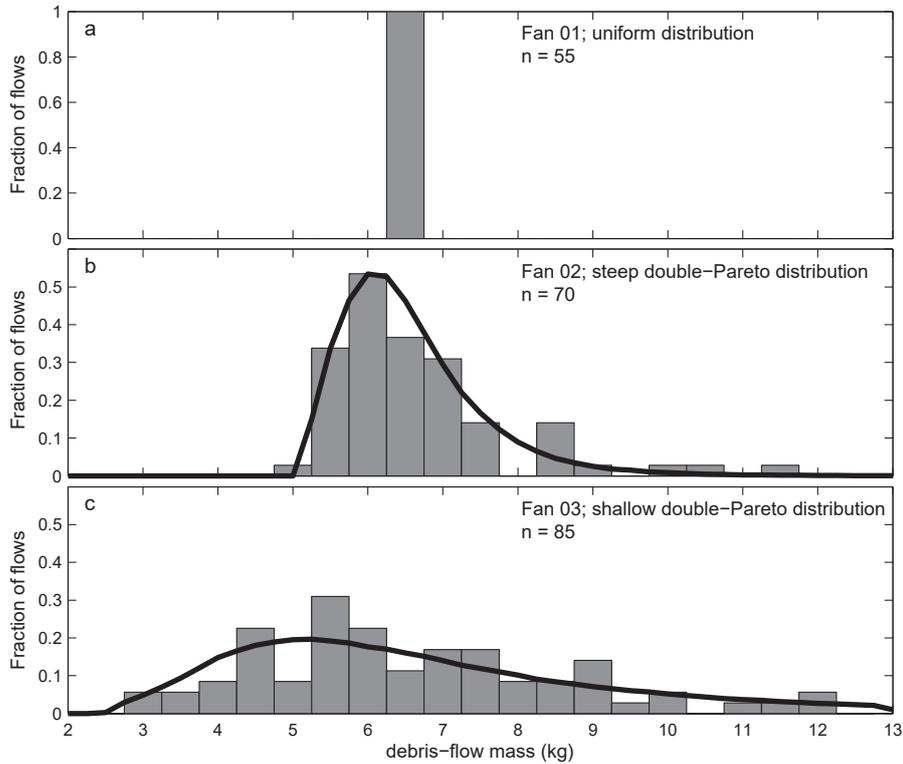


Figure 3: Magnitude-frequency distributions of the experimental debris-flow fans. The lines denote the double-Pareto distributions from which the debris-flow magnitudes were randomly extracted. The bars denote the actual number of events in each experiment, divided into 0.5 kg bins. The mean debris-flow mass is  $\sim 6.5$  kg for all experiments. (a) Fan 01 with a uniform distribution; (b) fan 02 with a steep-tailed double-Pareto distribution; (c) fan 03 with a shallow-tailed double-Pareto distribution.

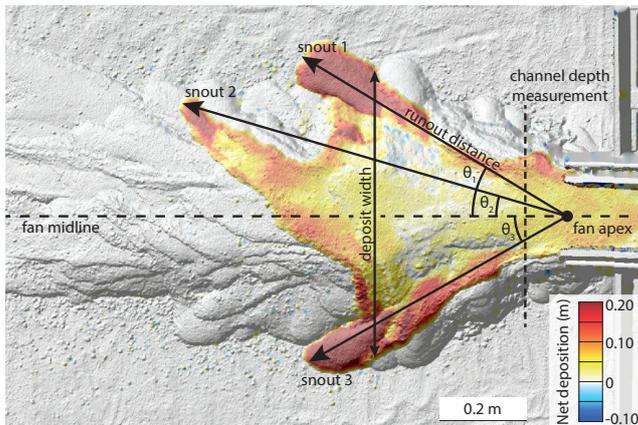


Figure 4: Depositional geometry and measurement of spatio-temporal patterns of debris-flow activity. For each snout, the runout distance is the distance between snout front and fan apex, and the flow angle is the angle between the fan midline and a straight line between debris-flow snout and apex. Deposit width is defined as the maximum width of the deposit, excluding individual snouts substantially outside of the main flow direction. Channel depth is measured 10 cm downstream of the fan apex.

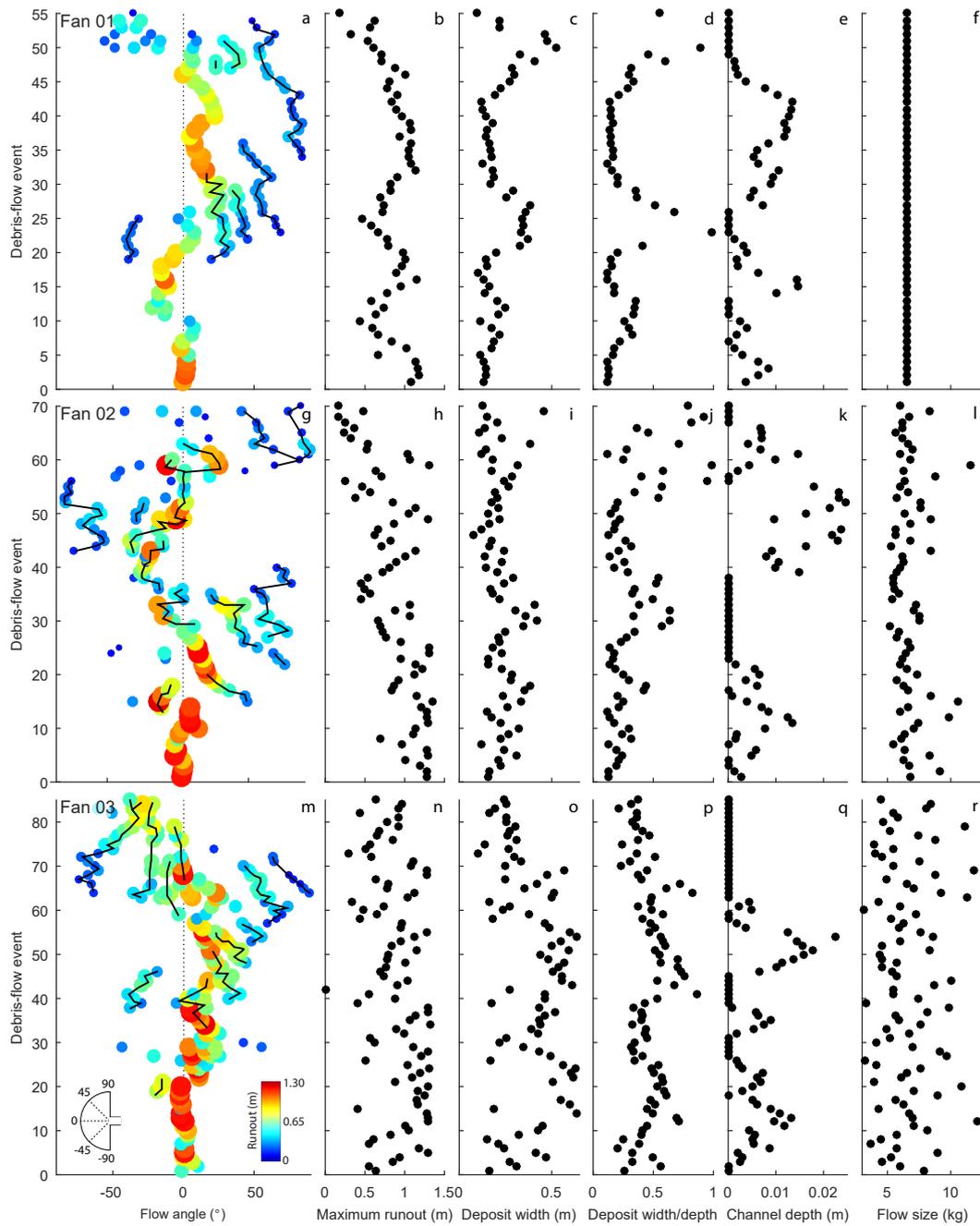


Figure 5: Summary of the spatio-temporal patterns of debris-flow activity on fans 01-03. (a) Flow angle and runout distance of the channels active during the debris-flow events that formed fan 01. Flow angle was previously published in *De Haas et al.* [2016]; the other variables are newly reported here. Solid line segments join successive flows in the same channel. (b) Maximum runout distance during each debris flow on fan 01. (c) Deposit width during each event on fan 01. (d) Deposit width/depth ratio for each debris flow on fan 01, defined as deposit width divided by maximum runout distance. (e) Channel depth after each debris-flow event on fan 01, measured 10 cm downstream of the fan apex. (f) Debris-flow mass in kg. (g-l) As above for fan 02. (m-r) As above for fan 03.

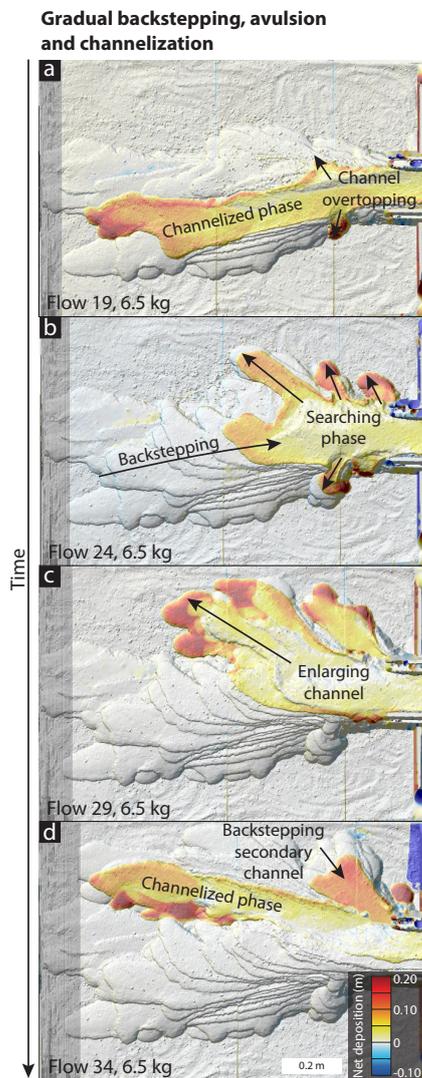


Figure 6: Typical avulsion sequence on fan 01, formed by a uniform flow magnitude-frequency distribution (Fig. 3a). The sequence shows evolution from a well-defined channel (panel a) through gradual backstepping of deposition toward the fan apex, followed by a searching phase (panel b), avulsion to a new channel pathway and enlargement (panel c), and channelization (panel d). Note that a secondary channel formed but was plugged and abandoned during this sequence.

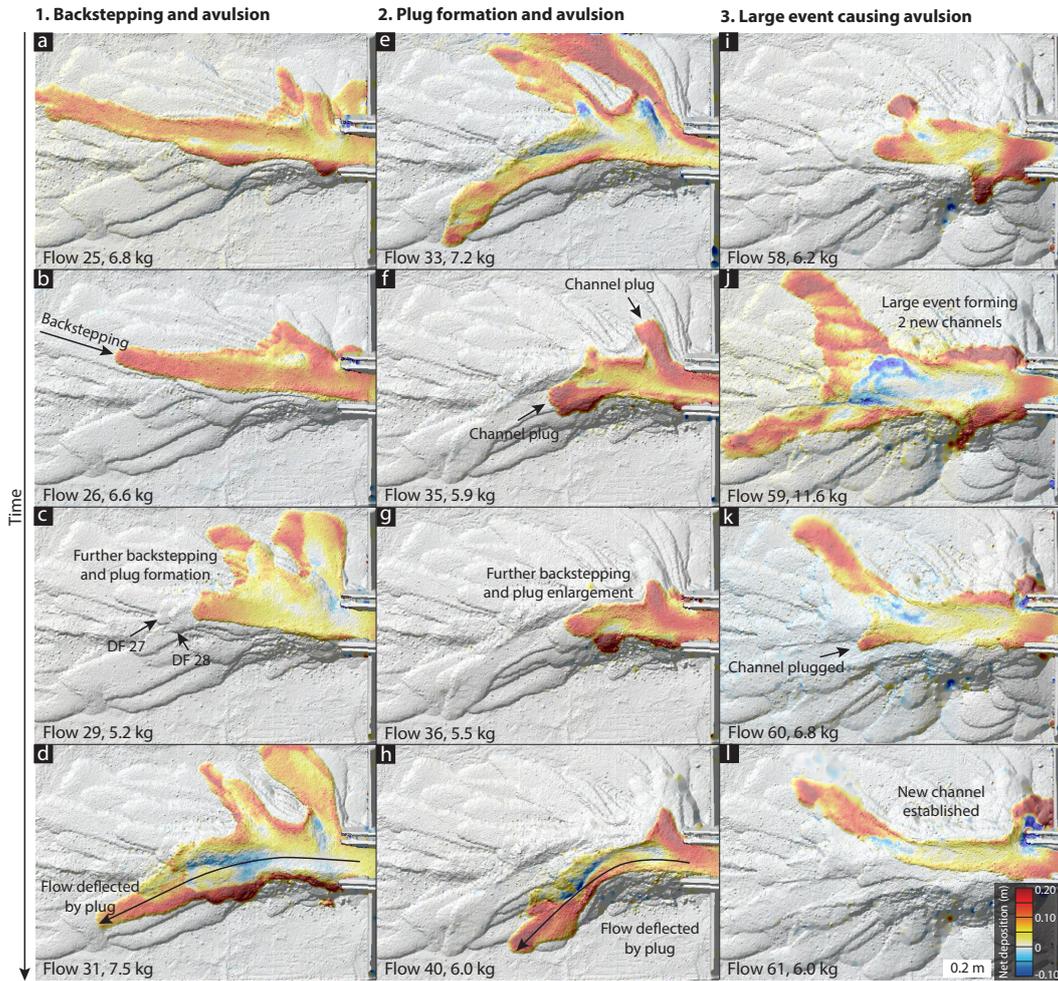


Figure 7: Common debris-flow-magnitude sequences leading to avulsion on fan 02, formed by a steep-tailed double-Pareto flow distribution (Fig. 3b). (a-d) Backstepping and avulsion sequence during debris flows 25-31. A sequence of small- to moderately-sized flows induced a sequence of backstepping deposition (panels b to c), which was followed by avulsion during the large flow 31 (panel d). (e-h) Channel plug formation by two small debris flows that blocked the main channel (panels f and g), followed by avulsion during a moderately-sized flow (panel h). (i-l) A very large debris flow created two new channels (panel j), one of which became blocked by a flow snout in the next, smaller flow (panel k). Avulsion then proceeded into the topographically-favored right-hand channel (panel l).

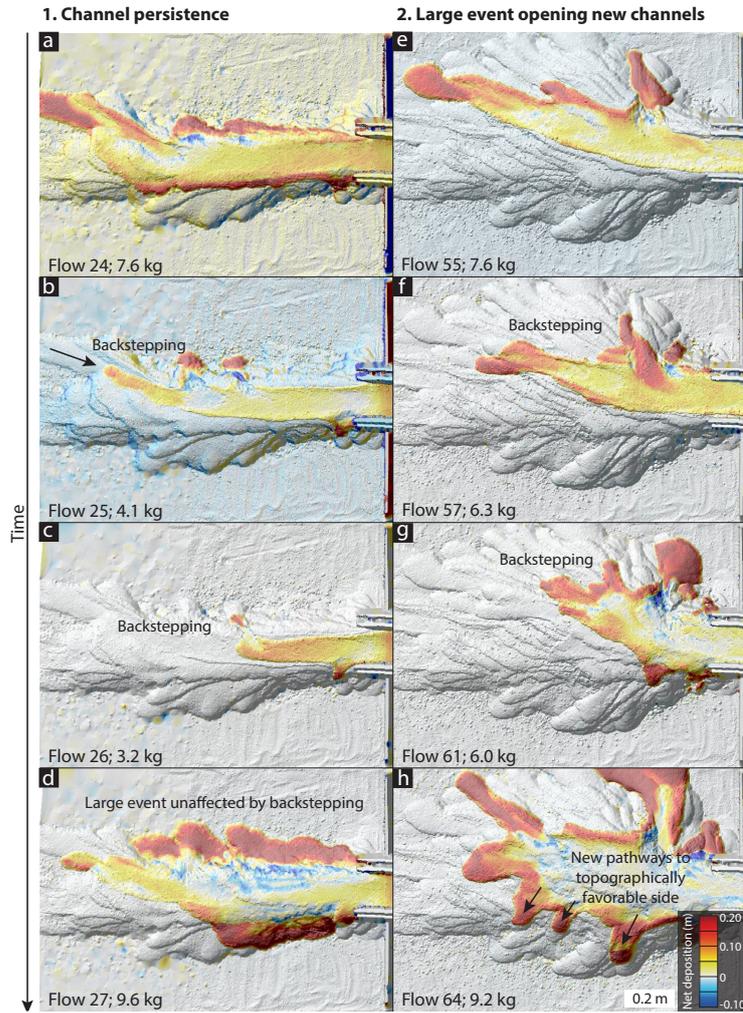


Figure 8: Common spatio-temporal patterns of debris-flow activity on fan 03, formed by a shallow-tailed double-Pareto flow distribution (Fig. 3c). (a-d) Persistent channel position during debris flows 24-27. The direction of the large debris flow 27 (panel d) was unaffected by the backstepping plug deposits from the small flows 25 and 26 (panels b and c). Flow 27 partly eroded the channel plug and filled the channel surrounding the plug. Overbank surges were widespread during the relatively large flows 24 and 27 (panels a and d). (e-h) After a partial backstepping sequence from debris flow 55 to 62 (panels e-g), large flows 63 and 64 opened up three new channel pathways on the left side of the fan (panel h, shown by arrows) that allowed subsequent avulsion towards the left.

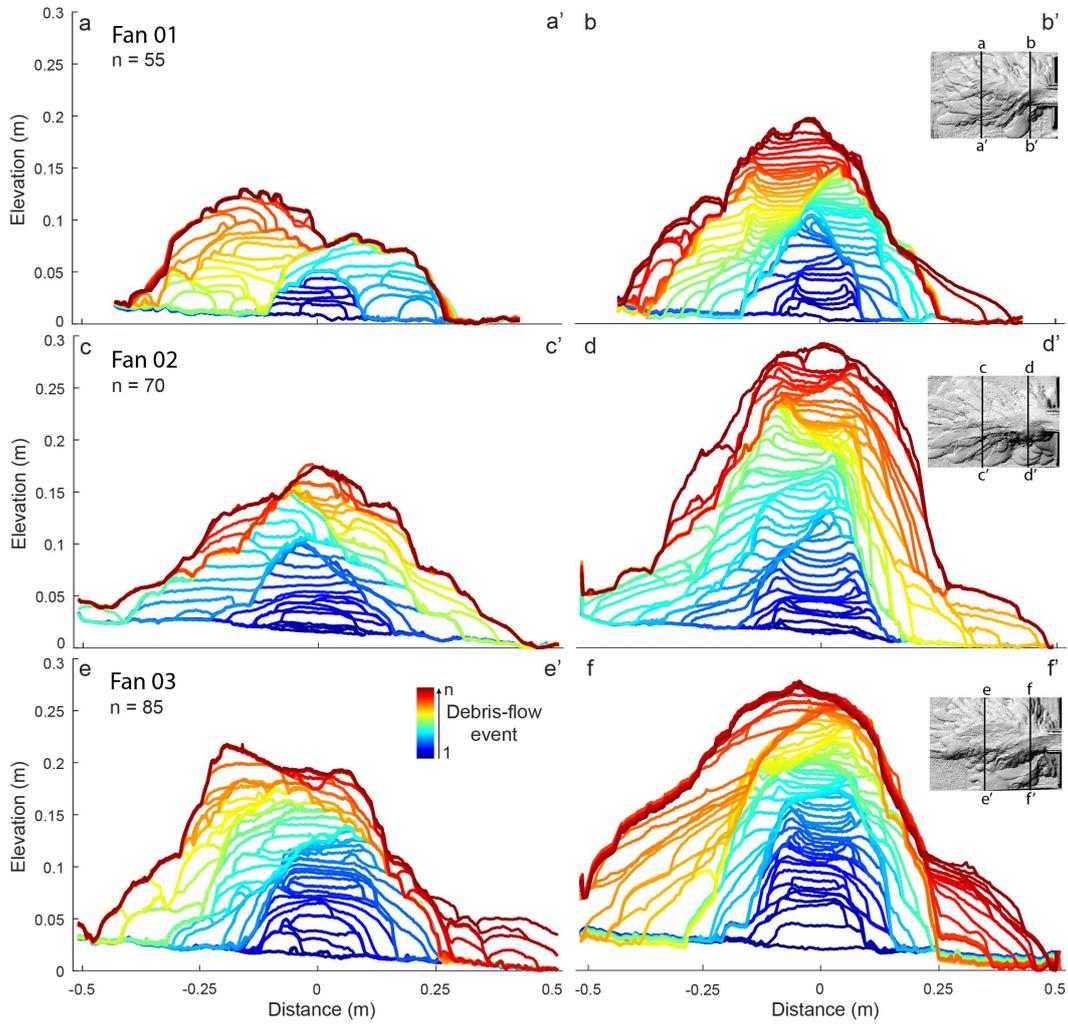


Figure 9: Cross-profiles through the experimental debris-flow fans at distances of 0.2 m (left-hand column) and 0.8 m (right-hand column) downstream of the fan apex. Colors show progressive flow sequence from cool to warm. (a-b) Fan 01, previously shown in *De Haas et al.* [2016]. (c-d) Fan 02. (e-f) Fan 03. Note how overbank deposition became increasingly important for fan construction and how large lateral shifts became less pronounced with increasing flow-magnitude variability from fan 01 to 03.

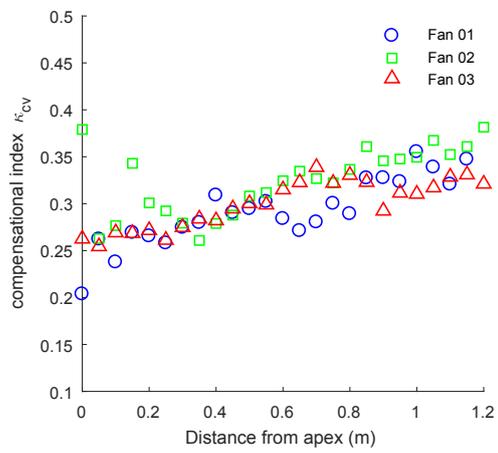


Figure 10: Compensation index from fan apex (left) to fan toe (right) for fans 01-03.

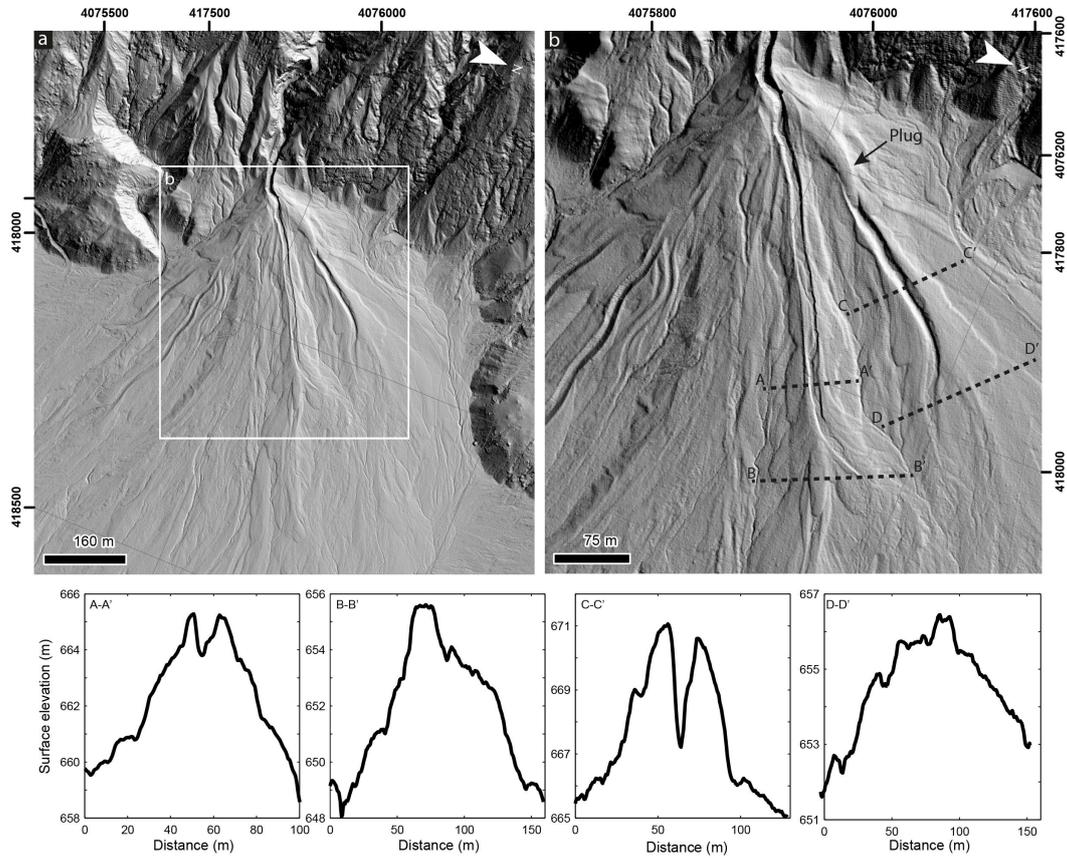


Figure 11: Channel superelevation on a debris-flow fan in southern Saline Valley, California, USA. Panels (a) and (b) show present-day topography of the fan surface. Data are from the EarthScope Southern & Eastern California Lidar Project ([www.opentopography.org](http://www.opentopography.org)) with a cell size of 0.5 m. The cross-sections below show the absolute superelevation of the two most recently-active channels on the fan. This example shows how deposits can act as attractors, mainly due to the presence of an incised apex channel, leading to superelevation of 2-8 m. Coordinates in UTM WGS 1984 11N.

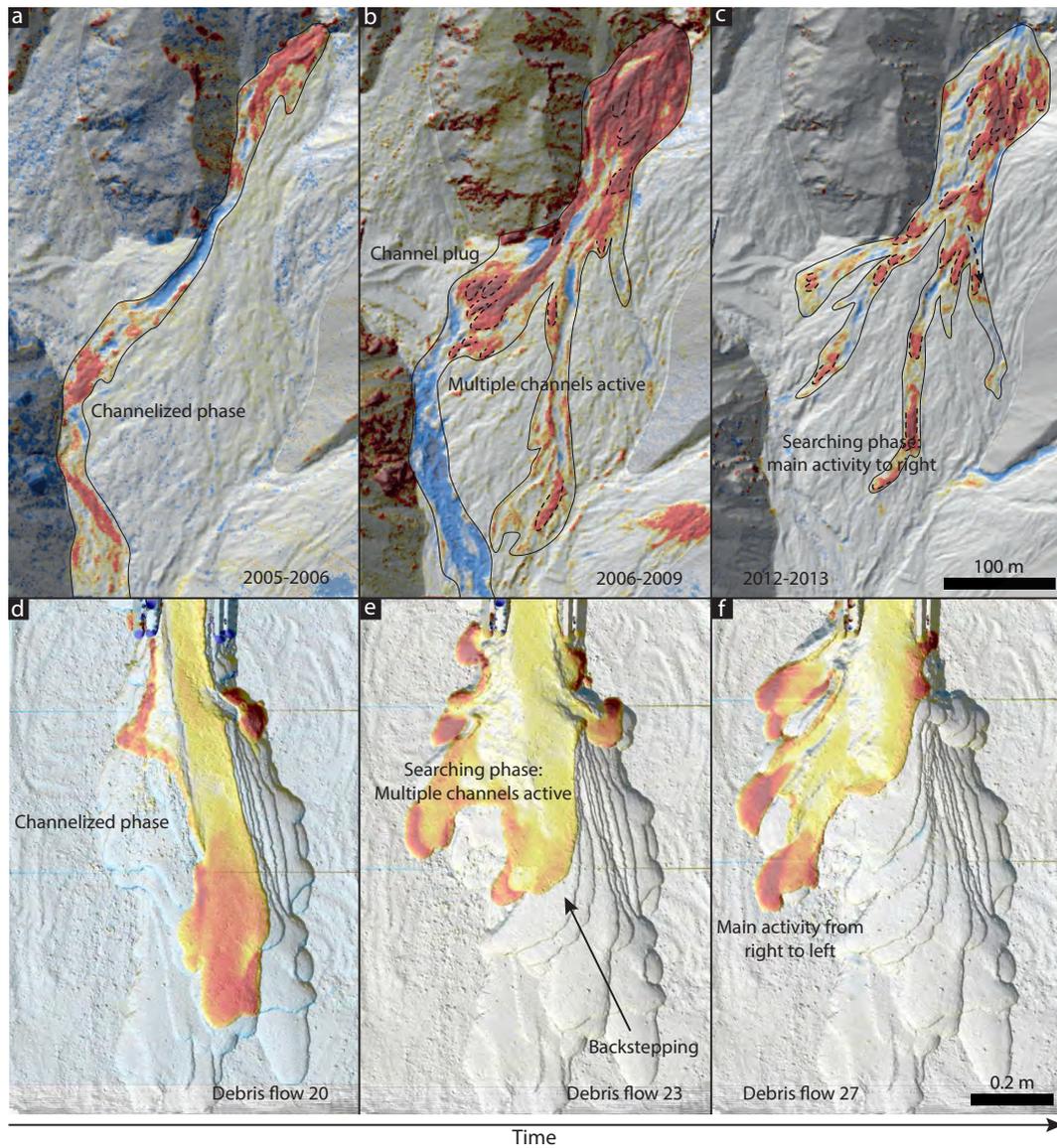


Figure 12: Examples of the transition from channelized to searching phases on (a-c) the Ohya debris-flow fan in Japan [images modified from *Imaizumi et al., 2016; De Haas et al., 2018*] and (d-f) experimental debris-flow fan 01. Flow in all panels was from top to bottom. On both the natural and experimental fans, activity during the searching phase was spread over multiple channels on the proximal fan, and the locus of activity shifted laterally across the fan over multiple debris flows. Warm colors indicate deposition and cool colors indicate erosion, although absolute scales differ.

501 **References**

- 502 Aslan, A., W. J. Autin, and M. D. Blum (2005), Causes of river avulsion: insights from  
503 the late Holocene avulsion history of the Mississippi River, USA, *Journal of Sedimen-*  
504 *tary Research*, 75(4), 650–664.
- 505 Bardou, E., and M. Jaboyedoff (2008), Debris flows as a factor of hillslope evolution  
506 controlled by a continuous or a pulse process?, *Geological Society, London, Special*  
507 *Publications*, 296(1), 63–78.
- 508 Beaty, C. B. (1963), Origin of alluvial fans, White Mountains, California and Nevada,  
509 *Ann. Assoc. Am. Geogr.*, 53(4), 516–535.
- 510 Bennett, G., P. Molnar, H. Eisenbeiss, and B. McArdell (2012), Erosional power in the  
511 Swiss Alps: characterization of slope failure in the Illgraben, *Earth Surface Processes*  
512 *and Landforms*, 37(15), 1627–1640.
- 513 Bennett, G., P. Molnar, B. McArdell, F. Schlunegger, and P. Burlando (2013), Patterns and  
514 controls of sediment production, transfer and yield in the Illgraben, *Geomorphology*,  
515 188, 68–82.
- 516 Bennett, G., P. Molnar, B. McArdell, and P. Burlando (2014), A probabilistic sediment  
517 cascade model of sediment transfer in the Illgraben, *Water Resources Research*, 50(2),  
518 1225–1244.
- 519 Blair, T. C., and J. G. McPherson (1994), Alluvial fans and their natural distinction from  
520 rivers based on morphology, hydraulic processes, sedimentary processes, and facies  
521 assemblages, *Journal of Sedimentary Research*, 64A, 450–489.
- 522 Blair, T. C., and J. G. McPherson (2009), Processes and forms of alluvial fans, in *Ge-*  
523 *omorphology of Desert Environments*, edited by A. Parsons and A. Abrahams, pp.  
524 413–467, Springer Netherlands.
- 525 Bollschweiler, M., M. Stoffel, and D. M. Schneuwly (2008), Dynamics in debris-flow  
526 activity on a forested cone - a case study using different dendroecological approaches,  
527 *Catena*, 72(1), 67–78.
- 528 Brizga, S. O., and B. L. Finlayson (1990), Channel avulsion and river metamorphosis: the  
529 case of the Thomson River, Victoria, Australia, *Earth Surface Processes and Landforms*,  
530 15(5), 391–404.
- 531 Cannon, S. H., J. E. Gartner, R. C. Wilson, J. C. Bowers, and J. L. Laber (2008), Storm  
532 rainfall conditions for floods and debris flows from recently burned areas in southwest-  
533 ern Colorado and southern California, *Geomorphology*, 96(3), 250–269.

- 534 Cannon, S. H., E. M. Boldt, J. L. Laber, J. W. Kean, and D. M. Staley (2011), Rainfall  
535 intensity–duration thresholds for postfire debris-flow emergency-response planning,  
536 *Natural Hazards*, 59(1), 209–236.
- 537 Clague, J. J., C. Huggel, O. Korup, and B. McGuire (2012), Climate change and haz-  
538 ardous processes in high mountain, *Revista de la Asociación Geológica Argentina*, 69(3),  
539 328–338.
- 540 Clarke, L., T. A. Quine, and A. Nicholas (2010), An experimental investigation of auto-  
541 genic behaviour during alluvial fan evolution, *Geomorphology*, 115(3), 278–285.
- 542 Dai, F., and C. Lee (2001), Frequency–volume relation and prediction of rainfall-induced  
543 landslides, *Engineering geology*, 59(3), 253–266.
- 544 D’Arcy, M., D. C. Roda Boluda, A. C. Whittaker, and A. Carpineti (2015), Dating alluvial  
545 fan surfaces in Owens Valley, California, using weathering fractures in boulders, *Earth  
546 Surface Processes and Landforms*, 40(4), 487–501.
- 547 De Haas, T., and T. Van Woerkom (2016), Bed scour by debris flows: experimental inves-  
548 tigation of effects of debris-flow composition, *Earth Surface Processes and Landforms*,  
549 41(13), 1951–1966, doi:10.1002/esp.3963, esp.3963.
- 550 De Haas, T., D. Ventra, P. E. Carbonneau, and M. G. Kleinhans (2014), Debris-flow dom-  
551 inance of alluvial fans masked by runoff reworking and weathering, *Geomorphology*,  
552 217, 165–181.
- 553 De Haas, T., M. G. Kleinhans, P. E. Carbonneau, L. Rubensdotter, and E. Hauber (2015a),  
554 Surface morphology of fans in the high-Arctic periglacial environment of Svalbard:  
555 Controls and processes, *Earth-Science Reviews*, 146, 163–182.
- 556 De Haas, T., L. Braat, J. F. W. Leuven, I. R. Lokhorst, and M. G. Kleinhans (2015b),  
557 Effects of debris-flow composition and topography on runout distance, depositional  
558 mechanisms and deposit morphology, *Journal of Geophysical Research: Earth Surface*,  
559 120, 1949–1972.
- 560 De Haas, T., W. Berg, L. Braat, and M. G. Kleinhans (2016), Autogenic avulsion, chan-  
561 nelization and backfilling dynamics of debris-flow fans, *Sedimentology*, 63, 1596–1619.
- 562 De Haas, T., A. L. Densmore, M. Stoffel, H. Suwa, F. Imaizumi, J. A. Ballesteros-  
563 Cánovas, and T. Wasklewicz (2018), Avulsions and the spatio-temporal evolution of  
564 debris-flow fans, *Earth Science Reviews*, 177, 53–75, doi:10.1016/j.earscirev.2017.11.  
565 007.

- 566 Dühnforth, M., A. L. Densmore, S. Ivy-Ochs, P. A. Allen, and P. W. Kubik (2007), Tim-  
567 ing and patterns of debris flow deposition on Shepherd and Symmes creek fans, Owens  
568 Valley, California, deduced from cosmogenic  $^{10}\text{Be}$ , *Journal of Geophysical Research:*  
569 *Earth Surface (2003–2012)*, 112(F3).
- 570 Dühnforth, M., A. L. Densmore, S. Ivy-Ochs, and P. A. Allen (2008), Controls on sed-  
571 iment evacuation from glacially modified and unmodified catchments in the eastern  
572 Sierra Nevada, California, *Earth Surface Processes and Landforms*, 33(10), 1602–1613.
- 573 Edmonds, D. A., D. C. Hoyal, B. A. Sheets, and R. L. Slingerland (2009), Predicting  
574 delta avulsions: Implications for coastal wetland restoration, *Geology*, 37(8), 759–762.
- 575 Ganti, V., Z. Chu, M. P. Lamb, J. A. Nittrouer, and G. Parker (2014), Testing morphody-  
576 namic controls on the location and frequency of river avulsions on fans versus deltas:  
577 Huanghe (Yellow River), China, *Geophysical Research Letters*, 41(22), 7882–7890.
- 578 Ganti, V., A. J. Chadwick, H. J. Hassenruck-Gudipati, and M. P. Lamb (2016), Avulsion  
579 cycles and their stratigraphic signature on an experimental backwater-controlled delta,  
580 *Journal of Geophysical Research: Earth Surface*, 121(9), 1651–1675.
- 581 Guthrie, R., and S. Evans (2004), Analysis of landslide frequencies and characteristics  
582 in a natural system, coastal British Columbia, *Earth Surface Processes and Landforms*,  
583 29(11), 1321–1339.
- 584 Hamilton, P. B., K. Strom, and D. C. Hoyal (2013), Autogenic incision-backfilling cycles  
585 and lobe formation during the growth of alluvial fans with supercritical distributaries,  
586 *Sedimentology*, 60(6), 1498–1525.
- 587 Harvey, A. (2011), Dryland alluvial fans, *Arid Zone Geomorphology: Process, Form and*  
588 *Change in Drylands, Third Edition*, pp. 333–371.
- 589 Helsen, M. M., P. J. M. Koop, and H. Van Steijn (2002), Magnitude-frequency relation-  
590 ship for debris flows on the fan of the Chalance torrent, Valgaudemar (French Alps),  
591 *Earth Surface Processes and Landforms*, 27(12), 1299–1307.
- 592 Hoefling, R. (2004), High-speed 3D imaging by DMD technology, in *Electronic Imaging*  
593 *2004*, pp. 188–194, International Society for Optics and Photonics.
- 594 Hooke, R. B., and W. L. Rohrer (1979), Geometry of alluvial fans: Effect of discharge  
595 and sediment size, *Earth Surface Processes*, 4(2), 147–166.
- 596 Hooke, R. L. (1967), Processes on arid-region alluvial fans, *The Journal of Geology*, pp.  
597 438–460.

- 598 Hovius, N., C. P. Stark, and P. A. Allen (1997), Sediment flux from a mountain belt de-  
599 rived by landslide mapping, *Geology*, 25(3), 231–234.
- 600 Huang, R., and X. Fan (2013), The landslide story, *Nature Geoscience*, 6(5), 325–326.
- 601 Hungr, O., S. Evans, and J. Hazzard (1999), Magnitude and frequency of rock falls and  
602 rock slides along the main transportation corridors of southwestern British Columbia,  
603 *Canadian Geotechnical Journal*, 36(2), 224–238.
- 604 Imaizumi, F., D. Trappmann, N. Matsuoka, S. Tsuchiya, O. Ohsaka, and M. Stoffel  
605 (2016), Biographical sketch of a giant: Deciphering recent debris-flow dynamics from  
606 the Ohya landslide body (Japanese Alps), *Geomorphology*, 272, 102–114.
- 607 Iverson, R. M. (1997), The physics of debris flows, *Reviews of Geophysics*, 35(3), 245–  
608 296.
- 609 Iverson, R. M., M. E. Reid, M. Logan, R. G. LaHusen, J. W. Godt, and J. P. Griswold  
610 (2011), Positive feedback and momentum growth during debris-flow entrainment of wet  
611 bed sediment, *Nature Geoscience*, 4(2), 116–121.
- 612 Ma, C., Y. Wang, K. Hu, C. Du, and W. Yang (2017), Rainfall intensity–duration  
613 threshold and erosion competence of debris flows in four areas affected by the 2008  
614 Wenchuan earthquake, *Geomorphology*, 282, 85–95.
- 615 Malamud, B. D., D. L. Turcotte, F. Guzzetti, and P. Reichenbach (2004), Landslide in-  
616 ventories and their statistical properties, *Earth Surface Processes and Landforms*, 29(6),  
617 687–711.
- 618 Mohrig, D., P. L. Heller, C. Paola, and W. J. Lyons (2000), Interpreting avulsion process  
619 from ancient alluvial sequences: Guadalope-matarranya system (northern Spain) and  
620 wasatch formation (western Colorado), *Geological Society of America Bulletin*, 112(12),  
621 1787–1803.
- 622 Okano, K., H. Suwa, and T. Kanno (2012), Characterization of debris flows by rainstorm  
623 condition at a torrent on the Mount Yakedake volcano, Japan, *Geomorphology*, 136(1),  
624 88–94.
- 625 Okuda, S., S. Suwa, K. Okunishi, and K. Yokoyama (1981), Depositional processes of  
626 debris flow at Kamikamihori fan, Northern Japan Alps, *Trans. Japan. Geomorph. Union*,  
627 2, 353–361.
- 628 Paola, C., K. Straub, D. Mohrig, and L. Reinhardt (2009), The unreasonable effectiveness  
629 of stratigraphic and geomorphic experiments, *Earth-Science Reviews*, 97(1), 1–43.

- 630 Pederson, C. A., P. M. Santi, and D. R. Pyles (2015), Relating the compensational stack-  
631 ing of debris-flow fans to characteristics of their underlying stratigraphy: Implications  
632 for geologic hazard assessment and mitigation, *Geomorphology*, 248, 47–56.
- 633 Powell, E. J., W. Kim, and T. Muto (2012), Varying discharge controls on timescales of  
634 autogenic storage and release processes in fluvio-deltaic environments: Tank experi-  
635 ments, *Journal of Geophysical Research: Earth Surface*, 117(F02011).
- 636 Rebetez, M., R. Lugon, and P.-A. Baeriswyl (1997), Climatic change and debris flows in  
637 high mountain regions: the case study of the Ritigraben torrent (Swiss Alps), *Climatic  
638 change*, 36(3-4), 371–389.
- 639 Reed, W. J. (2001), The pareto, zipf and other power laws, *Economics Letters*, 74(1),  
640 15–19.
- 641 Reed, W. J., and M. Jorgensen (2004), The double pareto-lognormal distribution a new  
642 parametric model for size distributions, *Communications in Statistics-Theory and Meth-  
643 ods*, 33(8), 1733–1753.
- 644 Reitz, M. D., D. J. Jerolmack, and J. B. Swenson (2010), Flooding and flow path selection  
645 on alluvial fans and deltas, *Geophysical Research Letters*, 37(6), L06,401.
- 646 Schumm, S., M. Mosley, and W. Weaver (1987), *Experimental Fluvial Geomorphology*,  
647 John Wiley and Sons, New York.
- 648 Schürch, P. (2011), Debris-flow erosion and deposition dynamics, Ph.D. thesis, Durham  
649 University.
- 650 Schürch, P., A. L. Densmore, N. J. Rosser, and B. W. McArdell (2011), Dynamic controls  
651 on erosion and deposition on debris-flow fans, *Geology*, 39(9), 827–830.
- 652 Schürch, P., A. L. Densmore, S. Ivy-Ochs, N. J. Rosser, F. Kober, F. Schlunegger,  
653 B. McArdell, and V. Alfimov (2016), Quantitative reconstruction of late Holocene  
654 surface evolution on an alpine debris-flow fan, *Geomorphology*, 275, 46–57.
- 655 Shieh, C.-L., Y. Chen, Y. Tsai, and J. Wu (2009), Variability in rainfall threshold for de-  
656bris flow after the Chi-Chi earthquake in central Taiwan, China, *International Journal of  
657 Sediment Research*, 24(2), 177–188.
- 658 Slingerland, R., and N. D. Smith (1998), Necessary conditions for a meandering-river  
659 avulsion, *Geology*, 26(5), 435–438.
- 660 Stark, C. P., and N. Hovius (2001), The characterization of landslide size distributions,  
661 *Geophysical Research Letters*, 28(6), 1091–1094.

- 662 Stoffel, M. (2010), Magnitude–frequency relationships of debris flows—a case study based  
663 on field surveys and tree-ring records, *Geomorphology*, 116(1), 67–76.
- 664 Stoffel, M., D. Conus, M. A. Grichting, I. Lièvre, and G. Maître (2008), Unraveling the  
665 patterns of late Holocene debris-flow activity on a cone in the Swiss Alps: chronol-  
666 ogy, environment and implications for the future, *Global and Planetary Change*, 60(3),  
667 222–234.
- 668 Straub, K. M., and D. R. Pyles (2012), Quantifying the hierarchical organization of com-  
669 pensation in submarine fans using surface statistics, *Journal of Sedimentary Research*,  
670 82(11), 889–898.
- 671 Straub, K. M., C. Paola, D. Mohrig, M. A. Wolinsky, and T. George (2009), Compensa-  
672 tional stacking of channelized sedimentary deposits, *Journal of Sedimentary Research*,  
673 79(9), 673–688.
- 674 Suwa, H., and S. Okuda (1983), Deposition of debris flows on a fan surface, Mt.  
675 Yakedake, Japan, *Zeitschrift fur Geomorphologie NF Supplementband*, 46, 79–101.
- 676 Suwa, H., and T. Yamakoshi (1999), Sediment discharge by storm runoff at volcanic tor-  
677 rents affected by eruption, *Zeitschrift fur Geomorphologie NF Supplementband*, 114,  
678 63–88.
- 679 Suwa, H., K. Okano, and T. Kanno (2009), Behavior of debris flows monitored on test  
680 slopes of Kamikamihorizawa Creek, Mount Yakedake, Japan, *International Journal of*  
681 *Erosion Control Engineering*, 2(2), 33–45.
- 682 Van Dijk, M., G. Postma, and M. G. Kleinhans (2009), Autocyclic behaviour of fan  
683 deltas: an analogue experimental study, *Sedimentology*, 56(5), 1569–1589.
- 684 Van Dijk, M., M. G. Kleinhans, G. Postma, and E. Kraal (2012), Contrasting morphody-  
685 namics in alluvial fans and fan deltas: effect of the downstream boundary, *Sedimentol-*  
686 *ogy*, 59(7), 2125–2145.
- 687 Wang, Y., K. M. Straub, and E. A. Hajek (2011), Scale-dependent compensational stack-  
688 ing: an estimate of autogenic time scales in channelized sedimentary deposits, *Geology*,  
689 39(9), 811–814.
- 690 Wasklewicz, T., and C. Scheinert (2016), Development and maintenance of a telescoping  
691 debris flow fan in response to human-induced fan surface channelization, Chalk Creek  
692 Valley Natural Debris Flow Laboratory, Colorado, USA, *Geomorphology*, 252, 51–65.
- 693 Whipple, K. X., and T. Dunne (1992), The influence of debris-flow rheology on fan mor-  
694 phology, Owens Valley, California, *Geological Society of America Bulletin*, 104(7),

695 887–900.

696 Zaginaev, V., J. Ballesteros-Cánovas, S. Erokhin, E. Matov, D. Petrakov, and M. Stoffel  
697 (2016), Reconstruction of glacial lake outburst floods in northern Tien Shan: Implica-  
698 tions for hazard assessment, *Geomorphology*, 269, 75–84.