1 Crystallisation in basaltic magmas revealed via *in situ* 4D synchrotron X-ray 2 microtomography

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Magma crystallisation is a fundamental process driving eruptions and controlling the style of 15 16 volcanic activity. Crystal nucleation delay, heterogeneous and homogeneous nucleation and 17 crystal growth are all time-dependent processes, however, there is a paucity of real-time experimental data on crystal nucleation and growth kinetics, particularly at the beginning of 18 19 crystallisation when conditions are far from equilibrium. Here, we reveal the first in situ 3D 20 time-dependent observations of crystal nucleation and growth kinetics in a natural magma, 21 reproducing the crystallisation occurring in real-time during a lava flow, by combining a 22 bespoke high-temperature environmental cell with fast synchrotron X-ray microtomography. 23 We find that both crystal nucleation and growth occur in pulses, with the first crystallisation 24 wave producing a relatively low volume fraction of crystals and hence negligible influence on 25 magma viscosity. This result explains why some lava flows cover kilometres in a few hours from 26 eruption inception, highlighting the hazard posed by fast-moving lava flows. We use our 27 observations to quantify disequilibrium crystallisation in basaltic magmas using an empirical 28 model. Our results demonstrate the potential of in situ 3D time-dependent experiments and 29 have fundamental implications for the rheological evolution of basaltic lava flows, aiding flow 30 modelling, eruption forecasting and hazard management.

31 Magmas are multiphase mixtures consisting of liquid, crystals and bubbles. The arrangement and 32 evolution of these textures during magma ascent or residence in a magma chamber reflects the 33 physico-chemical history of the magma and thermodynamic properties of each component. Together, 34 these textures control magma rheological behaviour and permeability for volatile flow (e.g., ref. 1, 2). 35 These, in turn, dictate magma crystallisation and volatile exsolution, providing powerful feedback 36 mechanisms and complex non-linear behaviour. This richness makes volcanic processes fascinating, 37 as there is a very wide range of possible behaviours with rapid transitions between styles of activity. 38 However, these complex, interdependent processes must be accurately described if we are to succeed 39 in producing accurate models of volcanic systems. Such models are the long-term goal for many 40 volcanologists, as they will open the possibility of aiding volcano observatories in the interpretation of their observations, improving eruption forecasting and thereby allowing risk managers to keep 41 42 populations safe whilst minimising costly evacuations.

43 The combined effects of crystal and vesicle textures on magma rheology and their relationship to 44 eruptive style and intensity is a very active field of investigation. It has long been established that 45 viscosity, the most important rheological property governing magma transport processes, is a function 46 of magma composition, temperature, volatile content, crystal content, size, shape, distribution and orientation^{2,3} and vesicle content and arrangement^{4,5}. Recently, earth scientists have shown that crystal 47 shapes, besides their abundances, have a strong effect on magma rheological response^{3,6}. It is for this 48 49 reason that accurate knowledge of the processes determining the formation (e.g., nucleation, growth, 50 phase changes) of crystal phases with time is of critical importance to realistically describe dynamic 51 processes occurring during magma transport and syn- and post-eruptive emplacement.

52 Experimental work on crystallisation kinetics has been conducted through the study of 2D textures 53 (e.g., refs. 7-10); however, the texture of a volcanic rock is the final product of a dynamic process and 54 it is difficult to quantify with snapshot experiments and 2D measurements. This is particularly 55 relevant when studying low-viscosity systems such as basalts where quench effects may alter the 56 sample texture and chemistry. In addition, degassing and crystallisation are time-dependent processes 57 that introduce strong non-linearity in magmatic and volcanic systems. Crystallisation occurs in two 58 steps: nucleation and growth. Crystal growth is the process of the evolution of a single crystal 59 nucleated in a melt with a regular structure into the characteristic arrangement of a crystalline Bravais lattice¹¹. Growth could be related to a sequence of processes, from Ostwald ripening to crystal 60 aggregation¹² or dissolution, which significantly complicates the understanding of crystal textural 61 62 evolution in both space and time. Although previous studies have attempted to investigate such processes on *ex situ* samples, as demonstrated in other systems^{13,14}, only *in situ*, real-time 63 quantification of 3D crystal sizes and morphologies can provide a realistic description of kinetics 64

65 parameters such as crystal nucleation and growth rates, which will improve our knowledge of 66 crystallisation and, consequently, magma rheological properties.

With a few notable exceptions, the volcanology community has conventionally assumed that the processes of magma degassing and crystallisation occur as an equilibrium response to depressurisation during magma ascent and eruption. However, it is now recognised that the timescales required to achieve equilibrium for both crystal growth¹⁵⁻¹⁸ and volatile exsolution¹⁹⁻²¹ are similar to, or longer than, ascent times for erupting basaltic magmas, and therefore disequilibria are expected to be ubiquitous in such systems. Neglecting disequilibrium crystallisation and degassing therefore limits quantitative modelling and our understanding of volcanic processes.

74 Disequilibrium processes are challenging to study because the P, T, volatile content, melt composition 75 and rate-of-ascent parameter space is huge and because the traditional experimental techniques used 76 to explore disequilibrium processes are challenging and time-consuming. Until now, the principle 77 experimental approach has been 2D post-mortem examination of samples from laborious, 78 conventional petrological experiments requiring interruption and quenching (e.g., refs. 22, 23). By 79 using a high-temperature mossainite cell, in ref. 12 (and successive papers, see ref. 24 and references 80 therein) Schiavi *et al.* were able to optically observe *in situ* crystallisation of a basaltic andesitic melt 81 at atmospheric pressure over time, but their results are limited to a 2D space and may be significantly 82 influenced by surface effects.

In this study we report the first *in situ*, real-time 3D crystallisation experiments which constrain timescales of crystal kinetics in basaltic melts. We combine a high-temperature resistance furnace¹⁴ with the fast X-ray microtomographic capabilities of beamline I12 at the Diamond Light Source to directly quantify 4D (i.e. space and time) crystal nucleation and growth in a natural basaltic melt from Mt. Etna at atmospheric pressure. This allowed a full 3D picture with 3.2 micron³ voxel size to be captured every three minutes of the growing crystals, whose individual number and volumes could be quantified using post-processing segmentation techniques (see Methods for details).

90 Our experiments were intended to help inform part of the crystallisation process that occurs during 91 lava emplacement. The Etna 1981 eruption produced a lava flow that extended up to 6 km in ~3 92 hours²⁵, a similar timescale to our experiments. Whilst cooling would occur on the border of the flow, 93 the core would have insufficient time to cool significantly in 3.5 hours, and so our isothermal 94 experiments can capture the natural crystal dynamics. We cannot apply a shear force in these 95 experiments, and therefore our results are only representative of the core of the lava flow where strain 96 rate is minimal. 97 Cooling rates of basaltic lavas, measured at the surface and within active lava channels during 98 emplacement at a variety of locations range from 0.01 to 15 °C/min^{26,27}. The conditions investigated 99 in the presented experiments, starting at super-liquidus temperature (1250 °C), imposing a single-step 100 cooling and maintaining the temperature constant for 4 hours, likely represent the slowest cooling 101 rates (~0.01 °C/min) of the interior of large lava flows and lava channels.

102 Our 4D experiments on basaltic melts at realistic magmatic temperatures are, to our knowledge, the 103 first that have been conducted on natural basalts. We capture disequilibrium crystallisation pathways 104 by investigating the size, shape, orientation and evolution of crystals as a function of both time and 105 space. This powerful in situ experimental technique opens a new frontier in the research of complex 106 multiphase magmas, which were previously limited to ex situ quench studies, which have major 107 challenges in terms of repeatability and workload. Each experimental in situ 3D image we capture is 108 the equivalent of a single ex-situ quench experiment, and during a four hour experiment we capture 109 \sim 80 images. This is a step-change in how experimental petrology can be performed. The 4D X-ray 110 microtomographic method represents therefore a paradigm shift in how experimental petrology and 111 volcanological studies can be conducted.

112 **Results**

113 General observations. In our cooling crystallisation experiments (see Methods for details) a natural 114 anhydrous basaltic melt sampled from the 2001 Mt. Etna eruption was heated at atmospheric pressure 115 to 1250 °C on beamline I12 at the Diamond Light Source (Supplementary Table S1), and then cooled to 1150 °C or 1170 °C over a dwell time of 4 h (Supplementary Table S2). The experiments 116 117 crystallised augitic clinopyroxene (Supplementary Fig. S1, Supplementary Table S3), hereafter 118 referred to as simply pyroxene. We also observed abundant precipitation of Fe-Mg oxides during the 119 entire duration of the experiments, generated primarily by exposure to the atmosphere, which 120 produced uncontrolled redox conditions within our samples, favouring the nucleation of oxides. 121 However, in this paper we focus primarily on pyroxene crystallisation. In fact, we could not perform a 122 thorough quantitative image analysis on oxides owing to a fundamental resolution issue: a fraction of 123 oxides is close to or below the resolution limit of our system (pixel size=3.2 microns), producing 124 artefacts and showing most of the oxides connected to each other and forming oxide clusters and/or 125 chains (see Supplementary movie S2 in Methods). Indeed, their real spatial distribution consists of 126 euhedral isolated oxide crystals, which are arranged in layers. Because of this resolution issue, it is 127 very challenging to separate them and we could not reliably measure oxide number density, but we 128 were instead able to quantify oxide volume fraction with time.

During the entire dwell time, full 3D datasets were acquired every 3 minutes for quantitative imageanalysis (Supplementary Table S4). In the following, we focus on crystal nucleation and growth

kinetics results from an experiment cooling to 1150 °C (named ET1150 hereafter), and use an
experiment cooling to 1170 °C (named ET1170 hereafter) to provide additional information on
nucleation delay for pyroxenes and oxides.

MELTS calculations^{28,29} performed at ambient pressure and ambient fugacity (MH redox buffer), 134 indicate that the oxide liquidus temperature is 1230 °C (which is the liquidus phase of the bulk 135 136 composition), in agreement with Orlando et al. [ref. 30]. The clinopyroxene liquidus temperature is 137 1188 °C for the composition of our starting material. This estimate is similar to experimental determinations of clinopyroxene liquidus temperatures from Etnean basalt compositions^{16,30}. The 138 single-step cooling method was adopted in order to induce basaltic melt crystallisation in response to 139 140 an instantaneously applied thermodynamic driving force (i.e., undercooling, $\Delta T = T_{\text{liquidus}} - T_{\text{experimental}}$). 141 On the basis of the calculated clinopyroxene liquidus temperature, the experiments were performed at 142 nominal $\Delta T = 18$ °C and 38 °C for ET1170 and ET1150 respectively. The oxide crystals were formed at nominal $\Delta T = 60$ °C and 80 °C during experiments ET1170 and ET1150, respectively. 143

144 We choose to focus on experiment ET1150, because experiment ET1170 crystallised only a very 145 small volume fraction of pyroxene crystals (crystal volume fraction < 0.001) after 4 h at the 146 melt/sample holder interface at the bottom of the sample, and $\sim 0.11\pm 0.01$ of oxide crystals. This implies that pyroxene crystallisation kinetics are very slow at $\Delta T = 18$ °C, in agreement with ref. 16. 147 148 The nucleation, growth and textural development of pyroxene crystals with time is captured for the 149 first time in Fig. 1 (see Supplementary Movie S1 for full experiment). Each frame illustrates 150 successive crystal nucleation and/or growth, with a dwell time of 15 minutes between consecutive 151 frames. Each frame represents a 3D volume rendering of the same representative digital segmented 152 volume of interest (VOI) selected from the centre of each tomographic scan within experiment 153 ET1150 for quantitative analysis (see Methods).

154 Pyroxene crystals in ET1150 are dominantly euhedral and their habits change from blocky to more 155 elongate as crystallisation proceeds (Fig. 1). They are initially distributed as individual crystals 156 (Fig.1a, b, c); however, both crystal branching and crystal aggregates become ubiquitous as 157 crystallisation continues (Fig. 1e, f and successive frames). Oxide crystals are spherical and mostly 158 arranged in layers (see Methods, supplementary Movie S2). In our 3D tomographic images, we did 159 not find pyroxene crystals that nucleated from the bottom of our VOI. We observe what appears to be 160 both heterogeneous and homogeneous nucleation. Pyroxene and oxide crystals nucleate 161 heterogeneously from the top of the sample holder at the melt/air interface within the first 12 minutes 162 (Fig.1a, b, Table 1) and 6 minutes of the dwell time (Table 2), respectively. Because of our nominal 163 resolution of 3.2 μ m, we cannot however rule out the possibility that we are missing submicron 164 crystals forming at the onset of dwell time, implying that the nucleation delay found in our 165 experiments and reported in the following sections may be shorter. Heterogeneous nucleation at the 166 melt/air and melt/sample holder interfaces represents the first nucleation event, although the latter 167 becomes apparent in Fig. 1 only after 1 h of dwell time. The first pyroxene crystal nucleates in the 168 melt after 60 minutes of dwell time (Fig.1d). Because in the tomographic images we do not see any 169 surface upon which this crystal could have nucleated, we assume that it nucleated homogeneously. 170 We cannot however exclude that the nucleation of this first crystal in the bulk was heterogeneously 171 promoted by a small pore or an oxide that cannot be resolved at the 3.2 µm pixel resolution used in our experiments³¹. This crystal nucleus promotes a spherulite-like crystal growth, consisting mostly of 172 173 crystalline branches and crystal aggregates, visible near the bottom right-hand corner of our volume 174 renderings (Fig.1e, f, g and successive frames). After 60 minutes of dwell time crystallisation is 175 dominated by heterogeneous nucleation events, generating branched polycrystalline morphologies.

176 It is important to highlight that, in this study, the total crystal volume in each 3D frame reflects a 177 combination of both nucleation and crystal growth, and is the sum of the volume of individual 178 crystals, crystal branches and crystal aggregates. The calculated crystal number density with time also 179 includes all types of crystal morphologies (individual, branched and aggregates). For this reason, in 180 the following discussion we use the term *crystallisation rate* (expressed as crystal volume fraction 181 over time), which more faithfully illustrates our combined nucleation and growth processes, rather than using *crystal growth rate*, which conventionally relates only to individual crystal growth^{7,32}. 182 183 Specifically, we use *average crystallisation rate* to describe the crystallisation rate averaged over the 184 entire crystal nucleation and growth period, and *instantaneous crystallisation rate* to describe the 185 crystallisation rate between two successive temporal frames (see Methods for further details).

186 Pyroxene and oxide crystal volume fraction and crystallisation rate. During experiment ET1150 187 pyroxene crystal volume fraction varies between 4.3E-05 after the first 15 minutes of dwell time and 188 0.084 at the end of the experiment (Fig. 2a, Table 1). Pyroxene crystal volume fraction grows slowly 189 in the first 60 minutes of dwell time; it then increases almost linearly until the end of the experiment 190 (Fig. 2a). We observe a rapid increase in the volume fraction, from 0.032 to 0.053 (see red circles in 191 Fig. 11, m), between 150 minutes and 165 minutes of the dwell time. Pyroxene average crystallisation rate displays values from 4.75E-08 sec⁻¹ to 5.94E-06 sec⁻¹ (Fig. 2b). As a first order approximation, 192 the average crystallisation rate (Fig. 2b) increases rapidly in the first hour (up to 2E-06 sec⁻¹), 193 plateauing to an almost constant rate in the latter part of the experiment (from 3E-06 sec⁻¹ to 6E-06 194 195 sec⁻¹), meaning that equilibrium conditions are not reached. The instantaneous crystallisation rate 196 shows, in the first hour, a similar behaviour to the average crystallisation rate, increasing up to 4E-06 197 sec^{-1} (see Fig. 2b). After the first hour, although the average crystallisation shows an almost constant rate, the instantaneous crystallisation displays several fluctuations ranging from 7E-07 sec⁻¹ to 3E-05 198

sec⁻¹, implying that the crystallisation rate is not constant with time as it might appear from a first
order approximation.

The crystal volume fraction of Fe-Mg oxides varies between 0.03 after the first 15 minutes of dwell time and 0.17, the latter values reached at the end of the 4 h dwell time in experiment ET1150 (Fig. 2a, Table 2). Oxide crystal volume fraction grows quickly in the first 150 minutes of dwell time, then plateauing in the latter part of the experiment (Fig. 2a). Oxide average crystallisation rate is the highest at the beginning of dwell time (Fig. 2b, Table 2), and it then decreases very slowly from 3.33E-05 sec⁻¹ to 1.20E-05 sec⁻¹ until the end of the experiment.

207 Pyroxene and oxide crystal nucleation delay, pyroxene number density and nucleation rate. The 208 nucleation delay for heterogeneous pyroxene and oxide nucleation is, respectively, within 12 minutes and 6 minutes for experiment ET1150 (ΔT_{Px} = 38 °C; ΔT_{Ox} = 80 °C), and 30 minutes and 12 minutes 209 for experiment ET1170 (ΔT_{Px} = 18 °C; ΔT_{Ox} = 60 °C), demonstrating that the nucleation delay is 210 inversely proportional to ΔT . This is consistent with the nucleation behaviour of plagioclase observed 211 during isothermal decompression experiments¹⁰, and of olivine in cooled basaltic melts at ambient 212 pressure³³. These results suggest that nucleation delay for silicate and oxide crystals in basaltic melts 213 214 is strongly controlled by ΔT , whether crystallisation is induced by ambient cooling or isothermal decompression. The pyroxene crystal number density ranges between about 4 and 100 mm⁻³ over the 215 216 course of experiment ET1150 (Fig. 3a, Table 1). We observe three clear spikes in the crystal number 217 density as time progresses: the first and second occur after 60 minutes and 150 minutes in the dwell time, while the third occurs at the end of the experiment (Fig. 3a). These spikes suggest that pyroxene 218 nucleation occurs in pulses. Pyroxene average crystal nucleation rates range between 4.70E-03 mm⁻³ 219 sec⁻¹ and 1.13E-02 mm⁻³ sec⁻¹ (Fig. 3b, Table 1). Owing to the choice of VOI, our crystal nucleation 220 221 results are challenging to interpret. Pyroxene crystals nucleated heterogeneously at the melt/sample 222 holder interface within 12 minutes of the dwell time, at the same time as pyroxene crystals nucleated 223 at the melt/air interface. However, the former were initially not included in our VOI. In fact, in the 224 choice of the VOI, we excluded regions in the proximity of the sample holder walls (Supplementary 225 Fig. 2, see Methods), where crystal segmentation was highly subjective owing to sample surface 226 roughness and the presence of many highly phase-contrasted small bubbles. Therefore, pyroxene 227 crystals at the melt/sample holder interface entered and were counted in our selected VOI only after 1 228 h of the dwell time, with a delay of 45 minutes after they nucleated, implying that the nucleation trend 229 that we observe from 1 h onwards could be potentially shifted backwards in time.

230 Discussion

Our 4D *in situ* synchrotron X-ray tomographic results capture for the first time the kinetics of pyroxene nucleation and growth and of oxide crystallisation, revealing distinct stages of the process 233 quantitatively. Past experiments on crystal nucleation and growth have mostly focussed on the study 234 of feldspars in rhyolitic and basaltic melts (e.g., ref. 10.34). Studies on pyroxenes are very few and are 235 limited to conventional petrological experiments longer than 12 h⁹. Oxide crystallisation kinetics are also poorly experimentally investigated^{3,30}, however, they could have an important role in 236 237 heterogeneous nucleation of pyroxene in basaltic melts promoting, for example, crystal aggregation³⁵. 238 Because longer experimental durations provide lower apparent kinetic rates, previous studies may 239 therefore have underestimated pyroxene nucleation and growth rates. This has been already 240 highlighted, for example, in decompression-induced crystallisation experiments of a hydrous basaltic melt¹⁰, where they found that plagioclase nucleation and growth rates estimated in their 1 to 8 h 241 242 experiments are up to one order of magnitude higher than those of plagioclase obtained from longer 243 duration studies⁹. Recently, shorter *in situ* temporal observations of pyroxene crystallisation in a high-K basaltic melt were optically conducted at atmospheric pressure with a high-temperature moissanite 244 cell²⁴, but only 2D surface measurements were performed, which might prevent identifying textural 245 features related to crystal nucleation and growth kinetics that are instead apparent in 3D (e.g., growth 246 247 in pulses etc.). Our experiments uniquely measure the complex 3D crystal geometries in situ, allowing 248 the first quantitative investigation of crystallisation kinetics in a natural magma as a function of time. 249 For these reasons, the only kinetic data that can reliably be compared with our results are those measured in 3D trachytic feldspar spherulites³⁶. In ref. 36, Arzilli et al. report feldspar average 250 crystallisation rates between 1E-07 sec⁻¹ and 1E-08 sec⁻¹ in cooling crystallisation experiments of 251 252 trachytic melts lasting 4 and 6 h. These crystallisation rates are up to one order of magnitude lower than rates obtained for our pyroxene crystals (1E-06 sec⁻¹/1E-08 sec⁻¹, see Fig.2b and Table 1), 253 254 suggesting that the kinetics of nucleation and crystal growth are potentially faster in basaltic than trachytic melts owing to their lower viscosity³⁷. 255

256 In experiment ET1150 we observe three distinct pyroxene nucleation events. The first heterogeneous 257 nucleation event occurs at the melt/air and melt/sample holder interfaces and starts just a few minutes 258 after the beginning of the dwell time (Figs. 1a and 3a). When magma reached sub-liquidus conditions, crystal nucleation was firstly favoured by the presence of pre-existing surfaces such as melt/air and 259 260 melt/sample holder interfaces³⁸, which will be cooler and may have very small oxides ($<3.2 \mu m$) that 261 act as heterogeneous nuclei, although the latter cannot be resolved at the resolution used in our 262 experiments. Bubble surfaces (melt/gas interfaces) in a magma can induce heterogeneous crystal nucleation^{36, 39-41}. Our results therefore suggest that nucleation of pyroxene may be faster in the 263 264 presence of bubbles, which has important implications for conduit dynamics during magma 265 degassing. Because magma rheology is profoundly affected by the presence of crystals, bubble 266 enhanced crystal nucleation will increase magma viscosity, which, in turn, affects the overall magma fluid dynamics in the conduit⁴². Finally, it is important to highlight that the initial oxide crystallisation 267

268 may affect the initial pyroxene crystal nucleation but it does not affect pyroxene crystal growth269 kinetics as the two phases appear to follow distinct crystallisation trends (Fig. 2).

270 We observe that the first crystal (termed homogeneous) nucleated in the bulk after 1h initiated the 271 generation of the large spherulite-like texture, mostly consisting of crystal aggregates and branched 272 polycrystalline pyroxenes. From Fig.1e onwards, it is apparent that homogeneously nucleated 273 pyroxene crystals and pyroxene crystals nucleated on pre-existing surfaces start to develop 274 polycrystalline morphologies by branching. This implies that a second type of heterogeneous 275 nucleation occurred on pre-existing pyroxene crystals, i.e. at solid/solid interfaces. This type of 276 nucleation dominates from 75 minutes to the end of the dwell time, while homogeneous nucleation 277 and growth of individual crystals are secondary processes during this time interval. Indeed, when the 278 system is far from equilibrium, the pre-existing crystal surface can be perturbed by heterogeneities and instabilities, resulting in the nucleation of new grains at the growth front^{36,43,44}, and yielding a rich 279 280 variety of polycrystalline growth patterns. Heterogeneous nucleation leads to a decrease in the 281 interfacial surface energy, reducing the work of cluster formation, W. The work of cluster formation 282 on a pre-existing surface can be defined by 45 :

283
$$W_{het} = \frac{1}{4} W_{hom} \left(2 - 3\cos\theta + \cos^3\theta\right) \tag{1}$$

where θ is the dihedral angle at the contact between two crystals and the melt. The dihedral angle (θ) is related to the ratio of interfacial free energies (solid-liquid interfacial energy/solid-solid grain boundary energy)⁴⁵ expressed as:

$$\frac{\sigma_{sl}}{\sigma_{ss}} = \frac{1}{2\cos\frac{\theta}{2}}$$
(2)

288 where σ_{sl} and σ_{ss} are the solid-liquid and solid-solid interfacial free energies, respectively. For low θ , 289 the interfacial energy ratio will be low and heterogeneous nucleation will be energetically favoured^{38,46}. In this study, dihedral angles (θ) between a pre-existing crystal and the new grain were 290 291 measured, using the Avizo 3D software (FEI Visualization Sciences Group), as indicators of the 292 interfacial energy ratio. Dihedral angles range between 30° and 54°, implying that heterogeneous 293 nucleation on pre-existing crystals was strongly promoted in comparison with homogenous 294 nucleation. Finally, we find that pyroxene nucleation and crystallisation rates (Figs. 2b, 3b) in 295 experiment ET1150 differ from those that would be predicted by either the classical crystal size distribution theory⁴⁷⁻⁴⁹ or successive revisions of this theory (see, for example, ref. 24). The former 296 297 states that crystal nucleation and growth occur simultaneously over a large temperature and time 298 interval at a constant rate, while the latter predicts instead that the nucleation of each crystal phase is 299 limited to a short event, followed by a long period of crystal growth and annealing. Our data on 300 pyroxene crystals demonstrate the occurrence of at least three nucleation events over a period of 4 h

(Fig. 3a). In addition, the evolution of instantaneous crystallisation allows us to identify three growth dominated crystallisation events at 75, 165 min and at the end of the experiment (Fig. 2b),

303 highlighting that crystallisation is not constant but proceeds through pulses over time.

Multiple crystal nucleation events have been ascribed by Armienti *et al.* (1994) to sudden temperature and pressure perturbations during magma ascent and storage⁵⁰. We believe instead that our multiple pyroxene crystal nucleation events and growth through pulses at high temperature and atmospheric pressure are likely to reflect different interfacial free energies and delays in nucleation under disequilibrium conditions.

309 Experiment ET1150 displays a pyroxene crystallisation trend (Figs. 2a and 2b) that differs from that 310 assumed by numerical models of magma crystallisation in volcanic conduits. We find that pyroxene volume fraction increases slowly in the first 60 minutes, then almost linearly over the remaining 3 h 311 312 of the experiment (Fig. 2a). The model used by La Spina et al. in refs. 51, 52 instead assumes the 313 opposite trend, where the crystal volume fraction increases rapidly at the onset of crystallisation and 314 slows with time. In order to derive a new formulation for disequilibrium crystallisation in basaltic 315 magmas, we fitted our empirical data on crystal volume fraction of pyroxene versus time. In our 316 experiments, however, the equilibrium crystal volume content of pyroxene is not achieved; therefore we cannot describe the complete pathway up to the equilibrium value. We have calculated this 317 number using MELTS^{28,29}. We also assumed that at a certain point, as soon as the actual crystal 318 content is very close to the equilibrium value, the crystallisation rate should decrease with time. 319 320 However, since in our experiment we do not reach the equilibrium value for pyroxene, we cannot 321 constrain exactly when the crystallisation rate starts to decrease to zero. With these assumptions, we 322 found that the best fitting for the disequilibrium crystallisation in our experiment can be obtained 323 from:

324
$$\Phi_{c}(t) = \Phi_{c,0} + \left(\Phi_{c}^{eq} - \Phi_{c,0}\right) \left[\arctan\left(\frac{\pi}{2}c_{1}t^{c_{2}}\right)\right]^{c_{3}},$$
 (3)

where $\Phi_{c,0}$ is the initial crystal volume fraction, Φ_c^{eq} is the equilibrium value, $\Phi_c(t)$ is the actual value at the time t, and c_1 , c_2 and c_3 are proper fitting parameters. Given an initial crystal volume fraction of 0 and the equilibrium value of 0.14 (obtained from MELTS calculations), the values of the fitting parameters c_1 , c_2 and c_3 for the pyroxene crystal volume fraction curve are 9.46e-7, 2.57 and 0.86 respectively (Fig. 4, blue line).

Using the same equation, we were able to fit also the oxide crystal volume fractions with time. Compared to pyroxene, we do not need to estimate the equilibrium oxide content, since it reaches the equilibrium value during the course of the experiment (i.e. 0.17). The red line in Fig. 4 shows the fitting curve of the oxide crystal volume fraction obtained assuming the fitting parameters c_1 , c_2 and c_3 equal to 5.4e-9, 3.84 and 0.22, respectively.

Eq. (3) is expressed with a parametric formulation, and therefore it can be applied to describe disequilibrium crystallisation pathways of different crystal phases, in different conditions and for different natural systems. Clearly, different conditions such as pressure, temperature, oxygen fugacity, etc., will affect the fitting parameters. However, the overall trend of the disequilibrium crystallisation will remain similar to that found in this study, suggesting that Eq. (3) will be valid also for different conditions.

341 In addition, Fig. 4 illustrates the comparison between the empirical crystal volume fractions obtained 342 from our experiment and the fitting curves obtained from Eq. (3). Although the oxide crystal content 343 seems to reach the equilibrium value within 240 minutes (4 h), using our equation, we determine that 344 the difference between the equilibrium value of pyroxene and the actual value becomes small after 345 1000 minutes (~17 h), and it reaches equilibrium after ~1500 minutes (~25 h). The trend of pyroxene 346 illustrated by Fig. 4 is similar to that obtained by Melnik and Sparks in ref. 53, where they model the 347 evolution of crystal content in the Montserrat magma over time. In their Fig. 2, they show that the 348 crystal content initially increases slowly and then grows linearly until it eventually reaches 349 equilibrium. This finding has implications for magma rheology and dynamics in volcanic conduits 350 and lava flows, whereby slower crystallisation implies a lower magma viscosity and a faster magma 351 flow rate. Our results strongly suggest that disequilibrium crystallisation processes in basaltic systems 352 may be more pronounced than has previously been predicted.

353 The results of this study are important to quantify the crystallisation kinetics of basaltic lava lakes and 354 lava flows in the first 4 hours from the emission. It is generally assumed that changes in mineral 355 assemblage, crystal content and rheological properties of lava flows occur instantaneously with 356 changes in temperature. However, transient crystallisation processes and delay of crystal nucleation 357 occur in isothermal conditions, meaning that crystallisation and melt viscosity can change with time^{3,54}. Crystal content controls the rheology of lava flows and our results show that crystallisation 358 359 kinetics may slow the crystallisation process, promoting lower viscosities and longer flow lengths. An 360 archetype for a very fast-moving lava flow was produced during the 1981 eruption of Mount Etna. This eruptive episode was characterised by the emplacement of a ~ 6 km-long lava flow in about 3 361 hours (17 March 1981)²⁵. Our isothermal experiment ET1170 at 1170 °C shows a negligible amount 362 of pyroxene crystals (<0.001) and ~0.10 of oxides after 3 hours, whereas, considering the same 363 364 duration in isothermal experiment ET1150 at 1150 °C the crystal fraction of Fe-Mg oxides is 0.16 and 365 that of pyroxene is 0.05. From a rheological point of view, the spherical shapes of Fe-Mg oxides have 366 a negligible effect on the viscosity of lava flows, whereas the elongated shapes of pyroxene crystals could affect lava viscosity³; however, in agreement with ref. 3, we suggest that the low pyroxene 367

368 crystal fraction in the first 3 hours was not enough to induce a dramatic viscosity change. In ref. 3, 369 Vona *et al.* performed isothermal experiments, studying the viscosity evolution of Etna basalt with 370 time. However, through their ex situ experiments it is not possible to quantify the amount of crystals 371 and which phases are produced during the first hours of the experiments, and therefore, 3D 372 information on the crystal textural evolution in real time cannot be obtained. Our 3D real-time 373 experiments illustrate that a small amount of pyroxene crystals formed during the first 3 hours, 374 implying that these crystals did not increase lava viscosity significantly during its emplacement. This finding explains the reason why lava can flow for long distances in a short time²⁵. The novelty of our 375 376 approach is that we can quantify in 3D the evolution of crystallisation in real time, showing exactly 377 the onset of crystallisation of each mineral phase. This aspect allows us to investigate the textural 378 evolution of basaltic melts at the beginning of the solidification process when the magma is still able 379 to move and flow, providing a fundamental improvement in comparison to ex situ crystallisation and 380 rheological experiments. This study provides the first in situ time-dependent measurements of 3D 381 disequilibrium crystal nucleation and growth in a natural magma and opens the possibility to study 382 volcanic melts under realistic conditions for volcanic systems. The in situ 4D technique provides 383 orders of magnitude increase in the number of experimental runs that can be conducted compared 384 with traditional ex situ experiments, opening a new frontier for experimental petrology and 385 volcanology. Our results provide a fundamental improvement of our knowledge of crystallisation in 386 basaltic magmas, with implications for the rheological behaviour of basaltic magma, which is strongly 387 dependent on crystal content and crystal morphology, at near-vent conditions, relevant to the flow of lava and the ascent of explosive magma in the shallow conduit. As such, our results should be used in 388 389 models of lava flow dynamics and emplacement, and of magma ascent in the shallow conduit, to 390 improve their capacity to forecast the evolution of volcanic eruptions, assess volcanic hazard and 391 reduce the related volcanic risk.

392 Methods

Choice of starting material, starting material preparation and composition, pyroxene 393 394 composition. The starting material used for our cooling-driven crystallisation experiments consists of volcanic products from the lower vents of the 2001 Mt. Etna eruption⁵⁵⁻⁵⁷. In our experiments, we do 395 396 not attempt to reproduce faithfully the petrological and textural characteristics of the natural erupted 397 products: rather, we use the eruption and its products as a case study to investigate crystallisation 398 processes in shallow basaltic systems. It is however worth highlighting that our experimentally 399 produced pyroxene crystals resemble both compositionally (see Supplementary Figure S1 and 400 Supplementary Table S3) and texturally (ubiquitous presence of crystal aggregates, Fig. 2a, b in ref. 401 56) those found in natural samples. We chose material from the 2001 eruption because i) in terms of 402 style and intensity, this is one of the most representative eruptions that Mt. Etna has produced in the

403 last couple of decades, and can therefore be considered an archetype eruption for scientists 404 investigating processes in basaltic systems; ii) this is one of the most well monitored and studied 405 eruptions at this volcano, and there are numerous studies in the literature addressing the eruption 406 chronology and dynamics^{55,57-59}, its geophysical⁶⁰ and geochemical and petrological features^{56,61}, and 407 numerical modelling of conduit magma ascent⁵².

The anhydrous, glassy starting material was obtained by melting crushed rock samples in a 100 ml large, thin-walled, Pt crucible. Melting was performed in a Nabertherm® MoSi₂ box furnace at 1400 °C in air and at atmospheric pressure by adding the crushed sample to the crucible every 15 minutes. Once the crucible was completely filled, the melt was left in the furnace for about four hours to allow the melt to fully degas. The melt was then quenched in air to glass by pouring it onto a steel plate, and was crushed and melted a second time. Finally, glassy cylinders 3 mm in diameter and 4 mm in length were drilled from the synthesized glass for synchrotron X-ray microtomography experiments.

415 The chemical composition of the glassy starting material and of pyroxene crystals have been analysed 416 with a Jeol JXA 8530F microprobe in the facilities of the School of Earth and Environmental, 417 Sciences, University of Manchester, UK, and are reported in Supplementary Table S1, Supplementary 418 Table S3 and Supplementary Fig. S1. Analyses were performed using a 15 kV accelerating voltage, 419 10 nA beam current and beam size of 10 μ m. Standards used for calibration were albite for Na, 420 periclase for Mg, corundum for Al, fayalite for Fe, tephroite for Mn, apatite for P, sanidine for K, 421 wollastonite for Ca and Si and rutile for Ti. Sodium and potassium were measured first to minimize 422 loss owing to volatilisation.

423 Synchrotron X-ray microtomography experiments of 4D crystallisation in the 2001 Mt. Etna

424 basalt. Cooling-driven crystallisation experiments are defined by the temperature decrement imposed during cooling⁹: either one large drop in T (named single-step cooling or SSC) or a continuous 425 lowering of T (named continuous cooling or CC). Our experiments ET1150 and ET1170 belong to the 426 427 former type of crystallisation experiments (Supplementary Table S2). All experiments were 428 performed at Diamond Light Source beamline I12. Small cylindrical chips from the glassy starting 429 material described above were put in an alumina sample holder and were heated to 1250 °C and then equilibrated at this temperature for 30 minutes before cooling. For these experiments, we used the 430 high-temperature resistance Alice furnace¹⁴, which was commissioned and successfully used up to 431 1460 °C on I12, with controlled cooling at 0.05 °C/sec to 0.5 °C/sec. Crystallisation was induced by 432 isobarically decreasing temperature from 1250 °C to 1170 °C or 1150 °C, and then holding at the final 433 434 temperature for a dwell time of 4 h. Each experiment lasted about 6 $\frac{1}{2}$ h in total. Both experiments 435 were conducted in phase-contrast mode with a detector-sample distance of 2300 mm, using 436 monochromatic 53 keV light at a temporal resolution of 3 min per scan and a pixel size of 3.2 µm. 437 The detector was a high-resolution imaging PCO.edge camera with optical module 3, corresponding 438 to a field of view of 8.0 mm x 7.0 mm. Because we wanted to check the conditions of the starting 439 material both before and after the experiments, we acquired one scan of the sample in static mode at 440 the beginning and end of each experiment. Continuous scanning started before the end of sample 441 melting/equilibration, covered cooling and lasted for the entire duration of the dwell time. In each 442 scan, 1800 tomographic projections were acquired by the detector with equiangular steps over a full 443 rotation angle of 180° and an exposure time/projection of 0.05 s. These experimental conditions were 444 sufficient to capture pyroxene nucleation and growth and oxide crystallisation processes and their 445 textural evolution in 3D through time.

446 Tomographic data processing and quantitative analysis. Tomographic projections were 447 reconstructed into 2D slices by using Diamond I12 in-house python codes. The pre-processing 448 pipeline includes centre of rotation calculation, zinger removal, blob removal, regularisation-based 449 ring removal, then the GRIDEC algorithm is used for reconstruction 450 (http://confluence.diamond.ac.uk/display/I12Tech/Reconstruction+scripts+for+time+series+tomograp 451 hy, and refs. 62,63). The resulting 2D reconstructed tomographic slices were converted to 8-bit raw format and stacked into the freeware ImageJ software⁶⁴ to produce 3D digital volumes where the 452 isotropic voxel size has an edge length of 3.2 µm. 3D visualisation (volume rendering) of the 453 454 reconstructed volumes was obtained with the commercial software VGStudio 3.0 (Volume Graphics), 455 which allowed us to make qualitative textural observations on the evolution of pyroxene and oxide 456 nucleation and growth kinetics with time (Fig. 1, Supplementary Movie S1, Supplementary Movie 457 S2). While the whole scan time series consisting of 80 frames and covering the dwell time was 458 visualised carefully for qualitative textural investigation, quantitative image analysis was performed 459 on 16 frames with a time window of 15 minutes between the selected frames. Reconstructed volumes 460 of SSC experiment ET1150 were then cropped with ImageJ to select a volume of interest (VOI) for 461 quantitative image analysis of pyroxene and oxide crystals. The VOI was selected carefully in order to 462 i) discard the original imaged volume edges, which in the reconstructed volumes included the 463 experimental apparatus, and ii) specifically avoid the two entrained air bubbles at the bottom and top 464 of the sample holder, as well as small bubbles next to the sample holder walls that would have 465 complicated image segmentation owing to their highly phase-contrasted walls, and iii) avoid regions 466 where thermal gradients were the highest, i.e. the top and bottom of the sample holder. The VOI in 467 each frame therefore corresponds to the central part of the sample (Supplementary Figure S2), which 468 represents the furnace hotspot and the largest available melt volume that remains unaffected by surface effects. In experiment ET1170 pyroxene crystals nucleated in negligible numbers (crystal 469 470 volume fraction < 0.001) and only at the bottom of the sample holder at the melt/sample holder 471 interface, together with oxide crystals (0.11 ± 0.01) . In this experiment we analysed pyroxene and 472 oxide textures only qualitatively, and used this information to provide pyroxene and oxide nucleation 473 delay times for comparison with experiment ET1150.

474 We performed image processing of the selected VOIs following the protocol described in ref. 65 475 (Supplementary Table S4). Segmentation is the process that allows separation of objects from the 476 background to obtain binary volumes containing only the feature of interest. After using the 477 brightness-contrast function in ImageJ to increase the intensity of crystals and decrease that of the 478 matrix, segmentation of pyroxene crystals was operated in the 3D domain with the Pore3D software library⁶⁶ by using manual bi-level greyscale thresholding based on the greyscale histogram of the 479 480 selected VOIs and visual inspection of the slices in different directions. This allowed high sensitivity 481 to the presence of noise and artefacts and proved favourable for our purpose compared with automatic 482 thresholding algorithms. To avoid biasing from texturally complicated microtomographic images, pre-483 and post-segmentation smoothing filters were required to both ease and refine the segmentation 484 procedure (Supplementary Table S4, and refs. 67-69). The 3D bilateral filter in Pore3D was applied to 485 smooth the greyscale input images prior to segmentation, while preserving object edges⁷⁰. Because 486 oxide crystals have a near-identical greyscale contrast to pyroxene crystals in our VOIs, a significant 487 fraction of oxides were segmented together with pyroxenes in the same binary VOIs. To obtain 488 pyroxene crystal segmentation, we removed the segmented oxides from the volumes using a 489 combination of post-segmentation image processing steps in both ImageJ and Pore3D. This consisted 490 in applying a series of consecutive binary operations in ImageJ such as open, erosion, remove outliers 491 and dilate on both sagittal and coronal planes, which removed oxides and only a negligible part of the 492 smallest pyroxene crystals, as well as using the 3D minimum volume filter (MVF) in Pore3D, in the range 500-1500 pixel³ (corresponding to a volume of $25^3 \mu m^3$ and $37^3 \mu m^3$). An example of the 493 494 procedure is illustrated in Supplementary Movie S2 and Supplementary Movie S3. Here, a volume 495 rendering of frame 40 after manual greyscale thresholding and displaying both pyroxene and oxide 496 crystals (Supplementary Movie S2) is compared with a volume rendering of the same frame 497 containing only pyroxenes after the post-segmentation image processing protocol described above 498 (Supplementary Movie S3). The whole image processing protocol, including segmentation, pre- and 499 post-segmentation processing, lasted up to 3-4 hours per VOI. The resulting segmented pyroxene crystals were counted in each VOI and their volumes computed with the blob analysis in Pore3D⁷¹. 500 501 Pyroxene crystals numbers and individual volumes were used to obtain the textural and kinetic 502 parameters reported in Table 1, and Figs. 2 and 3. Oxide volume fractions were computed using the 503 plugin BoneJ in ImageJ, while their average crystallisation rate was obtained following the same 504 equation reported in the following for pyroxenes. Specifically, for each frame n acquired at the time 505 $t^{(n)}$, given the volume of pyroxene $V_{nx}(t^{(n)})$ and the number of pyroxene crystals $N_{nx}(t^{(n)})$, we computed:

$$\phi(t^{(n)}) = \frac{V_{px}(t^{(n)})}{VOI};$$
(4)

$$N_{\nu}(t^{(n)}) = \frac{N_{px}(t^{(n)})}{VOI};$$
(5)

. . . .

$$I_{v}(t^{(n)}) = \frac{N_{v}(t^{(n)})}{t^{(n)}};$$
(6)

$$Y_{v}(t^{(n)}) = \frac{\phi(t^{(n)})}{t^{(n)}};$$
(7)

$$Y_{i\nu}(t^{(n+1)}) = \frac{\phi(t^{(n+1)}) - \phi(t^{(n)})}{t^{(n+1)} - t^{(n)}};$$
(8)

where ϕ is the crystal volume fraction, *Nv* is the crystal number density, *Iv* is the average crystal nucleation rate, *Yv* is the average crystallisation rate, and *Yiv* is the instantaneous crystallisation rate. The uncertainty associated with image processing and analysis was calculated at 6% for ϕ and 7% for *Nv*, *Ivt*, and *Yvt*, and the standard deviation of the mean value is reported for each parameter in Table 1 and Table 2.

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687 Figure Captions

Figure 1 Crystallisation in a cooling basaltic melt at 1atm. Volume renderings showing pyroxene crystallisation kinetics in single-step cooling experiment ET1150.Vertical field of view is 2 mm. Red circles in l) and m) indicate region of pyroxene crystal volume fraction increase in Fig. 2a. See text for further details.

Figure 2 Crystal volume fraction and crystallisation rate in single-step cooling experiment
ET1150 with time. Pyroxene (red squares) and oxide (black squares) crystal volume fraction (a),
pyroxene average (closed red circles) and instantaneous (open red circles) crystallisation rate, and
oxide average crystallisation rate (closed black circles) (b). See text for further details.

Figure 3 Crystal number density and crystal nucleation rate in single-step cooling experiment
ET1150 with time. Pyroxene crystal number density (a) and pyroxene average crystal nucleation rate
(b). Red, blue and green lines in (a) indicate onset of heterogeneous nucleation, homogeneous
nucleation and heterogeneous nucleation branching. See text for further details.

700 Figure 4 Disequilibrium crystallisation pathway in single-step cooling experiment ET1150 with

time. Pyroxene (in blue) and oxide (in red) model of the disequilibrium crystallisation pathway
 obtained from the fitting of the experimental data. See text for further details.

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706 Data availability

The data that support the findings of this study are available within the article, its SupplementaryInformation files and from the corresponding author upon request.

709 Acknowledgements

710 This research has been funded by the RCUK NERC DisEqm project (NE/N018575/1) and project 711 NE/M013561/1. We gratefully thank D. Andronico, Osservatorio Etneo, Catania (Italy), for logistic 712 support during collection of samples from the 2001 Mt Etna eruption, and C. Romano for access to 713 the University of Roma3, Roma (Italy), analytical facilities. J. Fellowes, SEES, University of 714 Manchester, Manchester (UK), is acknowledged for help with microprobe analysis, and F. Brun, 715 INFN-Sezione di Trieste, Trieste (Italy), for support with the use of the Pore3D library. We gratefully 716 acknowledge Diamond Light Source, UK, for beamtimes EE12392/2 and EE16188/1 at I12, and the 717 Research Complex at Harwell, UK, for laboratory and office space.

718 Author contribution

MP, MB, FA and PDL conceived the research project. All authors, MP, FA, GLS, NLG, BC, MEH,
DDG, NTV, SN, RCA, EWL, PDL and MRB, contributed to the beamline experiments. FA collected
the volcanic rocks for the starting material. DDG prepared the starting material. MP performed image
reconstruction, processing and analysis, together with FA and GLS. GLS applied the crystallisation
model. MP wrote the manuscript, with contributions from all other authors.

724 Competing interests

- 725 The authors declare no competing interests.
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Frame#(time)	φ	<i>Nv</i> (mm ⁻³)	$Iv (mm^{-3}sec^{-1})$	$\frac{Yv}{(\sec^{-1})}$	Yiv (sec ⁻¹)
5 (15 min)	4.3E-05(2.6E-06)	4.27 (0.3)	4.70E-03 (3.3E-04)	4.75E-08 (3.30E-09)	4.75E-08
10 (30 min)	7.3E-04 (4.4E-05)	17.31 (1.2)	9.60E-03 (6.7E-04)	4.02E-07 (2.81E-08)	7.64E-07
15 (45 min)	2.3E-03 (1.4E-04)	28.04 (2.0)	1.00E-02 (7.0E-04)	8.51E-07 (6.0E-08)	1.74E-06
20 (60 min)	5.6E-03 (3.4E-04)	40.83 (2.8)	1.13E-02 (8.0E-04)	1.55E-06 (1.10E-08)	3.66E-06
25 (75 min)	0.011 (6.6E-04)	30.96 (2.2)	6.88E-03 (5.0E-04)	2.50E-06 (1.74E-07)	6.25E-06
30 (90 min)	0.016 (9.6E-04)	34.50 (2.4)	6.40E-03 (4.5E-04)	2.97E-06 (2.08E-07)	5.40E-06
35 (105 min)	0.021 (0.003)	49.98 (3.5)	7.93E-03 (5.5E-04)	3.31E-06 (2.32E-07)	5.36E-06
40 (120 min)	0.024 (0.001)	69.48 (4.9)	9.65E-03 (6.7E-04)	3.41E-06 (2.04E-07)	4.11E-06
45 (135 min)	0.025 (0.001)	84.11 (5.9)	1.04E-02 (7.3E-04)	3.12E-06 (2.18E-07)	7.51E-07
50 (150 min)	0.032 (0.002)	100.08 (7.0)	1.11E-02 (7.7E-04)	3.60E-06 (2.51E-07)	7.81E-06
55 (165 min)	0.053 (0.003)	90.08 (6.3)	9.10E-03 (6.4E-04)	5.30E-06 (3.70E-07)	2.32E-05
60 (180 min)	0.056 (0.003)	71.92 (5.0)	6.60E-03 (5.0E-04)	5.20E-06 (3.63E-07)	3.45E-06
65 (195 min)	0.065 (0.004)	78.01 (5.5)	6.70E-03 (5.0E-04)	5.60E-06 (3.92E-07)	1.04E-05
70 (210 min)	0.075 (0.004)	69.97 (4.9)	5.55E-03 (4.0E-04)	5.94E-06 (4.15E-07)	1.05E-05
75 (225 min)	0.079 (0.005)	71.67 (5.0)	5.31E-03 (4.0E-04)	5.84E-06 (4.10E-07)	4.51E-06
79 (237 min)	0.084 (0.005)	96.54 (6.7)	6.80E-03 (5.0E-04)	5.91E-06 (4.14E-07)	5.73E-06

730 Table1 Experimental results of pyroxene crystallisation kinetics in single step-cooling experiment ET1150

 ϕ = crystal volume fraction, Nv= Crystal number density, Iv= average crystal nucleation rate, Yv= average

732 crystallisation rate, Yiv= instantaneous crystallisation rate. Values in parentheses are the standard deviation of

the mean value.

Table2 Experimental results of oxide crystallisation

Frame#(time)	ϕ	Yv (sec ⁻¹)	
5 (15 min)	0.030 (0.002)	3.33E-05 (2.31E-06)	
10 (30 min)	0.04 (0.002)	2.22E-05 (1.55E-06)	
15 (45 min)	0.06 (0.004)	2.22E-05 (1.55E-06)	
20 (60 min)	0.08 (0.005)	2.22E-05 (1.55E-06)	
25 (75 min)	0.09 (0.005)	2.00E-5 (1.40E-06)	
30 (90 min)	0.10 (0.006)	1.85E-05 (1.30E-06)	
35 (105 min)	0.10 (0.006)	1.60E-05 (1.11E-06)	
40 (120 min)	0.14 (0.008)	1.94E-05 (1.4E-06)	
45 (135 min)	0.14 (0.008)	1.73E-5 (1.21E-06)	
50 (150 min)	0.16 (0.009)	1.77E-05 (1.24E-06)	
55 (165 min)	0.16 (0.009)	1.62E-05 (1.13E-06)	
60 (180 min)	0.16 (0.009)	1.50E-05 (1.04E-06)	
65 (195 min)	0.16 (0.009)	1.40E-05 (9.60E-07)	
70 (210 min)	0.17 (0.01)	1.35E-05 (9.44E-07)	
75 (225 min)	0.17 (0.01)	1.26E-05 (8.82E-07)	
79 (237 min)	0.17 (0.01)	1.20E-05 (8.40E-07)	

kinetics in single step-cooling experiment ET1150

 ϕ = crystal volume fraction, Yv= average crystallisation rate.

Values in parentheses are the standard deviation of the mean value.











b



























