



# Effects of chemical pretreatment and intra- and inter-specimen variability on $\delta^{18}\text{O}$ of aquatic insect remains

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**Abstract** Oxygen isotope ( $\delta^{18}\text{O}$ ) measurements on the exoskeletons of aquatic insects can be used to reconstruct changes in the  $\delta^{18}\text{O}$  of ambient water and, indirectly, to infer the climate and environmental conditions at the time of tissue synthesis. Prior to stable isotope analysis, it is often necessary to chemically pretreat insect remains to remove allochthonous organic and inorganic compounds without altering the  $\delta^{18}\text{O}$  signature. We tested the effectiveness and impact of duration of exposure to a buffered 2 M ammonium chloride ( $\text{NH}_4\text{Cl}$ ) solution

for removing carbonates at neutral pH from chironomid head capsules, water beetle sclerites and marine crab remains prior to stable isotope analysis. Immersion in  $\text{NH}_4\text{Cl}$  for 24 h efficiently removed the effect of carbonates with no long-term effects of prolonged exposure observed. Furthermore, we assessed the variability in  $\delta^{18}\text{O}$  values within and between individual sclerites (exoskeleton parts) of both modern and fossil water beetle remains. Both modern and fossil specimens had similar intra-sclerite variability in  $\delta^{18}\text{O}$  values ( $\sim 2\%$  range). In contrast, modern specimens had much smaller inter-sclerite variability ( $< 0.9\%$  range) compared with fossil specimens from the same sample (up to 10% range). The high inter-sclerite variability observed in fossil material likely results from the nature of fossil material: a mix of sclerites from a 1–2 L sample bin, originating from different individuals that may have existed at different times and under different environmental conditions. We therefore recommend that material to be analysed for stable isotopes be sampled at high temporal resolution to reduce uncertainties in paleotemperature estimates derived from water beetle  $\delta^{18}\text{O}$  records.

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## Introduction

Oxygen isotope ratios in the water of hydrologically open lakes ( $\delta^{18}\text{O}_{\text{lakewater}}$ ) dominantly reflect the isotopic composition of precipitation ( $\delta^{18}\text{O}_{\text{precipitation}}$ ) received by the lake via rainfall, streams and groundwater and thereby, indirectly, record a climate signal (Dansgaard 1964; Clark and Fritz 1997; Leng and Marshall 2004). The  $\delta^{18}\text{O}$  of the exoskeleton of aquatic insects has been shown to reflect the  $\delta^{18}\text{O}_{\text{lakewater}}$  at the time of cuticle formation (Wooller et al. 2004; Verbruggen et al. 2011), with around 70% of the oxygen required for biosynthesis acquired from the host lake water itself (Wang et al. 2009; Soto et al. 2013). Highly sclerotized chitinous remains of aquatic insects are generally abundant and well-preserved within lake sediments, with  $\delta^{18}\text{O}$  values on their fossil remains having the potential to provide information on paleoclimate (Wooller et al. 2004; Verbruggen et al. 2010b; Lasher et al. 2017; van Hardenbroek et al. 2018). In lacustrine environments, stratigraphic changes in insect  $\delta^{18}\text{O}$  values can be used to infer changes in  $\delta^{18}\text{O}_{\text{lakewater}}$  and, indirectly, past climate and environmental conditions. Biosynthesis of the chitinous cuticles does not seem to be affected by temperature-dependent fractionation processes (Schilder et al. 2015), in contrast to other lacustrine material, such as biogenic carbonate (Leng and Barker 2006). Moreover, the ability to identify insect remains under a microscope minimizes the possibility of including terrestrial contamination commonly faced by records derived for aquatic cellulose and algal lipids (Sauer et al. 2001; Leng and Barker 2006).

Chemical pretreatment of chitinous remains prior to stable isotope analysis is needed to eliminate endogenic or authigenic carbonates and allochthonous organic matter that may affect the  $\delta^{18}\text{O}$  signal. Conventional pre-treatment methods use acid (such as HCl or HF) to remove carbonates (Wooller et al. 2004; Brooks et al. 2007; Chambers et al. 2011), but at low pH oxygen can be exchanged between organic materials and water used in the pretreatment procedure (Wedeking and Hayes 1983). Moreover, the use of such acids was found to alter the chemical composition of chironomid cuticles by selective removal of chitin relative to proteins, resulting in decreased  $\delta^{18}\text{O}$  values compared to untreated controls (Verbruggen et al. 2010a). By contrast, pretreatment with a 2 M ammonium chloride ( $\text{NH}_4\text{Cl}$ ) solution buffered with

sodium hydroxide (NaOH) can be undertaken to remove sedimentary carbonates at a circum-neutral pH (Hieltjes and Lijklema 1980) in a way that does not promote oxygen exchange or alter  $\delta^{18}\text{O}$  values of chironomid cuticles (Verbruggen et al. 2010b). During pre-treatment, the  $\text{NH}_4\text{Cl}$  reacts with calcium carbonate ( $\text{CaCO}_3$ ), if present within the sample, to form calcium chloride ( $\text{CaCl}_2$ ) and ammonium carbonate ( $(\text{NH}_4)_2\text{CO}_3$ ), which rapidly degrades to gaseous ammonia ( $\text{NH}_3$ ) and  $\text{CO}_2$ . As such, the pretreatment removes carbonate at neutral pH, minimising exchange of OH-groups between water and insect cuticle and preserving the original  $\delta^{18}\text{O}$  signal of insect cuticles (Verbruggen et al. 2010b). It remains unclear, however, how efficient  $\text{NH}_4\text{Cl}$  pretreatment is in removing all residual carbonates in a sample, and whether it affects  $\delta^{18}\text{O}$  values of individual insect groups differently.

A second part of this study investigates variability of  $\delta^{18}\text{O}$  values of insect exoskeleton fragments. Fossil insect remains preserved within sediments are rarely intact specimens and instead, often comprise various parts of the exoskeleton from different individuals. It is not known how variable  $\delta^{18}\text{O}$  values of sclerite fragments from the same individual are, but this is important to consider when evaluating correlations between  $\delta^{18}\text{O}$  values of sclerite fragments and  $\delta^{18}\text{O}$  of precipitation (Gröcke et al. 2006, 2011; van Hardenbroek et al. 2013).

To address these issues, we designed a series of experiments to test (1) the efficiency of  $\text{NH}_4\text{Cl}$  to remove artificially added carbonate (2) the impact of  $\text{NH}_4\text{Cl}$  exposure duration on the  $\delta^{18}\text{O}$  of chironomid larvae, water beetle sclerites and a widely used purified chitin standard derived from marine crab shells; and (3) the extent to which there is natural variability in  $\delta^{18}\text{O}$  values within and between water beetle sclerites of modern and fossil material. Experiments 1 and 2 tested the effect of  $\text{NH}_4\text{Cl}$  on chironomid larvae, water beetle sclerites and the chitin standard derived from marine crab shells. Chironomid larvae (*Chironomus riparius*) were obtained from a commercial source (Discusfarm Marsilea, Lelystad, The Netherlands). In the third and final experiment, only beetle sclerites were tested, as their large size enabled an investigation of within-sclerite variability, unlike those of chironomids.

## Materials and methods

Soft tissues were left to decay in water for 3 months and remaining material was manually removed from head capsules with fine forceps. For the beetle chitin, a specimen of the genus *Dytiscus* was extracted from the surface sediments of Lake Nimetön (61.23°N, 25.19°E, Finland). The elytron was ground into a fine powder using an agate pestle and mortar to homogenize the sample. As a third sample, we used a marine crab chitin standard, which was purified by the supplier (Sigma-Aldrich, St. Louis, MO) using the method of Skujins et al. (1965).

For experiment 3, which tested intra- and intersclerite variability in  $\delta^{18}\text{O}$  values, we analysed modern and fossil specimens of North American water beetles from the genera *Dytiscus*, *Hydroporus* and *Helophorus*. Adults and their larvae of these three genera are fully aquatic, with *Dytiscus* and *Hydroporus* (predaceous diving beetles) inhabiting the pelagic zone of ponds and shallow lakes, whilst intermittently diving down deeper into the water column to catch their prey (Larson et al. 2000). Adults of the water scavenger beetle *Helophorus* inhabit the benthic zone of ponds and shallow lakes, and their larvae are known to be predaceous (Angus 1973; Landin 1980). In North America, the larvae of these three genera typically emerge within 6–8 days in spring or summer; *Hydroporus* in late July to early August (Larson et al. 2000), whereas *Helophorus* and *Dytiscus* species typically emerge weeks to months earlier in the year (Smetana 1988; Hilsenhoff 1995). All of the beetle specimens analysed in this study are fully aquatic and were collected from perennially wet ponds, lakes or associated organic deposits representing tundra lakes, peatland and/or mire sediments. The sample material consisted of heads, pronota, elytra, metasterna, hind legs and abdomens (Figs. 3 and 4). For details of collection location and age of the modern and fossil beetle specimens, see Electronic Supplementary Material (ESM) Tables 1 and 2.

## Experimental design

Experiment 1 tested the effectiveness of a buffered 2 M  $\text{NH}_4\text{Cl}$  solution in removing artificially added carbonate. All samples, including controls, were immersed in 10% KOH for 2 h at room temperature

as an initial cleaning step to remove any remaining soft tissues, following Verbruggen et al. (2010b), and rinsed with demineralized water over a 41- $\mu\text{m}$ -mesh sieve. Next, 1 g of calcium carbonate ( $\text{CaCO}_3$ ) powder was added to 1 mg of chironomid, beetle and crab sample batches. These samples were then soaked in 5 ml of a 2 M  $\text{NH}_4\text{Cl}$  solution buffered with NaOH at a pH of 7.5 (Verbruggen et al. 2010a) for either 2 h or 24 h. All samples were rinsed with deionised water and freeze-dried before  $\delta^{18}\text{O}$  analysis. A subset of each  $\text{CaCO}_3$ -treated batch received no pretreatment with  $\text{NH}_4\text{Cl}$  and acted as a positive control. Measured  $\delta^{18}\text{O}$  values of treated samples were compared with those obtained for untreated samples, which received no  $\text{CaCO}_3$  or pretreatment with  $\text{NH}_4\text{Cl}$ .

Experiment 2 tested whether the duration of exposure to  $\text{NH}_4\text{Cl}$  had an effect on the  $\delta^{18}\text{O}$  of aquatic insect remains. Batches of samples were drawn after 2, 12, 24 and 48 h of exposure to 5 ml of the buffered  $\text{NH}_4\text{Cl}$  solution at room temperature and rinsed with deionised water. Initial batches of each group were measured as an untreated control group receiving no pretreatment with  $\text{NH}_4\text{Cl}$ .

We tested for significant differences between mean  $\delta^{18}\text{O}$  values of control samples and the mean  $\delta^{18}\text{O}$  values for each experiment using the non-parametric Kruskal–Wallis test and between treatments in each experiment with Mann–Whitney tests (Past software version 2.14, Hammer et al. 2001).

## Analytical methods

All samples except the three modern *Dytiscus* beetle specimens were analysed for  $^{18}\text{O}/^{16}\text{O}$  in the Stable Isotope Biogeochemistry Laboratory at Durham University. Oxygen isotope ratios were determined using a High Temperature Conversion Elemental Analyzer (Thermo Finnigan TC/EA) coupled to a continuous flow Delta V Advantage isotope ratio mass spectrometer (IRMS) (Thermo Scientific) with no-helium dilution aspect in the ConFlo III, because of the small sample weights obtained for chitin samples. Results were calibrated against three international reference standards (IAEA-600, -601, -602) and two in-house standards (TU-2 and C9752). Data filtering followed a strict protocol to assess the quality of the electrical signal retrieved for samples, including reference standards, on the mass spectrometer. In total, three measurements of standards were removed from the

analysis because of their likely low weight and/or incomplete combustion. Replicate measurements of standards indicated a mean analytical precision of 0.4‰ (1 $\sigma$ ) for batch 1 and 0.3‰ (1 $\sigma$ ) for batch 2 (ESM Table 3).

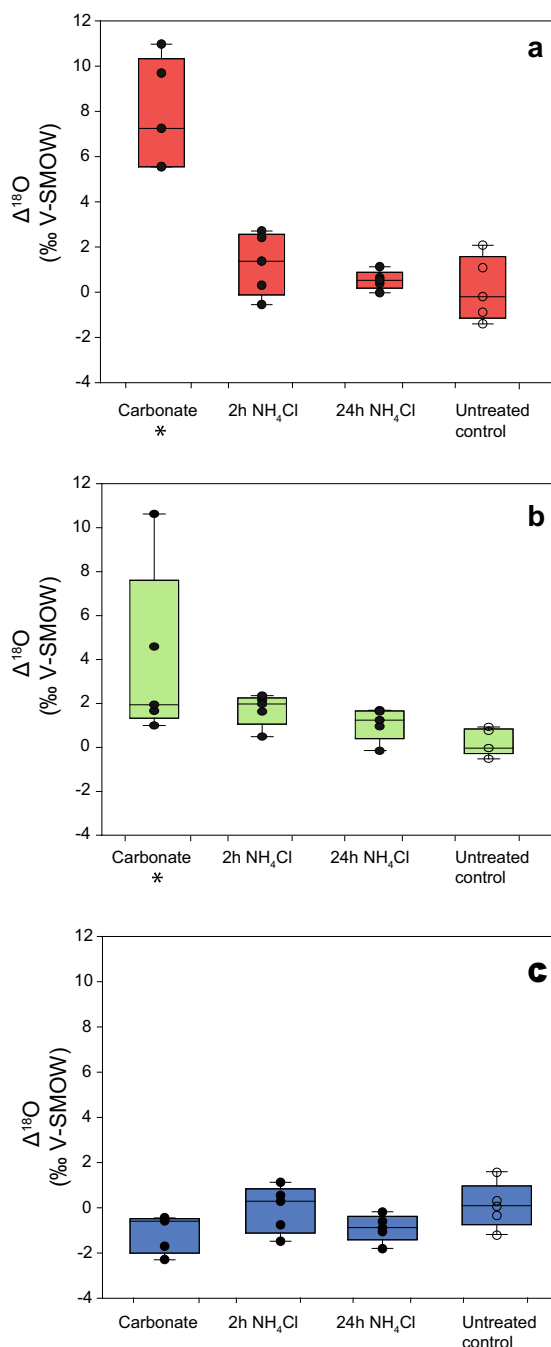
Modern *Dytiscus* beetle specimens were analysed in the Stable Isotope Biogeochemistry Laboratory at McMaster University on a Costech TC/EA connected to a Thermo Finnigan Delta Plus XP IRMS without helium dilution. The results were calibrated against in-house chitin and benzoic acid reference standards, and ANU sucrose as international reference material (ESM Table 3). Repeated measurement of standards indicated a mean analytical precision of < 0.4‰ (1 $\sigma$ ).

## Results

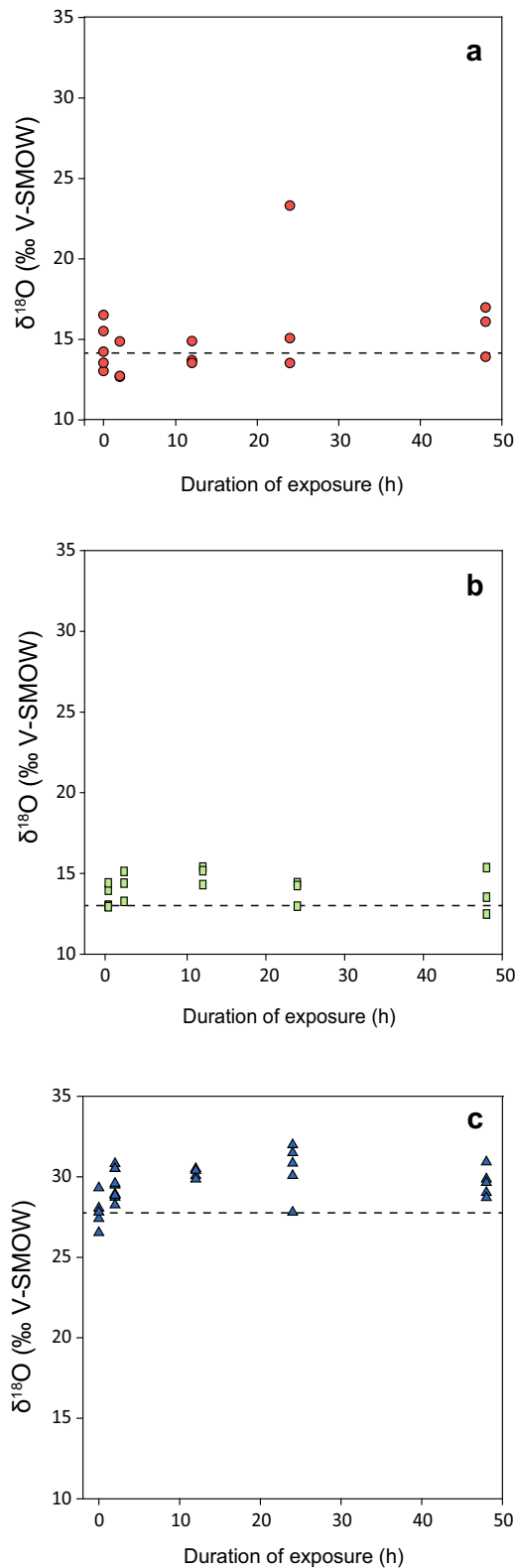
### Effect of NH<sub>4</sub>Cl pretreatment on invertebrate $\delta^{18}\text{O}$

Experiment 1 tested the efficiency of NH<sub>4</sub>Cl pretreatment for removing residual carbonates from chironomid, beetle and marine crab chitin prior to stable isotope analysis. The artificial addition of CaCO<sub>3</sub> was found to significantly change the  $\delta^{18}\text{O}$  of chironomid and beetle samples only compared to controls (Fig. 1), whereas the CaCO<sub>3</sub> and crab chitin samples had similar  $\delta^{18}\text{O}$  values (results of Mann–Whitney tests in ESM Tables 4b and 4c). After 2 h and 24 h of treatment with NH<sub>4</sub>Cl, the  $\delta^{18}\text{O}$  values of the aquatic insect samples did not indicate an effect of CaCO<sub>3</sub> anymore, as the samples were not statistically different from the untreated controls (ESM Tables 4b and 4c).

Experiment 2 tested the impact of prolonged exposure to NH<sub>4</sub>Cl on chironomid, beetle and crab  $\delta^{18}\text{O}$  (Fig. 2). One outlier in the chironomid batch exposed to NH<sub>4</sub>Cl for 24 h, possibly the result from contamination by CaCO<sub>3</sub> used in experiment 1, was excluded from further statistical analyses. No significant differences in  $\delta^{18}\text{O}$  values were found for samples exposed to NH<sub>4</sub>Cl for 2 h and those obtained for samples exposed for 12, 24 or 48 h ( $H > 2.6$ ,  $p > 0.2$ , ESM Table 4a). A significant difference was found, however, for untreated crab controls and those exposed to NH<sub>4</sub>Cl for any duration, possibly resulting from removing calcium carbonate that is an integrated part of crustacean exoskeletons (Vincent 2002; Gupta 2011). The effect of removing the internal carbonate



**Fig. 1**  $\Delta^{18}\text{O}$  values ( $\delta^{18}\text{O}$  untreated–treated) for **a** chironomid, **b** beetle and **c** crab chitin samples, having undergone pretreatment experiments (filled symbols), and untreated control samples (open symbols). Pretreatment experiments included the artificial addition of CaCO<sub>3</sub> with no NH<sub>4</sub>Cl exposure (carbonate), artificial addition of CaCO<sub>3</sub> with either 2 h exposure in NH<sub>4</sub>Cl (2 h NH<sub>4</sub>Cl) or 24 h exposure in NH<sub>4</sub>Cl (24 h NH<sub>4</sub>Cl). Pretreatment experiments found to be statistically different from untreated control samples are indicated (\*)



**Fig. 2** Impact of the duration of exposure to a buffered 2 M  $\text{NH}_4\text{Cl}$  solution on **a** chironomid, **b** beetle and **c** crab  $\delta^{18}\text{O}$  values. Dotted line indicates the mean  $\delta^{18}\text{O}$  of untreated controls for each invertebrate group

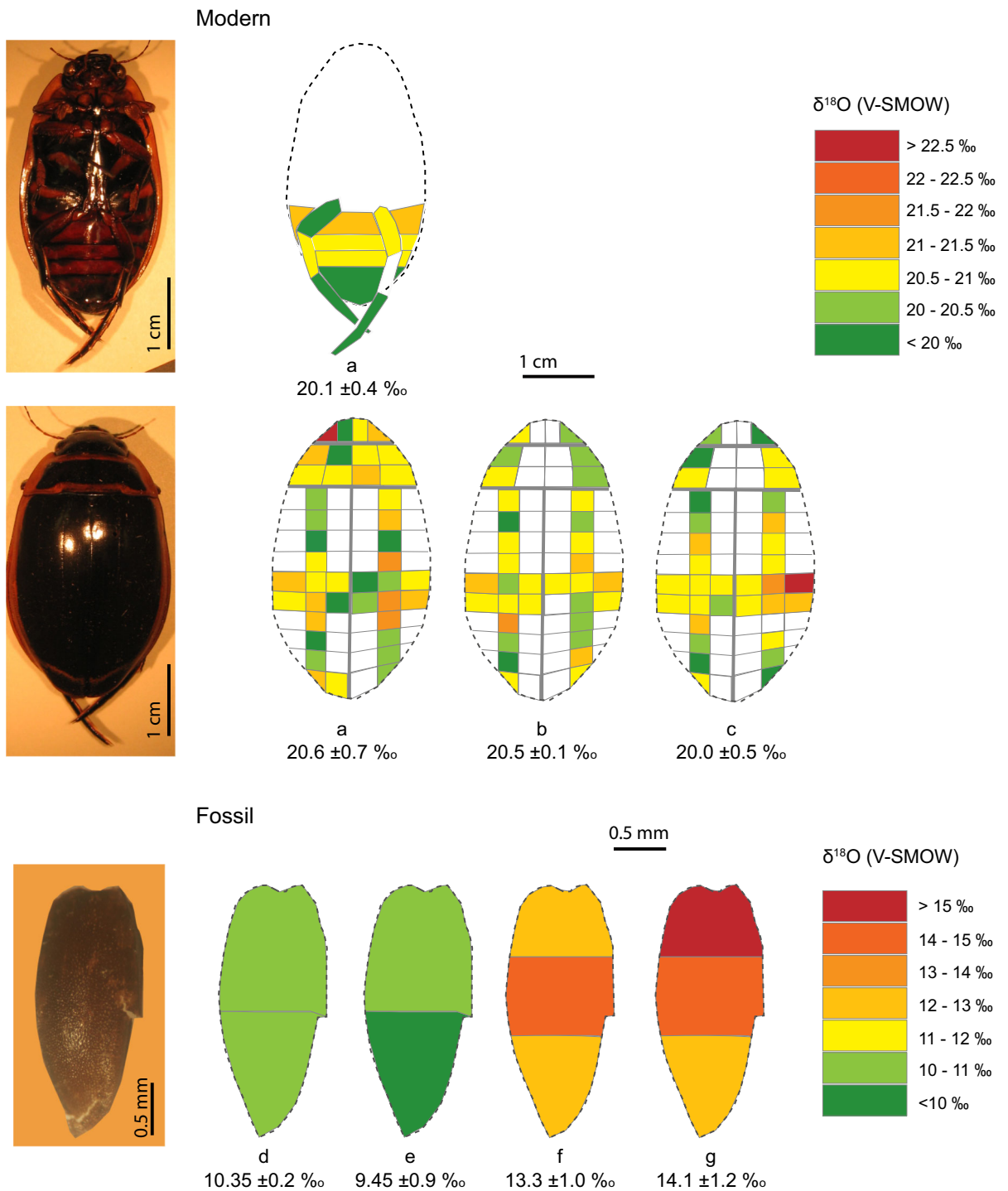
on  $\delta^{18}\text{O}$  was significant when comparing control and exposure to  $\text{NH}_4\text{Cl}$  for 2 h, but no further change in  $\delta^{18}\text{O}$  was found upon prolonged exposure to  $\text{NH}_4\text{Cl}$  (ESM Table 4e).

#### Intra- and inter-sclerite variability in beetle $\delta^{18}\text{O}$

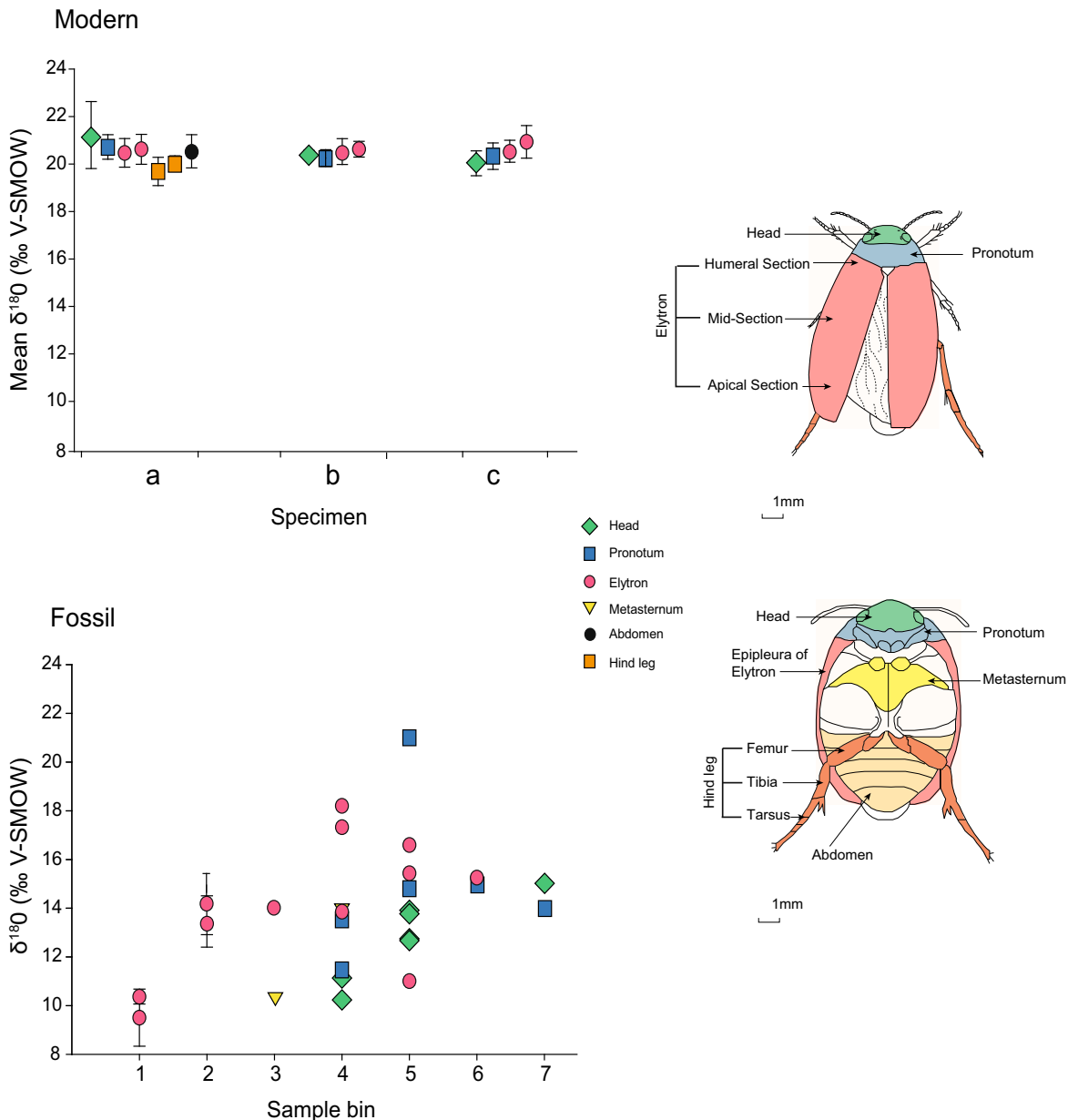
We investigated the extent of intra- and inter-sclerite variability in stable oxygen isotope values in both modern and fossil water beetles from North America. Within the exoskeletons of three modern *Dytiscus harrisii* diving beetles, we observed a variability in  $\delta^{18}\text{O}$  values (standard deviation) of 0.7, 0.4, and 0.6‰, with a range of  $\sim 2\%$  in any body part (Fig. 3a–c). Moreover, a difference in mean  $\delta^{18}\text{O}$  values of 0.3‰ was observed between the left and the right elytron. For the specimen presented in Fig. 3a, some of the ventral sclerites were also measured. The abdomen and hind legs displayed a standard deviation of 0.7‰ and 0.4‰, respectively.

The intra-sclerite variability in  $\delta^{18}\text{O}$  values within intact fossil remains of elytra from the genus *Hydroporus* are presented in Fig. 3d–g. On average for the four fossil specimens analysed,  $\delta^{18}\text{O}$  values were 1.05‰ higher for the humeral (top) section of the elytron compared to the apical (base) section. The intra-sclerite variability of fossil material was similar to that of modern material (ranges of  $\sim 2\%$  in the fossil elytra), suggesting that fossil elytra contain a similar stable isotope ‘signal’ as modern elytra. Modern beetle specimens were more than 10 × larger and sampled in more detail than fossil specimens, limiting the comparison of intra-sclerite variability in modern and fossil material.

Standard deviations of 0.2 to 0.5‰ were observed for  $\delta^{18}\text{O}$  values of individual sclerites of the modern *Dytiscus harrisii* specimens (e.g. head, elytron, pronotum, abdomen), with the elytra displaying the smallest standard deviation in  $\delta^{18}\text{O}$  values compared with other body parts (Fig. 4a). By contrast, high inter-sclerite



**Fig. 3** Intra-sclerite (exoskeleton part) variability in  $\delta^{18}\text{O}$  of modern *Dytiscus harrisii* (specimen a–c) and fossil *Hydroporus* (specimen d–g) water beetle specimens from Barter Island, Alaska. Note different scales for modern and fossil specimens



**Fig. 4** Variation in  $\delta^{18}\text{O}$  values for individual sclerites originating from both modern (museum) and fossil specimens of water beetles from North America. Fossil specimens from the same 1–2 L sample bin originate from the same site and same age fossil insect assemblage. All fossil specimens originated

from a different individual sclerite fragment. Error bars indicate variability within a single body part (intra-sclerite variability). See ESM Tables 1 and 2 for details on the collection location and age of modern and fossil beetle specimens, respectively

variability was observed for fossil specimens (Fig. 4b). Sclerites obtained from fossil assemblages of 1–2 L samples of the same age and collection

location displayed standard deviations of 0.1 to 3.2‰ (total range up to 10‰).

## Discussion

### Effect of $\text{NH}_4\text{Cl}$ pretreatment on invertebrate $\delta^{18}\text{O}$

We confirm the findings of Verbruggen et al. (2010b) for chironomids and further demonstrate that duration of exposure to  $\text{NH}_4\text{Cl}$  also has no significant effect on the  $\delta^{18}\text{O}$  of water beetles and a commonly used purified chitin standard derived from marine crab shells. Combined, the two experiments presented here indicate that, within the 0.4‰ analytical error of our measurements, a 24–48 h exposure to a buffered  $\text{NH}_4\text{Cl}$  solution can effectively remove all carbonates from aquatic insect remains prior to stable isotope analysis, without affecting the  $\delta^{18}\text{O}$  values of the sample material. The buffered  $\text{NH}_4\text{Cl}$  method was originally designed to measure carbonate-bound phosphorus in sediments and further developed by Hieltjes and Lijklema (1980) for lake sediments. Hieltjes and Lijklema (1980) used 50 ml of 1 M  $\text{NH}_4\text{Cl}$  solution buffered to pH 7 and 50 mg of dry sediment from Lake Brielle, The Netherlands, containing between 8.5 and 18.2% calcium carbonate by wt. They observed nearly complete carbonate removal ( $93.5 \pm 6.7\%$ , average  $\pm$  standard deviation) after 2 h of exposure to 1 M  $\text{NH}_4\text{Cl}$ , which is half as concentrated as the 2 M  $\text{NH}_4\text{Cl}$  solution used in our study and by Verbruggen et al. (2010b). No study to date has specifically tested the removal efficiency of the  $\text{NH}_4\text{Cl}$  method on different types of sedimentary and authigenic carbonates (e.g. calcite, aragonite, dolomite), but based on Hieltjes and Lijklema (1980) and Verbruggen et al. (2010b) and the current study, a buffered 2 M  $\text{NH}_4\text{Cl}$  will effectively remove carbonates from soils and lake sediments, especially when left to react for 24 h, which is 12 times as long as in the study by Hieltjes and Lijklema (1980).

### Intra- and inter-sclerite variability in beetle $\delta^{18}\text{O}$

Intra-sclerite variability ( $\sim 2\%$  range) in  $\delta^{18}\text{O}$  values for both modern and fossil water beetle sclerites (Fig. 3) was relatively low compared to the high inter-sclerite variability observed between fossil specimens originating from the same sample (up to 10‰ range; Fig. 4b). In general, the modern elytra displayed the most consistent  $\delta^{18}\text{O}$  values compared with other sclerites, but variability in  $\delta^{18}\text{O}$  does not seem to be systematic by region or body part, based on the

sampling locations and specimens analysed here (Fig. 3). This small within-sclerite variability could result from remaining soft tissues (e.g. inside the cavities of legs and heads) with different  $\delta^{18}\text{O}$  values, thereby altering the  $\delta^{18}\text{O}$  signature within sclerites originating from the same individual.

As there is also relatively low inter-sclerite variability in  $\delta^{18}\text{O}$  values for modern beetle specimens ( $< 0.9\%$  range; Fig. 4a), the high inter-sclerite variability observed in fossil material (up to 10‰ range) likely reflects either real differences in recorded source-water isotopes or interspecies differences in vital effects and/or biological fractionation. Typically, the paleoecological analysis of fossil beetle assemblages requires between 1 and 2 L of organic sediment to yield sufficient material to constrain paleoenvironmental interpretations (Elias 2001). The resulting fossil assemblages can therefore span thousands of years during which large shifts in the  $\delta^{18}\text{O}$  of precipitation and lake water  $\delta^{18}\text{O}$  are likely to have occurred. Furthermore, specimens retrieved from fossil beetle assemblages are often composed of a mix of sclerites from different individuals and/or species within the genera of interest, each with slightly different diet, habitat preferences, metabolism or utilisation of body water (deBruyn and Ring 1999). Differences in the seasonality of growth and/or scleral synthesis within and between the individuals and species present within the fossil assemblage could also explain some, or even most, of the observed variability in beetle  $\delta^{18}\text{O}$  values. No studies have investigated the seasonal variability of  $\delta^{18}\text{O}$  values in water beetle sclerites, but it is thought that aquatic beetle species mostly synthesize chitin during and shortly before emergence. For the three genera in this study, the larvae emergence typically takes place within a week in late spring or summer; *Hydroporus* in late July to early August (Larson et al. 2000), whereas *Helophorus* and *Dytiscus* species typically emerge weeks to months earlier in the year, usually in spring or early summer (Smetana 1988; Hilsenhoff 1995). Generally, emergence happens close to the warmest month, especially at higher latitudes (David Larson pers. commun.). Previous work comparing  $\delta^{18}\text{O}$  values of modern water beetles in North America and monthly  $\delta^{18}\text{O}$  values of modelled precipitation, found the strongest correlations for July and August (van Hardenbroek et al. 2013), also suggesting that variability of  $\delta^{18}\text{O}$  values in water beetle sclerites is



limited to variability of lake water  $\delta^{18}\text{O}$  values during the warmest months.

The low intra-sclerite and high inter-sclerite variability in  $\delta^{18}\text{O}$  from fossil beetle specimens has important implications for reconstructing paleoclimate. If it is assumed that the  $\delta^{18}\text{O}$  of environmental waters varies with temperature at a slope of 0.7‰ per °C (Dansgaard 1964), the mean observed variability in  $\delta^{18}\text{O}$  of fossil specimens derived from the same fossil assemblage (~ 1.5‰) corresponds to an uncertainty in paleotemperature estimate of  $\pm 2.6$  °C. Analysis of multiple beetle specimens from the same stratigraphic interval or assemblage will thus provide a larger, but more realistic range of beetle  $\delta^{18}\text{O}$  values and may give valuable information on short-term variability (i.e. within the time period represented by the fossil sample) in lake water isotope composition, as has been demonstrated with ostracod valves (Escobar et al. 2010).

In an attempt to test its utility in paleoenvironmental analysis, Motz (2000) measured stratigraphic changes in  $\delta^{18}\text{O}$  values of fossil bark beetle fragments and other tree-feeding insects from ca. 18,000-year-old sediment at the Gardena site, Illinois, and compared them to measured  $\delta^{18}\text{O}_{\text{cellulose}}$  values from the same horizons. High variability (standard deviation 0.62 to 1.76‰) was observed for fossil bark beetle fragments analysed from the same stratigraphic interval, likely because of (1) real differences in the fractionation of stable oxygen isotopes in different (parts of) trees as a result of differential evapotranspiration and/or (2) the large sample volumes (28–34 kg) used, limiting the temporal resolution possible for the fossil insect  $\delta^{18}\text{O}$  record. It would be expected that analysing material sampled at higher temporal resolution would reduce the number of sclerites in a sample, and thus the inter-sclerite variability. For example, high-resolution aquatic insect  $\delta^{18}\text{O}$  records have been produced using chironomid head capsules, which are generally more abundant in lake sediments than remains of water beetles (Verbruggen et al. 2010b; Arppe et al. 2017; Lombino 2014; Lasher et al. 2017).

## Conclusions

We demonstrated that pretreatment with buffered  $\text{NH}_4\text{Cl}$  efficiently removes the effect of artificially

added calcium carbonate on the  $\delta^{18}\text{O}$  values of chironomid and beetle chitin, without affecting their  $\delta^{18}\text{O}$  measure, even after prolonged (48 h) exposure. Moreover, we assessed the extent to which there is internal and within-sample variability in  $\delta^{18}\text{O}$  of modern and fossil specimens of water beetles. The variability in  $\delta^{18}\text{O}$  values within sclerites (exoskeleton parts) of modern specimens was similar to that observed for fossil specimens. Standard deviations of  $\delta^{18}\text{O}$  values between individual modern water beetles was only 0.2 to 0.5‰, and more than six times greater between individual fossil water beetle remains (0.1 to 3.2‰). Most likely, the large variability in fossil samples is caused by the low temporal resolution of the samples studied here, which span thousands of years. We therefore recommend that sediment records be sampled at higher temporal resolution for stable isotope studies of beetles, instead of large-volume sampling, which is often necessary for the paleoecological study of beetle assemblages.

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