Petrogenesis of the Triassic granitoids from the East Kunlun Orogenic Belt, NW China: Implications for continental crust growth from syn-collisional to post-collisional setting

Juanjuan Kong [a, b*], Yaoling Niu [b, c, d*], Yan Hu [e], Yu Zhang [b], Fengli Shao [f]

[a] College of Ocean Science and Engineering, Shandong University of Science and Technology, Qingdao 266590, China

[b] Laboratory for Marine Geology, Qingdao National Laboratory for Marine Science and Technology, Qingdao 266061, China

[c] Department of Earth Sciences, Durham University, Durham DH1 3LE, UK

[d] State Key Laboratory of Geological Processes and Mineral Resources, and School of Earth Science and Mineral Resources, China University of Geosciences, Beijing 100083, China

[e] Geology Survey Institute of Gansu Province, Lanzhou 730000, China

[f] Linyi University, Linyi 276000, China

***Corresponding authors:**

Miss Juanjuan Kong (juanjuan0317@foxmail.com)

Professor Yaoling Niu (yaoling.niu@durham.ac.uk)

Abstract

The Triassic granitoids are widespread in the eastern section of the East Kunlun Orogenic Belt (EKOB) on the northern Tibetan Plateau. These granitoids well record the evolution of the Paleo-Tethys oceans (named as A'nyemaqen Ocean in the EKOB). Our new zircon U-Pb data together with ages in literature show that these granitoids represent long-lasting magmatism from the early (T_1, ~ 251-248 Ma), middle (T_2, ~ 247-238 Ma) to late (T_3, ~ 234-214 Ma) Triassic. The Triassic granitoids display calc-alkaline I-type granite affinities and hybrid mantle-crust geochemical signatures. The T_1 granitoids possess andesitic to felsic bulk continental crust (BCC)-like chemical composition (e.g., enriched in Rb, K and Pb, depleted in Nb, Ta, Sr, P and Ti), coupled with high I_{Sr} (0.7067-0.7148), negative $\varepsilon_{Nd(t)}$ (-7.32 to -1.66) and negative to positive $\varepsilon_{Hf(t)}$ (-5.11 to 3.59) as well as $(Dy/Yb)_N$ = 1.1, suggesting that the T_1 granitoids were formed by melting of the subducted A'nyemaqen oceanic crust with terrigenous sediments under the amphibolite facies conditions in a syn-collisional setting. The T_2 and T_3 granitoids may be originated from a relatively homogeneous source with almost consistent mean values of I_{Sr} (0.7136 [T_2], 0.7094 [T_3]), $\varepsilon_{Nd(t)}$ (-5.83 [T_2], -5.97 [T_3]) and $\varepsilon_{Hf(t)}$ (-3.52 [T_2], -3.58 [T_3]). They present garnet signature of adakitic rocks and can be explained by partial melting of the juvenile mafic lower continental crust and mixing with upper crustal components during magma ascent. This process is considered to be associated with post-collisional extension which induced by asthenosphere upwelling and mantle melting, providing heat for mafic lower crust melting to form T_2

and T_3 granitoids. The T_1 granitoids with mantle signatures (e.g. $\varepsilon_{Hf(t)} > 0$) as well as BCC-like compositions represent a net flux of juvenile dioritic to granitic materials adding to the continental crust, in support of the hypothesis of "continental collision zones are primary sites for net continental crust growth" along the EKOB. The genetic link between T_2 and T_3 granitoids means the EKOB had transformed to post-collisional setting since the middle Triassic (~ 247 Ma). All these hypotheses are conceptually important for understanding the origin of the juvenile crust and continental crustal growth through magmatism from syn-collisional to post-collisional settings.

Keywords: East Kunlun Orogenic Belt; syn-collisional granitoids; post-collisional granitoids; continental crust growth; A'nyemaqen Ocean

1. Introduction

Granitoid batholiths are the most abundant constituent of continental crust, and their origin is essential for understanding the evolution and differentiation of the continental crust (Luo *et al.*, 2015). The East Kunlun Orogenic Belt (EKOB) as one of the major magmatic belts on the Tibetan Plateau is considered to undergo multi-cycle tectonic evolution and record the evolution of the Proto- and Paleo-Tethys oceans from the Early Paleozoic to the Early Mesozoic (Mo *et al.*, 2007; Xiong *et al.*, 2014). Widespread granitoid batholiths of two episodes, i.e. Early Paleozoic and Late

Paleozoic-Triassic in the EKOB formed in volcanic-arc, syn-collisional and post-collisional settings have supplied much information to help understand the geological evolution of the Proto- and Paleo-Tethys oceans over the past three decades (Jiang *et al.*, 1992; Huang *et al.*, 2014; Shao *et al.,* 2017). We focus on the Triassic granitoids in the eastern section of the EKOB where previous studies have revealed the history of Paleo-Tethys orogenic events of 270-195 Ma (Chen *et al.*, 2007; Sun *et al.*, 2009; Xiong *et al.*, 2012; Zhang *et al.*, 2012; Xu *et al.*, 2015; Ren *et al.,* 2016). Despite the large number of recent studies on the tectonic evolution of the Paleo-Tethys, many uncertainties and controversies remain (Mo *et al.,* 2007*;* Pan *et al.,* 2012*;* Xia *et al.,* 2015). Moreover, these studies concentrated on scattered/individual intrusions, lacking systemic chronology, geochemistry, and especially isotopic investigations.

Continental crustal growth is widely considered to be associated with the subduction-zone magmatism because of the arc-like incompatible element signature of the bulk continental crust (BCC) (e.g., enrichment in U, K and Pb, depletion in Nb, Ta and Ti; Niu *et al.*, 2013), which is termed “island arc model” (Taylor, 1967). However, the standard “island arc model” has many difficulties (Niu & O'Hara, 2009; Niu *et al.*, 2013), aiming at which Niu and co-workers proposed a testable hypothesis “continental collision zones are primary sites of net continental crustal growth”. Because of this and on the basis of their detailed studies of the Linzizong syncollisional volcanic sequence in southern Tibet (Mo et al., 2008), this hypothesis has been tested with success in several orogenic belts (Mo *et al.*, 2008; Niu & O'Hara, 2009; Niu *et al.*, 2013; Huang

84 *et al.*, 2014; Zhang *et al.,* 2016; Kong *et al.,* 2017; Shao *et al.,* 2017), including the
85 preliminary study of the EKOB syn-collisional granitoids. Furthermore, comparison
86 between the syn- and post-collisional granitoids in the Triassic offers important
87 information for the continental crust growth in different context of the orogenesis.
88 In this paper, we systemically present new geochemical and geochronological data
89 for the granitoids in the eastern section of the EKOB. Our new data, together with age
90 data from the literature, shed light on the petrogenesis of the granitoids from the syn-
91 to post-collisional settings and provide constraints on the evolution of the Paleo-Tethys
92 oceans and continental crustal growth in the EKOB.
93
94 **2. Geological setting and samples**
95 The EKOB is located south of the Kunlun-Qaidam terrane constrained between
96 the south Qilian suture (SQS) to the north and the A'nyemaqen-Kunlun-Mutztagh
97 suture (AKMS) in the south on the northern Tibet Plateau. The globally unique Tibetan
98 Plateau has been amalgamated through multiple continental collision events, resulting
99 in progressively younger sutures from northeast (Early Paleozoic) to southwest
100 (Cenozoic) (Fig. 1a; Niu *et al.*, 2013). The northern Tibet Plateau is thus the ideal site
101 for studying the processes of continental collision and geological consequences. The
102 EKOB is bounded by the Qaidam Basin to the north and Baryan Har-Songpan Ganze
103 terrane (BH-SG) to the south, extending W-E for ~ 1500 km (Fig. 1b; Jiang *et al.*, 1992;
104 Ding *et al.*, 2014). Three sub-units (the northern, middle and southern zone) have been

recognized in the EKOB, separated by two subparallel faults (Central and South Kunlun Fault) offset by the Altun Tagh sinistral strike-slip fault in the west (Fig. 1b; Jiang *et al.*, 1992; Liu *et al.*, 2004). Our study area is in the eastern section of the EKOB with middle and southern magmatic zones well exposed. The middle zone is dominated by Late Paleozoic and Triassic granitoids. It also contains Precambrian metamorphic basement, Devonian sandstones, conglomerates and Carboniferous marine limestones and clastic sedimentary rocks. Comparatively, the southern zone is a more complex fold belt, which is divided into western Late Paleozoic fold group, middle Precambrian uplift and eastern Paleozoic and Triassic fold group. The basement is dominated by the Mesoarchean-Mesoproterozoic Jinshuikou Group in the middle zone and the Paleoproterozoic Kuhai Group in the southern zone (Fig. 1c). The Jinshuikou Group, which comprises the lower Mesoarchean-Paleoproterozoic Baishahe Formation and the upper Mesoproterozoic Xiaomiao Formation (Wang *et al.*, 2004; Chen *et al.,* 2011). The Jinshuikou Group is consist of gneisses, marbles, greenschists, amphibolites, migmatites and quartzites and had been involved into later granulite-facies metamorphism and anatexis during 460-402Ma (Zhang *et al.*, 2003; Wang *et al.*, 2007). The Kuhai Group is a metamorphic rock series of amphibolite facies forming during 2330-1441 Ma, which consist of gneiss, plagioclase amphibolite, quartz schist and migmatites (Wang *et al.*, 2007). Two ophiolite belts distributed along the Central and South Kunlun Fault, the Nuomuhong-Qingshuiquan ophiolite belt (NQO) extending along the south of the Central Kunlun Fault and the A'nyemaqen ophiolite belt (AMO)

126 in the southern margin of the EKOB (Bian *et al.*, 2001; Jiang *et al.*, 1992; Zhang *et al.*,
127 2012). It is accepted that the NQO formed in the Early Paleozoic, indicating the closure
128 of Central Kunlun Ocean (Yang *et al.*, 1996; Li *et al.*, 2012). While the formation age
129 of the AMO is relatively complex. Many studies have confirmed that the AMO
130 represent two stages of ocean closing, i.e. Early Paleozoic (~467Ma) Qinling-Qilian-
131 Kunlun Ocean (Bian *et al.*, 2004); Late Paleozoic to Middle Triassic (~308-260Ma)
132 Paleo-Tethys Ocean (Jiang *et al.*, 1992; Yang *et al.*, 2009).

133 The EKOB preserves geological records of the Early Paleozoic Caledonian cycles
134 and the Late Paleozoic to Early Mesozoic Variscan-Indosinian cycle. The Early
135 Paleozoic (500-400 Ma) granitoids are comparable with those in the North Qilian
136 orogenic belt (Mo *et al.,* 2007). The Permian-Triassic granitoids are dominant in the
137 EKOB (~25,000 km^2) and account for 50% of the total outcrop area of the granitoids
138 (Fig. 1c, Jiang *et al.,* 1992; Zhang *et al.,* 2014).

139 Samples of this study are collected from the eastern section of the EKOB, which
140 extend W-E for ~ 300 km and distribute in the Middle and Southern Zones (Fig. 1c).
141 They are intermediate- to coarse- grained granites to granodiorites (Fig.2) with mineral
142 assemblage of quartz + potash feldspar + plagioclase + amphibole + biotite + magnetite
143 and accessory minerals of apatite and zircon (see Table 1 for detail). Fine-grained
144 dioritic mafic magmatic enclaves (MMEs) are dispersed in the pluton (Fig.2a, b).
145

146 **3. Analytical methods**

147 *3.1. Zircon U-Pb dating*

148 Zircons from seven samples were separated for U-Pb dating using combined
149 methods of heavy liquid and magnetic techniques. The zircon internal structure was
150 examined using cathodoluminescence (CL) imaging on a FEI Quanta 450 FEG
151 scanning electron microscope (SEM) at the State Key Laboratory of Geological
152 Processes and Mineral Resources (GPMR), China University of Geosciences, Wuhan.
153 U-Pb dating was conducted by laser ablation inductively coupled plasma mass
154 spectrometry (LA-ICP-MS) at the same laboratory, using a 32 μm spot size. Zircon
155 91500 was used as the external standard (Wiedenbeck *et al.*, 1995). Each block of 6
156 unknowns was bracketed by analyses of the standards. Off-line selection and integration
157 of background and analyte signals, and time-drift correction and quantitative calibration
158 for trace element analyses and U-Pb dating were performed by ICP-MS-Data-Cal (Liu
159 *et al.*, 2010). Eighteen points of each sample were chosen for LA-ICP-MS U-Pb
160 analysis and their results are given in Appendix A. Concordia diagrams and weighted
161 mean calculations were processed using the Isoplot/Ex_version 4.15 program (Ludwig,
162 2012). Concordia diagrams and representative CL images of seven samples are shown
163 in Figure 3.

164

165 *3.2. Major and trace elements*

166 Twenty-five freshest samples were chosen for elemental analysis. Weathered
167 surfaces and pen saw marks were removed, cleaned, and ultrasonically cleaned with

Milli-Q water and dried before powdered using an agate mill into ~200-mesh in a clean environment.

Major elements were determined using a Leeman Prodigy inductively coupled plasma-optical emission spectroscopy (ICP-OES) system with high dispersion Echelle optics at China University of Geosciences, Beijing (CUGB). The precision (1σ) based on rock standards BCR-1, AGV-2 and GSR-3 is estimated as ~ 1.5% for TiO_2, ~ 2.0% for P_2O_5 and better than 1% for other major oxides. And trace elements were analyzed using an Agilent-7500a inductively coupled plasma mass spectrometry (ICP-MS) at CUGB. The BCR-1 and BHVO-1 were used to calibrate the elemental concentrations of the samples. The analytical precision was generally better than 5% for most trace elements, 10-13% for Cu, Sc, Nb, Er, Th, and U, and 10-15% for Ta, Tm and Gd. Full sample preparation techniques and other details are described by Song *et al.* (2010). Results are listed in Table 3.

3.3. Sr-Nd-Hf isotopes

The whole-rock Sr-Nd-Hf isotopic composition of twenty-two samples were determined at GPMR. Sr and Nd isotopic ratios were measured on a Thermo Finnigan Triton Ti thermal ionization mass spectrometer (TIMS), Analytical details are given in Gao *et al.* (2004). And Hf isotopic analysis was conducted using a multi-collector inductively coupled plasma mass spectrometer (MC-ICP-MS) with a Thermo Neptune Plus system. The procedures of chemical separation and analysis are following Yang *et*

189 *al.* (2010b). The isotopic data are presented in Table 4.

190

191 **4. Results**

192 *4.1. Zircon U-Pb ages*

193 Zircons from all the analytical samples are transparent, light brown euhedral
194 columnar crystals. They have faint (Fig. 3a), oscillatory (Fig. 3b, c, f, g), sector (Fig.
195 3e) and transitional (Fig. 3d) zoning, which is consistent with a magmatic origin (Corfu
196 *et al.*, 2003). The LA-ICP-MS U–Pb analysis gave variable Th (101-4815ppm) and U
197 (115-8483ppm) concentrations with Th/U ratios of 0.15-1.21, which are of a magmatic
198 origin (Hoskin & Schaltegger, 2003). Thus, the youngest U-Pb age group of the zircons
199 represents the crystallization age. All the data are given in Appendix A.

200 Zircons from sample AKDL12-01 and NSK12-09 yield weighted mean $^{206}Pb/^{238}U$
201 ages of 251.4 ± 6.8Ma (MSWD=4.1, n=9) and 247.6 ± 4.8Ma (MSWD=10.3, n=10),
202 respectively. And zircons from sample BLX12-09 and BLX12-03 yield concordia ages
203 of 250.5 ± 0.74Ma (MSWD=3.5, n=16) and 247.8 ± 2.1Ma (MSWD=1.12, n=13),
204 respectively. Inherited /Captured zircon cores plotted along the concordia yield age
205 group of ~ 900Ma and ~ 500-300Ma (Figs. 3a-d), closing to the age of crust basement
206 of the EKOB (Xiaomiao Formation) and later metamorphism as well as anatexis event
207 (Wang *et al.*, 2004). Zircons from sample HYC12-01 and HXNC12-01 yield weighted
208 mean $^{206}Pb/^{238}U$ ages of 245.5 ± 9.2Ma (MSWD=9.6, n=8) and 237.8 ± 3.8Ma
209 (MSWD=4.6, n=15), and the concordia age 227.1 ± 0.74Ma (MSWD=0.84, n=17) is

given by zircons from sample BLXD12-06, respectively. Our new data can be considered as approximating the emplacement ages of the granitoids in the EKOB, which, together with the high-quality zircon U-Pb ages reported in the recent literature (summarized in Table 1), indicate that these granitoids continually formed from early Triassic (~ 251Ma) to late Triassic (~ 214 Ma).

4.2. Major and trace elements

The Triassic granitoids from the EKOB show high SiO_2 (60.06-75.83 wt.%) and low Na_2O+K_2O (4.72-9.17 wt.%) with a relatively narrow compositional range from diorite, granodiorite to granite in total alkali (Na_2O+K_2O)-SiO_2 space (Fig. 4a). In the (Na_2O+K_2O-CaO) against SiO_2 discrimination diagram (Fig. 4b), most of these rocks are restricted in the calc-alkalic field with scattered ones falling in the alkali-calcic and calcic fields. The early Triassic granitoids varying from low-K calc-alkaline series to high-K calc-alkaline series (K_2O = 0.47-4.47 wt.%). Comparatively, the middle Triassic granitoids fall in the calc-alkaline and high-K calc-alkaline fields (K_2O = 1.18-4.94 wt.%). The late Triassic granitoids with higher K_2O of 2.59-5.55 wt.% belong to high-K calc-alkaline and shoshonite series (Fig. 4c). They are all weakly metaluminous to peraluminous (A/NK = 1.11-2.34, A/CNK = 0.87-1.13), which are equivalent to I-type granitoids with only two samples straddling the I- to S-type boundary (Fig. 4d). As expected, these Triassic granitoids display decreasing trends in major elements with increasing SiO_2 (Fig. 5).

231 The Triassic granitoids invariably show enrichment of light rare earth elements
232 (LREE) with $(La/Yb)_N$ of 6.37-49.23. Most samples show significant negative Eu
233 anomalies with a few showing weak positive Eu anomalies (Eu/Eu* = 0.09-1.47; Table
234 2 and Figs. 6a-c). They share similar primitive mantle normalized patterns with
235 enrichment in Rb (although one sample of the early Triassic granitoids is distinct), K,
236 Pb and depletion in Nb, Ta, P, Ti, resembling the composition of the bulk continental
237 crust (BCC; Rudnick & Gao, 2003), as well as variable Sr anomalies (Sr/Sr* of 0.19-
238 1.06 for T_1 granitoids, 0.08-1.98 for T_2 granitoids and 0.03-1.11 for T_3 granitoids)
239 (Table 2).
240
241 *4.3. Sr-Nd-Hf isotopes*
242 Bulk-rock Sr-Nd-Hf isotopic data (Table 3) are plotted in Figures 7-10. Three
243 samples (AKDL12-04, DGL12-05 and DGL12-07) gives very high $^{87}Sr/^{86}Sr$ (0.7555-
244 0.8559) because of the high Rb/Sr (13.8-24.9) (Table 3), due to significant plagioclase-
245 dominated (and alkali feldspars to some content) fractional crystallization, resembling
246 peralkaline rhyolites (Shao *et al.*, 2015). Such high Rb/Sr resulting in high radiogenic
247 ^{87}Sr ingrowth, which makes the calculated $I_{Sr}(t)$ unreliable (Wu *et al.*, 2000).
248 Eliminating these samples with inaccurate I_{Sr} (not shown in the Figs.7-10), the Triassic
249 granitoids have variable I_{Sr} of 0.7067-0.7327.The $\varepsilon_{Nd(t)}$ and $\varepsilon_{Hf(t)}$ of all samples range
250 from -9.21 to -1.66 and -7.50 to 3.59, respectively.
251

5. Discussion

5.1 FC (fractional crystallization) or AFC (assimilation and fractional crystallization)

The data shown in SiO_2-variation diagrams (Fig. 5) are to a first-order consistent with varying extent of fractional crystallization of hornblende, plagioclase, Fe–Ti oxides and apatite. However, these trends are also consistent with modal variations of these phases in the samples although the depletion in P, Ti, Sr and Eu emphasizes the significance of fractional crystallization.

Crustal assimilation and fractional crystallization (AFC) of mantle-derived mafic magma generally produce continuous compositions from basaltic to felsic and typical linear trends between SiO_2 and I_{Sr} / $\varepsilon_{Nd(t)}$, because of high $^{87}Sr/^{86}Sr$ and low $^{143}Nd/^{144}Nd$ in the upper continental crust. There is a narrow spectrum of rock composition (Fig. 4a) and lack of linear trends in plots of SiO_2 against I_{Sr} and $\varepsilon_{Nd(t)}$ (Fig. 7a-b). Additionally, they have lower K_2O/Na_2O than the melts produced by assimilation experiments simulating the reaction between basalts and felsic pelitic gneiss in the crust (Castro, 2001) (Fig. 7d). These suggest that the AFC from common parental mantle-derived mafic magmas is unlikely an important mechanism for the Triassic granitoids from the EKOB. Alternatively, they show variable REE patterns and a wide range of bulk-rock Sr-Nd-Hf isotopic compositions from the early to late Triassic (Figs. 6a-c and 7a-c), which may indicate different sources and/or different petrogenetic processes. However, any petrogenetic model must satisfy the observation that the Triassic granitoids have hybrid mantle-crust geochemical signatures, as evidenced by their Sr-Nd-Hf isotopic

273 compositions (Figs. 8, 9 and 10).

274

275 5.2. Petrogenesis

276 S-, I-, A- and M-type granites are widely used classifications on the basis and
277 sources and petrogenesis (Collins *et al.*, 1982; Whalen *et al.*, 1987; Chappell & White,
278 2001). Chemically, the Triassic granitoids from the EKOB have relatively low A/CNK
279 values (<=1.1; Figure 4d)and amphiboles and biotite are common (Fig. 2), which accord
280 with features of I-type granites (Chappell & White, 2001). Therefore, the Triassic
281 granitoids from the EKOB are I-type granites, to be exact, they are calc-alkaline I-type
282 granitoids as shown in Figure 4. Such rocks can form by (1) crustal assimilation and
283 fractional crystallization of mantle-derived mafic magma (AFC) (Chen & Arakawa,
284 2005; De Souza *et al.*, 2007); (2) partial melting of crustal materials at deep or shallow
285 (Laurent *et al.*, 2014; Liu *et al.*, 2015); (3) hybridization between mantle-derived mafic
286 magmas and crustal melts (Dong *et al.*, 2011); (4) partial melting of remained basaltic
287 ocean crust with sediment mélanges (Castro *et al.*, 2010; Huang *et al.*, 2014; Zhang *et*
288 *al.*, 2016; Shao *et al.*, 2017; Kong *et al.,* 2017).

289

290 5.2.1. Petrogenesis of the early Triassic (~ 251-248 Ma) granitoids

291 The early Triassic granitoids (T_1, ~ 251-248 Ma) are enriched in Rb, K, Pb and
292 depleted in Nb, Ta, Sr, P and Ti, which resembles the composition of the bulk
293 continental crust (BCC; Fig. 6d). Abnormally, the sample (AKDL12-03) display strong

294 negative Rb anomaly due to its extremely low content of Rb (7.7 ppm). It may be
295 attributed to some possible reasons such as strong alteration. The T_1 granitoids have
296 higher I_{Sr} (0.7067-0.7148), negative $\varepsilon_{Nd(t)}$ (-7.32 to -1.66) and negative to positive $\varepsilon_{Hf(t)}$
297 (-5.11 to 3.59) than the mature continental crust ($[^{87}Sr/^{86}Sr]_i = 0.73802$, $\varepsilon_{Nd}(t) = -17.0$,
298 $\varepsilon_{Hf}(t) = -15.5$; Shao *et al.,* 2017), suggesting significant mantle contribution (or juvenile
299 crustal material) in terms of isotopes. The Hf isotopic data of the T_1 granitoids and the
300 other Kunlun Triassic granitoids (Fig. 9; Ding *et al.,* 2015; Huang *et al.,* 2014; Xia *et*
301 *al.,* 2015; Zhang *et al.,* 2016; Shao *et al.,* 2017) are also indicative of significant mantle
302 contribution. To produce such andesitic to felsic BCC-like magmas with mantle
303 signature, it requires a basaltic source plus continental materials mentioned above. As
304 shown in many previous studies, the A'nyemaqen Ocean (the north branch of the Paleo-
305 Tethys Ocean recorded in the EKOB; Jiang *et al.*, 1992) had closed and recorded a syn-
306 collisional setting in the early Triassic (Huang *et al.*, 2014; Shao *et al.,* 2017), in which
307 setting the potential basaltic source may be the arc crust (island arc basalt, IAB) or the
308 subducted ocean crust (mid-ocean ridge basalt, MORB). Higher Sr/Sr* values of IAB
309 (~ 2.72) than that of BCC (0.933) and MORB (~ 0.1-2.0) have been used by Niu and
310 co-authors (Niu and O'Hara, 2009; Niu *et al.* (2013) to argue against IAB for the source
311 of T_1 granitoids (Sr/Sr* = 0.19-1.06). Therefore, the basaltic end-member for the source
312 of T_1 granitoids is most probably supplied by the subducted ocean crust (i.e. MORB).
313 Note that their flat HREE (mean $[Dy/Yb]_N$ = 1.1, the same as BCC) patterns and
314 constant $(Ce/Yb)_N$ with decrease Yb_N (Fig. 6a, d), requires melting under amphibolite

315 facies conditions (< 50 km; without garnet signature). Additionally, the trace element
316 patterns, as well as the isotopic composition, especially the Nd-Hf isotopes of the T_1(~
317 251-248 Ma) granitoids are consistent with the I-type syn-collisional granitoids derived
318 from melting of subducted oceanic crust with terrigenous sediments of upper
319 continental crust origin (Figs. 7a, b, c and 8, 9,10). Partial melting of the ocean crust
320 produces felsic melts and the ocean crust derived from the mantle not long ago imparts
321 the mantle isotopic signature (Niu *et al.,* 2013). Meanwhile, the addition of terrigenous
322 sediment can explain the crustal signatures of the T_1 granitoids. Many lines of evidence
323 above suggest the T_1 (~ 251-248 Ma) granitoids are derived from partial melting of
324 remaining fragments of the subducted A'nyemaqen ocean crust with recycled
325 terrigenous sediment under amphibolite facies conditions. The underthrusting cold
326 Anyemaqen Ocean crust evolves along a high T/P geothermal path and has longer time
327 to absorb heat from the prior hot active continental margin. The highly hydrated ocean
328 crust (along with terrigenous sediments) begins to melt when it reaches the hydrous
329 basaltic solidus (<700 °C) under amphibolite conditions (see details in Mo *et al.,* 2008;
330 Niu *et al.,* 2013). The melts subsequently underwent plagioclase-controlled
331 fractionation within crustal reservoirs supported by relatively large decrease in Al_2O_3,
332 CaO (Fig. 5b, e) and Sr, Eu (Fig. 7e, f) with small increase in SiO_2. It should be noted
333 that the initial isotopic ratio (Isr, $\varepsilon_{Nd}(t)$ and $\varepsilon_{Hf}(t)$) variations within individual outcrops
334 largely reflect small scale isotopic heterogeneity due to incomplete homogenization of
335 melts affected by crustal assimilation or modal variation or both. This is because

granitoid magmas under sub-liquidus conditions are “crystal mashes”, complete homogenization is thus restricted by the efficient diffusion (Ramos and Reid, 2005). Inherited/captured zircon cores plotted along the concordia yield age group of ~ 900Ma and ~ 500-300Ma (Figs. 3a-d), closing to the age of crust basement of the EKOB (Xiaomiao Formation) and later metamorphism (Wang *et al.,* 2004), implying that they were involved in the ~250 Ma magmatism. Terrigenous sediments melted alongside oceanic crust also result in the isotope heterogeneity.

Studies of the Linzizong volcanic succession (LVS) in southern Tibet have testified that juvenile continental crust is produced via partial melting of the ocean crust under the amphibolite facies conditions and preserved as ‘net crust growth’ in the collision zones (Niu *et al.,* 2013; Mo *et al.,* 2008; Niu and O'Hara, 2009). The syn-collisional East Kunlun granitoids provide more geochemical and isotopic data to support this hypothesis. Firstly, the bulk compositions of granitoids are similar to the bulk continental crust with almost identical Nb-Ta-Ti and Sr and Eu depletion (Fig. 6a, d). Secondly, the whole rock Sr-Nd-Hf isotopes can be explained by partial melting of subducted Paleo-Tethyan MORB with terrigenous sediment addition in the melting region under the amphibolite-facies conditions (see above). Though the granitoids in the EKOB have relatively unradiogenic Nd isotope ($\varepsilon_{Nd}(t)<0$) compared to the LVS, they have high $\varepsilon_{Hf}(t)$ values (up to 3.59), which is the more convincing evidence that they are juvenile crust newly formed from ocean crust melting with inherited mantle isotopic signatures and crustal/sediment contribution.

357

358 5.2.2. Petrogenesis of the middle (~ 247-238 Ma) and late Triassic (~ 234-214 Ma)

359 granitoids

360 Since the middle Triassic (T_2, ~ 247-238 Ma) and late Triassic (T_3, ~ 234-214 Ma)

361 granitoids display similar trace elemental and isotopic features (Figs. 6-10), we hence

362 discuss them together here. The T_2 and T_3 granitoids have almost consistent mean

363 values of $\varepsilon_{Nd(t)}$ (-5.83 [T_2], -5.97 [T_3]) and $\varepsilon_{Hf(t)}$ (-3.52 [T_2], -3.58 [T_3]), indicating a

364 similar source. REE patterns of the T_2 and T_3 granitoids show weakly concave upward

365 between middle and heavy REEs (Fig. 6b-c), suggesting amphibole as a residual phase

366 in the source due to its high partition coefficients for middle to heavy REEs, especially

367 the highest D(Dy) in intermediate to felsic melts (Rollinson, 1993). Meanwhile, the

368 multi-element patterns of the T_2 and T_3 granitoids show obvious decrease from Dy to

369 Yb (Fig. 6e-f), which imply garnet may also be a residual phase in their source ($Kd_{Yb}^{Grt/L}$

370 = 39.9; Arth, 1976). The existence of garnet as a residual phase may be responsible for

371 the adakitic signature (Castillo, 2012) of the T_2 and T_3 granitoids. Most of the T_2 and

372 T_3 granitoids have relatively high Sr/Y and La/Yb, plotting in the field restricted to

373 adakite (Figs. 7h-i). Petrogenesis of adakitic rocks are considered as follows: (1) related

374 to the melting of subducted slab directly or indirectly (Martin *et al.*, 2005); (2) melting

375 of the mafic lower continental crust (e.g. Atherton & Petford, 1993;Goss & Kay, 2009);

376 (3) high pressure fractionation of garnet-bearing normal arc magma (Macpherson *et al.*,

377 2006). Their high $^{87}Sr/^{86}Sr$ (0.7107-0.7523), low $^{143}Nd/^{144}Nd$ (0.5120-0.5123) are not a

typical signature of melting of subducted slab (Castillo, 2012). Considering the lithostratigraphic records, these granitoids cannot be interpreted as product of arc magmatism (*Jiang et al.,* 1992). Therefore, partial melting of the lower continental crust may account for the origin of the granitic rock. First, the T_2 and T_3 granitoids from the EKOB have high $(La/Yb)_N$ (Fig. 6b,c) and high K_2O/ Na_2O ratios (Fig. 7d), which are consistent with the composition of adakitic rocks inferred to be derived from partial melting of the lower continental crust. Second, the T_2 and T_3 granitoids from the EKOB have Sr ($^{87}Sr/^{86}Sr$, 0.7107-0.7523), Nd ($\varepsilon_{Nd(t)}$, -5.83 [T_2], -5.97 [T_3]) and Hf ($\varepsilon_{Hf(t)}$, -3.52 [T_2], -3.58 [T_3]) isotopes similar to the T_1 grantoids (Fig. 8, 9), are likely indicative of significant juvenile mafic continental crust. Meanwhile, according to the mass balance, the T_2 and T_3 granitoids requires some input of mature crustal material (Fig.10b). Finally, the overlapping Sr-Nd-Hf isotope compositions of the T_2 and T_3 granitoids are consistent with the I-type granitoids derived from the lower continental crust as well as the sub-volcanic rocks derived from crust-mantle mixing (Fig. 8; Zhang *et al.*, 2012; Ding *et al.*, 2011; Hu *et al.*, 2016), which also suggest the T_2 and T_3 granitoids could be produced by melting of the juvenile mafic lower continental crust, mixing with the upper continental crust components during ascent. In which a newly lower crust thickened (probably ~ 50km with lowest temperature gradient of 20°C km^{-1}; Atherton & Petford, 1993) by underplating under garnet stability conditions played an important role, resulting in adakitic features of the T_2 and T_3 granitoids. They likely experienced fractional crystallization of plagioclase (Figs. 5b, e; 7e-f). Furthermore, high partition

399 coefficient of Sc but low value of Th in biotite will result in lower Sc/Th of residual
400 melts during biotite fractionation (Bea *et al.*, 1994). Negative correlation between
401 Sc/Th and SiO_2 denotes the biotite fractionation in T_2 and T_3 granitoids (Fig. 7g). In
402 fact, melting the lower crust clearly need a heat source from below. Post-collisional
403 setting of the late Triassic has been widely accepted in the EKOB (e.g. Pan *et al.*, 2012;
404 Wang *et al.*, 2014a). Many mechanisms such as mantle plume (Chung & Jahn, 1995),
405 slab break-off (Maury *et al.*, 2000) or convective lithosphere removal (Hoernle *et al.*,
406 2006) can cause asthenospheric upwelling, decompression melting, triggering crustal
407 melting. Accordingly, we prefer the more reasonable post-collisional extension and
408 related orogenic collapse for inducing asthenosphere upwelling, and then leading to
409 mafic lower crust melting to form T_2 and T_3 granitoids. The age range of the T_3
410 granitoids (~ 234-214 Ma) are similar to the late Triassic mafic dikes and felsic volcanic
411 rocks (228-218 Ma) associated with post-collisional extension in the EKOB (Hu *et al.*,
412 2016) suggest these post-collisional magmatism is an important tectono-magmatic
413 event which has influenced both the lower crust and the upper mantle. More importantly,
414 the identical genetic link between T_2 and T_3 granitoids has pushed this event back to
415 ~247 Ma, that is the EKOB had transformed to post-collisional setting since the middle
416 Triassic (~ 247 Ma).
417
418 5.3 An integrated model from syn-collisional to post-collisional settings
419 Our integrate model implies a process of partial melting of the subducted

A'nyemaqen ocean crust with terrestrial sediments in a syn-collisional setting (~ 251-248 Ma) and partial melting of juvenile mafic lower crust mixing with the upper crust components in a post-collisional setting (~ 247-214 Ma). Figure 10 shows mixing calculations for the Triassic granitoids in the EKOB. The T_1 (~ 251-248 Ma) granitoids represent mixing of at least ~ 50% A'nyemaqen MORB (Guo *et al.*, 2007) and ~ 50% terrigenous sediments represented by the Jinshuikou granites derived from the Proterozoic basement of Qaidam terrane (Yu *et al.*, 2005). Comparably, the T_2 (~ 247-238 Ma) and T_3 (~ 234-214 Ma) granitoids are best explained as hybrid magmas of ~ 55% juvenile mafic lower crust (Hu *et al.*, 2016) and ~ 45% upper continental crust materials (also logically represented by the Jinshuikou granitoids; Yu *et al.*, 2005). Melting of A'nyemaqen ocean crust fragments with recycled terrigenous sediment under amphibolite facies conditions resulted in the T_1 granitoids with mantle signatures (e.g. $\varepsilon_{Hf(t)} > 0$), which share similar compositions with the BCC. This process added a net flux of juvenile dioritic to granitic materials to the continental crust, pointing to the significance of the oceanic crust melting for continental crust accretion (Fig.11a). T_2 and T_3 granitoids are associated with post-collisional magmatism involving participation of juvenile continental crust. Asthenosphere upwelling and decompression melting would provide heat for juvenile mafic lower crust melting, then mixing with upper crust during post-collisional extension (Fig.11b). These hypotheses are conceptually important for understanding the origin of the juvenile crust and continental crustal growth through magmatism from syn-collisional to post-collisional settings.

441 Further research is needed to test these hypotheses.

442

443 **6. Conclusion**

444 1. The zircon U-Pb dating yields ages of ~ 251-248 Ma, ~ 247-238 Ma to ~ 234-214
445 Ma for the granitoids of the East Kunlun Orogenic Belt. These age data, together with
446 the literature data, suggest that the granitoids are products of syn-collisional (~ 251-248
447 Ma) and post-collisional (~ 247-238 Ma to ~ 234-214 Ma) magmatism during or shortly
448 after the closure of the A'nyemaqen Ocean.

449 2. The early Triassic (~ 251-248 Ma) granitoids are best explained by partial melting of
450 the subducted A'nyemaqen ocean crust fragments with recycled terrigenous sediment
451 under amphibolite facies conditions. The contribution of the ocean crust at least ~ 50%
452 to the parental melts of the T_1 granitoids, which subsequently underwent fractionation.

453 3. The middle (~ 247-238 Ma) and late (~ 234-214 Ma) Triassic granitoids with adakitic
454 features could be produced by melting of the juvenile mafic lower continental crust,
455 mixing with the upper continental crust components during ascent, in which the relative
456 proportion of juvenile mafic lower crust may be up to ~ 55% to their parental melts.
457 This process is more reasonably associated with post-collisional extension for inducing
458 asthenosphere upwelling, leading to mafic lower crust melting to form the granitoids.

459 4. The early Triassic (~ 251-248 Ma) granitoids with mantle signatures (e.g. $\varepsilon_{Hf(t)} > 0$)
460 as well as BCC-like compositions represent a net flux of juvenile dioritic to granitic
461 materials adding to the continental crust, pointing to the significance of melting of

462 ocean crust for continental crust accretion in the EKOB. And the identical genetic link
463 between T_2 (~ 247-238 Ma) and T_3 (~ 234-214 Ma) granitoids means the EKOB had
464 transformed to post-collisional setting since the middle Triassic (~ 247 Ma).
465
466
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477
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Figure Captions

735

736 **Fig. 1** (a) Geological framework of the Greater Tibetan Plateau showing the major
737 tectonic units and sutures as follows (from northeast to southwest): NQS, North Qilian
738 suture; DHS, Danghe Nan Shan suture; SQS, South Qilian suture; AKMS,
739 A'nyemaqen-Kunlun-Mutztagh suture; JS, Jinsha suture; BNS, Bangong-Nujiang
740 suture; IYS, Indus-Yarlung Zangbo suture, revised after (Hu *et al.*, 2016). (b) Outline
741 of the East Kunlun Orogenic Belt with two major faults (Central Kunlun Fault and
742 South Kunlun Fault) and three magmatic zones (Northern, Middle and Southern Zone)
743 (Hu *et al.*, 2016). (c) Simplified geological map of east section of the East Kunlun
744 Orogenic Belt (modified after XACGS, 2009). U-Pb ages shown for the granitic plutons
745 are new data of this study and from the recent literature indicated with superscript
746 numerals: 1. Ding *et al.* (2014), 2. Xiong *et al.* (2012), 3. Xia *et al.* (2014a), 4. Xia *et*
747 *al.* (2014b), 5. Sun *et al.* (2009), 6. Xu *et al.* (2015) and 7. Ding *et al.* (2011).

748

749 **Fig. 2** (a, b) An outcrop of the Triassic granitoids with some mafic magmatic enclaves
750 (MMEs) from the East Kunlun Orogenic Belt. Photomicrographs of (c) biotite
751 monzogranite, (d) syenogranite, (e) biotite moyite, (f) quartz diorite, (g) granodiorite,
752 (h) granite porphyry. Kfs = potash feldspar; Qtz = quartz; Pl = plagioclase; Bt = biotite;
753 Amp = amphibole.

754

755 **Fig. 3** Concordia diagrams of zircon U-Pb age data and cathodoluminescence images

of representative zircons for (a) AKDL12-01, (b) BLX12-09, (c) BLX12-03, (d) NSK12-09, (e) HYC12-01, (f) HXNC12-01 and (g) BLXD12-06 from the Triassic granitoids in the East Kunlun Orogenic Belt. Various amounts of inherited/captured old zircons of ~ 900 Ma and ~ 500-300 Ma in the early Triassic granitoids may indicate the presence and involvement of old crustal materials.

Fig. 4 Plots of (a) Na_2O+K_2O against SiO_2 (Middlemost, 1994); (b) Na_2O+K_2O-CaO against SiO_2 (Frost *et al.*, 2001); (c) K_2O against SiO_2 (Rickwood, 1989); (d) A/NK (molar $Al_2O_3/[Na_2O+K_2O]$ against A/CNK (molar $Al_2O_3/[CaO+Na_2O+K_2O]$) (Maniar & Piccoli, 1989; Chappell & White, 2001) for the Triassic granitoids from the East Kunlun Orogenic Belt.

Fig. 5 SiO_2 variation diagrams of the granitoids from the East Kunlun Orogenic Belt.

Fig. 6 (a-c) Normalized rare earth element (REE) and (d-f) multi-element patterns of the Triassic granitoids from the East Kunlun Orogenic Belt. Bulk continental crust (BCC; Rudnick & Gao, 2003) composition is also plotted for comparison. Chondrite and primitive mantle data are from Sun and McDonough (1989). The shaded fields in (a) and (d) are of I-type syn-collisional granitoids derived from subducted oceanic crust in the East Kunlun Orogenic Belt (Huang *et al.*, 2014).

Fig. 7 SiO_2 variation diagrams of (a) I_{Sr}, (b) $\varepsilon_{Nd(t)}$, (c) $\varepsilon_{Hf(t)}$, (d) K_2O/Na_2O (purple filled circles restricted in the purple dash line are products of the assimilation experiments that simulate the reaction between basalts and felsic pelitic gneiss in the crust; (Castro, 2001), (e) Sr, (f) Eu and (g) Sc/Th for the Triassic granitoids from the East Kunlun Orogenic Belt. Plots of (h) Sr/Y vs. Y and (i) La/Yb vs. Yb for these samples are used to distinguish adakitic rocks from normal arc andesite, dacite and rhyolite (Castillo, 2012).

Fig. 8 Diagrams of $\varepsilon_{Nd(t)}$ vs. I_{Sr} (a) and $\varepsilon_{Hf(t)}$ vs. $\varepsilon_{Nd(t)}$ for the Triassic granitoids from the East Kunlun Orogenic Belt. The field of I-type syn-collisional granitoids derived from subducted oceanic crust is based on Huang *et al.* (2014) and Zhang *et al.* (2016). The data for the sub-volcanic rocks derived from crust-mantle mixing are from Ding *et al.* (2011) and Hu *et al.* (2016). And the I-type granitoids derived from the lower crust and S-type granitoids derived from the upper crust are from Zhang *et al.* (2012) and Ba *et al.* (2012), respectively.

Fig. 9 The $\varepsilon_{Hf}(t)$ vs. Ages of the Triassic granitoids of the East Kunlun orogenic belt. Our new data and the literature data are indicated as the following: light blue diamonds (whole rock Hf data, West Kunlun; Zhang *et al.,* 2016); green triangles (whole rock Hf data, East Kunlun; Huang *et al.,* 2014); purple crosses and blue crosses (zircon Hf data, East Kunlun; Ding *et al.,* 2015; Xia *et al.,* 2015) and gay circles (whole rock Hf data,

798 East Kunlun; Shao *et al.,* 2017).

799

800 **Fig. 10** Mixing trends calculated with $\varepsilon_{Nd(t)}$ and $^{87}Sr/^{86}Sr$ composition for (a) the early
801 (~251-248Ma) and (b) middle (~247-238Ma) & late (~234-214Ma) Triassic granitoids
802 from the East Kunlun Orogenic Belt. T_1 granitoids could be explained by mixing
803 between A'nyemaqen MORB (average composition: $^{87}Sr/^{86}Sr = 0.707818$, $\varepsilon_{Nd(t)} = 12.9$,
804 Sr = 191.5 ppm, Nd = 8.9 ppm) (Guo *et al.*, 2007) and terrigenous sediments
805 (represented by the Jinshuikou S-type granites) (average composition: $^{87}Sr/^{86}Sr$ =
806 0.738282, $\varepsilon_{Nd(t)} = -11.3$, Sr = 260.1 ppm, Nd = 23.3 ppm) (Yu *et al.*, 2005). Comparably,
807 T_2 & T_3 granitoids could be interpreted as mixture of Juvenile mafic lower continental
808 crust (average composition: $^{87}Sr/^{86}Sr = 0.709589$, $\varepsilon_{Nd(t)} = -2.6$, Sr = 525.0 ppm, Nd =
809 33.2 ppm) (Hu *et al.*, 2016) and upper continental crust (represented by the Jinshuikou
810 S-type granites again).

811

812 **Fig. 11** Schematic illustration for the generation of the granitoids in EKOB during the
813 Triassic (modified from Shao *et al.,* 2017). See text for explanation.

Fig. 1

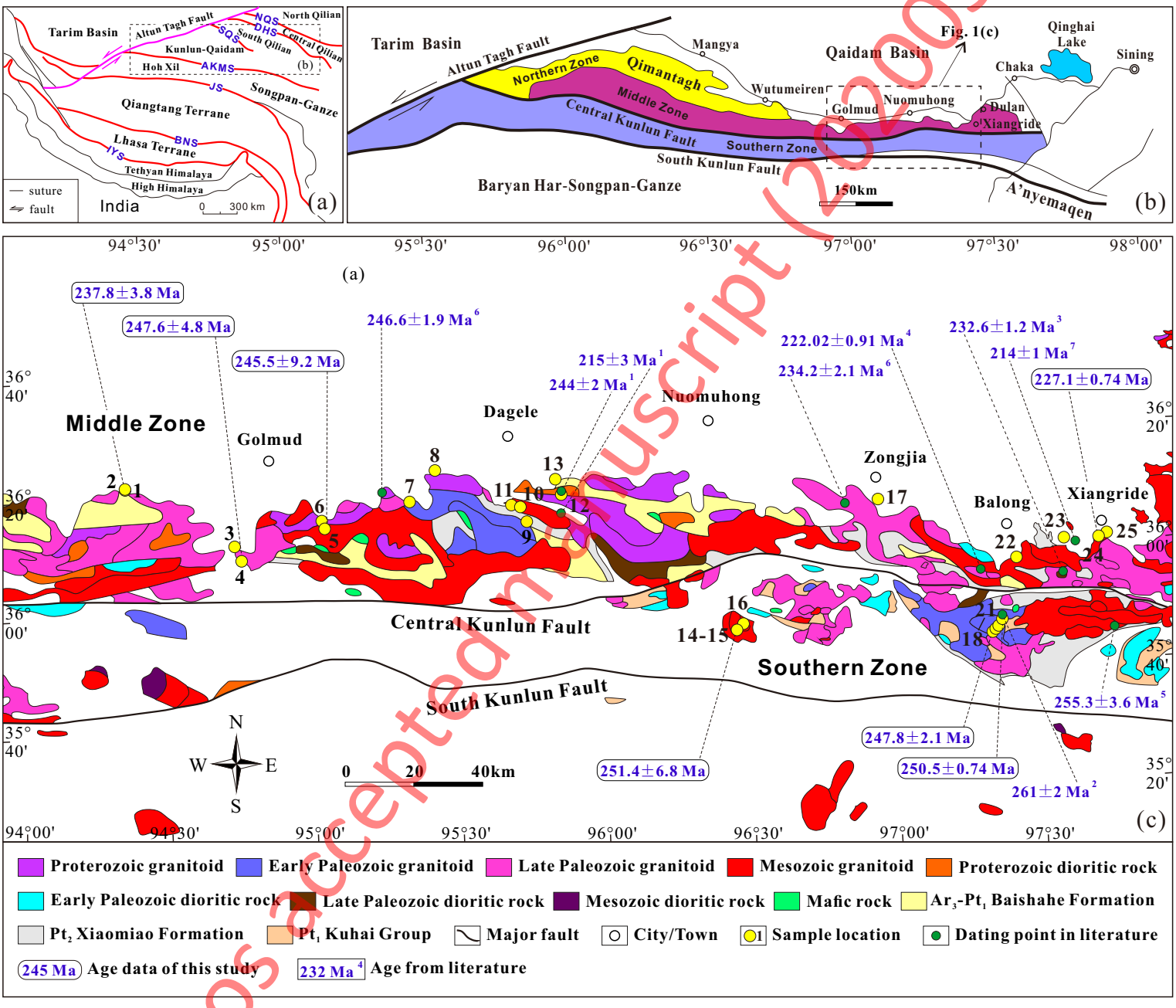
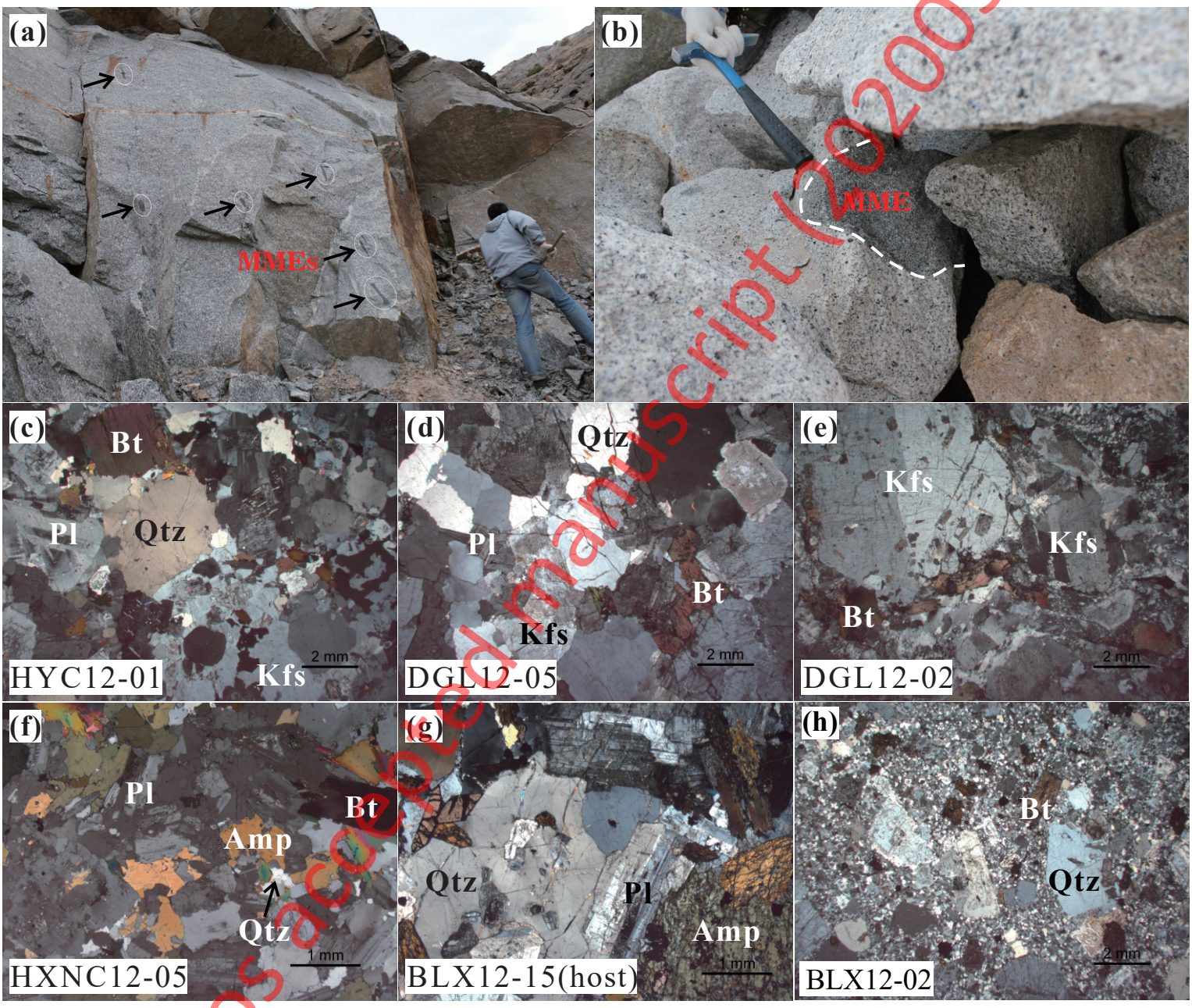


Fig. 2



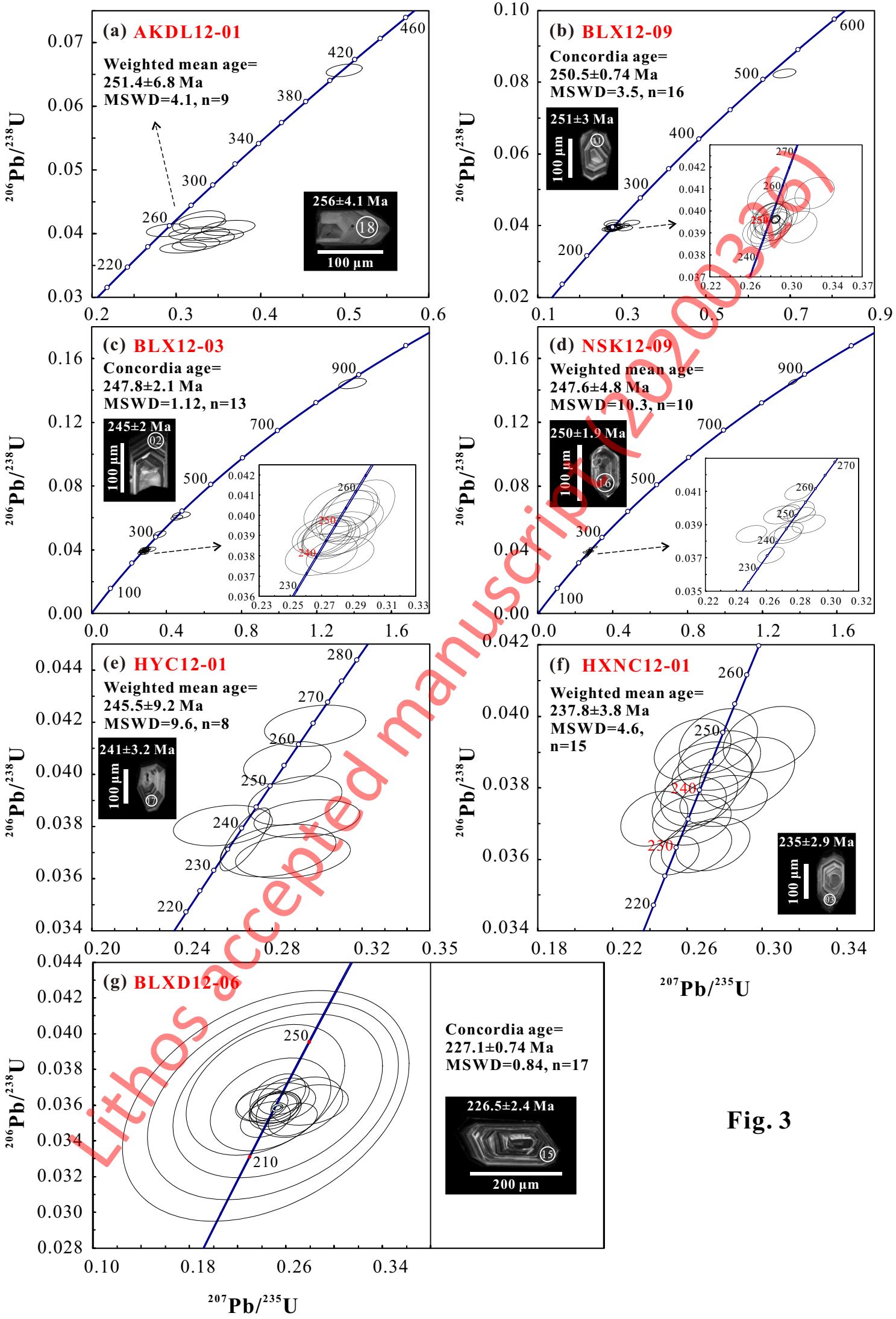
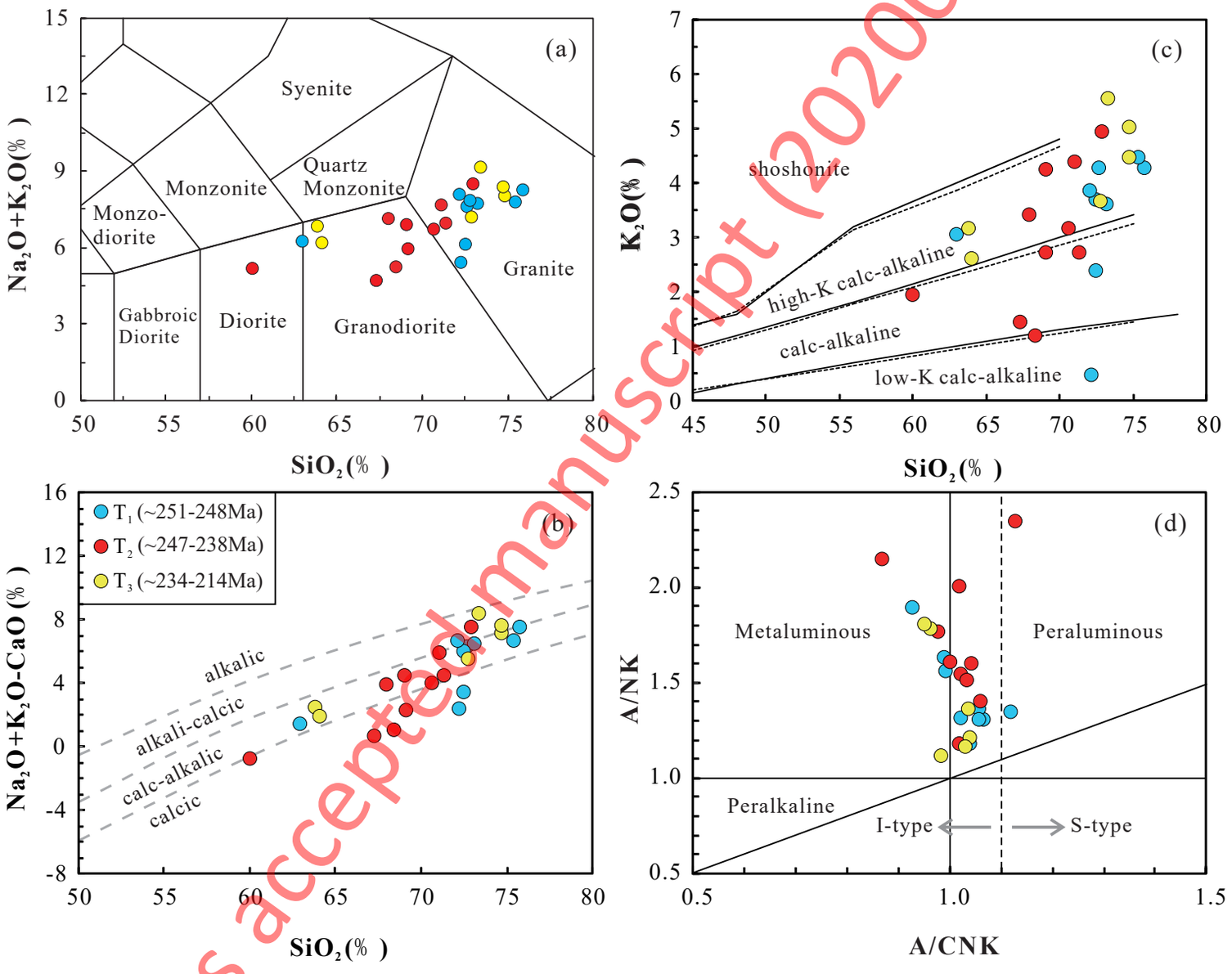


Fig. 3

Fig. 4



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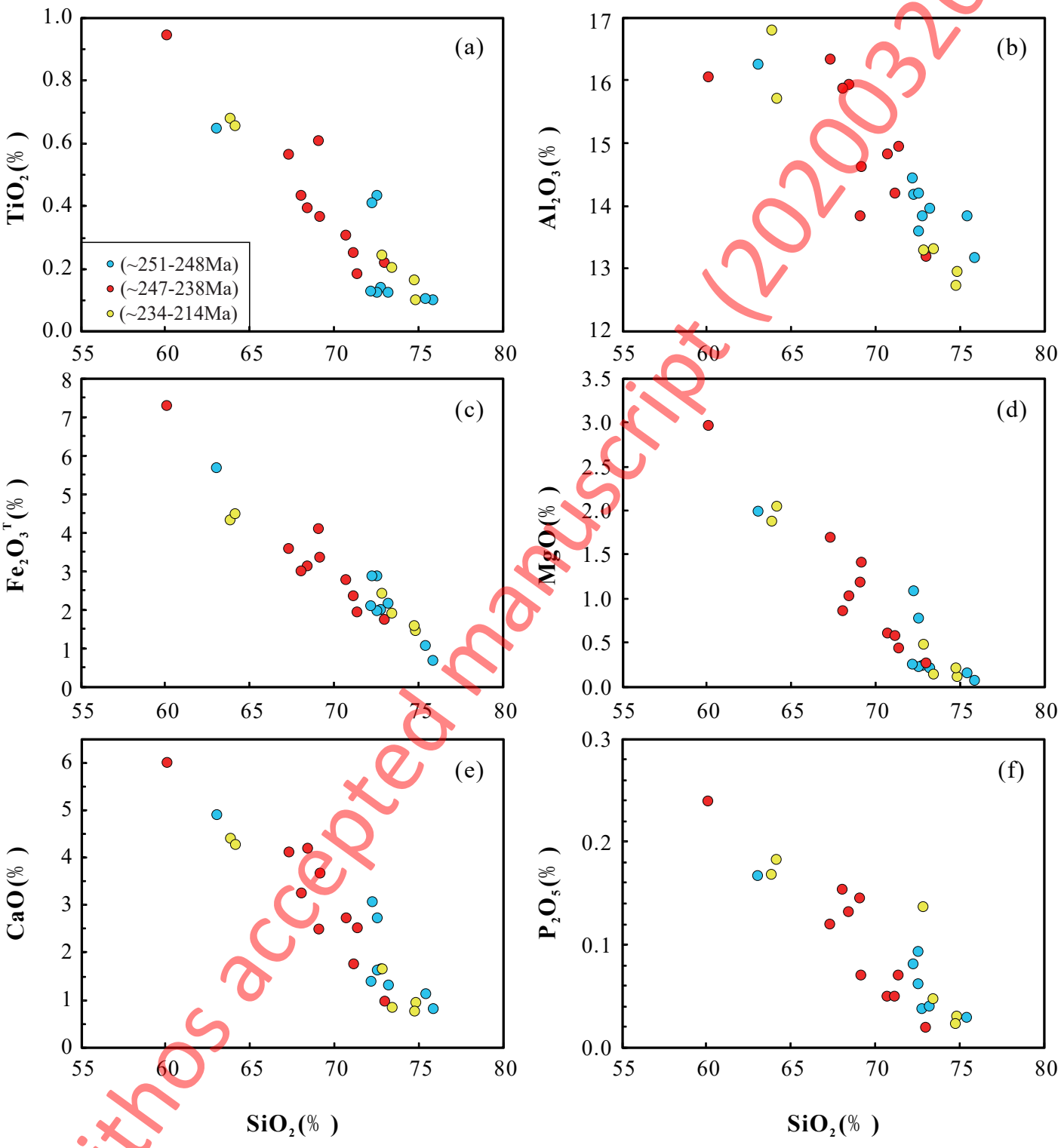

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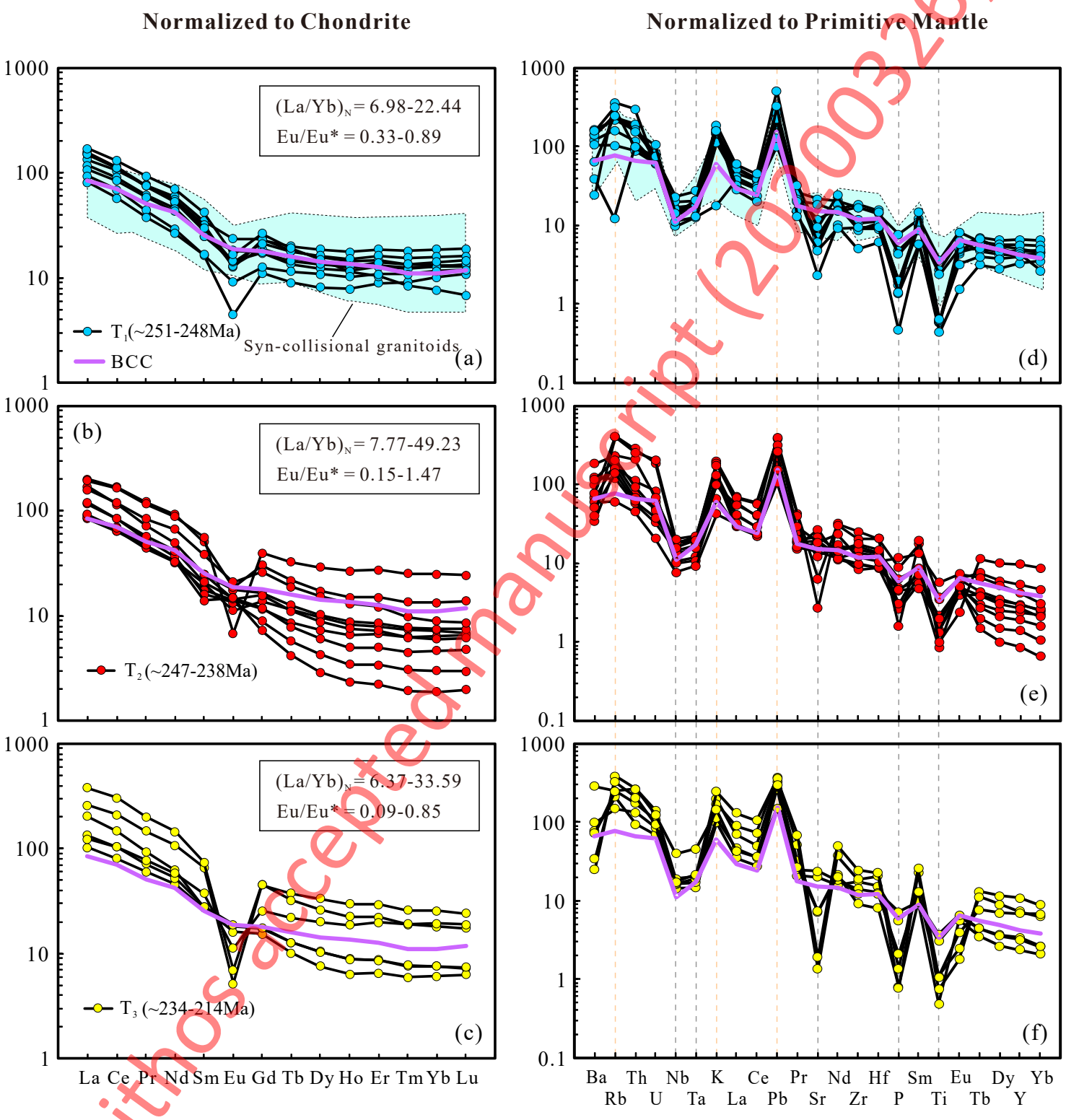
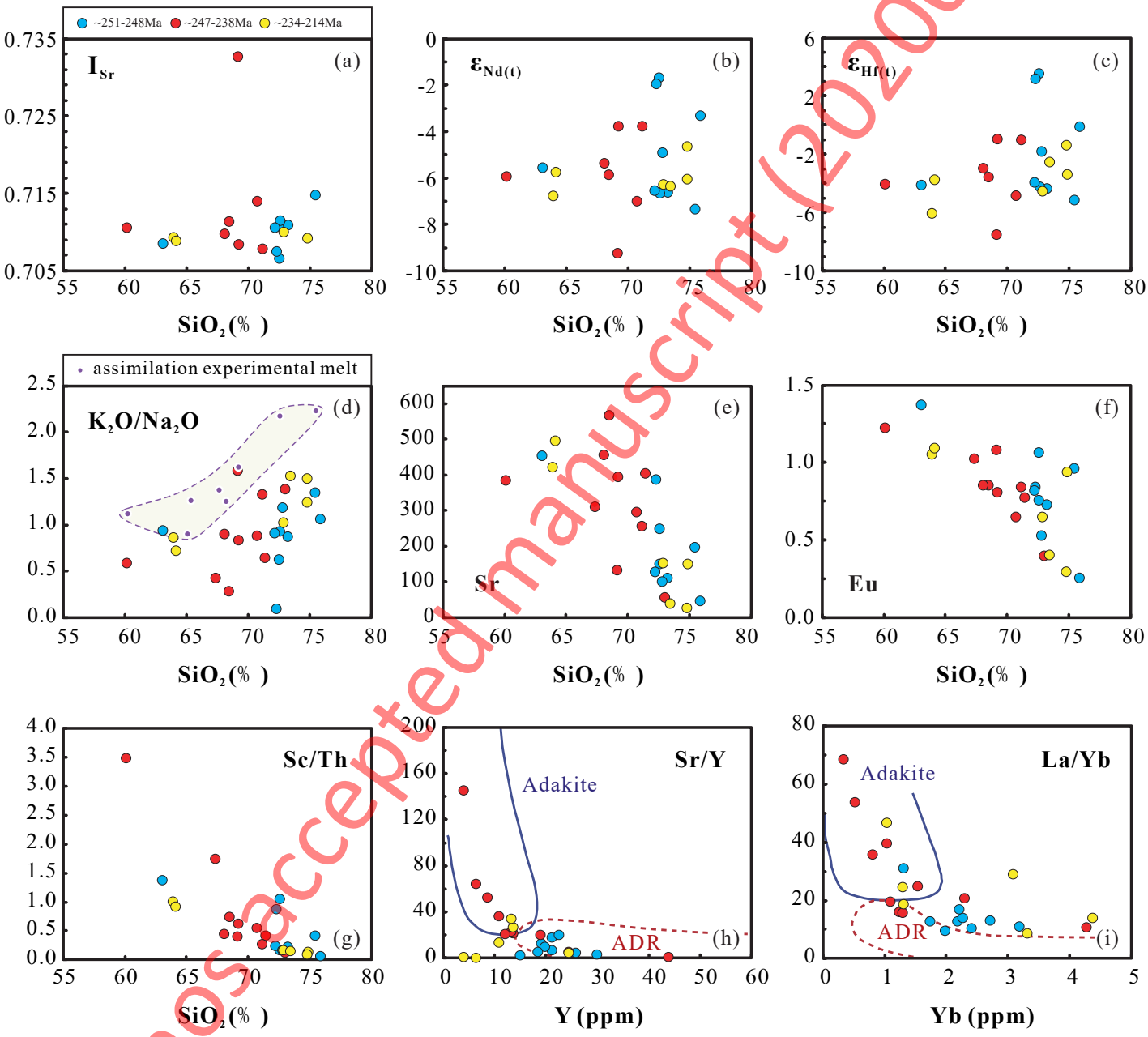


Fig. 7



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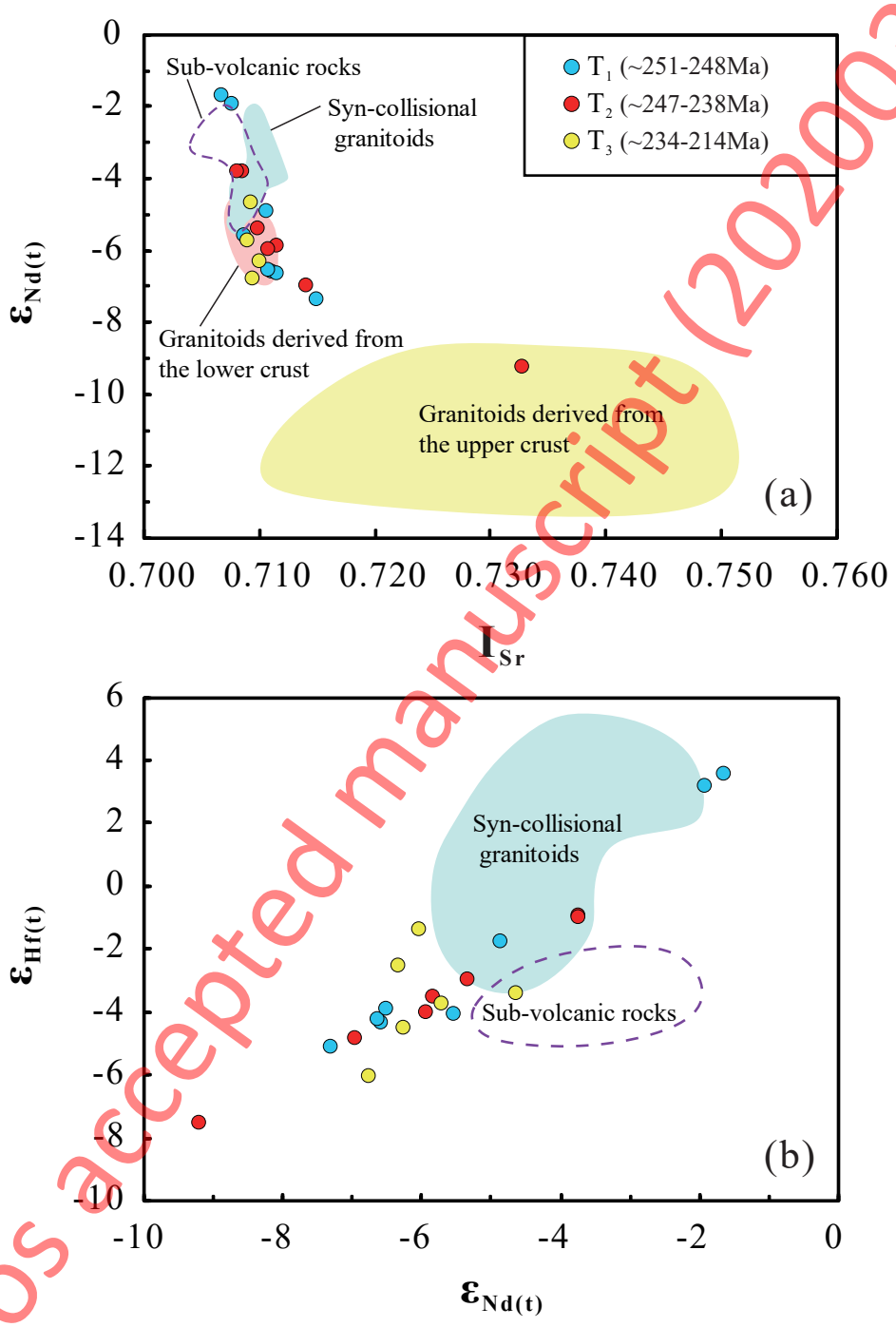


Fig.9

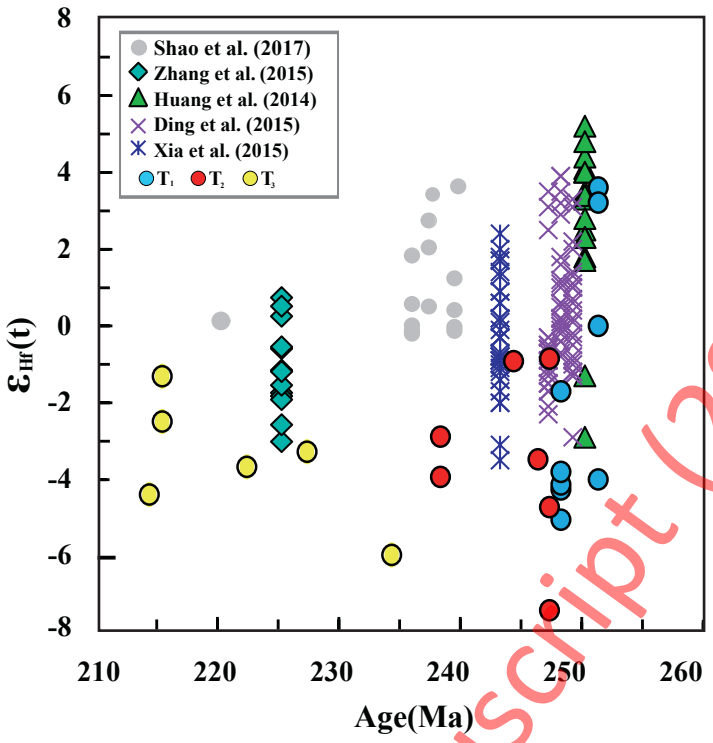
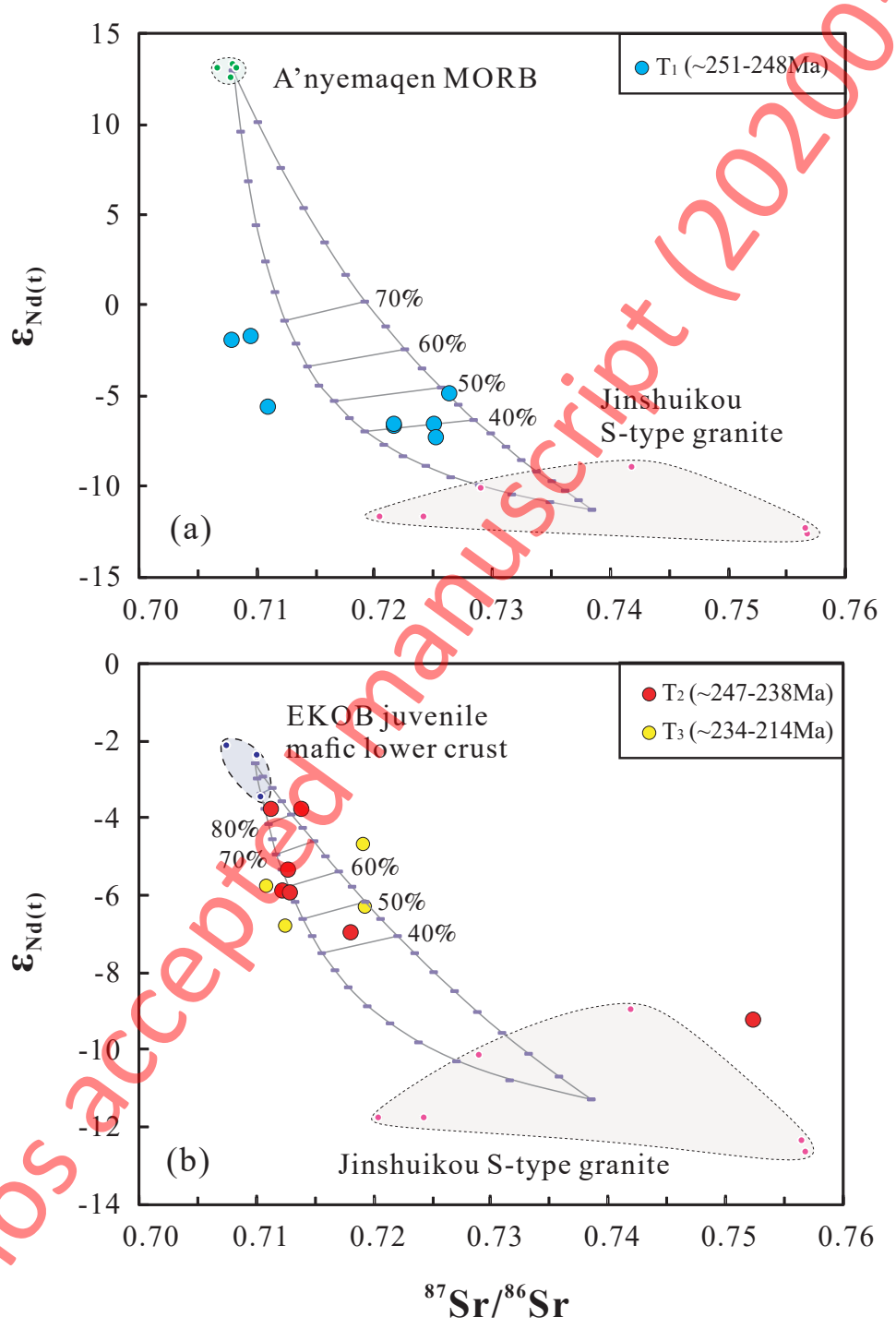


Fig. 10



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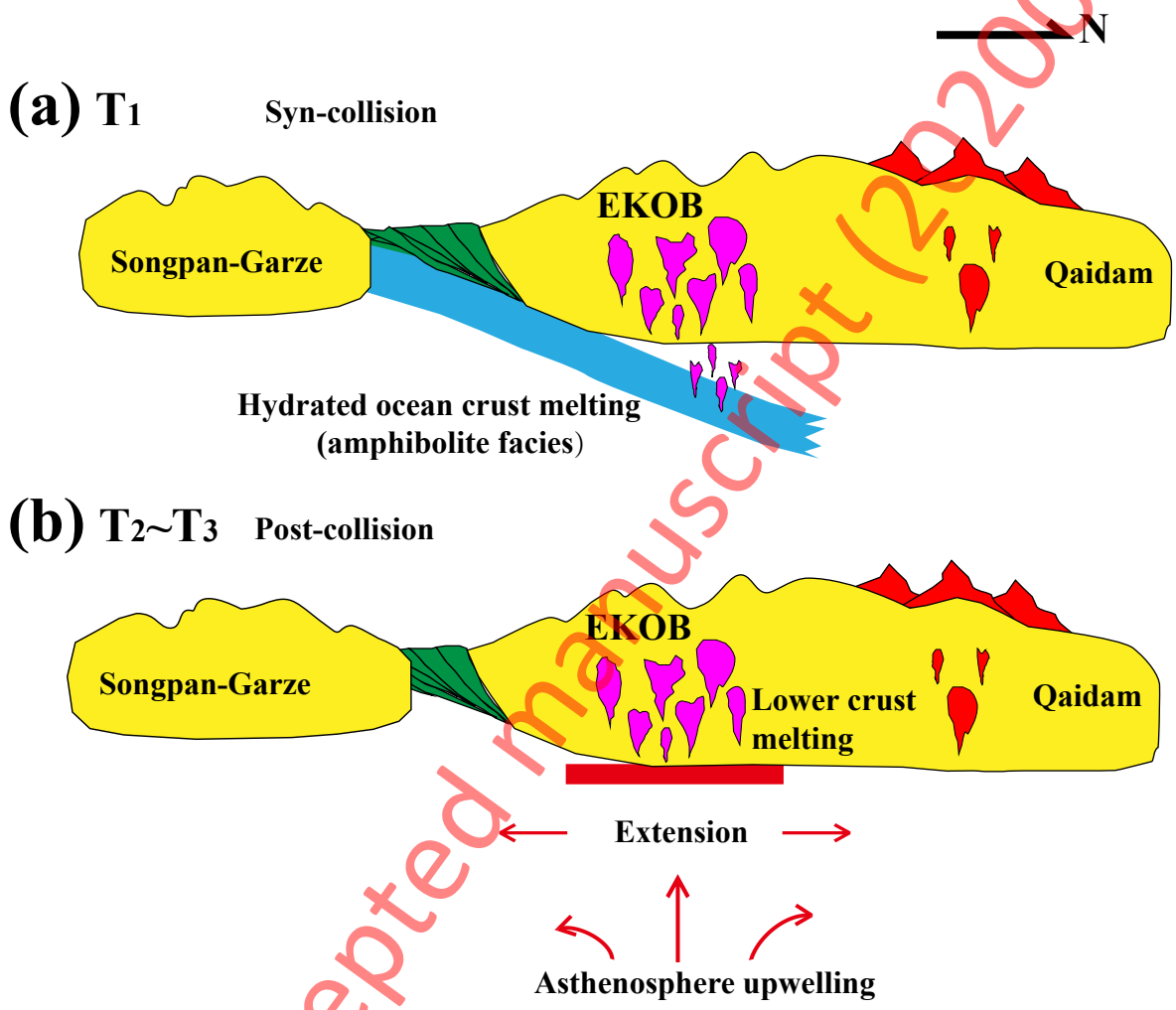

Table 1

Sample loctions and zircon U-Pb ages of the Triassic granitoids in the East Kunlun Orogenic Belt.

Sample	Point	GPS		Lithology	Petrology	Age (Ma)
AKDL12-01	14	N35°50'28.3"	E96°29'26.2"	syenogranite	Kfs (45%), Qtz (27%), Pl (15%), Mag (8%), Bt (5%); Kfs with kaolinization.	251.4±6.8
AKDL12-03	15	N35°50'28.2"	E96°29'32.1"	moyite	Kfs (50%), Qtz (30%), Mag (7%), Amp (5%), Bt (4%), Pl (4%).	251.4±6.8
AKDL12-04	16	N35°51'27.3"	E96°31'5.2"	moyite	Kfs (55%), Qtz (30%), Pl (5%), Mag (5%), Bt (5%).	251.4±6.8
BLX12-09	21	N35°48'34.9"	E97°24'25.8"	granodiorite	Pl (50%), Qtz (20%), Bt (10%), Amp (5%); accessory mineral (Ap)	250.5±0.74
BLX12-03	18	N35°46'38.5"	E97°22'11.2"	monzogranite	Qtz (35%), Kfs (30%), Pl (25%), Bt (10%); Kfs with kaolinization.	247.8±2.1
BLX12-06	19	N35°46'58.9"	E97°22'51.1"	granite porphyry	Phenocrysts of Qtz (40%), Kfs (30%), Pl (25%), Bt (5%), Qtz and Kfs with melt corrosion shapes; groundmass with the same mineral assemblage.	247.8±2.1
BLX12-08	20	N35°47'37.2"	E97°23'43.0"	monzogranite	Qtz (30%), Kfs (30%), Pl (25%), Bt+Mag (15%).	247.8±2.1
HXNC12-07	3	N36°18'53.8"	E94°35'21.4"	granite	Qtz (40%), Pl (35%), Kfs (10%), Bt (10%), Mag (5%); accessory minerals (Ap+Zrn); Pl with sericitization.	247.6±4.8
NSK12-09	4	N36°8'14.6"	E94°47'57.7"	syenogranite	Kfs (45%), Qtz (28%), Pl (15%), Bt (8%), Mag (4%); Kfs with kaolinization.	247.6±4.8
GYK12-05	7	N36°16'19.3"	E95°23'56.6"	biotite granite	Qtz (35%), Pl (30%), Kfs (15%), Bt (20%).	246.6±1.9[6]
DGLX12-01	8	N36°21'23.1"	E95°29'25"	granodiorite	Pl (35%), Qtz (25%), Kfs (10%), Bt+Amp+Mag (30%).	246.6±1.9[6]
DGL12-01	9	N36°11'33.5"	E95°47'39.2"	biotite granite	Qtz (30%), Pl+Kfs (45%), Bt (20%), Mag (5%).	246.6±1.9[6]
DGL12-02	9	N36°11'33.5"	E95°47'39.2"	biotite moyite	Kfs (45%), Qtz (30%), Bt (25%); porphyroid texture.	246.6±1.9[6]
HYC12-01	5	N36°12'55"	E95°05'36"	biotite monzogranite	Qtz (30%), Kfs (30%), Pl (30%), Bt (10%).	245.5±9.2
HYC12-06 (host)	6	N36°14'6.3"	E95°05'6.8"	biotite granite	Qtz (28%), Pl (35%), Kfs (15%), Bt (20%), Amp (2%).	245.5±9.2
WLG12-05	12	N36°15'40.5"	E95°55'12.2"	syenogranite	Kfs (50%), Qtz (30%), Pl (10%), Chl (10%).	244±2[1]
WLG12-07	13	N36°18'17"	E95°54'11.4"	syenogranite	Kfs (40%), Qtz (30%), Pl (15%), Bt (15%); accessory mineral (Zrn)	244±2[1]
HXNC12-01	1	N36°21'31.4"	E94°24'37.6"	syenogranite	Kfs (48%), Qtz (32%), Pl (10%), Bt (10%).	237.8±3.8
HXNC12-05	2	N36°21'39.8"	E94°24'37.8"	quartz diorite	Pl (40%), Qtz (20%), Bt+Amp (40%).	237.8±3.8
ZJX12-01	17	N36°10'27.8"	E97°01'10.6"	quartz diorite	Pl (40%), Qtz (10%), Kfs (10%), Bt+Amp (35%), Mag (5%).	234.2±2.1[6]
BLXD12-06	24	N36°01'8.7"	E97°46'17"	amphibole quartz monzonite	Kfs (30%), Pl (30%), Qtz (15%), Bt+Amp (25%).	227.1±0.74
WLS12-01	25	N36°01'33.5"	E97°48'7.3"	syenogranite	Kfs (50%), Qtz (30%), Pl (12%), Bt (5%), Mag (3%).	227.1±0.74
BLX12-15 (host)	22	N35°58'39.6"	E97°28'32.1"	granodiorite	Pl (45%), Qtz (25%), Kfs (5%), Bt+Amp (20%); Mag (5%).	222.02±0.91[4]
DGL12-05	10	N36°14'4.3"	E95°46'35.7"	syenogranite	Kfs (57%), Qtz (30%), Pl (8%), Bt (5%).	215±3[1]
DGL12-07	11	N36°14'28"	E95°44'47.6"	moyite	Kfs (56%), Qtz (28%), Bt (8%), Pl (5%), Mag (3%).	215±3[1]
BLXD12-02	23	N36°01'34.3"	E97°39'4.1"	granite porphyry	Phenocrysts of Qtz and feldspar, Qtz and feldspar with melt corrosion shapes; groundmass of Kfs (50%), Qtz (30%), Bt (10%), Pl (5%), Mag (5%).	214±1[7]

Kfs = potash feldspar; Qtz = quartz; Pl = plagioclase; Bt = biotite; Amp = amphibole; Mag = magnetite; Ap = apatite; Zrn = zircon.

Ages with superscript are from literature being the same as that in Fig. 1.

Table 2

Whole-rock major and trace element data of the Triassic granitoids in the East Kunlun Orogenic Belt.

Sample	AKDL12-01	AKDL12-03	AKDL12-04	BLX12-09	BLX12-03	BLX12-06	BLX12-08	HXNC12-07	NSK12-09
Age (Ma)	*(Early Triassic)*	~ 251			~ 248				
Major elements (wt.%)									
SiO_2	72.50	72.24	75.83	63.01	73.19	72.55	72.13	75.40	72.74
TiO_2	0.44	0.41	0.10	0.65	0.13	0.13	0.13	0.11	0.14
Al_2O_3	13.60	14.17	13.18	16.27	13.96	14.20	14.45	13.84	13.85
$Fe_2O_3^T$	2.87	2.89	0.67	5.67	2.16	1.98	2.09	1.06	2.01
MnO	0.06	0.09	0.02	0.11	0.07	0.04	0.07	0.02	0.06
MgO	0.78	1.09	0.07	1.99	0.22	0.23	0.25	0.16	0.24
CaO	2.73	3.08	0.81	4.90	1.31	1.63	1.40	1.14	1.65
Na_2O	3.74	4.97	4.00	3.23	4.12	3.93	4.20	3.30	3.58
K_2O	2.37	0.47	4.26	3.05	3.60	3.69	3.86	4.47	4.28
P_2O_5	0.09	0.08		0.17	0.04	0.06		0.03	0.04
LOI	0.71	0.99	0.44	0.49	0.52	0.67	0.53	0.52	0.54
Total	99.89	100.49	99.38	99.52	99.32	99.12	99.12	100.04	99.11
A/NK	17.09	14.07	76.26	7.33	40.97	33.73	37.25	50.11	33.42
A/CNK	0.99	0.99	1.04	0.93	1.07	1.06	1.06	1.12	1.02
$Mg^{\#}$	0.37	0.45	0.19	0.44	0.18	0.21	0.21	0.25	0.21
Trace elements (ppm)									
Li	4.47	9.60	2.64	21.8	40.6	14.8	39.9	33.9	50.0
Sc	7.90	7.34	1.95	13.8	3.46	2.88	3.40	3.59	2.28
V	54.6	44.4	2.07	119	4.49	4.66	4.96	3.67	4.88
Cr	4.44	3.78	17.2	10.0	4.82	21.6	4.75	3.00	3.85
Co	5.81	3.39	0.29	11.5	0.93	0.94	0.99	0.53	1.14
Ni	2.51	2.49	8.88	11.8	0.97	8.42	1.63	1.27	0.50
Cu	2.51	0.75	0.24	3.68		2.37		0.08	1.56
Zn	33.5	28.6	6.97	59.1	33.4	21.7	27.6	19.1	27.6
Ga	16.2	15.5	13.9	17.8	17.3	15.8	16.6	21.3	14.6
Rb	65.5	7.73	231	102	157	150	139	200	156
Sr	249	389	48.5	455	113	152	128	198	101
Y	19.0	21.1	15.0	22.5	29.8	21.1	25.7	19.6	18.3
Zr	206	187	56.3	189	119	101	131	96.5	105
Nb	6.94	7.95	14.2	9.77	11.9	9.81	11.4	16.3	10.0
Cs	0.52	0.86	2.60	2.22	3.31	2.72	3.06	4.41	3.42
Ba	734	268	168	845	974	995	1131	1098	447
La	28.0	25.6	22.2	31.8	35.7	37.5	35.6	40.7	19.2
Ce	55.2	51.5	42.2	64.7	68.8	69.7	67.8	80.5	35.4
Pr	5.89	5.57	4.20	7.29	7.34	7.17	7.16	8.85	3.63
Nd	21.3	20.2	13.6	28.0	25.8	24.6	25.0	32.7	12.4
Sm	3.81	3.81	2.53	5.35	4.96	4.34	4.61	6.43	2.55
Eu	1.07	0.85	0.26	1.37	0.73	0.76	0.82	0.96	0.53
Gd	3.53	3.60	2.28	4.85	4.73	3.91	4.36	5.45	2.59
Tb	0.51	0.54	0.34	0.68	0.76	0.57	0.67	0.72	0.43
Dy	3.13	3.43	2.08	4.01	4.76	3.49	4.16	3.81	2.75
Ho	0.66	0.74	0.45	0.81	1.00	0.71	0.87	0.65	0.58
Er	2.12	2.32	1.48	2.45	3.12	2.18	2.69	1.71	1.80
Tm	0.31	0.35	0.23	0.34	0.46	0.32	0.40	0.22	0.28
Yb	2.18	2.40	1.72	2.25	3.18	2.21	2.70	1.30	1.97
Lu	0.34	0.37	0.28	0.33	0.48	0.33	0.41	0.17	0.30
Hf	4.82	4.48	1.88	4.33	3.43	2.79	3.54	3.26	2.96
Ta	0.52	0.52	0.85	0.67	0.75	0.79	0.74	1.11	0.86
Pb	9.12	7.08	15.7	13.6	20.9	14.5	19.1	36.5	22.9
Th	7.45	8.34	25.4	10.0	14.7	16.3	13.9	8.36	13.2
U	1.26	1.38	1.29	1.34	1.72	1.35	1.59	1.59	2.22
$(La/Yb)_N$	9.22	7.65	9.25	10.13	8.05	12.20	9.46	22.44	6.98
Eu/Eu*	0.89	0.70	0.33	0.82	0.46	0.57	0.56	0.50	0.63
Sr/Sr*	0.65	1.06	0.19	0.92	0.24	0.33	0.28	0.34	0.44

(continued)

Table 2 *(continued)*

Sample	GYK12-05	DGLX12-01	DGL12-01	DGL12-02	HYC12-01	HYC12-06 (host)	WLG12-05	WLG12-07	HXNC12-01
Age (Ma)	*(Middle Triassic)*	~ 247			~ 246		~ 244		~ 238
Major elements (wt.%)									
SiO_2	70.68	69.16	67.32	69.08	71.35	68.42	71.12	72.95	68.01
TiO_2	0.31	0.37	0.57	0.61	0.19	0.39	0.25	0.22	0.43
Al_2O_3	14.83	14.63	16.33	13.83	14.96	15.93	14.20	13.19	15.87
$Fe_2O_3^T$	2.79	3.35	3.60	4.11	1.95	3.13	2.37	1.76	3.02
MnO	0.06	0.07	0.05	0.06	0.06	0.06	0.09	0.02	0.06
MgO	0.62	1.41	1.70	1.19	0.44	1.04	0.58	0.28	0.87
CaO	2.72	3.67	4.11	2.50	2.52	4.21	1.77	0.96	3.27
Na_2O	3.55	3.24	3.29	2.66	4.21	4.05	3.28	3.55	3.76
K_2O	3.16	2.72	1.43	4.23	2.72	1.18	4.37	4.94	3.40
P_2O_5	0.05	0.07	0.12	0.15	0.07	0.13	0.05	0.02	0.15
LOI	0.49	0.61	0.88	1.12	0.51	0.51	1.54	1.48	0.48
Total	99.25	99.30	99.40	99.53	98.97	99.06	99.64	99.37	99.33
A/NK	17.83	10.97	9.28	14.82	20.59	10.91	24.67	48.54	13.68
A/CNK	1.04	0.98	1.13	1.02	1.03	1.02	1.06	1.02	1.00
$Mg^{\#}$	0.33	0.48	0.51	0.39	0.33	0.42	0.35	0.26	0.39
Trace elements (ppm)									
Li	65.0	41.2	52.2	57.9	38.4	16.9	15.0	18.7	33.2
Sc	3.70	5.97	10.0	9.92	2.09	2.89	4.92	3.00	3.67
V	10.5	46.8	60.8	35.4	10.9	22.0	20.8	7.99	27.9
Cr	2.75	10.1	31.1	11.1	4.06	5.31	5.74	5.15	3.97
Co	2.04	6.36	5.90	6.43	1.57	3.28	2.03	1.13	3.39
Ni	0.70	5.07	8.68	2.78	0.90	2.17	2.72	1.45	1.62
Cu	0.37	1.33	1.38	4.75		0.16	3.61	0.44	0.54
Zn	58.1	50.5	50.0	80.5	53.2	56.8	50.3	43.5	48.4
Ga	19.6	16.5	20.3	21.9	20.0	16.9	16.8	20.6	19.0
Rb	117	105	102	260	76.0	38.4	149	257	131
Sr	297	397	311	134	406	569	257	57.1	458
Y	13.5	10.8	13.3	24.4	6.27	3.91	12.0	43.7	8.60
Zr	153	107	184	281	92.1	111	167	227	197
Nb	13.1	7.56	7.13	13.0	10.4	5.45	12.9	14.4	9.50
Cs	3.96	3.18	10.3	10.1	2.55	0.72	2.99	5.04	1.91
Ba	471	697	240	546	852	420	1298	282	818
La	20.0	21.4	20.6	47.9	27.6	22.1	40.6	46.2	28.5
Ce	39.2	39.4	43.6	104	51.4	41.8	70.8	102	51.7
Pr	4.34	4.19	5.00	11.6	5.16	4.47	6.89	11.1	5.49
Nd	15.7	15.0	19.4	43.2	17.1	15.5	23.0	41.5	19.4
Sm	3.15	2.70	3.83	7.76	2.47	2.14	3.84	8.64	3.20
Eu	0.65	0.81	1.03	1.09	0.78	0.86	0.84	0.40	0.86
Gd	2.84	2.39	3.43	6.28	1.83	1.50	3.25	8.15	2.46
Tb	0.43	0.32	0.47	0.81	0.22	0.16	0.42	1.22	0.30
Dy	2.46	1.86	2.59	4.37	1.10	0.74	2.25	7.39	1.55
Ho	0.47	0.38	0.50	0.85	0.20	0.13	0.43	1.52	0.29
Er	1.32	1.12	1.41	2.49	0.57	0.37	1.20	4.57	0.83
Tm	0.19	0.16	0.20	0.35	0.08	0.05	0.16	0.65	0.11
Yb	1.23	1.09	1.28	2.29	0.51	0.32	1.02	4.27	0.79
Lu	0.17	0.17	0.19	0.35	0.08	0.05	0.16	0.62	0.12
Hf	3.86	2.70	4.24	6.40	2.70	2.74	4.01	6.32	4.59
Ta	0.80	0.50	0.44	0.88	0.70	0.38	0.79	0.92	0.64
Pb	19.2	14.6	13.3	26.9	16.2	7.41	28.2	22.8	18.6
Th	6.64	9.42	5.71	24.3	4.89	3.83	17.7	21.9	8.04
U	1.04	1.75	1.25	3.96	0.69	0.43	1.45	4.33	0.99
$(La/Yb)_N$	11.67	14.14	11.49	15.00	38.72	49.23	28.46	7.77	25.83
Eu/Eu*	0.67	0.98	0.87	0.48	1.12	1.47	0.73	0.15	0.94
Sr/Sr*	1.04	1.45	0.91	0.17	1.25	1.98	0.59	0.08	1.28

(continued)

Table 2 *(continued)*

Sample	HXNC12-05	ZJX12-01	WLS12-01	BLX12-15 (host)	DGL12-05	DGL12-07	BLXD12-02
Age (Ma)	~ 238	*(Late Triassic)* ~ 234	~ 227	~ 222	~ 215		~ 214
Major elements (wt.%)							
SiO_2	60.06	63.86	74.78	64.10	74.75	73.38	72.82
TiO_2	0.95	0.68	0.10	0.66	0.17	0.21	0.24
Al_2O_3	16.06	16.81	12.94	15.72	12.72	13.32	13.29
$Fe_2O_3^T$	7.29	4.33	1.47	4.50	1.60	1.90	2.44
MnO	0.10	0.07	0.04	0.07	0.01	0.03	0.04
MgO	2.96	1.88	0.12	2.05	0.21	0.15	0.48
CaO	6.02	4.40	0.94	4.27	0.76	0.85	1.66
Na_2O	3.26	3.65	3.58	3.59	3.34	3.62	3.54
K_2O	1.94	3.16	4.47	2.59	5.02	5.55	3.65
P_2O_5	0.24	0.17	0.03	0.18	0.02	0.05	0.14
LOI	1.60	0.84	0.76	1.27	0.70	0.56	1.08
Total	100.48	99.86	99.24	99.01	99.31	99.61	99.39
A/NK	5.27	8.11	61.55	8.00	63.95	62.73	28.13
A/CNK	0.87	0.96	1.04	0.95	1.03	0.98	1.04
$Mg^{\#}$	0.47	0.49	0.15	0.50	0.22	0.15	0.30
Trace elements (ppm)							
Li	49.0	48.4	10.7	15.9	15.9	21.3	24.5
Sc	18.3	8.15	2.11	10.3	2.18	2.83	3.83
V	153	60.4	2.19	100	3.15	6.55	19.5
Cr	17.9	14.7	18.4	17.5	2.65	10.3	8.11
Co	15.4	7.04	0.64	10.9	0.55	0.89	2.59
Ni	5.91	5.05	8.16	6.52	0.48	5.47	4.69
Cu	6.08	1.61	2.03	1.66	0.59	0.17	2.37
Zn	68.1	58.9	23.3	49.6	42.1	45.6	44.6
Ga	20.8	19.7	17.4	20.3	20.7	22.5	16.9
Rb	87.7	132	158	94.1	244	205	157
Sr	387	422	153	497	28.4	39.9	153
Y	18.8	15.4	10.8	14.4	49.5	33.1	31.4
Zr	136	196	102	159	208	272	136
Nb	11.3	10.2	11.7	11.9	13.4	12.4	28.5
Cs	2.40	10.8	2.51	3.71	7.24	6.04	4.66
Ba	357	534	2012	684	173	235	507
La	37.9	24.2	48.3	31.9	61.6	89.7	29.4
Ce	72.9	49.4	89.8	63.5	129	186	63.8
Pr	8.04	5.73	8.85	6.83	14.0	19.1	7.39
Nd	31.4	21.4	29.2	24.5	50.1	68.0	27.3
Sm	5.88	4.16	4.24	4.28	10.1	11.4	5.76
Eu	1.22	1.06	0.94	1.09	0.30	0.41	0.65
Gd	5.37	3.54	3.18	3.63	9.29	9.38	5.28
Tb	0.71	0.48	0.38	0.48	1.40	1.21	0.82
Dy	3.89	2.68	1.94	2.63	8.53	6.64	5.12
Ho	0.74	0.51	0.36	0.50	1.69	1.28	1.06
Er	2.02	1.44	1.08	1.43	4.89	3.67	3.28
Tm	0.25	0.19	0.15	0.20	0.66	0.48	0.49
Yb	1.52	1.30	1.03	1.29	4.36	3.07	3.31
Lu	0.22	0.18	0.16	0.19	0.61	0.44	0.48
Hf	3.23	4.68	2.48	3.90	6.38	6.96	3.73
Ta	0.81	0.60	0.69	0.78	0.84	0.88	1.86
Pb	10.8	18.7	23.3	10.7	26.4	25.2	20.9
Th	5.23	8.00	14.7	11.2	22.5	17.4	22.5
U	0.80	1.46	2.23	1.79	2.97	1.96	2.61
$(La/Yb)_N$	17.84	13.41	33.59	17.77	10.12	20.94	6.37
Eu/Eu*	0.67	0.84	0.78	0.85	0.09	0.12	0.36
Sr/Sr*	0.71	1.11	0.28	1.11	0.03	0.03	0.31

A/NK = molar $Al_2O_3/(Na_2O+K_2O)$, A/CNK = molar $Al_2O_3/(CaO+Na_2O+K_2O)$, $Mg^{\#}$ = molar $Mg/(Mg+Fe^{2+})$; LOI, loss on ignition. $(La/Yb)_N$ refers to the value normalized against chondrite. $Eu/Eu^* = Eu_{PM}/[Sm_{PM}\times Gd_{PM}]^{1/2}$, $Sr/Sr^* = Sr_{PM}/[Pr_{PM}\times Nd_{PM}]^{1/2}$, where subscript $_{PM}$ denotes normalized values against primary mantle.

Table 3

Whole rock Sr-Nd-Hf isotopic composition of the Triassic granitoids from the East Kunlun Orogenic Belt.

Sample	Rb (ppm)	Sr (ppm)	$^{87}Rb/^{86}Sr$	$^{87}Sr/^{86}Sr$	±2σ	I_{sr}	Sm (ppm)	Nd (ppm)	$^{147}Sm/^{144}Nd$	$^{143}Nd/^{144}Nd$	±2σ	$^{143}Nd/^{144}Nd$	$\varepsilon_{Nd}(t)$	Lu (ppm)	Hf (ppm)	$^{176}Lu/^{177}Hf$	$^{176}Hf/^{177}Hf$	±2σ	$^{176}Hf/^{177}Hf_i$	$\varepsilon_{Hf}(t)$	t (Ma)
T_1 group																					
BLX12-09	102	455	0.65	0.710934	6	0.708626	5.35	28.0	0.12	0.512222	3	0.512030	-5.55	0.33	4.33	0.01	0.282553	2	0.282501	-4.06	251
AKDL12-01	65.5	249	0.76	0.709400	5	0.706684	3.81	21.3	0.11	0.512409	4	0.512230	-1.66	0.34	4.82	0.01	0.282765	2	0.282718	3.59	251
AKDL12-03	7.73	389	0.06	0.707823	3	0.707617	3.81	20.2	0.11	0.512405	6	0.512216	-1.92	0.37	4.48	0.01	0.282762	2	0.282707	3.21	251
HXNC12-07	200	198	2.92	0.725152	5	0.714843	6.43	32.7	0.12	0.512138	5	0.511944	-7.32	0.17	3.26	0.01	0.282509	3	0.282474	-5.11	248
NSK12-09	156	101	4.46	0.726310	6	0.710573	2.55	12.4	0.12	0.512272	4	0.512069	-4.88	0.30	2.96	0.01	0.282636	3	0.282569	-1.74	248
BLX12-03	157	113	4.01	0.725105	5	0.710972	4.96	25.8	0.12	0.512171	3	0.511981	-6.59	0.48	3.43	0.02	0.282588	3	0.282496	-4.31	248
BLX12-06	150	152	2.87	0.721635	3	0.711511	4.34	24.6	0.11	0.512153	3	0.511979	-6.64	0.33	2.79	0.02	0.282577	3	0.282499	-4.20	248
BLX12-08	139	128	3.13	0.721689	5	0.710650	4.61	25.0	0.11	0.512167	3	0.511985	-6.52	0.41	3.54	0.02	0.282584	3	0.282508	-3.88	248
AKDL12-04	231	48.5	13.80	0.755502	5	0.706213	2.53	13.6	0.11	0.512332	4	0.512146	-3.30	0.28	1.88	0.02	0.282713	4	0.282615	-0.05	251
T_2 group																					
GYK12-05	117	297	1.15	0.718030	4	0.714006	3.15	15.7	0.12	0.512161	5	0.511963	-6.97	0.17	3.86	0.01	0.282512	2	0.282483	-4.81	247
DGLX12-01	105	397	0.77	0.711228	5	0.708529	2.70	15.0	0.11	0.512304	4	0.512127	-3.76	0.17	2.70	0.01	0.282633	3	0.282592	-0.93	247
DGL12-02	260	134	5.60	0.752345	4	0.732679	7.76	43.2	0.11	0.512025	3	0.511848	-9.21	0.35	6.40	0.01	0.282443	3	0.282407	-7.50	247
HYC12-06 (host)	38.4	569	0.20	0.712135	5	0.711452	2.14	15.5	0.08	0.512157	4	0.512022	-5.85	0.05	2.74	0.00	0.282532	3	0.282520	-3.51	246
WLG12-05	149	257	1.68	0.713781	3	0.707954	3.84	23.0	0.10	0.512294	3	0.512131	-3.76	0.16	4.01	0.01	0.282619	3	0.282594	-0.96	244
HXNC12-01	131	458	0.83	0.712664	5	0.709853	3.20	19.4	0.10	0.512214	3	0.512058	-5.35	0.12	4.59	0.00	0.282558	2	0.282541	-2.94	238
HXNC12-05	87.7	387	0.66	0.712885	4	0.710665	5.88	31.4	0.11	0.512205	4	0.512027	-5.94	0.22	3.23	0.01	0.282554	3	0.282511	-4.01	238
T_3 group																					
ZJX12-01	132	422	0.91	0.712378	4	0.709362	4.16	21.4	0.12	0.512172	4	0.511990	-6.76	0.18	4.68	0.01	0.282481	3	0.282456	-6.04	234
WLS12-01	158	153	2.99	0.718922	4	0.709265	4.24	29.2	0.09	0.512239	5	0.512108	-4.65	0.16	2.48	0.01	0.282575	3	0.282536	-3.36	227
BLX12-15 (host)	94.1	497	0.55	0.710661	5	0.708930	4.28	24.5	0.11	0.512213	3	0.512059	-5.73	0.19	3.90	0.01	0.282557	2	0.282529	-3.74	222
DGL12-05	244	28.4	24.95	0.855875	5	0.779601	10.1	50.1	0.12	0.512224	3	0.512051	-6.05	0.61	6.38	0.01	0.282655	3	0.282601	-1.34	215
DGL12-07	205	39.9	14.90	0.800046	4	0.754482	11.4	68.0	0.10	0.512180	4	0.512036	-6.34	0.44	6.96	0.01	0.282604	3	0.282567	-2.53	215
BLXD12-02	157	153	2.97	0.719073	4	0.710046	5.76	27.3	0.13	0.512221	4	0.512041	-6.27	0.48	3.73	0.02	0.282586	3	0.282512	-4.49	214

$I_{sr}=[(^{87}Sr/^{86}Sr)-(^{87}Rb/^{86}Sr)(e^{\lambda t}-1)]$; $^{143}Nd/^{144}Nd_i=[(^{143}Nd/^{144}Nd)-(^{147}Sm/^{144}Nd)(e^{\lambda t}-1)]$; $\varepsilon_{Nd}(t)=[(^{143}Nd/^{144}Nd_i)/(^{143}Nd/^{144}Nd_{CHURi})-1]\times10^4$; $^{176}Hf/^{177}Hf_i=[(^{176}Hf/^{177}Hf)-(^{176}Lu/^{177}Hf)(e^{\lambda t}-1)]$; $\varepsilon_{Hf}(t)=[(^{176}Hf/^{177}Hfi)/(^{176}Hf/^{177}Hf_{CHURi})-1]\times10^4$.
$^{147}Sm/^{144}Nd_{CHUR}=0.1967$; $^{143}Nd/^{144}Nd_{CHUR}=0.512638$; $^{176}Lu/^{177}Hf_{CHUR}=0.0332$; $^{176}Hf/^{177}Hf_{CHUR}=0.282772$; $\lambda(^{87}Rb)=1.42\times10^{-11}yr^{-1}$; $\lambda(^{147}Sm)=6.54\times10^{-12}yr^{-1}$, $\lambda(^{176}Lu)=1.865\times10^{-11}yr^{-1}$.