## **1 SUPPLEMENTARY MATERIAL FOR:**

# The fate of fluvially-deposited organic carbon during transient floodplain storage

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## 23 S1. Estimating the contribution from POC<sub>petro</sub> in river and floodplain sediment

- 24 Oxidation of petrogenic organic carbon (POC<sub>petro</sub>) and biospheric organic carbon
- 25 (POC<sub>bio</sub>) have different influences on the geologic carbon cycle (e.g., Hilton and West, 2020),
- 26 and the presence of POC<sub>petro</sub> can influence the isotopic signature of POC (e.g., Hilton et al.,
- 27 2010), obscuring detection of isotopic changes caused by floodplain oxidation. While we do not
- 28 use the relative proportions of POC<sub>petro</sub> and POC<sub>bio</sub> in Rio Bermejo sediment to estimate the
- 29 potential for allochthonous POC oxidation in floodplain storage, our data does allow for
- 30 estimating the presence of POC<sub>petro</sub> in our samples. Given the interest in separating contributions
- 31 of POC<sub>bio</sub> and POC<sub>petro</sub> to the geologic carbon cycle (e.g., Berner, 1999; Blair and Aller, 2012;
- 32 Hilton and West, 2020; Horan et al., 2019), we provide such estimates in this supplement. These
- 33 results do not influence the findings presented in the main text, except for the fact that our
- 34 analyses suggest the presence of POC<sub>petro</sub> in Rio Bermejo sediments, thereby making our

36 minimum bound on the total amount of oxidation of allochthonous POC in floodplain storage.

### 37 S1.1. Methods for isolating POC<sub>petro</sub> contribution

38 We isolate POC<sub>petro</sub> in floodplain and river sediment samples by solving for POC<sub>petro</sub> 39 weight percent ( $C_{org petro}$ ) using the Galv et al. (2008a) method, as well as a simple mixing model 40 of Fm vs.  $1/C_{org}$  (sensu Wang et al., 2019), which is free from autocorrelation present in the Galy 41 et al. (2008a) method. Both the Galy et al. (2008a) method and simple mixing model method 42 assume that  $C_{org \ petro}$  is constant for all samples, such that variations in the total amount of POC 43 and Fm among samples is due exclusively to variations in the POC<sub>bio</sub> weight percent ( $C_{org\_bio}$ ). In 44 some cases, we measured a total POC concentration less than the calculated POC<sub>petro</sub> 45 concentration, indicating the constant  $C_{org\_petro}$  assumption was violated. For such cases, we set 46 Corg petro to its maximum possible value by assuming a binary mixture of POC<sub>bio</sub> (assumed to 47 have Fm = 1.07, the highest measured Fm in this study, Table S1) and POC<sub>petro</sub> (Fm = 0) such 48 that

49 
$$C_{org\_petro} = C_{org} \left( 1 - \frac{Fm}{1.07} \right)$$
(S1).

Note that *C*<sub>org\_petro</sub> could exceed the value calculated in Eq. (S1) if samples contain POC
biosynthesized after nuclear weapons testing when atmospheric *Fm* values exceeded 1.07.
Additionally, we independently estimated *C*<sub>org\_petro</sub> by attempting to remove POC<sub>bio</sub> from
a subset of samples with a H<sub>2</sub>O<sub>2</sub> rinse prior to radiocarbon analysis. We leached ~5 g aliquots of
select bedload and active bar deposits at room temperature in 10% H<sub>2</sub>O<sub>2</sub> on a shaker table for
>24 h prior to sample crushing and decarbonation (sensu Galy et al., 2008b). Following H<sub>2</sub>O<sub>2</sub>leaching, samples had *Fm*>0 indicating incomplete removal of POC<sub>bio</sub>. Assuming the leaching

57	did not remove POC <sub>petro</sub> , we estimated a maximum $C_{org\_petro}$ value for these samples following
58	Eq. (S1) using $C_{org}$ and Fm measured on the H <sub>2</sub> O <sub>2</sub> -leached aliquots.

#### 59 **S1.2. POC**<sub>petro</sub> content

60 <i>C</i> <sub>org_petro</sub> for all river and floodplain samples overlapped within error, and we observed a	60	<i>C</i> <sub>org_petro</sub> for all	river and floodplair	n samples overlapped	d within error, and	d we observed n
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- 61 systematic change in Corg\_petro content between samples collected in the upstream versus
- 62 downstream extent of the lowland Rio Bermejo (Figure S4). We estimated
- 63  $0.03\% < C_{org\_petro} < 0.04\%$  and  $0.007\% < C_{org\_petro} < 0.01\%$  using the method of *Galy et al.* (2008a)
- 64 and the simple mixing model, respectively (Figure S4), which bound the range from the  $H_2O_2$ -
- for rinsed samples ( $0.007\% < C_{org\_petro} < 0.03\%$ , Table S3). To estimate the fraction of POC<sub>petro</sub> in the
- floodplain and river sediment samples, we followed Eq. (S1) to calculate a maximum  $C_{org\_petro}$ .
- For samples in which  $C_{org\_petro}$  calculated in Eq. (S1) was >0.04%, we reduced  $C_{org\_petro}$  to 0.04%
- 68 following the results of the Galy et al. (2008a) method, the simple mixing model, and the H<sub>2</sub>O<sub>2</sub>-
- 69 rinsed samples. Using these estimates, the fraction of POC<sub>petro</sub> contributing to the total POC in
- 70 our samples (i.e.,  $C_{org\_petro}/C_{org}$ ) ranged from 0.006–0.6 (Figure S4c); samples with lower  $C_{org}$
- 71 generally contained greater proportions of *C*<sub>org\_petro</sub> (Figure 5).
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## 73 Figure Captions

Figure S1: Example photos of cored floodplain deposits, labeled by floodplain ID and
minimum and maximum deposit ages (Table 1)

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**Figure S2:** (a)  $\mathbb{R}^2$  values indicating the fit of linear regressions of the fractions of grains finer than a given value versus mineral specific surface area (SSA). (b) Fraction of grains finer than 2  $\mu$ m (*f*<sub>2</sub>) versus SSA.

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Figure S3: Particulate organic carbon weight percent ( $C_{org}$ ) (top row), stable carbon isotopic composition ( $\delta^{13}C_{org}$ ) (middle row), and radiocarbon fraction modern (*Fm*) (bottom row) for

- 84 actively-transported suspended and bedload sediment collected in the Rio Bermejo (a and c) and
- floodplain deposits (b and d) versus median particle size ( $D_{50}$ ) (a and b) and Al/Si ratio (c and d).
- 86 In panels (a and c), color and symbol groupings indicate distance downstream from the junction
- 87 with the Rio San Francisco, while in panels (b and d) color and symbol show floodplain

88 depositional age. Error bars show standard deviation from replicate measurements and are

- smaller than the symbol size where not shown.
- 90

91 **Figure S4:** Estimate of  $C_{org\_petro}$  from suspended and bedload sediments following (a) *Galy et al.* 

- 92 (2008a) and (b) with a simple mixing model. Insets show enlarged version of the gray-shaded
- area in the main plot.  $C_{org}$  error bars denote standard deviation of multiple measurements, Fm94 error is analytical uncertainty, and error on the product  $C_{org} \ge Fm$  is propagated assuming random
- and uncorrelated error in  $C_{org}$  and Fm (Table S1). Error bars are smaller than the symbol where
- 96 not shown. (c) Estimate of the fraction of petrogenic organic carbon to total organic carbon
- 97  $(C_{org\_petro}/C_{org})$  and (d)  $C_{org\_petro}$  calculated following Eq. (S1) for samples with  $C_{org} < 0.04\%$ . In
- 98 all panels, squares are floodplain sediment, circles are river sediment, and samples are split
- 99 between upstream portions of the Rio Bermejo (<300 km straight-line distance from the Rio San
- 100 Francisco junction) and downstream portions of the Rio Bermejo (>300 km straight-line distance
- 101 from the Rio San Francisco junction).
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- **Figure S5:** Floodplain depth profiles of median grain size  $(D_{50})$  and Al/Si ratio as a function of
- 104 depth below the surface. (a and b) Show all profiles of  $D_{50}$  and Al/Si, respectively, on the same 105 plot, color-coded by floodplain age. (c and d) Highlight individual profiles of  $D_{50}$  and Al/Si,
- 105 plot, color-coded by floodplain age. (c and d) Highlight individual profiles of  $D_{50}$  and Al/Si, 106 respectively, for each floodplain core (black line) with profiles from other floodplain cores in
- gray. Plot axis extent of individual profiles in (c) and (d) match extent shown in (a) and (b),
- 108 respectively. Error bars are removed for clarity, but are reported in Table S1. Box and whisker
- plots in (a) and (b) show median, inter-quartile range, and full extent of values observed inactively transported river sediments.
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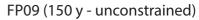
**Figure S6:** Comparison of POC weight percent ( $C_{org}$ ) versus (a) Al/Si ratio and (b) median grain diameter ( $D_{50}$ ) for actively transported river sediment (gray circles) and floodplain deposits (squares). Floodplains deposits are color-coded by depositional age and symbol size indicates sample depth below surface. Solid squares show measured  $C_{org}$ , and open squares show calculated allocthonous POC ( $C_{allo}$ ) in floodplain samples. In cases where  $C_{org} = C_{allo}$ , only solid

- squares are shown. Error bars show standard deviation from replicate measurements, and are
- squares are shown. Error bars show standard deviation from repricate measurements, and are smaller than the symbol size when not shown. Floodplain deposit FP09 is ommitted from the
- 119 figure as we have only a minimum a constraint (150 y) on its age.
- 120
- Figure S7: Comparison of particle size distributions for the two oldest floodplain deposits (FP14 and FP15).
- 124 Tables (available as a single .xlsx file with tables in inidividual tabs)
- 124 125
- 126 **Table S1:** Particulate radiocarbon fraction modern (*Fm*), organic carbon weight percent (*C*<sub>org</sub>),
- 127 total nitrogen (TN) and stable carbon isotope values ( $\delta^{13}C_{org}$ ) for river and floodplain sediment.
- 128 Replicate columns of  $C_{org}$ ,  $\delta^{13}$ C, and TN show measurements collected in individual runs. We
- 129 use the mean and standard deviation of these measurements in all figures. River distance refers to
- 130 the straight line distance downstream from the junction of the Rio San Francisco. Specific
- 131 Surface Area measurements from Repasch et al. (2020).
- 132

133 134 135 136	<b>Table S2:</b> Optically stimulated luminesce and radiocarbon dating results for floodplain deposits (as reported in Repasch et al. (2020)). OSL analysis used quartz of 63 - 90 $\mu$ m, 2 mm aliquots and the central age model (CAM (Galbraith et al., 1999)).
130 137 138 139 140 141 142 143	<b>Table S3:</b> Comparison of organic carbon weight percent ( $C_{org}$ ) and stable carbon isotopes before and after rinsing samples in H <sub>2</sub> O <sub>2</sub> . We calculate maximum possible petrogenic organic carbon weight percent, $C_{org\_petro}$ , after H <sub>2</sub> O <sub>2</sub> rinsing following Eq. (S1). River distance indicates the distance downstream from the junction with the Rio San Francisco junction. Samples with negative river distances are bedload from the Rio Bermejo and Rio San Francisco upstream of the junction of the two rivers. $D_{50}$ indicates median grain size.
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145 146	Supplemental References
147 148 149 150 151 152 153 154 155 156 157 158 159 160 161 162 163 164 165 166 167 168 169 170 171 172 173 174 175	<ul> <li>Berner, R.A., 1999. A new look at the long-term carbon cycle. GSA Today 9.</li> <li>Blair, N.E., Aller, R.C., 2012. The Fate of Terrestrial Organic Carbon in the Marine Environment. Annual Review of Marine Science, Vol 4 4, 401-423.</li> <li>Galbraith, R.F., Roberts, R.G., Laslett, G.M., Yoshida, H., Olley, J.M., 1999. Optical dating of single and multiple grains of quartz from jinmium rock shelter, northern Australia, part 1, Experimental design and statistical models. Archaeometry 41, 339-364.</li> <li>Galy, V., Beyssac, O., France-Lanord, C., Eglinton, T., 2008a. Recycling of Graphite During Himalayan Erosion: A Geological Stabilization of Carbon in the Crust. Science 322, 943-945.</li> <li>Galy, V., France-Lanord, C., Lartiges, B., 2008b. Loading and fate of particulate organic carbon from the Himalaya to the Ganga-Brahmaputra delta. Geochimica Et Cosmochimica Acta 72, 1767-1787.</li> <li>Hilton, R.G., Galy, A., Hovius, N., Horng, MJ., Chen, H., 2010. The isotopic composition of particulate organic carbon in mountain rivers of Taiwan. Geochimica Et Cosmochimica Acta 74, 3164-3181.</li> <li>Hilton, R.G., West, A.J., 2020. Mountains, erosion and the carbon cycle. Nature Reviews Earth and Environment 1, 284-299.</li> <li>Horan, K., Hilton, R.G., Dellinger, M., Tipper, E., Galy, V., Calmels, D., Selby, J., Gaillardet, J., Ottley, C.J., Parsons, D.R., Burton, K.W., 2019. Carbon dioxide emissions by rock organic carbon oxidation and the net geochemical carbon budget of the Mackenzie River Basin. Am J Sci 319, 473-499.</li> <li>Repasch, M., Wittmann, H., Scheingross, J.S., Sachse, D., Szupiany, R., Orfeo, O., Fuchs, M., Hovius, N., 2020. Sediment transit time and floodplain storage dynamics in alluvial rivers revealed by meteoric 10Be. J. Geophys. Res. Earth Surf.</li> <li>Wang, J., Hilton, R.G., Jin, Z.D., Zhang, F., Densmore, A.L., Grocke, D.R., Xu, X.M., Li, G., West, A.J., 2019. The isotopic composition and fluxes of particulate organic carbon exported from the eastern margin of the Tibetan Plateau.</li></ul>
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## FP01 (1 y - unconstrained)

FP05 (4 y - unconstrained)







FP14 (1970 y - 4080 y)

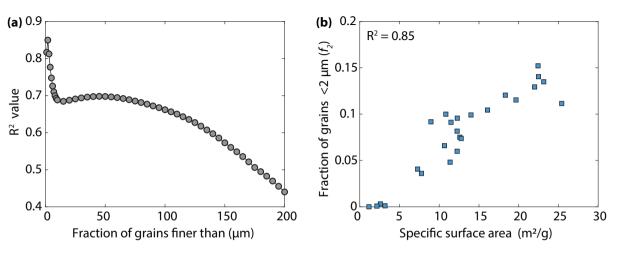


FP13 (380 y - 1960 y)

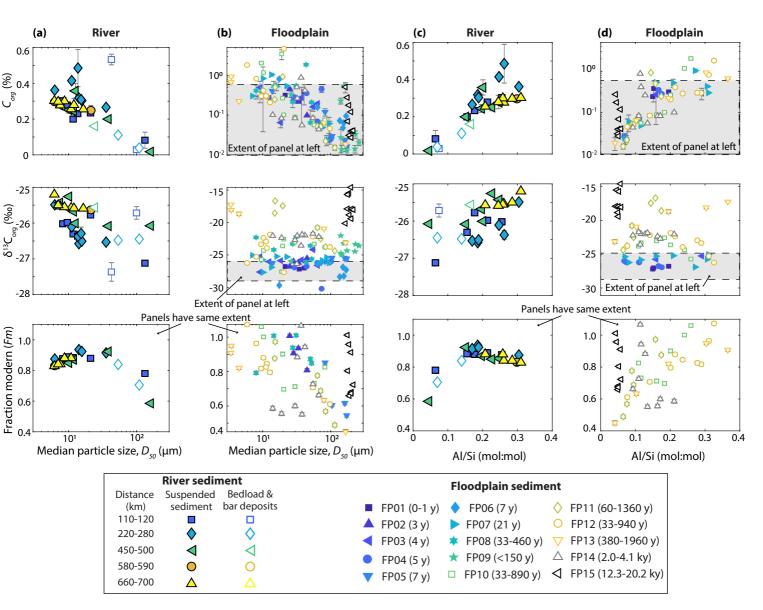
FP15 (12.3 ky - 20.2 ky)



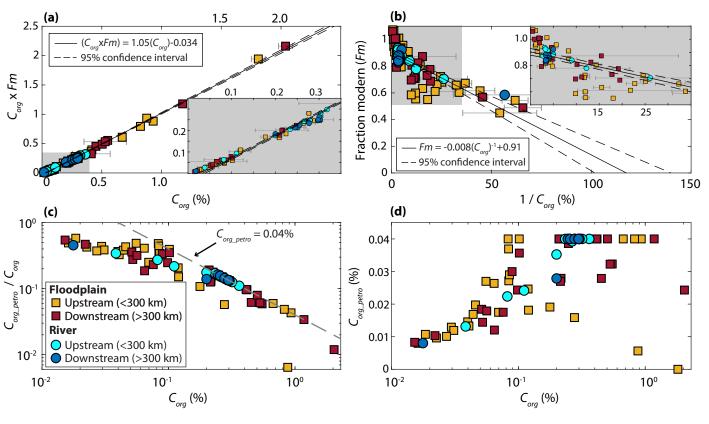
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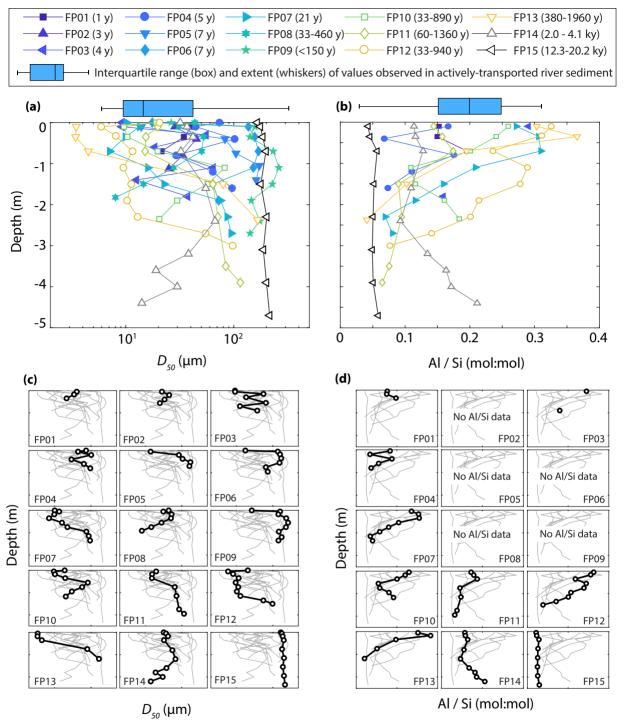
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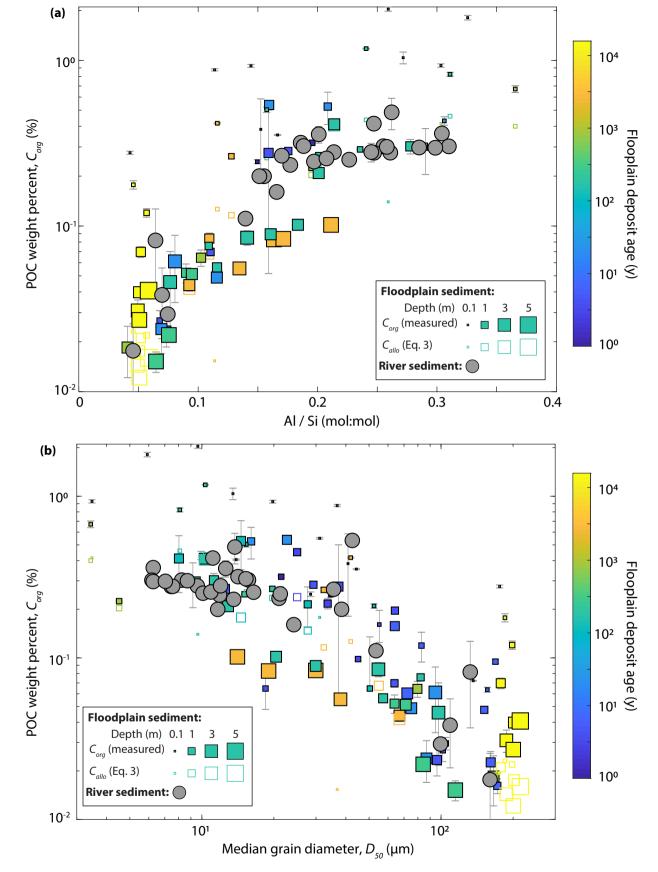
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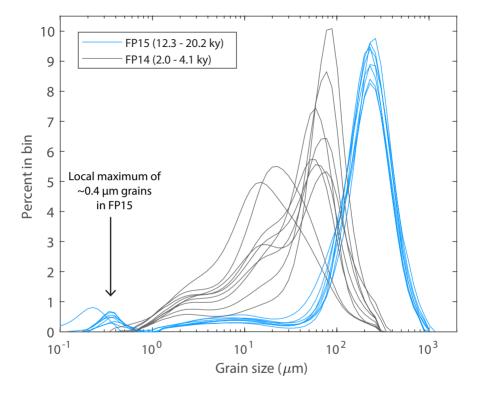
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